6. SITE 569¹

Shipboard Scientific Party²

HOLE 569

Date occupied: 12 February 1982, 2045 hr.

Date departed: 15 February 1982, 0135 hr.

Time on Hole: 52 hr., 50 min.

Position: 12°56.31'N; 90°50.35'W

Water depth (sea level; corrected m, echo-sounding): 2744

Water depth (rig floor; corrected m, echo-sounding): 2754

Bottom felt (m, drill pipe): 2799.7

Penetration (m): 250.7

Number of cores: 27

Total length of cored section (m): 250.7

Total core recovered (m): 134.83

Core recovery (%): 53.7

Oldest sediment cored: Depth sub-bottom (m): 250.7 Nature: mudstone Age: late Oligocene Measured velocity (km/s): 1.8

Basement: not reached

HOLE 569A

Date occupied: 15 February 1982, 0320 hr.

Date departed: 17 February 1982, 0050 hr.

Time on Hole: 46 hr., 30 min.

Position (latitude; longitude): 12°56.22'N; 90°50.81'W

Water depth (sea level; corrected m, echo-sounding): 2795

Water depth (rig floor; corrected m, echo-sounding): 2805

Bottom felt (m, drill pipe): 2814

Penetration (m): 364.9

Number of cores: 11

Total length of cored section (m): 100.0

Total core recovered (m): 16.32

Core recovery (%): 16.3

Oldest sediment cored: Depth sub-bottom (m): 351.4 Nature: mudstone Age: early Eocene Measured velocity (km/s): 2.4

Basement:

Depth sub-bottom (m): 351.4 Nature: amphibolites (altered gabbros and diabases) Velocity range (km/s): 4.982-5.232

Principal results: (Holes 569 and 569A): At Site 569, situated in the midslope area, drilling penetrated a 351-m Recent-late Pleistocene through Eocene sedimentary sequence overlying a basement of gabbros and diabases, of which 14 m were cored. The sequence recovered in Hole 569 is:

Unit I. 0-57 m, Pleistocene and late Pliocene green mud.

Unit II. 57-250 m, Miocene green mudstone including 57-77.9 m late Miocene, 77.9-87.4 m middle Miocene, 87.4-231.4 m early Miocene, and 231.4-250 m, late Oligocene light green calcareous mudstone.

In Hole 569A the sequence recovered is:

Unit II (continued). 246.0-332 m late Oligocene light green calcareous mudstone.

Unit III. 332-351 m, early to late Eocene green and black siliceous mudstone.

Hard-rock basement. 351-365 m, gabbro and diabase metamorphosed to amphibolite and greenschist facies.

Unconformities are suspected between the upper Pliocene and Miocene, the upper Oligocene and Eocene, and between the Eocene and the igneous basement rock.

Together with the results from Sites 566 on the lower slope, and 567 at the base of the slope, the results from this site show that the landward slope of the Middle America Trench is constructed of a basement of ophiolitic rocks beneath a cover of sediments. The oldest material overlying ophiolitic rocks is late Miocene at Site 566, early Miocene at Site 567, and early Eocene at Site 569; the minimum age of the tectonics that emplaced the original ophiolitic rock is, at least, pre-early Eocene.

These results strongly suggest an analogy with on-land geology and confirm the fact that off Guatemala the whole margin is constructed of an igneous basement belonging to the continental framework of Central America.

BACKGROUND AND OBJECTIVES

Site 569, situated in about 2800 m of water, 3200 m above and 32 km from the Middle America Trench axis, is in the middle of the slope where the generally thick sediment becomes thin enough to be penetrated by the *Glomar Challenger*. The seismic records show an irregular mantel of slope deposits burying a rough basement topography. The maximum length of drill string available on board the *Glomar Challenger* (3300 m), and the area where the base of the sediment is above the BSR (bottom simulating reflector) were important considerations in order to avoid the gas hydrate problem in getting to the basement. Site 569 is situated above a bump in the basement perhaps related to faulting. Two strong

¹ von Huene, R., Aubouin, J., et al., *Init. Repts. DSDP*, 84: Washington (U.S. Govt. Printing Office).
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reflectors, at a sub-bottom depth of about 400 m (assuming an average sound velocity of 1.8 km/s in the sediments), are the indication of basement upon which rests first a conformable sediment cover and then prograding slope deposits.

The main objectives at this site were (1) to sample the basement so that it would be possible to compare it with the basement at Sites 566 and 567, which are igneous rocks of an ophiolitic suite; (2) to recovering a complete section of the sediments above the basement, in order to compare this sequence with the results at Sites 494 (Leg 67), 566, and 567 (Leg 84), and the sites abandoned before basement was reached for safety reasons related to the gas hydrates problem (Leg 67: Sites 496, 497, 498: Leg 84: Site 568).

OPERATIONS

Glomar Challenger departed Site 568 at 1615L (local time) 12 February, under way to Site 569 where a 16-kHz beacon was dropped at 1950L. The ship was in position at 2045L. Running in began at 2330L after making up BHA (bottom hole assembly) and sorting 27 joints of premium pipe.

The hole was spudded at 0852L, 13 February, and the first core recovered at 1052L.

After having recovered Core 27 at 2107L, 14 February, it appeared that there was annular packing around collars in overpressured formation (250 psi static pressure on pipe), which caused us to abandon the hole.

At 0136L, 15 February, with four joints retrieved above the mud line, *Glomar Challenger* moved about 900 m west, 200 m south from the beacon to Hole 569A, the ship was positioned at 0320L.

After spudding at 0445L, drilling first washed down to 246 m, below which continuous coring was begun. The wash core, 569A-H1, was retrieved at 1010L, and the first core recovered at 1226L. The coring was somewhat difficult on February 16, with repeated mud flushes in the hole. In order to try to keep on drilling the hole and to reach basement, it was decided to drill two joints before retrieving the core barrel: recovery had been very low, so further recovery was not jeopardized, which reduced the tendency to stick during the wireline trips. This was done for Cores 8 and 9; at 1352L, 16 February, Core 10 was recovered with gabbros and dolerites below the Eocene mudstone. Only 4 m were cut of Core 11, because the main target of the hole had been reached, and further drilling was abandoned.

After filling the hole with heavy mud, running out began at 1615L, 16 February, and ended at 0045L, 17 February.

At 0050L, 17 February, *Glomar Challenger* departed for Site 570.

Table 1 shows the coring summary for Site 569.

LITHOSTRATIGRAPHY

Site 569 is located between Sites 568 and 567 on the landward side of the Middle America Trench at a water depth of about 2800 m. The site is 32, km from the Trench axis and 125 km south of the coast of Guatemala (Fig. 1).

Twenty-seven cores were drilled in Hole 569 with good recovery to a sub-bottom depth of 251 m, when caving in the hole became a problem. The hole was abandoned and our location offset about 1 km to drill Hole 569A. After washing down to 246 m at Hole 569A, continuous coring was resumed for eight cores, at which point the consistently poor recovery and the imminent danger of caving promoted the decision to drill two cores' length between core retrievals. This procedure was continued for two cores (8 and 9) when the drilling encountered hard rock, and continuous coring was resumed (Cores 10 and 11).

We divide the sedimentary lithology into three units on the basis of color, texture, constituents, and seismic data; below Unit III, hard rock was drilled (Fig. 2).

Hole 569

Unit I

Unit I, composed of (Cores 1 to 7, 0 to 57 m sub-bottom depth (Recent-late Pleistocene to late Pliocene), is a generally massive, dark olive gray (5Y 3/2) to grayish olive (10Y 4/2) mud that easily deforms into swirled layers during drilling. Average sand-silt-clay composition of the unit is 6, 17, and 77%, respectively. Siliceous biogenic remains (radiolarians, sponge spicules, silicoflagellates, and, in particular, diatoms) are common and locally constitute 15% of the total sediment composition, and sandy horizons occur in Cores 1 to 3 and in Core 7. The sand in Core 7 is coarse, dark, and angular, and contain a high percentage of ash. Sedimentary layering is weakly preserved in Cores 2 to 4, and is generally less than 10 cm thick.

Twelve ash or muddy-sandy ash layers occur in Cores 1 to 5, and numerous ash mottles (which may be drilling-disturbed layers) and pumice clasts occur throughout the unit. Colors range from light gray olive (10Y 4/2) to olive black (5Y 2/1) (containing quartz, plagioclase, pyroxene, and orthopyroxene), and very light gray to white (N8-N9, usually glass-shard-rich).

Unit II

Unit II comprises Cores 7 to 27, 57-250 m, sub-bottom depth (Miocene-late Oligocene). The contact between Unit I and Unit II is placed at an abrupt color change below a 15-cm-thick dark sand layer in Core 7, Section 2. Mud below the contact is firmer. The horizon roughly corresponds to an acoustically defined contact between downslope-dipping reflections of sedimentary lobes and sediments draping over these (see the Geophysics section) and for this reason was favored as the contact over another color change observed in Core 8. This unit comprises ash-bearing mud-mudstone that is olive gray (10Y 4/2) to grayish blue green (5BG 5/2) in color, mottled and bioturbated particularly from Core 12 to the bottom of the hole. Small silty-sandy horizons are common, and burrows are commonly filled with coarser sediment. The average sand-silt-clay composition of the mud is 4, 21, and 75%, respectively.

Thin limestone layers occur in Cores 13 and 18 that are yellowish gray (5Y 7/2) and greenish gray (5GY 6/1),

Table 1. Coring summary, Site 569.

	222.00		Depth from drill floor	Depth below seafloor	Length	Length	P	
Core	Date Feb. 1982	Time	(m) Top Bottom	(m) Top Bottom	cored (m)	recovered (m)	Recovery (%)	
Hole 5	69							
1	13	0952	2799.7-2800.9	0.0-1.2	1.2	1.18	98	
2	13	1055	2800.9-2810.8	1.2-11.1	9.9	7.63	77	
3	13	1200	2810.8-2820.3	11.1-20.6	9.5	5.73	60	
4	13	1259	2820 3-2830 0	20.6-30.3	9.7	9.62	99	
5	13	1356	2830 0-2839 7	30 3-40 0	97	5 41	56	
6	13	1457	2839 7-2849 4	40 0-49 7	97	8 90	92	
7	13	1555	2849 4-2858 6	49 7-58 9	9.2	9.55	100	
8	13	1705	2858 6-2868 2	\$8 0 68 5	9.6	4 22	44	
0	13	1828	2858 2-2877 6	69 5 77 0	9.0	3 13	33	
10	13	1027	2000.2-2077.0	77 0 97 4	0.5	3.15	34	
11	13	2050	2877.0-2887.1	87 4 07 1	9.5	3.13	37	
12	13	2030	2007.1-2090.0	07.1 106.7	9.1	9.07	02	
12	13	2204	2090.0-2900.4	97.1-100.7	9.0	5.12	55	
15	15	2324	2900.4-2915.7	100.7-110.0	9.5	3.12	40	
14	14	0037	2915.7-2925.5	110.0-125.0	9.0	3.61	40	
15	14	0200	2925.3-2934.8	125.6-135.1	9.5	4.40	47	
10	14	0308	2934.8-2944.5	135.1-144.8	9.7	1.27	13	
17	14	0415	2944.5-2954.1	144.8-154.4	9.6	1.27	10	
18	14	0520	2954.1-2963.8	154.4-164.1	9.7	4.29	44	
19	14	0645	2963.8-2973.5	164.1-173.8	9.7	0.37	4	
20	14	0820	2973.5-2983.1	173.8-183.4	9.6	2.15	22	
21	14	0940	2983.1-2992.6	183.4-192.9	9.5	3.54	37	
22	14	1115	2992.6-3002.3	192.9-202.6	9.7	3.14	32	
23	14	1331	3002.3-3011.8	202.6-212.1	9.5	5.23	55	
24	14	1515	3011.8-3021.5	212.1-221.8	9.7	8.24	85	
25	14	1643	3021.5-3031.1	221.8-231.4	9.6	3.82	40	
26	14	1815	3031.1-3040.7	231.4-241.0	9.6	6.85	71	
27	14	2107	3040.7-3050.4	241.0-250.7	9.7	4.60	47	
Total					250.7	134.83	54	
Hole 5	69A							
HI	15	Washed	2814.0-3060.0	0.0-246.0	6.12	\sim		
1	15	1226	3060.0-3069.5	246.0-255.5	9.5	6.47	68	
2	15	1436	3069.5-3079.9	255.5-265.1	9.6	2.47	26	
3	15	1616	3079.1-3088.5	265.1-274.5	9.4	0.00	0	
4	15	1805	3088.5-3098.2	274 5-284 2	9.7	0.97	10	
5	15	2012	3098.2-3107.9	284 2-293 9	9.7	0.06	1	
6	15	2325	3107 9-3117 5	203 0-303 5	9.6	1.00	10	
7	15	0230	3117 5-3127 1	303 5-313 1	9.6	1 14	12	
8	16	0430	3127 1-3136 8	313 1-322 8	9.7	0.68	1	
.0	16	Drilled	3136 8-3146 3	377 8_337 3	2.1	0.00		
0	16	0947	3146 3-3156 0	332 3-342 0	97	1.96	20	
2	16	Drilled	3156 0-3165 4	342.0-351.4	2.1	1.50	20	
10	16	1352	2165 4 2174 0	342.0-331.4	0.5	1.25	14	
11	17	0008	3174 0 3179 0	360.0.364.0	4.0	0.22	6	
	17	0000	51/4.7-51/8.9	300.9-304.9	4.0	0.22	0	
Total					100.0	16.32	16	

Note: H1 designates wash core. "Drilled" in Time column indicates that recovery was combined with recovery of the preceding core, that is, two core lengths were drilled for every instance of recovery.

respectively, and bioturbated. Sandy layers occur in Cores 13, 18, and 21.

Bedding inclinations of 30 to 60° are observed in Cores 14 through 16 and 27. Veining and small-scale fractures are first observed in Core 12, where the mudstone is also stiff enough to break into biscuitlike chunks. These features are similar to the veining and fracturing observed at Site 568, except that they range in size to include greater dimensions (lengths observed up to 10 cm, thickness of parallel-type veins up to 3 mm); see Figure 3. Offset along fractures indicates extension.

Hole 569A

Unit II

Unit II comprises Cores 1 to 8, 246 to 332.3 m subbottom depth, (late Oligocene). After washing down to 246 m, coring was resumed in the Unit II sedimentary facies encountered in Hole 569. The major lithology is composed of heavily burrowed and mottled olive gray (5Y 3/2) to grayish olive (10Y 4/2) mudstone. Veining and fracturing as described at Site 568 and Hole 569 are common in this hole. Two ash layers occur in Cores 2 and 7; these are pale reddish purple (5NP 6/2) and medium dark gray (N4), respectively.

Cores 6 to 8 make up the minor lithology of this unit and consist of interbedding calcareous mudstone and claystone (pale blue green 5BG 6/6), which exhibit the same striking mottled, bioturbated, fractured, and veined appearance. In the core catcher of Section 8, a coarse blue gray (5G 6/1), sandy limestone identifies the base of Unit II.

Unit III

Unit III is Core 9, 322 to 351 m sub-bottom depth, (late-early Eocene). The contact between Units II and



Figure 1. Bathymetry of the Guatemala margin and San José Canyon, showing the UTMSI survey tracks and the location of Legs 67 and 84 sites off Guatemala.

III is placed at the base of Core 8 on the base of a textural and color change.

Unit III comprises heavily mottled and bioturbated radiolarian-rich mudstone. The top of Core 9, Section 1, is a blue gray mudstone that occurs in swirled, drilling-disturbed contact with darker (light olive gray 5Y 5/2 to grayish olive 10Y 4/2) mudstone. The lowermost 30 cm of the unit consists of a faintly mottled, brownish black mudstone.

Hard Rock Petrology

Gabbroic and diabasic rocks were recovered from the base of Hole 569A; amphibolite (hornblende) and greenschist facies minerals replace the primary minerals. Crosscutting veins filled with zeolite minerals and prehnite are common features in the rock.

If these rocks are of an igneous provenance similar to that of the hard rocks recovered at Hole 567A, they have been subjected to nearly complete alteration during greenschist metamorphism. Later mineralization under zeolite metamorphic conditions is identified by the abundant zeolite-filled veins that crosscut the rock.

Discussion

The dipping beds in Unit II lend themselves to the same suite of interpretations proposed in the discussion of the Site 568 lithology.

The presence of late Oligocene ash layers in Holes 569 and 569A has interesting implications for the regional tectonic history. One of the reasons for assuming that the initiation of subduction along the Middle America Trench took place in the Miocene is the first appearance of ash horizons during this time. The recovery of still older ash layer may permit extension of the age of the present subduction setting back into the late Oligocene. The altered metamorphosed igneous rocks are similar to the hard crystalline rocks of the ophiolitic suite recovered from Site 567.

BIOSTRATIGRAPHY

Introduction

Two holes drilled at Site 569 recovered 361 m of Pleistocene through early Eocene sediments overlying metabasite basement. Calcareous nannofossils are age-diagnostic and moderately preserved within most of the Pleistocene through late Eocene mudstones, but less frequent and nondiverse within much of the early Miocene and probable early Eocene sediments. Below the Pleistocene, diatoms are commonly low in abundance and moderately to poorly preserved. A barren interval exists until the early middle Miocene, when diatoms again become useful for biostratigraphy. Benthic foraminifers are moderately well preserved and abundant in the Pleistocene, however, they decrease in number and quality of preservation through the Miocene and Eocene.

In Hole 569 Sections 569-1-2 through 569-6,CC are Pleistocene and 569-7,CC is late Pliocene, although some reworked Miocene specimens were noted. Sample 569-8-1, 18-22 cm is late Miocene, on the basis of benthic foraminifers and diatoms; 569-10,CC is middle Miocene, based on nannofossils and benthic foraminifers; 569-11-2, 71 cm through 569-12,CC are early Miocene based on diatoms and nannofossils, whereas 569-13-4 through 569-26-4, 44 cm are early Miocene undifferentiated, based on nannofossils, diatoms, and benthic foraminifers. Nannofossil Samples 569-27-1, 19 cm through 569-27,CC are late Oligocene (Fig. 4).

In Hole 569A, Samples 569A-1-1, 99 cm through 569A-7,CC are late Oligocene, based on nannofossils, whereas 569A-8,CC is late Eocene. And 569A-9-1, 18 cm through 569A-9,CC are late middle Eocene, based on nannofossils and, in part, benthic foraminifers and diatoms. A muddy limestone recovered just above basement in 569A-10-1 is probably late early Eocene (possibly early middle Eocene), in the basis of nannofossils (Fig. 4).

Paleoecologic analysis of Holes 569 and 569A suggests deposition occurred at abyssal depths during the Miocene through middle Eocene. Additionally, the early Miocene through middle Eocene deposition was probably below of near the foraminiferal CCD (calcite compensation depth). Calcareous benthic foraminiferal tests were occasionally preserved (569A-8 and -9,CC) at these great abyssal depths, and planktonic foraminiferal tests only rarely (569A-10-1). Transported material in the abyssal biofacies was primarily from the lower and upper middle bathyal biofacies. In assemblages deposited below the CCD, transported outer shelf and upper bathyal species are often the only components of the fauna.

Sediment accumulation rates uncorrected for compaction (Fig. 5) vary from 7 m/m.y. to 33 m/m.y. for both holes and illustrate at least three (probably more) unconformities.

The accumulation rate for the Pleistocene (Hole 569, 1-55 m, Cores 1 through 7) is approximately 33 m/m.y. Below this interval is a sequence (56–85, Cores 7 through

10) of siliceous mudstones that contain late Pliocene nannofossils and diatoms at the top, and early middle Miocene diatoms at the bottom. This interval represents one, and possibly several early Pliocene-Miocene unconformities.

Rates of sediment accumulation are approximately 7 m/m.y. for the early middle and late early Miocene (Hole 569, 86–107 m, Cores 10 through 12) and increase to approximately 25 m/m.y. for the remainder of the recovered early Miocene section (108–231 m, Cores 13 through 25). The sedimentation rate for the late Oligocene is approximately 7 m/m.y. (Hole 569, 232–251 m, Cores 26 through 27, and Hole 569A, Cores H1 through 1); this rate was averaged through two zones that bracket a barren interval.

The late Oligocene sedimentation rate for Hole 569A (250–324 \pm 10 m, Cores through 10) was also averaged across a barren interval bracketed by two zones, and is approximately 23 m/m.y.

Late Eocene sediments are present from 332 to 352 m, with accumulation rates of 7 m/m.y. An unconformity near the base of the late Oligocene is therefore suggested somewhere between 315 and 330 m, which is an interval of poor sample recovery.

A few centimeters of probably early Eocene limestones are present in Section 569A-10-1, which is indicative of another unconformity at 355 m (\pm 5 m) near the base of the late Eocene.

Nannofossils

Recovery at Hole 569 include 250.7 m of Pleistocene to late Oligocene (Fig 4) mudstones. Hole 569A was washed to 246 m and cored through late Oligocene probably early Eocene sediments upon basement.

Calcareous nannofossils are frequent to common and moderately preserved within most of the Pleistocene through late Eocene intervals, except for the early Miocene interval from 125 to 231 m in Hole 569 where assemblages diversity and numbers drop. Early Eocene nannofossils are also rare and heavily recrystallized within Sample 569-10-1, 1-2 cm.

Reworking is minimal, and it is largely restricted to the Pleistocene, where several rare Cretaceous species were observed.

Hole 569

Sections 569-1-2 through 569-4, CC are assigned to the Recent-late Pleistocene *Emiliania huxleyi* to *Gephyrocapsa oceanica* Zones and contain common *G. oceanica, Helicosphaera kamptneri*, small *Gephyrocapsa* spp., and minute placoliths with affinities to *Emiliania huxleyi*. Samples 569-5, CC and 569-6, CC (40.0-49.7 m) are also Pleistocene, but species diversity and abundances are unfavorably low for detailed zonation. *Pseudoemiliana lacunosa* is present, which is indicative of an age older than the *G. oceanica* Zone. Very rare specimens of the Cretaceous species *Watznaueria barnesae* are present in 569-6, CC. Samples 569-7-2, 10 cm through 569-7-6, 10 cm are barren of nannofossils.

Sample 569-7, CC (58.9 m) is assigned to the late Pliocene *Discoaster pentaradiatus* Subzone. It contains frequent *Discoaster pentaradiatus*, and rare *D. brouwe*-



Figure 2. Lithostratigraphic summary of Site 569.

ri and D. decorus. A change in sediment color, which may be the Pleistocene/Pliocene boundary, is present near 569-7-6, 49 cm, but samples that bracket the contact are barren of nannofossils. Samples 569-8,CC and 569-9,CC are also barren of nannofossils. Sample 569-10-2, 20 cm through 569-10, CC are middle Miocene and assigned to the *Sphenolithus heteromorphus* Zone. Diagnostic species include *S. heteromorphus*, *Discoaster exilis* s.l., *Coccolithus miopelagicus*, and *Discoaster signus*.



Figure 2. (Continued).

Sample 569-11-2, 71 cm through 569-12,CC are assigned to the early Miocene Helicosphaera ampliaperta Zone and contain H. ampliaperta, S. heteromorphus and Discoaster deflandrei.

Sample 569-13-4, 3 cm through 569-22, CC (116-202.6 m) are early Miocene, based on rare to frequent abundances of moderately to poorly preserved specimens of *D. deflandrei*, *Cyclicargolithus floridanus*, *Sphenolithus moriformis*, *S. conicus*, *S. cf. belemnos*, and *Discoaster saundersi*. Sample 569-17, CC also contains rare *Triquetrorhabdulus milowii*, which is restricted to early Miocene or younger. The absence of *Discoaster druggii* indicates that this entire interval may be within the *Sphenolithus belemnos* Zones, but low assemblage abundance and diversity prohibit definite zonal assignment. Samples 569-23, CC and 569-24, CC are dominated by siliceous microfossils and are barren of nannofossils.

Samples 569-26-3, 62 cm and 569-26-4, 44 cm were selected from the light gray calcareous mudstones that fills the burrows of a highly bioturbated sediment. They contain *Discoaster deflandrei*, *Dictyococcites scrippsae*, *Helicosphaera recta*, and *H. bramlettei* and lack *Cyclicargolithus abisectus* and *Dictyococcites bisectus*, which is most indicative of the early Miocene *Discoaster deflandrei*, Subzone.

Sample 569-27-1, 19 cm through 569-27, CC (250.7 m) are late Oligocene and assigned to the *Dictyococcites bisectus* Subzone. Age-diagnostic species include *Cyclicargolithus floridanus, Discoaster deflandrei, Triquetrorhabdulus carinatus, Dictyococcites bisectus, Helico-*

sphaera recta, Cyclicargolithus abisectus, Sphenolithus ciperoensis, and Helicosphaera bramlettei.

Hole 569A

Sample 569A-1-1, 99 cm is probably early Miocene (Discoaster deflandrei Subzone) and contains Discoaster deflandrei, Helicosphaera bramlettei, and Dictyococcites scrippsae.

Samples 569A-2-1, 147 cm through 569A-7-1, 43 cm are late Oligocene and are assigned to the Dictyococcites bisectus Zone. They contain the species Cyclicargolithus floridanus, Discoaster deflandrei, Helicosphaera intermedia, C. abisectus, and Helicosphaera recta. Sample 569A-7,CC contains well developed Dictyococcites bisectus and Sphenolithus predistentus, along with the above assemblages, and is assigned to the late Oligocene Sphenolithus perdistentus Zone.

Sample 569A-8, CC (322.7 m) is late Eocene and is assigned to the Discoaster barbadiensis Zone. Discoaster barbadiensis, Reticulofenestra umbilica, and Dictyococcites scrippsae are present along with rare Calcidiscus formosus and common C. reticulatus. Sample 569A-9-1, 18-19 cm contains the contact of a heavily burrowed and possibly disturbed base of a light gray calcareous mudstone that overlies the dark gray mudstone sampled from 569A-9, CC. Rare specimens of Chiasmolithus grandis are present, which, if not reworked, would suggest a slightly older age (late middle Eocene Discoaster saipanensis Zone) for all of Core 569A-9.



Figure 3. Photograph of parallel veins in sediment at Site 569, Sample 569-15-1, 0-30 cm.

Core 569A-10 (360.9 m) penetrate dolerite basement and contains only several centimeters of disturbed, muddy, light gray limestone fragments at the top of Section 1. This limestone contains highly recrystallized specimens of *Discoaster lodoensis*, *Coccolithus crassus*, *C. gammation*, *Chiasmolithus solithus*, *D. barbadiensis*, and *Sphenolithus radians*. This assemblage is most indicative of the late early Eocene *Discoaster lodoensis* Zone, but extremely poor preservation may have destroyed key species that would place it within the early middle Eocene *Discoaster sublodoensis* Zone. Core 569A-11 recovered metabasite mixed with a piece of late middle Eocene limestone displaced from upsection

Diatoms

Diatomacerous material occurs throughout practically the entire length of Hole 569; only Core 19 is entirely barren of diatoms. There is a hiatus, however, from Cores 14 to 23, where no stratigraphically useful taxa were observed.

The Pleistocene is represented by a diverse flora in core catchers 569-1 through -6. *Pseudoeunotia doliolus* is present throughout this entire interval as well as in Section 569-7-2 and 569-7-4 which should put the Pliocene/Pleistocene boundary somewhere within this core.

Rhizosolenia praebergonii, a lower Pleistocene indicator, is present only in 569-5, CC.

The Pliocene is represented by all, or at least part of, Core 7. Rossiella tatsunokuchiensis, extant only during the Miocene, appears at the end of Cores 569-7 and -8. Thalassiosira convexa, a marker fossil for the late Miocene to early Pliocene, is common in 569-7.CC. Section 569-8.CC contains the last appearence of Thalassiosira oestrupii, a marker fossil for the interval from latest Miocene to mid-Pliocene. Core 10 contains definite late early to early middle Miocene material. First appearing in Core 10 is Actinocyclus radionovae, restricted to early Miocene (Barron, in press); Coscinodiscus lewisianus (middle Miocene); Craspedodiscus coscinodiscus (middle Miocene); Denticulopsis nicobarica (late middle Miocene); and Synedra jouseana (middle Miocene). Although the Miocene continues through Core 27, Cores 13 and 23 contain the late Oligocene taxon Coscinodiscus oblongus. Section 569-25, CC contains Melosira clavigera fragments (late Eocene) and Cestodiscus pulchellus (early to mid-Oligocene), which suggests that there was some Eocene and Oligocene reworking during the early Miocene.

The first cores of Hole 569A contain an assemblage again representing reworked Eocene and Oligocene at an early Miocene depositional site. This is confirmed by the presence of the early Miocene taxa Actinocyclus radionovae, Coscinodiseus rhombicus, and Craspedodiscus elegans. This interval also includes Xanthiopyxis acrolopha (late Eocene-early Oligocene). The section from 569A-2 through 569A-8,CC is barren of stratigraphically useful diatoms. In 569A-9,CC the following species appear: Hemiaulus polymorphus (late Eocene), Monobrachia simplex (late Eocene), melosira architecturalis (late Eocene-early Oligocene), and Archaeomonas dubia (late Eocene). This suggests an Eocene age for 569A-



Figure 4. Biostratigraphic and paleoecologic summary, Site 569. Hachures indicate barren intervals.



Figure 5. Sediment accumulation rates for Holes 569 and 569A, uncorrected for compaction.

9 through 569A-11,CC, after which dolerite is encountered to the base of the hole.

Benthic Foraminifers

Benthic foraminiferal assemblages in Holes 569 and 569A are highly variable in number and preservation. Well-preserved assemblages are encountered in the Pleistocene through late early Miocene, whereas the early Miocene and older sediments are poorly preserved and of low diversity. Paleoecologic analysis of the benthic foraminiferal assemblages indicates deposition occurred within the abyssal biofacies (\geq 4000 m) from the Pleistocene through the middle Eocene. The scarcity of calcareous tests in the early Miocene to middle Eocene interval suggests that deposition occurred near the foraminiferal CCD. Shallowing to present lower bathyal depths occurred in the latest Pleistocene.

Samples 569-1-1, 91-95 cm to 569-6,CC, contain Pleistocene assemblages similar to the Holocene of this area (Smith, 1964). Ecologic analysis of these samples indicates deposition occurred at abyssal depths, that is, below 4000 m. This interpretation indicates a decrease in water depths from the Pleistocene to the present depth of 2700 m. Uvigerinids (*Uvigerina senticosa, U. rustica*, and *U. peregrina*) are the most abundant component. Transported material was primarily from the lower middle bathyal biofacies.

Benthic foraminiferal assemblages in Core 569-7, may be Pliocene. This assemblage, through less diverse, contains many of the same species as the Pleistocene assemblage and some Miocene species. These Miocene species are rare and may have been reworked into younger sediments. Water depths of greater than 4000 m are again suggested by the benthic foraminifers. Transported material is principally from the lower middle bathyal biofacies.

Cores 569-8 to -25 are Miocene: Samples 569-8-1, 18-22 cm through 569-10, CC are middle Miocene, and Cores 11 through 25 are early Miocene. In this hole the Pliocene/Miocene and middle/early Miocene boundaries are marked by extensive faunal changes. Miocene benthic foraminifers are also present in Cores 569A-H1 and -1, but probably indicate downhole contamination rather than a Miocene age. Cores 569A-2 through -5 are barren or poorly preserved and contain no age-diagnostic benthic foraminifers. Paleoecologic analysis indicates that deposition continued in the abyssal biofacies $(\geq 4000 \text{ m})$ during the Miocene. Upper middle bathyal benthic foraminiferal species are most common in the transported material. This change suggests a change in source area from the Pliocene-Pleistocene and Miocene. The poorly preserved early early Miocene Samples 569-20,CC, 569-22,CC, 569-25,CC as well as the barren samples below these appear to have been deposited either below the foraminiferal CCD or were subjected to corrosive bottom waters because only the most solution-resistant forms remain.

Reworked middle Eocene benthic foraminifers are present in Cores 569-12 and -13. Species of *Lenticulina*, *Nodosaria*, and *Vaginulinopsis* are rare, poorly preserved, diagnostic of the middle Eocene, and indicate warm, shallow marine conditions. Thus their association with well-preserved abyssal Miocene assemblages makes them easily identifiable as reworked.

Benthic foraminiferal assemblages in Cores 569A-6 through -9, are poorly preserved and of low deversity. These cores are characterized by rapid faunal changes and intervening barren samples. Late Oligocene species are present in Cores 569A-6 and -7, and late Eocene species occur in Cores 569A-8 and -9. Samples from Core 569A-10 are barren or too poorly preserved for age interpretation. Paleoecologic analyses suggest deposition occurred in the abyssal biofacies (\geq 4000 m) and below the foraminiferal CCD, because most of the calcareous tests are siliceous.

PHYSICAL PROPERTIES

Methods

Measurements of bulk density, porosity, water content, compressive wave velocity, shear and unconfined compressive strength, and thermal conductivity were conducted as described for earlier sites and in Boyce (1976). All tests were run routinely except for thermal conductivity, which was run on a limited number of samples.

Results

Index properties for the sedimentary column at this site resemble those obtained at Sites 568 and 496. Bulk densities show a gradual increase from about 1.4 Mg/ m³ near subsurface to 1.8 Mg/m³ in the overall trend, with some higher densities reflecting cementation in calcareous mudstones. Corrected wet-water contents, porosities, and bulk densities are presented in Figure 6, where the former two inversely mimic the distribution of bulk densities described. Three features in index property trends stand out from the main tendencies at sub-bottom depths of approximately 60, 115, and 352 m. The 352-m depth represents the sediment/crystalline rock (amphibolite) contact, whereas the 60-m depth can be correlated to the Pliocene/Miocene boundary. As at Sites 568 and 496, the 60- and 115-m depth show a more abrupt change in index properties, which is best appreciated in the wet-water content profile.

Shear and unconfined compressive strengths were measured routinely on Cores 1 through 12 beyond which lithified mudstone biscuits were recovered. Figure 7 shows the vertical distribution of strengths measured and the close correlation. The section shows a slow increase of strength in the upper 35 m, below which the section is more variable. Strength data showing deviation from a downhole trend are found at about 35, 58, and 70 m sub-bottom. The Pliocene/Miocene boundary can also be appreciated from this curve as well as from index properties.

Compressional wave velocity and acoustic impedance are displayed in Figure 8. The overall trend is gradual. Velocities increase from 1.5 km/s at subsurface to 1.9 km/s at 310 m. Deviations occur at the Pliocene/Miocene boundary or unconformity (60 m) and at about 115 m sub-bottom. Bulk densities, velocities, and impedances for igneous rock samples are detailed in Table 2.

Limited thermal conductivity measurements were made and are displayed in Table 3. These values are not corrected to *in situ* conditions and the variability is attributed, in part, to degassed structure in unlithified sediments and questionable contact between probe and slabbed samples (569-24-4 and 569-27,CC are mudstones and 569A-10-1 is amphibolite).

Discussion

The sedimentary column drilled resembles other slope sites in this area, showing a slow and gradual diagenetic modification except for some limited horizons. Two zones are outstanding, however, showing rapidly changing characteristics typical of an increased state of consolidation relative to the remaining section. The upper of these events is correlated to the Pliocene/Miocene boundary and may represent one or more nondepositional or erosional unconformities. At about 115 m sub-bottom a similar horizon occurs. Through no clear paleostratigraphic boundary is distinguished at this deeper horizon, some evidence of one or more hiatuses is suggested from microfossil assemblages, which leads us to believe this zone may also represent some type of unconformity.

GEOPHYSICS

Site 569 is near the juncture of the middle and lower slope of the Trench. The middle slope area is distinguished by thick sediment and a rough acoustic basement surface. The lower slope is steeper has a thin cover of slope deposits conforming to the topography, and a relatively smooth acoustic basement surface except along seismic line GUA-13 (Fig. 9) where the lower slope has three steps. Site 569 is located on GUA-13, about 8 km upslope from the highest step. The site is also 1.5 km southeast of a canyon that traverses the lower slope of the Trench.

Although the topography of the slope appears relatively simple on GUA-13, variation in on-board bathymetric records was seen during the attempts to find the site and drop a beacon. The ship made five passes up and down the slope within about 1 km of each other,







Figure 7. Strength measurements for sediments of Hole 569.

Table 2. Physical properties of amphibolites at Site 569.

Sample	Sub-bottom depth (m)	Bulk density ^a (Mg/m ³)	Sonic velocity (km/s)	Acoustic impedance ^a (× $10^5 \text{ g/cm}^2 \cdot \text{s}$)
569A-10-1	351.67	2.739	5.232	14.330
569A-10-1	352.18	(2.591)	4.983	(12.911)
569A-10-1	352.38	2.581	5.125	13.228
569A-10-2	352.91	(2.847)		

^a Bulk density and acoustic impedance values obtained from gravimetric data. Values in parenthesis are based on 2-min. GRAPE counts.

Table 3. Thermal	conductivity	measurements at Site 569.	

Sample	Sub-bottom depth (m)	Thermal conductivity (mcal/cm°C s)
569-4,CC	30.30	1.713
569-6,CC	49.70	1,160
569-9,CC	77.90	1.826
569-13,CC	116.00	1.856
569-24-4	216.89	4,192
569-27,CC	250.70	3.068
569A-10-1 (amphibolite)	351.50	9.301

and all had some different features at the scale of 1 to 2 km. The style of the topography fits that along the side of a canyon system or in a highly channeled area.

Magnetic anomaly contours wander aimlessly around the site and have no distinct trend, nor do they show much variation of the magnetic field. This is consistant with the magnetic susceptibility measured on the recovered rocks (see Paleomagnetism section).

Across the site, seismic record GUA-13 shows first a seaward-dipping reflective sequence, then a sequence similar to the prograding sequence at Sites 496 and 568, which is underlain by a zone of very faint reflections



Figure 8. Acoustic characteristics of recovered materials at Site 569.

conforming to the basement topography, and finally a diffracted reflection of relatively high amplitude marking the top of acoustic basement (Fig. 10). Applying laboratory measurements of velocity (see Physical Properties section) to intercepts on the seismic record indicates that the top and bottom of the prograding sequence are 50 and 245 m, and the basement 340 m subbottom. This corresponds with lithology as follows:

Core o an	lepth (m) d age	Seismic record	Seismic sub-bottom depth (m)
0–50	Pleisto- cene	Hemipelagic cover	0-50
50-250	Pliocene- late Oligocene	Seaward dipping reflections, prograding sequence	50-245
250-351	Eocene- Oligocene	Faint reflections conforming with basement	245-340
351	Igneous rock	Top of basement reflection	340

At Hole 569, sudden failure was attributed to elevated pore pressure from the response of pump volume and circulation pressures at the rig floor. When the drill stem became stuck, the pressure was vented, and twice after going to 0 it built up again to 250 lb. without pumping. After freeing the drill stem, pressures returned to normal, indicating no unusual back-flow condition. This is similar to the conditions at Sites 566 and 567 where overpressures also appeared to be present.

PALEOMAGNETISM

Oriented samples of sediment were taken from each section of all cores that were sufficiently consolidated for paleomagnetic measurement. Alternating field demagnetization results for selected samples are plotted on Figure 11. All of samples except one show fairly stable behavior with little change in the inclinations. Figure 12 is stratigraphic plot of inclinations and intensities after alternating field demagnetization at 150 to 175 Oe. The inclinations are very scatter and do not define any pattern of normal and reversed intervals. However, the scattered does appear to decrease in the lower half of the hole where the sediment was more cohesive, suggesting that drilling disturbance was responsible for at least some of the scatter.

It is interesting to note the large variations in intensity of magnetization that occur in Figure 12. The highest intensity occurs in the sample from Core 5, which was taken close to an ash layer. It is possible that variations in volcanic activity occurring close to this site may have been responsible for the variations in intensity of magnetization of these sediments.

Results of A. F. demagnetization of two samples of crystalline basement obtained from Hole 569A are plotted in Figure 13. Both samples are fairly stable, although one (569A-10-1, 98 cm) contained a large vis-



Figure 9. Seismic record GUA-13 showing position of Site 569 (from Ladd et al., 1982). This display, at a vertical exaggeration of about 10 times, is processed as a 24-fold stacked section with a vertical axis in units of time.



Figure 10. A part of seismic record GUA-13 showing detail of reflections. This display has been migrated and thus the highly diffractive basement now appears as a series of short discontinuous reflections. The vertical exaggeration is about 2.4 times and the vertical axis is in units of time.

cous component that is responsible for the erratic variations in both intensity and direction. The interpreted stable inclinations are about 55° for 569A-10-1, 27 cm, and 75° for 569A-10-1, 98 cm. Sample 569A-10-1, 98 cm was taken from a more foliated section of the dolerite. The nonuniformity of its magnetization may explain the 20° difference between the inclinations of the two samples. The difference could also have been caused by overprinting during the tectonic events that produced the foliation. Even an inclination of 55° is considerably higher than the axial dipole inclination for this site, which is about 20°. The result therefore suggests either a tilting of at least 25° about an east-west axis, or that the dolerite was emplaced at a latitude of 35°N. It is interesting that these two inclinations are close to the inclinations of the two serpentinite samples from Hole 556C (84° and 47°). It is possible that the same two episodes of magnetization were sampled in Hole 569.

Magnetic susceptibilities for the two serpentinite samples and the resulting Königsberger ratios (Q_n) appear in the table below.

Sample	NRM (× 10^{-4} emu/cm ³)	Susceptibility (× 10^{-4} cm · g · s)	Q _n
569-10-1, 27 cm	0.21	2.3	0.23
569-10-1, 98 cm	0.041	1.8	0.057

Figure 11. Alternating field demagnetization of selected sedimentary samples from Hole 569.

GEOCHEMISTRY

Gas Analyses

Hydrocarbon gases methane (C₁), ethane (C₂), propane (C₃), isobutane (i-C₄), normal butane (n-C₄), neopentane (neo-C₅), isopentane (i-C₅), and normal pentane (n-C₅) were measured in cores from Holes 569 and 569A. Hole 569 penetrated from the seafloor to a total subbottom depth of 251 m before it was necessary to abandon it. Hole 569A was washed to 246 m and then continually cored to a total sub-bottom depth of 365 m. Results of gas analyses are shown in Table 4.

The use of vacutainers for collection of gases requires that sufficient gas be present to be collected within the core liners, preferably in pockets formed as the sediment expands. Samples of gas were not collected from the first 30 m of core from Hole 569. Sediments in this interval did not show any expansion cracks because of gas but smelled strongly of H₂S. These cores came from the zone of sulfate reduction, and amount of hydrocarbon gas in these sediments undoubtedly was low. From 30 to 67.5 m sub-bottom the cores did not develop, at least while on

Figure 12. Stratigraphic plot of data from Hole 569, demagnetized at either 150 or 175 Oe.

deck, cracks or pockets resulting from gas expansion; therefore, in order to monitor the hydrocarbon gas composition, samples were taken through the end caps at the lower end of the cores (Cores 5 through 8). Below 67.5 m most cores developed gas pockets that could be sampled. Gas samples taken through the end caps usually had low concentration of hydrocarbons, because of the ample opportunity for air to dilute the gas mixture.

Although the concentrations of hydrocarbon gases varied and depended in part on where the samples were taken (end cap, gas pocket, annular space), the methaneethane ratios (C_1/C_2) were remarkably consistent throughout both Holes 569 and 569A (Fig. 14). The ratios

Figure 13. Alternating field demagnetization of two samples of basement from Hole 569A.

ranged between 5900 and 13,600 and averaged 8300 ± 1600 ; the ratios did not decrease exponentially with depth as they did at Sites 569 and 568.

The mixture of hydrocarbon gases throughout the sediments at this site is very simple, being composed mainly of C_1 and C_2 and much lower amounts of C_3 and i- C_4 . Other hydrocarbons are present but usually in amounts less than 0.1 ppm. This simple mixture of hydrocarbons could reflect that the gas is present as a gas hydrate because of the very low amounts of hydrocarbons larger than C_2 ; however, there was no visual evidence of gas hydrates at this site.

Interstitial Water Chemistry

The following inorganic parameters were measured: calcium, magnesium, chlorinity, salinity, alkalinity, and pH. The results are displayed in Figure 15. The slight decrease with depth of chlorinity and salinity may reflect the presence of gas hydrates, as discussed in the Sites 565 and 568 reports.

Summary

At Site 569 the C_1/C_2 ratios are remarkably consistent in all cores sampled and average 8300; the ratios do not decrease with depth as is the general rule. The gas mixture is simple and dominated by C_1 and C_2 , with lower amounts of C_3 and i-C₄. This distribution of hydrocarbons could reflect that the gas is present as gas

hydrate. The decreasing values with depth of chlorinity and salinity support the conclusion that gas hydrates are present. However, no visual evidence of gas hydrates was noted.

SUMMARY AND CONCLUSIONS

Site 569 is in the mid-slope area off Guatemala, in about 2800 m of water, 32 km upslope and 3200 m above the Middle America Trench axis. The objectives at this site were to sample the basement and the cover of sediment immediately overlying it to study the tectonic history of the mid-slope area.

Seismic record GUA-13 across the site shows first a seaward-dipping reflective sequence similar to the prograding sequence at Sites 496 and 568. The sequence is underlain by a zone of very faint reflections conforming to the basement topography, and beneath which a diffracted reflection of relatively high amplitude marks the top of acoustic basement. Application of the velocities measured on cores to the time intercepts in the seismic record indicates that the top and bottom of the prograding sequence are at 50 and 245 m, respectively, and the basement at 340 m. This corresponds well with the depth of lithologic changes found in the cores and shows that an Oligocene-Miocene prograding sequence overlies an Eocene and Oligocene sediment cover, which overlies in turn the igneous basement.

Two holes were drilled at the site: 569 penetrated the late Oligocene before a suspected overpressure zone caused the hole to fail, and 569A penetrated into basement composed of metamorphosed gabbro and diabase beneath Eocene mudstones. The sequence recovered in Hole 569 is:

- 0-58.9 m, Pleistocene and late Pliocene green mud;
 58.9-250.7 m, Miocene to late Oligocene green mudstone (including 58.9-77.9 m upper Miocene, 77.9-87.4 m middle Miocene, 87.4-240 m early Miocene, and 240-250.7 m, late Oligocene light green calcareous mudstone).
- In Hole 569A the sequence is:
 - 246.0-313.1 m, late Oligocene light green calcareous mudstone;
 - 313.1-351.4 m, late Eocene green and black siliceous mudstone, and a few pieces of blue gray mudstone overlying early Eocene mudstone;
 - 351.4–360.9 m, gabbro and diabases metamorphosed to amphibolite and greenschist facies.

Sediment accumulation rates for Hole 569 are approximately 33 m/m.y. for Pleistocene mudstones, 7 m/m.y. for the middle and late early Miocene mudstones, 25 m/ m.y. for early Miocene undifferentiated sediments, and approximately 7 m/m.y. for late Oligocene mudstones.

Unrecovered unconformities are suspected between the late Pliocene and Miocene, the late Oligocene and Eocene, and between the Eocene and the igneous basement rock.

The presence of late Oligocene ash layers in Holes 569 and 569A has interesting implications for the regional tectonic history. One reason for dating the initiation of subduction along the Middle America Trench at the beginning of the Miocene is the first appearance of ash horizons during this time. The recovery of still older

Section (core-section)	Sub-bottom depth (m)	C ₁ (%)	C2 (ppm)	C3 (ppm)	i-C4 (ppm)	<i>n</i> -C ₄ (ppm)	neo-C5 (ppm)	i-C5 (ppm)	<i>n</i> -C5 (ppm)
Hole 569									
5-4	38	0.6	0.96	0.39	0.07	0.02	-	0.07	0.09
6-6	49	1.2	1.60	0.34	0.07			0.45	0.19
7-7	59	4.0	4.80	0.16	0.18	0.05	0.77		0.33
8-3	63	26.0	30.00	0.38	0.41	0.04	0.01	0.01	0.22
9-1	69	9.2	11.00	0.29	0.18	0.09	0.04	0.05	0.12
10-1	78	15.0	17.00	0.16	0.23	0.02	0.08	0.01	0.03
11-2	90	55.0	67.00	0.38	0.75	0.07	0.09	0.06	0.02
12-6	105	87.0	130.00	0.22	1.30	0.04	0.06	0.06	-
13-3	110	88.0	120.00	0.28	1.30	0.04	0.06	0.03	\sim
14-1	117	65.0	94.00	0.27	1.00	0.06	0.09	0.05	0.22
15-3	129	84.0	140.00	0.21	1.20	0.07	0.10	0.09	—
16-1	136	0.7	0.48	0.07	0.05		0.01	-	-
17-6	153	0.5	0.61	0.09	0.10		0.04	0.05	\rightarrow
18-2	156	54.0	60.00	0.17	0.68	0.06	0.07	0.04	0.08
20-1	175	17.0	17.00	0.08	0.30		0.04		0.18
21-2	186	42.0	38.00	0.12	0.52	0.02	0.05	0.05	0.05
22-2	195	11.0	12.00	0.22	0.17	0.04	0.06	0.02	—
23-3	207	94.0	160.00	0.23	1.30	0.04		-	-
24-5	219	90.0	130.00	0.22	1.10	0.03	0.08	0.02	-
25-2	224	92.0	120.00	0.21	1.10	0.01	0.08	0.02	—
26-5	238	89.0	130.00	0.22	0.90	0.03	0.08		\sim
27-3	244	91.0	120.00	0.20	0.83	0.03	0.07		-
Hole 569A									
1-2	248	89.0	110.00	0.21	1.50	0.07	0.07	0.12	0.11
2-2	258	1.9	2.00	0.19	0.09	0.01	0.04	0.13	-
4-1	275	20.0	21.00	0.10	0.38	0.03	0.04		-
6-1	294	84.0	110.00	0.19	1.00	0.18	0.06	0.06	0.11
7-1	304	85.0	110.00	0.37	0.80	0.07	0.14	0.17	_
9-1	333	66.0	70.00	0.19	0.32	0.04	0.15	0.05	0.05

Table 4. Distribution of hydrocarbon gases at Site 569.

Note: - indicates not detected.

Figure 14. Ratios of methane/ethane (C_1/C_2) with depth at Site 569.

ash layers may permit the age of the present subduction setting to be moved back into the late Oligocene.

Paleoecologic analysis of Holes 569 and 569A suggests deposition occurred at abyssal depths during the Miocene through middle Eocene. Additionally, the early Miocene through middle Eocene deposition was probably below or near the CCD. Calcareous benthic foraminiferal tests are replaced with silica in Eocene intervals occasionally preserved at these great abyssal depths and planktonic foraminiferal tests only rarely. Transported material in the abyssal biofacies was primarily from lower and upper middle bathyal biofacies. In assemblages deposited below the CCD, transported outer shelf and upper bathyal species are often the only components of the fauna.

Physical properties for the sedimentary column at this site resemble those obtained at Sites 568 and 496. Bulk densities show a gradual increase from about 1.4 Mg/m³ near subsurface to 1.8 Mg/m³ in the overall trend, with some higher densities reflecting cementation in calcareous mudstone. Three deviations in the physical property trends stand out from the main tendencies at sub-bottom depths of approximately 60, 115, and 352 m. The latter represents the sediment/crystalline rock (gabbro) contact, whereas the uppermost can be correlated to the Pliocene/Miocene boundary. As at Sites 568 and 496, the first two deviations show a more abrupt change in index properties, which is best appreciated in the wetwater content profile.

At Hole 569, sudden failure was attributed to elevated pore pressure as indicated by the response of pump volume and circulation pressures at the rig floor. During sticking of the drill stem the pressure was vented, and twice after going to zero it built up again to 250 lbs. without pumping. After freeing the drill stem pressures

Figure 15. Chemistry of interstitial water at Site 569.

returned to normal, indicating that the pressure cannot be attributed to an unusual backflow condition. This is similar to the conditions at Sites 566 and 567 where overpressures also appeared to be present.

At Sites 569 the C_1/C_2 ratios are remarkably consistent in all cores sampled and average 8300; the ratios do not decrease with depth as is the general rule. The gas mixture is simple and dominated by C_1 and C_2 with lower amounts of C_3 and i- C_4 . This distribution of hydrocarbons could reflect that the gas is present as gas hydrate. The decreasing values with depth of chlorinity and salinity support the conclusion that gas hydrates are present. However, no visual evidence of gas hydrates was noted.

Together with the results from Sites 566 on the lower slope, and 567 at the base of the slope, the results from this site show that the landward slope of the Middle America Trench is constructed of a basement of ophiolitic rocks beneath a cover of slope sediment. The results from Site 569 show that the tectonic emplacement of ophiolitic rocks is at least pre-early Eocene.

This pre-Eocene age of ophiolitic rocks suggests an analogy with on-land geology and confirms the fact that off Guatemala the margin is constructed of an igneous basement belonging to the continental framework of Central America.

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SITE	569	-	HOL	.Ε		_	CC	RE	1 CORED	INTE	RVAL	0.0-1.2 m sub-bottom			
	PHIC		F	OSS	L	3				П	Π				
TIME - ROCH	BIOSTRATIGRA ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS		SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE DISTURBANCE SEDIMENTARY	STRUCTURES SAMPLES		LITHOLOGIC DES	CRIPTI	ON
istocene	leyi						1	0.5			•		Dominant lithology and ashy mottles. Color: grayish olive	r: mud (10Y 4	containing an ashy sand laye /2).
E.	hux		AM			CC		1.0			1.	Pumice clasts	SMEAR SLIDE SU	MMAR	Y (%):
bbb	2		1				-	-						1,77	CC
1	nill												Textures		В
ine.	ũ						1				- 1		Sand	1.27	6
ž	1 C 1						1				- 1		Sile		20
~													Class	-	25
							I .						Composition		15
											- 1		Quarty	25	
							1						Fairtenat	20	2
							L				- 1		Mica	10	i
	1 1				1		1						Heavy minerals	-	Tr
											- 1		Clay	20	82
							1				- 1		Volcanic glass	3	2
			11		L								Glauconite	10	1
											- 1		Pyrite	-	a
											- 1		Carbonate unspec.	4	5
													Foraminifers	5	2
													Calc. nannofossils	-	1
							L .						Diatoms	1	-
		2			1	1	1				- 4		Radiolarians	-	2
					L 1	I 1	L 1						Sponge spicules	-	1
							I 1						Rock fragments	20	Tr
													Wood fragment	± 1	Τr
													CARBONATE BON	18 (% C	aCO ₃):
													1, 100 cm = 5 (shis	board	analysis

SITE 569 HOLE CORE 2 CORED INTERVAL 1.2-11.1 m sub-bottom TIME - ROCK UNIT POSTIFIC - ROCK UNIT POSTIFIC - ROCK POSTIFIC - ROCK RAMINIFIC - ROCK RAMI FOSSIL SECTION DRILLING DISTURBANCE SEDIMENYARV STRUCTURES SAMPLES GRAPHIC LITHOLOGIC DESCRIPTION Dominant lithology: mud with high siliceous biogenic component. Several sandy ash layers, clasts of ash and pumice. Scattered foraminifera in base of Section 5. Color: grayish olive (10Y 4/2). SMEAR SLIDE SUMMARY (%): 1,66 4,120 5,140 M M Texture: Sand Silt 80 20 10 70 20 5 20 75 Clay Composition: Quartz 2 20 - 2 2 Tr 60 15 2 8 2 1 Feldspar Mica Heavy minerats Tr -Tr 81 20 3 8 Clay Volcanic glass 6 - - 2 Glauconite 10 Pyrite 5 Carbonate unspec. 10 Foraminifers 2 3 - 1 Tr 1 6 5 1 1 4 5 Calc. nannofossils -Diatoms 1 Radiolarians cene 3 oceanica Radiolarians – Sponge spicules – Rock fragments 15 Plets esck 1.1 Gephy 4 * Olive gray (5Y 3/2) mottles 5 CC

SITE 569

253

254

×	APHIC		CHA	OSS	TER			GONED		T					
TINU	BIOSTRATIGR	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DMILLING	SEDIMENTARY SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DES	CRIPT	FION
						ĩ	0.5	0.0.0.0					Dominant litholog Mottled in Section Color: olive gray SMEAR SLIDE SU	y: dia 3. (5Y 3 MMAF 1, 13 D	tomaceous mud, weakly layered V21 and grayish olive (10Y 4/2) IY (%): IY 4,50 M
stocene	bad					2	tititititititi	0 0 0 0 0	*****************			Void	Texture: Sand Sift Clay Composition: Feldspar Mica Heavy minerals Clay Volcanic glass Glauconite	17 10 73 Tr 1 73	65 36 - 2 - 8 - 8 -
rie!	szun					3	the set of the	0 0 0	*******************	C* 61		Moderate office brown (5Y 4/4) anottling	Pyrite Carbonate unspec. Foraminifers Diatoms Radiolarians Sponge spicules Silicoflageilates	2 2 Tr 15 1 3 1	
			RP			4	I TOTAL					Disturbed ash layer of light bluish gray (f very light gray (N8), and yellow gray (5Y 8/1)	58 7/1),		

SITE	569		HOL	E		CC	RE	5 CORED	INTER	RVA	40.0-49.7 m sub-bottom
	PHIC		F	RAC	IL						
TIME - ROCK	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE SEDIMENTARY STRINTINGS	SAMPLES	LITHOLOGIC DESCRIPTION
						1	0.5	Q Q	******		Scattered glassy adv, light gray (N7) Dominant lithology: mud, weakly layered as indicated color alternations. Color: grayish ofive (10Y 4/2) and light ofive (10Y 5/ SMEAR SLIDE SUMMARY (K): 1, 121 3, 76 5, 125 CC
						2		2	***************		D D M D Texture: Sand 1 5 Sitt 15 60 Chay 84 35 Composition: Outrtr 10 - Feldspar 2 20 40 1 Mitta 2 - 2 Tr Heavy minerals Tr 5 20 - Clay 85 59 - 89
						3	ind contrarts	1 <u>111</u>			Volcanic glass 1 10 Palagonite 1 Glacconice - Tr Pyrite 2 2 - 1 Dark greenish grav (5GY 5/1) ash Foraminiters - Tr Calc narmofosults Tr 1 Diatoms 3 1 - 1 Hadiolariant 1 1 - 1 Hadiolariant 2 1
Pleistocene						4					(Guayy an - Hgm1 Silicoftagillates 1 2 - 1 grav (N7) Rock fragments - 30 -
						5		6- 	*************		Light aith layers distorted by defiling disturbance
						6	a transferration	0			
			FM			7		Void			Color is 5Y 3/2

	PHIC		F	OSS	TER														PHIC		F
TIME - ROCI	BIOSTRATIGRA ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	STRUCTURIS SAMPLES			LITHOLOGIC DE	SCRIPT	ION			TIME - ROCK UNIT	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS
							-				Ι		Dominant litholo alternations about	igy: mu t 30 cm	id – wi thick.	eakly layered by color in					
						-	0.5			t	t	-Void	Minor lithology: volcanic material;	coarse rounde	sand co	ontaining shell fragments, Fangular material,					
						Ê	1.0						Color: dominant grayish olive (10)	- alt (4/2).	ernating	olive gray (5Y 3/2) and					
												Ashy mottles	Minor lithology:	lark gra	y (N3).						
						1					L		SMEAR SLIDE S	UMMA	RY (%):			0 0	()	11	
							- 3				L			1, 1	17 4, 14	10 7,49					
						1					L		T	D	D	м					
						2	1 -						Sand	-	-	70					
							1.4		1	1	L		Silt	15	15	20					
							1 2				L		Clay	85	85	10					
							1.5						Composition:	1	2						
							-					12/62	Feldsoar	3	2	30	k 8			11	
			2				1 2		11		1	Ash	Heavy minerals	-	1	5				11	
							1 3				h		Clay	85	89	10			1 0	11	
							-	المراجع المالية المالية			Ш		Volcanic glass	Б	-	-					
						1 3	-				Ш		Glauconita	2	-	16			1	11	
						1			1	1	11	Minor lithology	Distoms	2	1	15					
							-				11	Sario Deus	Radiolarians	-	2	-				11	
							1				Ш		Sponge spicules	-	1	-	1 0				
							1 5				11		Silicoflagellates	1.00	2	-	5				В
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stoc							1 2				1						1				
Plai							-														
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		1				F	-	1071270-757	.1		1	stanu as above					1				
								the first of the	11			Reworked, graded at	thy beds				1				
			0		11	7	1		1	1.							1				
			5 KAN					A DESCRIPTION OF A DESC													

×	APHIC		HA	RAC	TER									
TIME - ROC UNIT	BIOSTRATIGR	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SAMPLES			LITHOLOGIC DESI	CRIPTION
						,	0.5					Offive gray (5Y 3/2) eshy bend	Dominant litholog Major difference in cm. Firmer underly Color: grayish oliv erate: olive brown below.	y: mud with ashy layers and clasts n color and firmness in Section 2, 90 ing mud. (107 4/2) to Section 2, 80 cm; mod to grayish olive (5Y 4/4–10Y 4/5)
						2	terri muturu terri	Void			-	Lamineted ash bed Clast of ash Gravish olive (10Y 4/2) Moderate olive brown (5Y 4/4)	SMEAR SLIDE SU Texture: Sand Silt Clay Composition: Ouerz Feldspar Mica Volcanic glass Glauconite Carbonate unspec. Calc. canofosailis	MMARY (%): 1, 135 D 3 12 85 3 2 7 7 5 Tr 1 2
			в			CC	the second second					Dusky blue green (58G 3/2) spot	Diatoms Radiolarians Silicoflagellates	Tr 2 1

SITE 569 CORE 10 CORED INTERVAL 77.9-87.4 m sub-bottom HOLE FOSSIL TIME - ROCK UNIT FORAMINIFERS NANVOFOSSILS RADIOLARIANS SECTION GRAPHIC LITHOLOGY ZONE TARY LITHOLOGIC DESCRIPTION AADIOLAR DIATOMS BIOSTF -..... Pyrite nodules Dominant lithology: mud with several ashy layers. 0.5 Color: grayish olive (10Y 4/2). BUBO LTA AF DTA T Ash bands SMEAR SLIDE SUMMARY (%) 1.0 1, 73 2, 22 niddle I -----Texture Sand 50 30 20 ्य -----12 Silt 2 Sphe Clay 85 cc Composition RCM 5 Chuartz. 1 Feldspar 3 84 2 Clay 12 80 Volcanic glass Glauconite Tr Carbonate unspec -Calc. nannofossils Radiotarians 3 _

1

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Silicoflagellates

SITE 56	9 HOLE	CORE 12 CORED INTERVAL	. 97.1-106.7 m sub-bottom	SITE 569 HOLE CORE 13 CORED INTERVAL 106.7-116.0 m sub-bottom
TIME - ROCK UNIT BIOSTRATIGRAPHIC	FORAMINIFERS FORAMINIFERS FOR ANIANS FOR ANIANANANIANANANANANIANANANANANANANANAN	SB110000 SB111000000 SB110000000 SB110000000 SB110000000000	LITHOLOGIC DESCRIPTION	The character of the ch
lover Micconé lover Micconé			Dominant lithology: structureless mud.	Provide
94	CM	5 5 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	Gray ash bend	

SITE 569

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SITE	569		HOL	.Е		CC	RE	14 CORED	INT	ER	VAL	116.0-125.6 m sub	bottom
×	APHIC		F	OSS	TER								
TIME - ROC UNIT	BIOSTRATIGRU	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DESCRIPTION
						,	0.5		0 C D G D D D D			Dark gray (N3) ash 30° 30°	Dominant lithology: bioturbeted, mottled mudstone. Bedding is at an angle, Fracturing, Color; grayish olive (10Y 4/2).
Iower Miocene						2		Void	9 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0		•	Ashy layer 30 [°] ashy layer	SMEAR SLIDE SUMMARY (%): 2, 37 Texture: Sand – Sith
			RM			3	and and an		••• ۵ م د ۵ م				Glauconite IF Radiolarians Tr Biscuit from Section 3, 90 cm – typical structures application of the standard of the standard structures of extended 1 Non-sandy layers open dewatering 2 Normal faulting 3 Normal faulting
													1 Also cuts 2 and 3 type extensio

PHIC		F	RAC	TER				Π				
BIOSTRATIGRA ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC	DRILLING DISTURBANCE	STRUCTURES	SAMPLES		LITHOLOGIC DESCRIPTION
lover Micenne		RP			1 2 3 CC	0.5		· ·	**	•	40° N1 ashy layer Pale blue green (SBG 7/2) mottling Black (N1) ash, coaley clast	Dominant Lithology: mudstone. Mottled and bioturbation bedding an angle to horizontal, fractured and veined Color: grayish olive (10Y 4/2) to dusky vellow green (BY 5/2). SMEAR SLIDE SUMMARY (%), 3, 94 Texture: Sand – Siti 15 Clav 85 Composition: Ouart 5 Feldpar 3 Mica Tr Clay 85 Gamootien Tr Clay 99 Voicanic glass 5 Glauconite Tr Cation at 5 Feldpar 3 Mica Tr Clay 99 Voicanic glass 5 Glauconite Tr Cationate unspec: 2 Cationate unspec: 1 Spring spicoles 1 Stilooflagetters Tr

	APHIC		FO	SSI	L TER							
TINUT	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARV STRUCTURES	SAMPLES	LITHOLOGIC DE	SCRIPTION
er Middene						1	0.5				Dominant litholo	gy: mudstone.
10Mi			RP			cc	1.0		đ	•	SMEAR SLIDE S	UMMARY (%): 1, 94 D
											Texture: Sand Silt Clay Composition:	5 95
											Quartz Feldspar Ctay Volcanic dass	2 Tr 92 3
											Glauconite Calc. nannofossil Radiolarians	1 . Tr 1
											Sponge spicules Silicoflagellates	1 Tr

ITE	569	1	HO	E	_		co	RE 17 COREC	IN	TER	AVA	. 144.8-154.4 m sub-bottom
	PHIC		CHA	OSS	TER	1						
UNIT - ROCI	BIOSTRATIGRA ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS		SECTION	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
							1	0.5		1		Ash-filled burrow Yellowish gray (SY 8/11 ash Cotor: bluish green (58G 5/2) to pale green (10G 6/ SMEAR SLIDE SUMMARY (%)-
							2					5, 110 Texture: Sand 5 Silt 25 Clay 70 Composition: Quartz 10 Feldspar 5 Mica Tr Clay 71
							3	Void				Volcanic plass 6 Glauconite 1 Carbonate unspec. 1 Diatoms 1 Diatoms 1 Hadiotariani 3 Sponga spiculae 1 Silicoflagellates Yr
fower Miocene							4					
							5	Void				
							6					Ashy Ash Coaley clast Ash
		R	RN	F			7	e	1			Ash

SITE 569 HOLE CORE 19 CORED INTERVAL 164.1-173.8 m sub-bottom

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TIME - ROC UNIT	BIOSTRATIGR/ ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY	SAMPLES	LITHOLOGIC DESCRIPTION
ower Miacene		R	FM			1			0 4 4		Dominant lithology: mudstone, fractured, burrowed, mot- tied, Color: pale green (10G 6/2) to pale blue green (58G 7/2).

₽	<u> </u>	HOL	.t		C	ORE	20 CORED	INTER	VAI	173.8-185.4 m sub-bottom
1		F СНА	OSS	TER						
BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION
lower Miocene	F	RM			1	0.5				Dominant lithelogy: muditone, burrowed, mottled, fra tured and velned. Color: pale green (5BG 7/2) to gravith blue green (5BG 5/2 SMEAR SLIDE SUMMARY (%): 1, 123 Texture: Sand 6 Sitt 14 Clay 80 Composition: Quartz 1 Feddpar 1 Feddpar 1 Feddpar 1 Clay 75 Volcanic glass 10 Glauconitt 5 Diatoms 1
E 569	_	HOL	E	L.	cc	DRE	21 CORED	INTER		183.4-192.9 m sub-bottom
UNIT OSTRATIGRAPH ZONE	ORAMINIFERS	HA STISSOLONN	RAC SNAIAN	TER	NO	s		**		
m	L L	NA	RADIOL	DIATOMS	SECTI	METER	GRAPHIC LITHOLOGY	DRILLING DISTURBANC SEDIMENTAR STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION

	HIC		F	OSSI	L	Ť	T	LE CORED INTERVAL	192.0=202.0 m sub-bottom
TIME - ROCK UNIT	BIOSTRATIGRAP	FORAMINIFERS	MANNOPOSSILS	RADIOLARIANS	DIATOMS	CECTION	METERS	GRAPHIC LITHOLOGY SUNTINUNG SUNTINUN SUNTINUN SUNTINUN SUNTINUN SUNTINUN SUNTINUN SUNTINUN SUNTI	LITHOLOGIC DESCRIPTION
sr Miocene							0.5	Vegi 6	Deminant lithology: pebbly muditone. Class are rounded to subrounded, composed of limestone, shell fragments, greenish mudstone. Color: pale green (10G 8/2). SMEAR SLIDE SUMMARY (%): 2, 140
lowe			RP			cc /	2	Void 6	Téxture: Sand 5 Sit 40 Clay 55 Composition: Felóspar 2 Mica 1 Heavy minarab Clay 88 Volcaring das 3 Carbonate unspec. 5 Radiolarians Tr Sponge spiculas 1

	HIC		F	oss	L					Π			
TIME - ROCK UNIT	BIOSTRATIGRAP ZONE	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DESCRIPTION	
probable lower Miodene						2	0.5		2 4 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	*	Very light (N8) ash Light brownish gray (SYR 6/1) ash spot Light gray (N7) ash	Dominant lithology: mudstone, mottled with abundant fractures and veins, Burrow Color; grayish olive green (5GY 3/2) with olive gray (5Y 5/2). SMEAR SLIDE SUMMARY (%): 1, 20 Texture: Sand 5 Sill 25 Clay 70 Composition: Quartz 1 Feldspar 2 Mica Tr Heavy minerals Tr Clay 89 Volcanic glass 2 Pyrite 1 Diatoms Tr	and burrowe
						4		Void	100		Pumice clast white (N9)	Sponge spicules 2 Silicoffagellates Tr	

SITE 569

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TIME - ROC UNIT	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSIL'S	RADIOLARIANS		DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION
probable lower Micorne	11001	FORAM	NAMAG	RADIOL		04709	3	0.5		1011010 10 40 40 40 0 0 0 0 0 0 0 0 0 0		sample	Dominant lithology: mudstone, with elightly calcurate burrows, burrows, and winantion and fractures. Color: grayish olive green (5GY 3/2) to grayish oli (04 4/2). Salay structure: BMCAR SubMAARY (%): <u>5,98 CC</u> <u>M</u> D Texture: Sand <u>5</u> 1 Mice <u>5</u> 3 Mice <u></u>
							6			Do-ADA D			Brazek azu:

×	APHIC		F CHA	OSSI	L TER								
UNIT UNIT	BIOSTRATIGR	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DESCRIPTION
						1	0.5						Dominant lithology: burrowed and fractured mudston Color: grayish olive green (SGY 3/2).
arty Miodene						-		7.53636.44.94 	0 0 0 0 0	1	•	Light gray (N7) glassy ash Olive gray (SY 4/1) ash	SMEAR SLIDE SUMMARY (%): 1, 142 Texture: Sand 2 Siti 10 Citure 88
Probable 6						2	ALCONDUCT.		000				Composition: Feldspar 3 Mica Tr Heavy minerals Tr Clav 89 Volcanic gilas 1 Citocomic 1
						3	11111		а 6 4 7				Pyrite Tr Diatomi 1 RadioBarians 1 Sponge spicules 4

	FOSSIL								
APH	CHARACTER	_							
UNIT BIOSTRATIGR ZONE FORAMINIFERS	NANNOFOSSILS RADIOLARIANS DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING	SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DESCRIPTION
ou 18 19 19 10 10 10 10 10 10 10 10 10 10 10 10 10		3	0.5		80 > 0 8 8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	18 tt	SA	Very light gray (N8) sah Dusky blue green (68G 3/2)	Dominant lithology: mudstone, burrowed, fractured and veined. Color: gravish olive (10Y 4/2) with light olive brown (5Y 5/6) burrows.

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TIME - ROC UNIT	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES		LITHOLOGIC DESC	RIPTION
80	ctus					,	0.5				60° bedding 30° some veina	Dominant lithology: veined with slightly Color: gravish olive erate olive brown (1 and dusky blue (5PR	mudstone, burrowed, fractured an calcareous burrows, Zoophycus type (10Y 4/2), Burrows filled with mod Y 4/4), light elive brown (BY 5/6) 3/2).
upper Oligoce	Dictyococcites bise					2	and an advance					SMEAR SLIDE SU Texture: Sand Silt Clay Composition:	MMARY (%); 3,35 D
						3	treat treat					Duartz Feldspor Mica Clay Volcanic glass Glauconite Calc. nannofosails Diatoms Radiolarians Sponge spicules Sulice futeritates	2 1 Tr 85 1 10 - Tr 1 Tr
						C	c		8			Siliconagenates	

SITE 569 HOLE CORE 27 CORED INTERVAL 241.0-250.7 m sub-bottom

SITE 569 HOLE A	CORE H1	CORED INTERVAL	0.0-246.0 m sub-bottom (wash)	SITE	569	HO	LE A	CC	ORE	1 CORED	INTERVA	L 246.0-255.5 m sub-bottom
	SECTION	GRAPHIC SBIRLDIN BRAPHIC SBIRLDIN SCHOLOGY SHITLING STHOLOGY	LITHOLOGIC DESCRIPTION	TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FORAMINIFERS	ARADIOLARI ANDIOLARI	SECTION	METERS	GRAPHIC LITHOLOGY	OBILLING DISTURBANCE SEDIMENTARY SIRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION
probable forer Minoure unzoned	0.5 1 10 2 2 3 3		Wash core. Dominant lithology: mudstone with mixture of veined and burrowed biscuits and clasts and soupy to brecciated mud matrix. Color: olive gray (5Y 3/2) to gravish olive green (5GY 3/2). Olive gray (5Y 3/2) Color change Grayish olive green (5GY 3/2)		Leared			1 2 3 4 5 5	1.0			CaCO ₃ = 43% Dominant lithology: mottled, burrowed, fractured and veined mudstone. Burrows are calcureous. Color: olive gray (SY 3/2) with moderate olive brown (SY 4/4) mottles. SMEAR SLIDE SUMMARY (%): CC Texture: Sand 3 Sitt 15 Clay 82 Composition: Quartz 1 Feldspar 3 Mica 1 Heavy mineralt 1 Clay 89 Volcanic glas 3 Pivite 1 Radiolarians 1 Silicoflagellates Tr

SITE 569 HOLE A CORE 6 CORED INTERVAL 294.0-303.6 m sub-bottom FOSSIL CHARACTER ROCK METERS 22 NO DISTURBANCE SEDIMENTARY STRUCTURES BIOSTRATIGE GRAPHIC LITHOLOGIC DESCRIPTION TIME - R LITHOLOGY SECT 6 Dominant lithology: fine claystone; slightly calcareous. 0 1 0.5 a Minor lithology: massive mudstone. 0 Oligo narren Color: mudstone is olive gray (5Y 4/1). Claystone is pale 0 blue green (5BG 6/6). addr 110 SMEAR SLIDE SUMMARY (%): 1,82 Texture Sand Silt Clay 97 Composition Clay 70 Volcanic glass Tr Carbonate unspec. 20 Calc. nannofossils 10 Calc. nannofossils

×	APHIC		F	OSSI	TER								
TIME - ROC UNIT	BIOSTRATIGRA	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC D	SCRIPTI	ON
pocene	s bisectus					,	0.5				Light olive gray (5Y 5/2) modstone Color: pale blu	ogy: calc green (51	areous claystone. BG 6/6).
upper OG	ctyococciter		FM			-	1.0	Secretary L			(N4) ash SMEAR SLIDI pyrite	SUMMA 1, 1	RY (%): 8 1,84
	Dia										Texture: Sand Silt Class	- 5	25 30
											Composition; Quartz Feldspar	1 Tr	5 2
											Clay Volcanic glass Zeolite		68 20 Tr
											Carbonate uns Cale: nannofos	ec. 10 Is 6	5
											CARBONATE 1, 105 cm = 15	SOMB (% (shipboar	CaCO ₃)

SITE 569	9 F	IOL	E	A	C	ORE	8 CORED	INTER	VAL	2.3 m sub-bottom	
CK		FO	RAC	TER	_						
TIME - RO UNIT BIOSTRATIGR	FORAMINIFERS	NANNOFOSSIL	RADIOLARIANS	DIATOMS	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION	
upper Eodene Discoaster barbachensis		FP				0.5	Void	444	*	Dominant lithology: mottled, biotu tured. Color: pale blue green (5BG 7/2) wit (5BG 5/8) and light olive gray (5Y 5/ Minor lithology: medium bluish gray gray (5G 6/1).	rbated, veined, frac- h mottles of pale blue 2). r (58 5/1) to greenish
										SMEAR SLIDE SUMMARY (%): 1, 54 Texture: Sand 8 Sitt 22 Clay 70 Composition: Ouerr 3 Feldspair 1 Mica Tr Clay 90 Volcenic glass 5 Zeolite Tr Cale, namofossiht Tr Radiolarians 1 Sponge spiceles Tr	

TE	269a	Г	F	OSSI	L		RE	GORED	INTER		L 332.3-351.4 m sub-b	ottom		-		
UNIT UNIT	BIOSTRATIGRAF	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS	SECTION	METERS	GR APHIC LITHOLOGY	DISTURBANCE SEDIMENTARY	SAMPLES		LITHOLOGIC DESC	RIPTIO	N		
upper middle Eccene	Discoutler automentes		FP			1	0.5				Pate blue gray (5BG 7/2) Light olive gray (5Y 5/2) Olive gray (5Y 3/2) Grayish olive (10Y 4/2), calererous Brownish black (5YR 2/1) scally fabric (2) Pale olive (10Y 6/2)	Dominant Titholog turred mulations & slanted contact. Radiolarian rich. Minor fithology: (as indicated as an a	y: burrowed bithology Darker Inthe sen lith MMAR'1 1, 14 30 40 - 3 30 40 - 1 1 45 - 20 5 - - 20 5 - -	wed, n sove, ex y colum mudsto ory fabri sudston o facies 1, 54 8 22 70 3 1 Tr - 90 5 - Tr Tr Tr Tr Tr acOogli:	hottled, ve 60-80%, orpt with nn and b ne with high burrons and b with burrons are sharp. CC, top 20 555 22 1 2 2 - Tr 7 6 3 3 - 1 - 5 2 2 4 6 5 -	ined and frac 0 - > 80 cm different color nurrowed over lever burrow we (a) $0 - 0$ ad or silicor CC, 17 10 30 60 61 - 58 3 - - - 6 5 5 -
												1, 15 cm = 20 (ship	NB (% C board a	aCO ₃): nalysis)		

84-569A-10-1

0-5 cm: pair blue green calcarcous serpentinite modstone.

Pieces 1—10: dark greenish yray (5G 4/1) amphibolite, metamorphically altered to amphibolite facies then intro-graded to zeolite facies. Foliated and fracturied containing plagiculas veius that are also offset by the fractures that pervade the rock. Prehnite is present in Piece 3C and prehnite and pompetiyite are present in Pieces 6–10. Plagicolare veining is present in isome sampler and may be replaced by natorities.

84-569A-10-2

Greenish gray (5G 4/1). Foliated amphibolite with plagoclase veim.

84-5694-11-1

Depth 360.9 364.0 m

Depth 351 4-364,9 m

Depth 351.4 364.9 m

Foliated diabase with plagloclase(?) veins, dark greenish gray (5G 4/1).

268

-0 cm	7-4	7-5	7-6	7-7	7,CC	8-1	8-2	8-3	8,CC	9-1	9-2	9,CC

-0 cm 12,CC	13-1	13-2	13-3	13-4	13,CC	14-1	14-2	14-3	14,CC	15-1	15-2
-0 cm 12,CC 					13,CC		14-2	14-3	14,CC		15-2
-											1 Th
150	have .					-					

-0 cm 24-6	24,CC	25-1	25-2	25-3	25,CC	26-1	26-2	26-3	26-4	26-5	27-1
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-0 cm H1-1	H1-2	H1-3	H1-4	H1,CC	1-1	1-2	1-3	1-4	1-5	1,CC	2-1
-0 cm +11.1	H1-2		H1-4	H1,CC				1-4	1-5	1,CC	

