

5. SITE 568¹

Shipboard Scientific Party²

HOLE 568

Date occupied: 8 February 1982, 1730 hr.

Date departed: 12 February 1982, 1615 hr.

Time on hole: 94 hr., 24 min.

Position: 13°04.33'N; 90°48.00'W

Water depth (sea level; corrected m, echo-sounding): 2010

Water depth (rig floor; corrected m, echo-sounding): 2020

Bottom felt (m, drill pipe): 2031

Penetration (m): 417.7

Number of cores: 44

Total length of cored section (m): 417.7

Total core recovered (m): 308.4

Core recovery (%): 73.8

Oldest sediment cored:

Depth sub-bottom (m): 417.7

Nature: mudstone

Age: earliest Miocene

Measured velocity (km/s): 1.595

Basement: not reached

Principal results: Drilling ended at 417.7 m in order to stay above the base of gas hydrate as required by the Safety Panels.

The main results concern the evolution of gaseous components through the section and the problem of gas hydrates: drilling of Hole 568 was dedicated to this problem. Sequence of gas components is as follows:

—from 0 to 190 m sub-bottom, the composition of gas is normal for a biogenic source;

—from 190 to 345 m, the dominance of methane and ethane, disproportionately low or absent heavier molecules, and the low values of salinity and chlorinity suggest a dispersed hydrate;

—from 345 to 391 m, the composition returns to normal with a normal proportion of heavier hydrocarbons suggesting gas in the free state;

—from 391 to 410 m, hydrates were recovered in sediment with large quantities of methane and ethane, low quantities of heavier molecules, and low values of salinity and chlorinity;

—from 410 to 417.7 m, the section is again not hydrated.

Hydrates form only in relatively porous substratum, either porous sediments such as sandstones and ash layers, or fractured nonporous sediment. The recovered hydrate at Site 568 was in fractures in a tuffaceous mudstone; therefore, hydrates probably do not form a continuous layer at the site; no seal is formed. Temperature measurements and seismic records place the base of gas hydrates from 20 to 40 m below the level where drilling was abandoned.

Other results of the drilling at Site 568 confirm the results from Leg 67, Site 496. The sedimentary sequence consists of an upper underconsolidated green mud from 0 to 182 m of the middle to late Pleistocene, and a lower indurated green mudstone from 182 to 417 m, ranging from early Pliocene to earliest Miocene. The progressive increase of dip with depth reached 70° in Core 43. Seismic data indicate that the site is on a depositional lobe, and slumping at the front of the lobe may contribute to the steep dip seen locally in the core.

The comparison with the early Miocene sediment recovered at Site 567 shows the local origin of reworked material at that site, and suggests that the original unstable topography of the slope was buried during the early Miocene, whereas the sedimentation of overlying formations was more regular. If this is so, the prograded sections at the base of slope deposits could be everywhere considered early Miocene, which thus questions the origin of the paleotopography of the slope.

BACKGROUND AND OBJECTIVES

Site 568 is on the upper part of the Middle America Trench slope, in about 2031 m of water, 4000 m above and about 47 km landward from the Trench axis. It is located about 1 km upslope from Leg 67 Site 496, on the bearing 300°.

Site 496 had been abandoned at a depth of 378 m after a gas composition was encountered that suggested possible thermogenic components and perhaps gas hydrate. After Leg 67, a good base of gas hydrate reflection was found beneath the site. Therefore, the major objective at Site 568 was a detailed monitoring of the gas through the whole section in a study of the formation of gas hydrate.

Seismic record GUA-13 was reprocessed for this purpose to bring out the shallow structure and the base of gas hydrate seismic reflector. Within the Safety Panel recommendations, drilling at Site 568 could penetrate to 100 m above the reflector or the base of hydrate, as calculated from temperature data, and at least to the same depth as at Site 496.

OPERATIONS

Glomar Challenger departed from Site 567 at 1245L (local time), 7 February. At 1600L a 16-kHz beacon was dropped and the ship was in position at 1730L. From 1708, 8 February, to 0330L, 8 February, a new bottom-hole assembly was made up; then until 1000L, 117 joints were moved from between decks to the starboard casing rack. Running in the hole began at 1030L, while

¹ von Huene, R., Aubouin, J., et al., *Init. Repts. DSDP*, 84: Washington (U.S. Govt. Printing Office).

² Roland von Huene (Co-Chief Scientist), U.S. Geological Survey, Menlo Park, California; Jean Aubouin (Co-Chief Scientist), Département de Géotectonique, Université Pierre et Marie Curie, Paris, France; Miriam Baltuck, Scripps Institution of Oceanography, University of California, San Diego, La Jolla, California (present address: Department of Geological Sciences, Tulane University, New Orleans, Louisiana); Robert Arnott, Department of Geology, University of Oxford, Oxford, United Kingdom (present address: Shell International, The Haag, Holland); Jacques Bourgeois, Département de Géotectonique, Université Pierre et Marie Curie, Paris, France; Mark Filewicz, Union Oil Company, Ventura, California; Roger Helm, Institut Für Geologie, Ruhr-Universität Bochum, Bochum, Federal Republic of Germany; Keith A. Kvenvolden, U.S. Geological Survey, Menlo Park, California; Barry Leinert, Hawaii Institute of Geophysics, University of Hawaii, Manoa, Honolulu, Hawaii; Thomas J. McDonald, Department of Oceanography, Texas A&M University, College Station, Texas; Kristin McDougall, U.S. Geological Survey, Menlo Park, California; Yujiro Ogawa, Department of Geology, Kyushu University, Hakozaki, Fukuoka-Shi, Japan; Elliott Taylor, Department of Oceanography, Texas A&M University, College Station, Texas; Barbara Winsborough, Espey, Houston, and Associates, Austin, Texas (present address: Department of Geology, Princeton University, Princeton, New Jersey).

a new drill string was reconstructed and the remaining joints checked. After spudding the hole at 2100L, drilling began; the first core was recovered at 2236L.

In situ pore water and heat flow measurements were done after recovering Cores 10 (sub-bottom depth 89.4 m); 15 (137.2 m); 20 (185.7 m); 30 (282.3 m); and 38 (359.9 m). A pressure core barrel was used for Cores 11 (sub-bottom depth of 89.4–98.8 m); 21 (185.7–195.4 m); and 31 (282.3–292.0 m).

After recovering a piece of crystallized hydrate at less than 100 m from the base of hydrate reflection, drilling was stopped at 1615L, 11 February, to wait for the decision from DSDP on further drilling. This decision was negative and the hole was given over to logging.

The bit was released from 2030 to 2130L; the hole was filled with weighted mud from 2130 to 2215L, and logging began at 0030L, 12 February.

After sonic, waveform, and density-neutron logs were run, the drill string was retrieved from 1145 to 1600L.

Glomar Challenger departed Site 568 at 1615L, 13 February.

Table 1 shows the coring summary for Site 568.

LITHOSTRATIGRAPHY

Site 568 is located within 1.5 km of Leg 67 Site 496 at a depth of 2031 m (drill pipe) on the upper slope of the Middle America Trench, about 47 km from the Trench axis (Fig. 1). Forty-four cores were drilled with very good recovery to a sub-bottom depth of 418 m. We divide the section penetrated into two lithostratigraphic units.

The contact with the lower unit is identified in Core 20 on the basis of color change, a slightly coarser texture, and a difference in deformation style from swirling of soft mud to breaking of stiff mudstone into squarish "biscuits." Unit I is early-middle Pleistocene at its base in Core 20; Unit II is early Pliocene to middle Miocene in Core 21 (ages based on shipboard analyses of core

Table 1. Coring summary, Site 568.

Core	Date (Jan. 1982)	Time	Depth from drill floor (m)		Depth from seafloor (m)		Length cored (m)	Length recovered (m)	Recovery (%)
			Top	Bottom	Top	Bottom			
1	8	2236	2031.0	2034.4	0.0	3.4	3.4	3.43	100
2	8	2342	2034.4	2044.0	3.4	13.0	9.6	9.55	99
3	9	0055	2044.0	2053.7	13.0	22.7	9.7	8.55	88
4	9	0145	2053.7	2063.3	22.7	32.3	9.6	8.47	88
5	9	0230	2063.3	2073.0	32.3	42.0	9.7	9.57	99
6	9	0325	2073.0	2082.4	42.0	51.4	9.4	7.31	78
7	9	0410	2082.4	2092.1	51.4	61.1	9.7	7.72	80
8	9	0515	2092.1	2101.4	61.1	70.4	9.3	9.70	100
9	9	0610	2101.4	2111.1	70.4	80.1	9.7	7.61	78
10	9	0656	2111.1	2120.4	80.1	89.4	9.3	6.51	70
11	9	0955	2120.4	2129.8	89.4	98.8	9.4	1.05	11
12	9	1045	2129.8	2139.3	98.8	108.3	9.5	6.69	70
13	9	1150	2139.3	2148.8	108.3	117.8	9.5	3.87	41
14	9	1250	2148.8	2158.4	117.8	127.4	9.6	9.90	100
15	9	1358	2158.4	2168.2	127.4	137.2	9.8	6.05	62
16	9	1825	2168.2	2177.9	137.2	146.9	9.7	5.39	56
17	9	1925	2177.9	2187.5	146.9	156.5	9.6	6.46	67
18	9	2028	2187.5	2197.2	156.5	166.2	9.7	6.07	63
19	9	2132	2197.2	2207.0	166.2	176.0	9.8	8.50	87
20	9	2235	2207.0	2216.7	176.0	185.7	9.7	9.56	99
21	10	0205	2216.7	2226.4	185.7	195.4	9.7	1.61	17
22	10	0310	2226.4	2236.1	195.4	205.1	9.7	9.14	94
23	10	0420	2236.1	2245.8	205.1	214.8	9.7	4.62	48
24	10	0525	2245.8	2255.5	214.8	224.5	9.7	9.67	99
25	10	0626	2255.5	2265.2	224.5	234.2	9.7	9.34	96
26	10	0730	2265.2	2274.4	234.2	243.9	9.7	0.00	0
27	10	0830	2274.4	2284.4	243.9	253.4	9.5	5.83	61
28	10	0940	2284.4	2294.1	253.4	263.1	9.7	9.62	99
29	10	1110	2294.1	2303.6	263.1	272.6	9.5	8.51	90
30	10	1220	2303.6	2313.3	272.6	282.3	9.7	9.05	93
31	10	1555	2313.3	2323.0	282.3	292.0	9.7	1.16	12
32	10	1739	2323.0	2332.6	292.0	301.6	9.6	8.86	92
33	19	1900	2332.6	2342.4	301.6	311.3	9.7	9.63	99
34	10	2025	2342.3	2352.2	311.3	321.2	9.9	9.52	96
35	10	2134	2352.2	2362.0	321.2	331.0	9.8	9.03	92
36	10	2254	2362.0	2371.7	331.0	340.7	9.7	9.20	95
37	11	0025	2371.7	2381.2	340.7	350.2	9.5	4.58	48
38	11	0210	2381.2	2390.9	350.2	359.9	9.7	9.61	99
39	11	0550	2390.9	2400.6	359.9	369.6	9.7	3.03	31
40	11	0735	2400.6	2410.3	369.6	379.3	9.7	7.01	72
41	11	1100	2410.3	2420.0	379.3	389.0	9.7	8.83	91
42	11	1245	2420.0	2429.7	389.0	389.7	9.7	6.11	63
43	11	1413	2429.7	2439.4	398.7	408.4	9.7	5.71	59
44	11	1546	2439.4	2448.7	408.4	417.7	9.3	6.77	73
Total							417.7	308.40	74

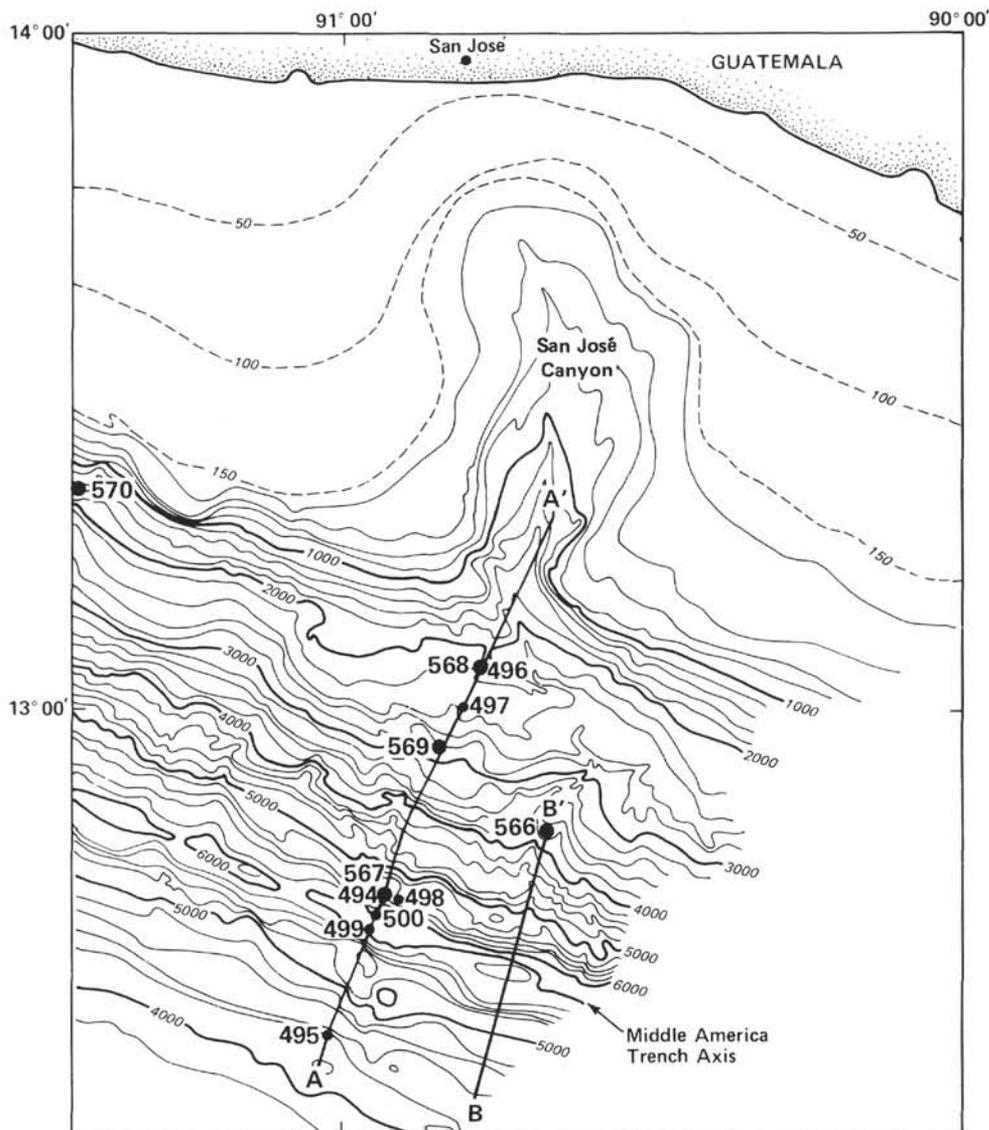


Figure 1. Bathymetry of the Guatemala margin and San José Canyon, showing the UTMSI survey tracks and the locations of Legs 67 and 84 sites off Guatemala.

catchers; see Biostratigraphy) and thus the contact probably represents a hiatus. Figure 2 shows a lithostratigraphic summary of Site 568.

Unit I

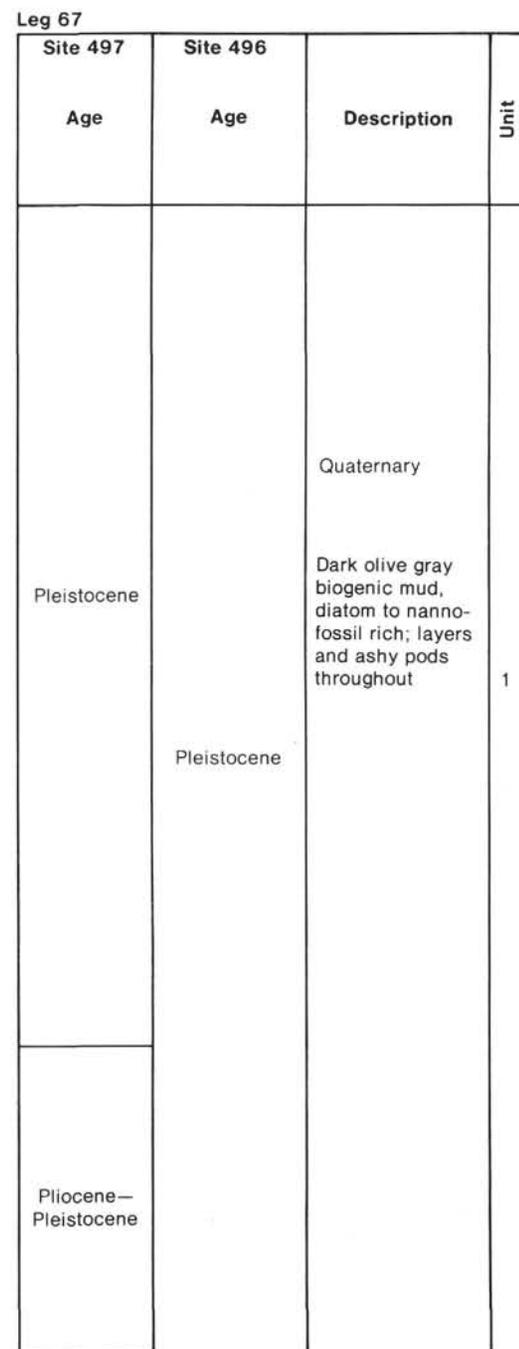
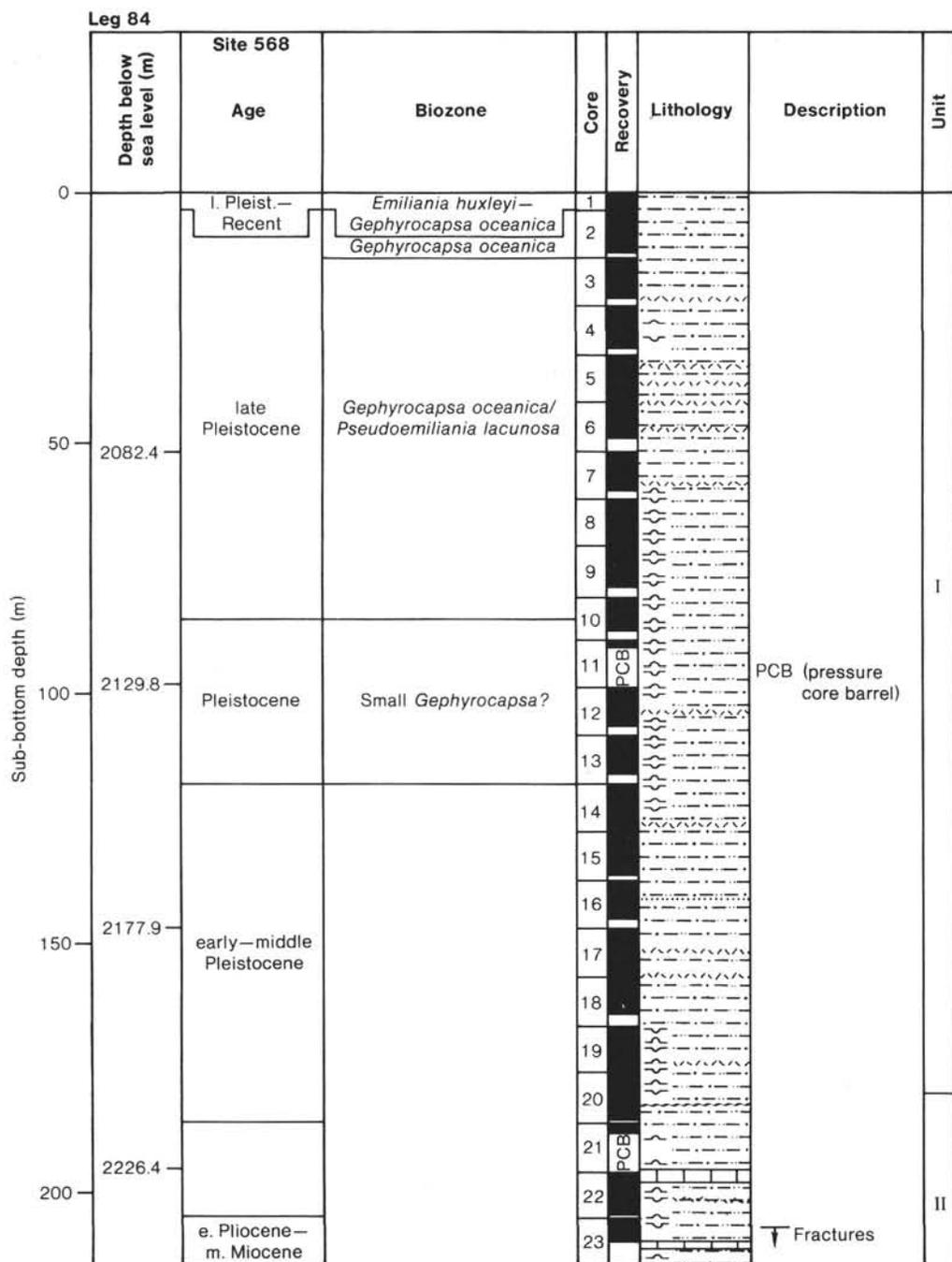
Unit I comprises Cores 1 to 20, 0 to 182.0 m sub-bottom depth, from Recent to late Pleistocene, and is composed of a massive, dark olive gray (5Y3/2) mud that is easily deformed into swirled layers during drilling. The average sand-silt-clay composition of the unit (based on shipboard smear-slide analyses) is 6, 29, and 65%, respectively, with a slight coarsening downhole. Siliceous biogenic remains (diatoms, radiolarians, silicoflagellates, and sponge spicules) make up more than 15% of the total sediment in Cores 8 to 17 and 19, and sandy horizons occur in Cores 1, 8, and 16. Foraminifers are dispersed in the sediment and fill a burrow in Core 19. Thin carbonaceous beds and clasts (including carbonaceous wood fragments) occur in Cores 8 and 9. Large (2–4 cm) bivalve shell fragments were found in Cores 1,

2, and 20; in Core 1 two valves are still attached by ligament.

Cores 1 through 3 smelled strongly of H_2S after splitting. Gas expansion cracks (mostly horizontal to sub-horizontal and about 1 cm long) occurred in the firmer parts of soft cores. Toward the top of the hole, gas expansion in more indurated mud resulted in coherent breaks across a core creating voids in the cross sections. This phenomenon resulted in a slight exaggeration of the figure for percent of core recovered.

Bedding laminations in otherwise structureless cores are weakly preserved in parts of Cores 1, 2, 8, 12 to 14, and 16 to 18.

Ash layers occur in Cores 3, 5 to 9, 12, 14, and 17 and are somewhat more common in the upper half of the unit. Twenty-three layers of ash or muddy ash and numerous ashy mottles and pumice clasts are discerned. These range in colors from light gray (N7), medium bluish gray (5B 5/1), olive black (5Y 2/1), to pale olive (10Y 4/2).



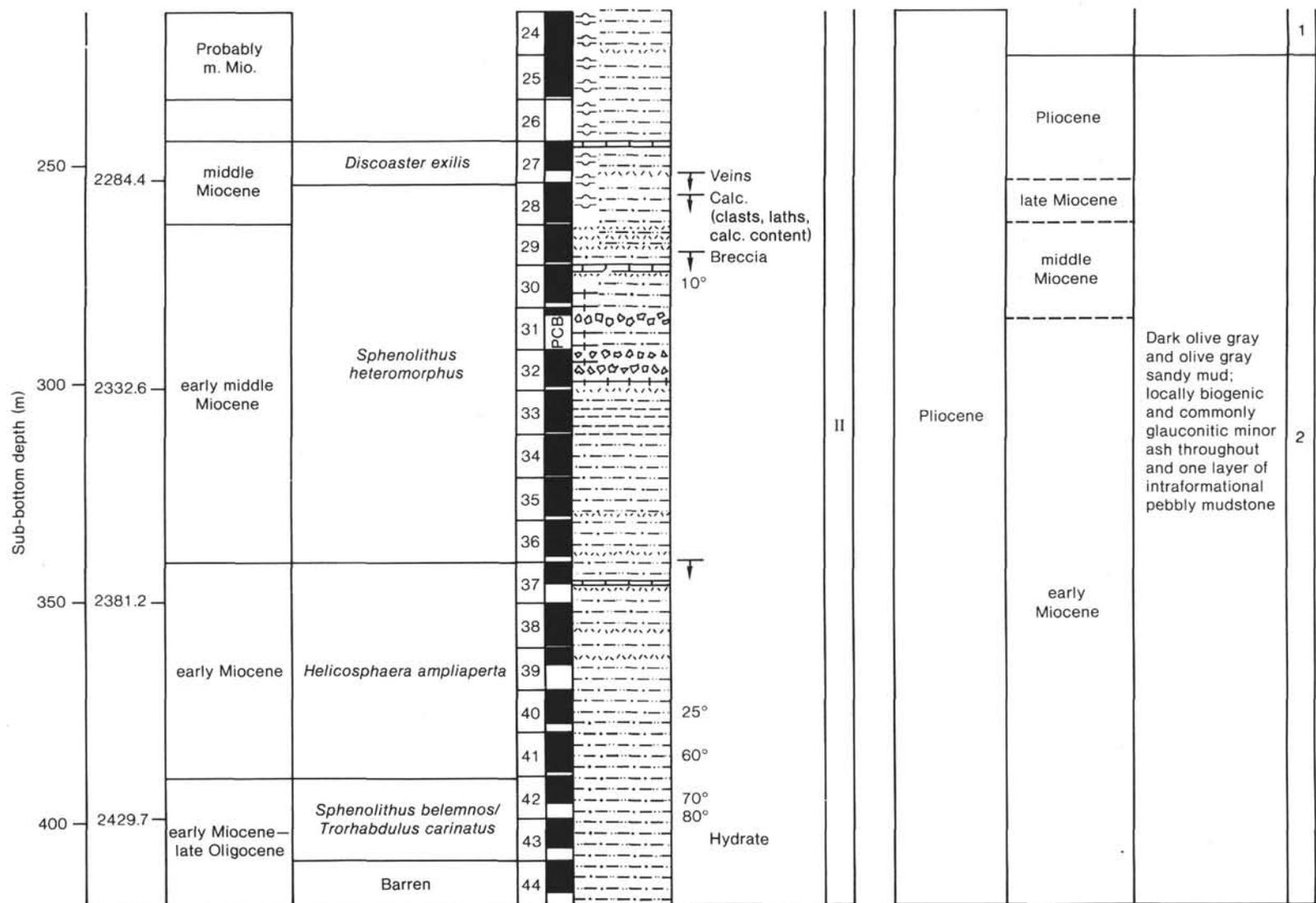


Figure 2. Lithostratigraphic summary of Site 568 and comparison with thicknesses of equivalent-aged sections at Sites 496 and 497. Apparent dip of bedding in Description column noted where observed in the hole adjacent to Lithology column.

Unit II

Unit II (Cores 20 to 44, 182.0–417.7 m sub-bottom depth from early Pliocene to earliest Miocene) is a mottled and bioturbated mudstone, generally grayish olive (10Y 4/2) to grayish blue green (5BG 5/2). These colors occur as mottles and pebble clasts and as matrix material. The average sand–silt–clay composition is 11, 35, and 54%, respectively (as was determined by shipboard smear-slide analyses). This unit contains at least 15% siliceous biogenic remains (diatoms, radiolarians, silicoflagellates and sponge spicules), and several horizons of slightly calcareous mudstone (Cores 20, 21, 22, 28, 29), nannofossil ooze (Cores 21, 31–33), and limestones and limestone conglomerate (Cores 20–23, 30, and 37), as well as isolated limestone clasts throughout the core. Thin mudstone breccia layers with angular to subangular mudstone clasts occur in Cores 31 through 32. The bottom half of Core 33 is very fine-grained calcareous mudstone–claystone composed almost entirely of clay and nannofossils.

Light to moderate olive brown mottling and clasts (5Y 5/6 to 5Y 4/4) occur in Cores 42 through 44. These contain reworked Eocene nannofossil and foraminifer fossils (see Biostratigraphy).

Bedding is roughly horizontal in the upper half of Unit II, but in Cores 29 and 30 the apparent dip of bedding is 10°, Cores 40 to 42 show some apparent dips of 30 to 60°, and in Core 43 locally preserved bedding exhibits apparent dips of 70 to 80°.

We distinguished 25 ash layers (Cores 23, 24, 27, 29, 30, 33, 35–37, and 39) as well as numerous ashy mottles and pumice clasts in the upper part of the unit (above Core 40).

Small-Scale Deformation in Unit II

Small-scale fractures such as those described from Site 566 were observed offsetting beds beginning in Core 22. Undeformed beds commonly overlie fractured beds (Fig. 3). The subhorizontal–horizontal scaly fabric observed at Sites 565, 566, and 567 is also observed at Site 568 from Core 20 to the bottom of the hole. As at previous sites, the development of scaly fabric is attributed to the alignment of clays during sediment mass movement. Possible slickenside-type lineation is observed along some of the scaly surfaces, particularly toward the bottom of the hole.

From Core 28 to hole base, deformation structures called “veining” are preserved in mudstone biscuits. Veins are about 1 mm thick and usually 1 to 3 cm long. The veining occurs in two main styles. Figure 4A shows an anastomosing network of small veins of horizontal to subhorizontal orientation relative to bedding. Figure 4B illustrates the second style of veining, where a series of subparallel veins are planar to curvilinear or even sigmoidal. These veins are usually only a few millimeters apart. Orientation of this second type of veining ranges from subhorizontal to vertical relative to bedding planes, with subvertical to vertical orientation increasing in proportion downhole. Cowan (1982) studied the

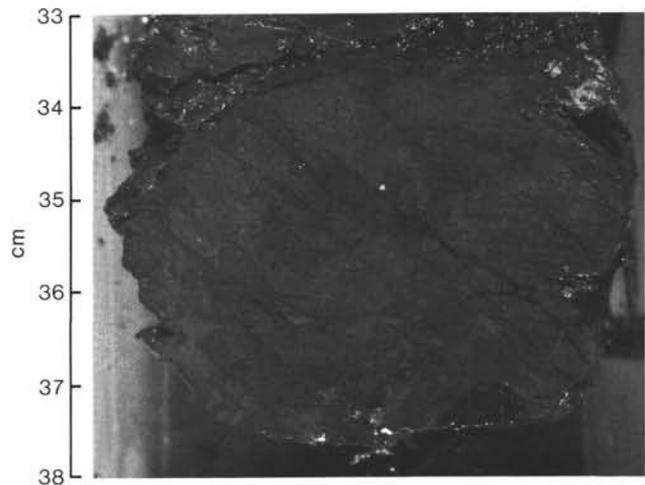


Figure 3. Fractured bed over- and underlain by undisturbed beds (Sample 568-34-1, 33–38 cm).

parallel “veining” of Site 496 in detail for mineralogy and petrographic fabric. The vein fillings were compositionally and texturally similar to the surrounding mud matrix except for very fine-grained phyllosilicates, which are preferentially oriented parallel to vein boundary. Cowan interprets these veins as dewatering conduits geometrically analogous to extension fractures. Deformation of veins may have occurred in response to extension during downslope mass movement of sediments on the upper slope of the Middle America Trench.

Post-Miocene Sedimentation at Site 568

Three observations can be interpreted in several ways to model sedimentation at Site 568. (1) There is a difference of nearly 200 m between thicknesses of Pliocene–Pleistocene section at Sites 568, 496, and 497, all within about 10 km of each other. (2) Seismic stratigraphy shows considerable relief along an acoustic basement that may represent an unconformity (see Geophysics section). (3) Bedding inclination shows an increase concentrated in Core 40 from only about 10° in Cores 29 and 30 to nearly vertical in Cores 42 and 43.

There is abundant indication of mass movement and small-scale deformation in the presence of scaly fabric and fractures, and of nonuniform horizontal pressure on the sediments in the dewatering veins and extension cracks in the Miocene sediments. Such slumping and/or folding of coherent units may account for the high-angle bedding dip observed at Site 568.

Finally, the presence of reworked Eocene material in clasts at the base of the hole can be accounted for by at least two very different origins. The microfossils in sediment from the bottom of the hole indicate an age of early Miocene–late Oligocene (see Biostratigraphy section). Either the reworked material is very locally derived from underlying Eocene basement (in which case the drilling was halted quite close to the unconformity—if that is what the acoustic reflection represents) or it was transported downslope from an upslope exposure

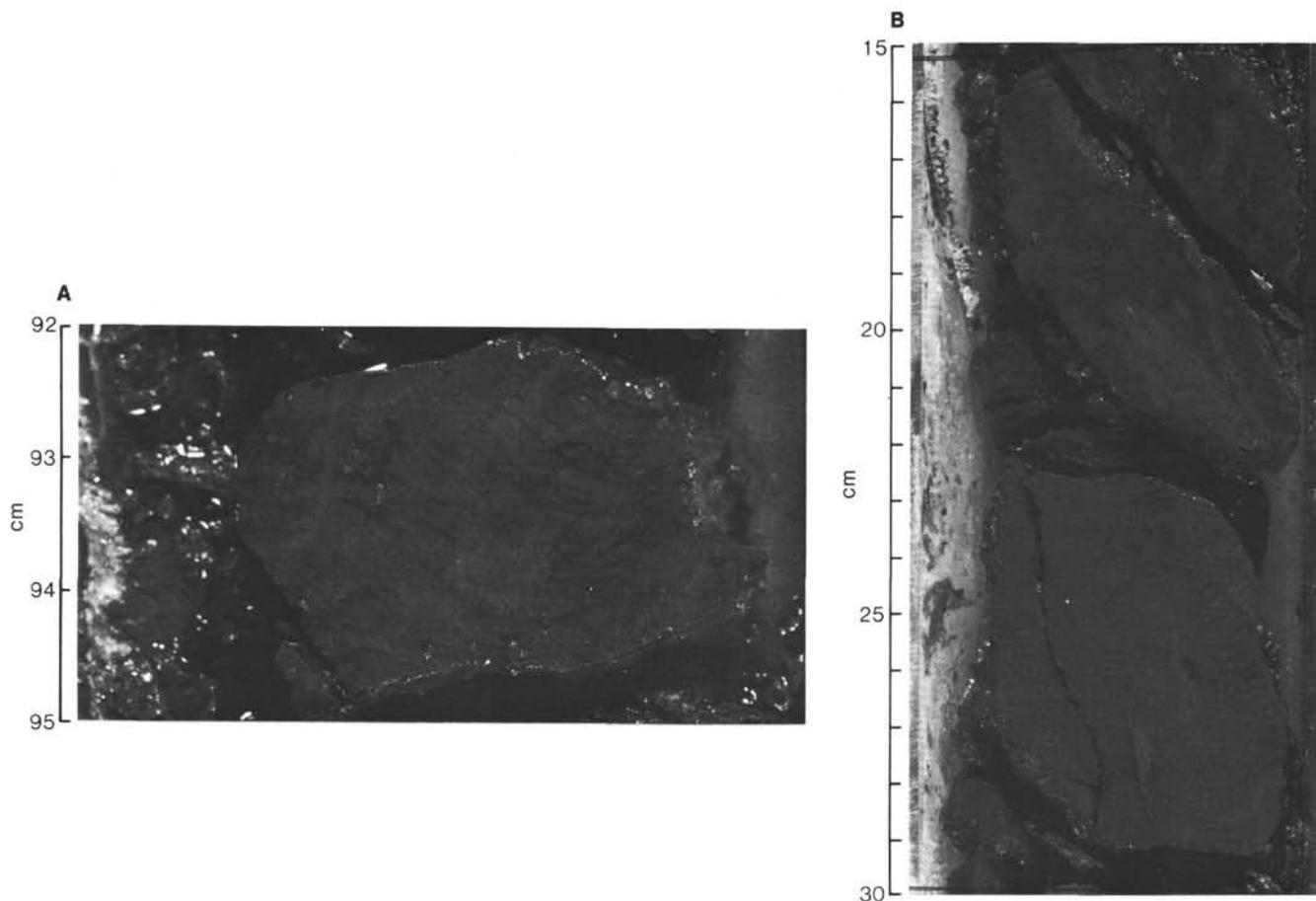


Figure 4. Photo of A. anastomosing (Sample 568-32-6, 92–95 cm) and B. subparallel veining (Sample 568-42-7, 15–30 cm) in sediments.

of Eocene material. The clasts contained pyritized radiolarians similar to those characteristic of the Eocene biostratigraphic facies recovered in the Esso Petrel Well.

BIOSTRATIGRAPHY

Introduction

At Site 568, 417.7 m of Pleistocene through early Miocene–late Oligocene mudstones were recovered (Fig. 5). Calcareous nannofossils, diatoms, and benthic foraminifers are in general age agreement, moderately preserved, and present in rare to common numbers to within 20 m of total depth.

Sections 568-1-2 to 568 21,CC are Pleistocene (1–195.4 m), as indicated by all microfossil disciplines; 568-22,CC is tentatively Pliocene, based on diatoms, whereas 568-23,CC, through 568-24,CC (214.8–224.5 m) are late Miocene, based on benthic foraminifers and diatoms. The barren and/or poorly preserved Samples 568-22-2, 32–36 cm through 568-25,CC delineate one long or several short Pliocene–late Miocene unconformities. Samples 568-25-4, 23 cm through 568-28-2, 11 cm are assigned to the middle Miocene *Discoaster exillis* Zone, 568-28-6, 10 cm through 568-38-2, 147 cm are assigned to the middle Miocene *Sphenolithus heteromorphus* Zone, and 568-38-4, 138 cm through 568-42-3, 126 cm are assigned to the early Miocene *Helicosphaera am-*

pliaperta Zone. Cores 42 through 44 (398.7–417.7 m) contain rare early Miocene–late Oligocene diatoms (to 568-44,CC), nannofossils (to 568-43,CC) and benthic foraminifers (to 568-44,CC).

Reworking of the microfossils is not significant until Cores 42 through 44, where Eocene diatom species first appear. Very rare Eocene nannofossils and a very sparse benthic foraminiferal assemblage associated with pyritized radiolarians is also present through this interval.

Paleoecologic analysis of Site 568 based on the benthic foraminifers suggests that deposition occurred in the upper middle bathyal (500–1500 m) to abyssal (≥ 4000 m) biofacies. The oldest sediments recovered in Cores 42 to 44 were deposited in the abyssal biofacies and below the foraminiferal CCD (calcite compensation depth). Foraminifers are extremely rare in these samples, whereas radiolarians are common and are preserved as both siliceous tests and pyritized molds. Benthic foraminiferal assemblages in Cores 21 through 41 suggest a gradual increase in water depths from upper middle bathyal (500–1500 m) to lower bathyal (2000–4000 m). Transported outer shelf material dominates the lower part of this interval. In upper part of the section transported material is principally from the upper slope.

Pleistocene deposition occurs in the lower middle bathyal and abyssal biofacies. In the lower part of this interval, Cores 19 through 7 were deposited in the lower

Core	Age	Biostratigraphy			Paleo-bathymetry					
		Nannofossil zones	Diatom zones	Benthic foraminifers	Neritic	Bathyal			Abyssal	
						Upper	Middle	Lower		
1	Pleistocene/ Recent	<i>Emiliana huxleyi</i> — <i>Gephyrocapsa oceanica</i>								
2										
3										
4										
5										
6										
7	Pleistocene	<i>Pseudoemiliana lacunosa</i>	<i>Pseudoeunotia doliolus</i>							
8										
9										
10										
11										
12										
13										
14										
15										
16										
17										
18										
19										
20										
21										
22	late Miocene	Indeterminate	Indeterminate							
23										
24	middle Miocene	<i>Discoaster exilis</i>								
25										
26										
27										
28										
29										
30										
31										
32										
33										
34										
35										
36										
37										
38	early Miocene	<i>Helicosphaera ampliaperta</i>								
39										
40										
41										
42										
43										
44	<i>Sphenolithus belemnos</i> — <i>Discoaster dellandrei</i>									

Figure 5. Biostratigraphic and paleoecologic summary, Site 568. Hachures indicate barren intervals.

middle bathyal biofacies (1500–2000 m), and paleobathymetry appears to decrease up section. Between Cores 7 and 6, paleodepths increase rapidly. Abyssal faunas dominate Sections 568-6, CC through 568-2-7. The core catcher of Core 1 contains a fauna indicative of depths intermediate between the present lower bathyal depth and the abyssal depths of Core 2.

Sediment accumulation rates uncorrected for compaction (Fig. 6) for the Pleistocene section (0–196 m) are approximately 117 m/m.y. This figure seems low in comparison to Site 496 (205 m/m.y.), unless the entire Pleistocene interval (225 m) from that site is averaged over a 1.65-m.y. time interval. A rate of 136 m/m.y. is then possible for Site 496, which compares favorably to Site 568 sedimentation rates.

Sedimentation rates for the early through middle Miocene section range from 66 m/m.y. (based on a constant sedimentation rate—straight line fit) to 33 m/m.y. (averaged over the entire time period encompassed by successive zones). Early Miocene sedimentation rates approximated at Site 496 were 60 m/m.y. and are in agreement with the higher figure. Paleontologic resolution is not sufficient to allow sedimentation rate calculations for the early Miocene to late Miocene and early Miocene–late Oligocene undifferentiated sediments.

Nannofossils

Abundant nannofossil assemblages are well preserved throughout the Pleistocene (568-1 through 568-21, CC) but lack consistent marker species for detailed zonation.

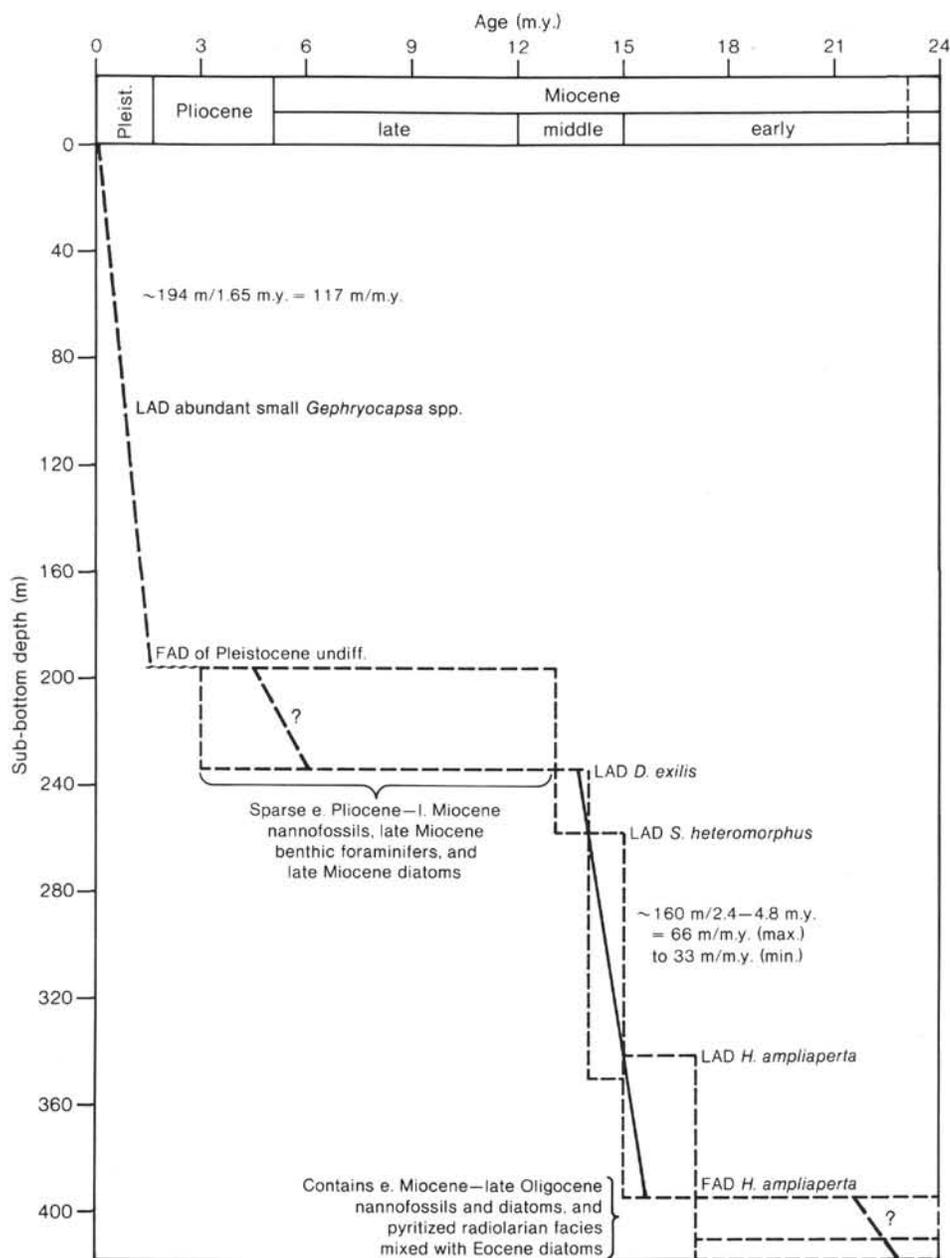


Figure 6. Sediment accumulation rates uncorrected for compaction, Site 568.

Gartner's (1977) Pleistocene zonation was tentatively applied from 568-1-2 through 568-21, CC, whereas Okada and Bukry's (1980) zonation is applicable for the early through middle Miocene interval.

Early Pliocene–middle Miocene nannofossils are rare to absent but are moderately preserved when present in Cores 23 through 25. Middle and early Miocene nannofossils encountered in 568-27 through 568-41, CC contain rare to common moderately preserved nannofossils. Early Miocene–late Oligocene nannofossils in Cores 42 and 43 are rare, and poorly to moderately preserved, whereas 568-44, CC (417.7 m, TD) was essentially barren.

Nannofossil reworking is minimal, with several early Pleistocene to middle Miocene species present in Cores 1 through 5. Very rare fragments of an Eocene *Micrantholithus* sp. are also present in Sample 568-43, CC.

Samples 568-1-2, 80 cm through 568-6-2, 43 cm are assigned to the Recent–late Pleistocene *Emiliania huxleyi* to *Gephyrocapsa oceanica* Zones on the basis of the occurrence of *Geophyrocapsa oceanica*, *Helicosphaera kamptneri*, small *Gephyrocapsa* spp., and *E. cf. huxleyi*. Samples 568-6-4, 43 cm through 568-11-1, 89 cm are assigned to the Pleistocene *Pseudoemiliania lacunosa* Zone, based on the occurrence of this same assemblage, with the addition of *P. lacunosa* and the absence of *E. cf. huxleyi*.

Core 11 through Sample 568-13-2, 61 cm are tentatively assigned to the small *Gephyrocapsa* Zone because of the dominant abundance of small *Gephyrocapsa* spp. and lack of *G. oceanica*.

Samples 568-13-4, 45 cm through 568-21, CC are assigned to the early Pleistocene. There is little change from the assemblage encountered upsection; and the sporadic, rare occurrence of *P. lacunosa* and *Helicosphaera sellii* makes zonal assignment difficult. Lack of *Calcidiscus macintyreii* suggests that the earliest Pleistocene may be absent.

Samples 568-22-2, 146 cm through 568-23-2, 90 cm are dominated by siliceous microfossils and barren of calcareous nannofossils. And 568-23, CC through 568-25-2, 148 cm are early Pliocene–middle Miocene, given the occurrence of *H. kamptneri*, *C. leptopora*, *Discoaster brouweri*, *Sphenolithus abies*, and *Reticulofenestra pseudoumbilica*.

Samples 568-25-4, 23 cm through 568-28-2, 11 cm are middle Miocene and assigned to the *Discoaster exilis* Zone. Rare *Discoaster exilis*, *D. challengerii*, and *Discoaster cf. bollii* are present in Sample 568-25-4, 23 cm whereas Sample 568-27, CC contains *Coccolithus miopelagicus* and *Discoaster adamanteus* with *D. exilis*.

Samples 568-28-6, 10 cm through 568-38-2, 147 cm are assigned to the middle Miocene *Sphenolithus heteromorphus* Zone. They contain a well-developed assemblage with *S. heteromorphus*, *D. exilis* s. 1., *Discoaster signus*, *C. miopelagicus*, *H. kamptneri*, and *Discoaster variabilis*.

Samples 568-38-4, 138 cm through 568-42-3, 126 cm are assigned to the early Miocene *Helicosphaera ampli-
aperta* Zone. They contain *H. ampli-
aperta*, *S. hetero-*

morphus, *H. intermedia*, *H. euphratis*, and *Discoaster deflandrei*.

Samples 568-42-7, 40 cm through 568-44-4, 70 cm (firm mudstone lithologies) are considered early Miocene and contain consistent occurrences of *Helicosphaera euphrates*, *D. deflandrei*, *Triquetrorhabdulus carinatus*, and *Sphenolithus moriformis*. Specimens of *Cyclicargolithus abisectus* or *Dictyococcites bisectus* are not present, which indicates that Oligocene sediments were not recovered at this depth. Section 568-44, CC is essentially barren of nannofossils.

Diatoms

Hole 568 contains the most diverse assemblage of diatoms to date on Leg 84—about 175 species were observed from the 44 cores.

The Pleistocene is represented by the reliable marker fossil *Pseudoeunotia doliolus*, which is present continuously down through Core 20. The *Nitzschia reinholdii* Zone spans Cores 14 through 20. The last appearance of *Rhizosolenia praebergonii* is in the Pliocene, which is penetrated by Core 22 (21 is a short pressure core), and *Thalassiosira oestrupii* became extinct in the Pliocene, as well.

Hole 23 contains *Rossiella praepaleacea*, which is restricted to the upper Miocene.

Cores 24 to 44 contains diatoms whose ranges are restricted to the middle and lower Miocene (Barron, 1981, and Barron, in press). These include *Actinocyclus ellipticus* var. *javanicus*, *Annellus californicus*, *Cestodiscus peplum*, *Actinocyclus radianovae*, *Coscinodiscus lewisianus*, *Craspedodiscus coscinodiscus*, *Cymatogonia amblyoceros*, *Denticulopsis nicobarica*, *Goniothecium odontella*, *Stephanopyxis grunowii*, and *Synedra jouseana*. According to Barron (in press), *Craspedodiscus coscinodiscus*, *Coscinodiscus lewisianus*, and *Annellus californicus* have their upper range limits in the lower middle Miocene. These species first occur in Core 24. *Synedra jouseana*, which became extinct in the middle Miocene, ranges from Cores 26 through 44.

Benthic Foraminifers

Benthic foraminiferal assemblages from Site 568 are abundant and well preserved in the majority of samples examined. Poorly preserved assemblages occur particularly in the intervals associated with a hiatus of where deposition occurred below the foraminiferal CCD. Biostratigraphic interpretation indicates that Core 1 through 568-21, CC are Pleistocene, and Samples 568-22-2, 34–36 cm through 568-23-3, 92–96 cm are probably late Miocene. Samples 568-23, CC through 568-37, CC are early Miocene, and Cores 42 through 44 are earliest Miocene. Paleoecologic analysis indicates a complex bathymetric history with depths ranging from upper middle bathyal to abyssal.

The Pleistocene interpretation (Cores 1 through 21) is based on diagnostic Pleistocene species, the similarity of faunas to Holocene assemblages (Smith, 1964), and the absence of any extinct species. *Uvigerina senticosa*, *Laticarinina pauperata*, *Planulina weullerstorffi*, and

various bolivinids dominate the upper cores in this interval. This faunal composition indicates deposition occurred in the abyssal biofacies (4000 m) with transport predominately from the upper and lower bathyal biofacies. The benthic foraminiferal fauna similarly changes in Cores 7 through 21 and is characterized by a larger component of lower middle bathyal species and rare abyssal species. Samples 568-22-5, 32-36 cm through 568-25-6, 51-55 cm are poorly preserved and dominated by resistant species or barren. The common occurrence of *Siphogenerina basispinata* in Sample 568-23, CC and subsequent samples suggests that the overlying barren or dissolved interval represents one or more of the Pliocene to middle Miocene hiatuses. Well preserved middle Miocene assemblages are mixed with dissolved or barren samples down to Sample 568-37, CC. In 568-38, CC *Siphogenerina transversa* and nannofossils indicate the early Miocene. Additional diagnostic early Miocene species occur down to 568-41, CC.

Paleoecologic analysis of the Miocene interval at Site 568 suggests that from Cores 21 to 41 there was a transgression with water depths progressing from upper middle bathyal (500-1500 m) in early Miocene to lower bathyal (2000-4000 m) by late Miocene. Transported material is derived from the outer shelf biofacies in the early Miocene. Lenticulinids and various species of *Marginulina* are common components of the transported material. As the water depths increased in the later part of the Miocene, shelf materials were no longer deposited in this area. In the late and middle Miocene, benthic foraminifers were transported principally from the upper slope.

Benthic foraminiferal species in Cores 43 and 44 are rare, poorly preserved, and probably early Miocene. These assemblages could also be diagnostic of the latest Oligocene. Because radiolarians dominate these samples, deposition is believed to have taken place in the abyssal biofacies (4000 m) and below the foraminiferal CCD.

PHYSICAL PROPERTIES

Methods

The suite of measurements performed to determine physical properties consists of bulk density, porosity, water content, sonic velocity, shear and compressive strength, and thermal conductivity. The techniques employed are the same as for previous sites of this leg. Thermal conductivity tests were made on only a limited number of samples.

Results

Bulk densities and their inversely related wet-water contents and porosities are shown in Figure 7. Bulk density increases slowly from 1.31 Mg/m³ at sub-surface to approximately 1.4 Mg/m³ at 230 m sub-bottom. At this depth all three index properties show a more rapid change to a sub-bottom depth around 240 m; bulk densities increment in this interval to about 1.55 Mg/m³. Below this point, the trends once again demonstrate an overall gradual and slow shift to the depth interval between 360 and 375 m, which once again shows a faster rate of

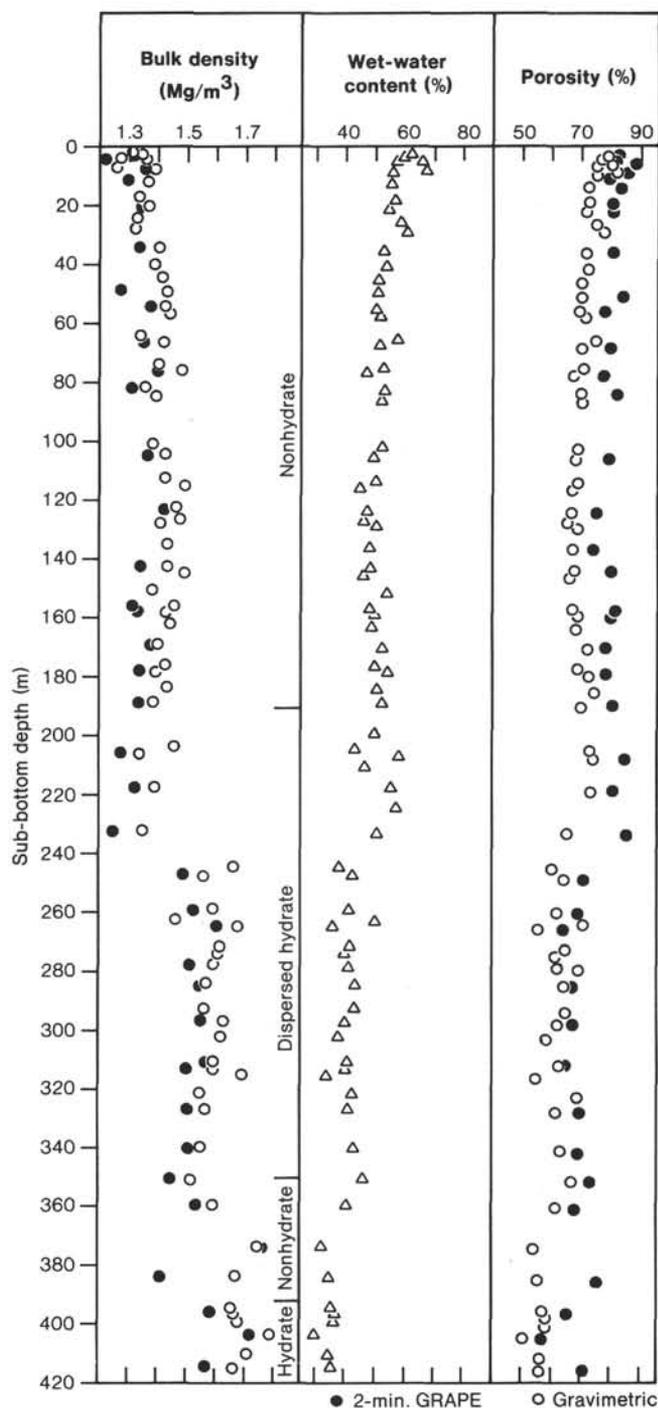


Figure 7. Index properties of Site 568 sediments.

change in index properties, that is, bulk density increases from 1.55 Mg/m³ to roughly 1.70 Mg/m³. The remaining data resume the gradual trend of the overall section. Variability occurring around the tendencies may be related to the disturbance caused from degassing of the sediment, drilling, and/or the fractured nature of sediment below 166 m.

Sonic velocities and acoustic impedance are plotted in Figure 8. Sonic velocity data show an overall increase with depth, velocities varying from 1.56 km/s near the

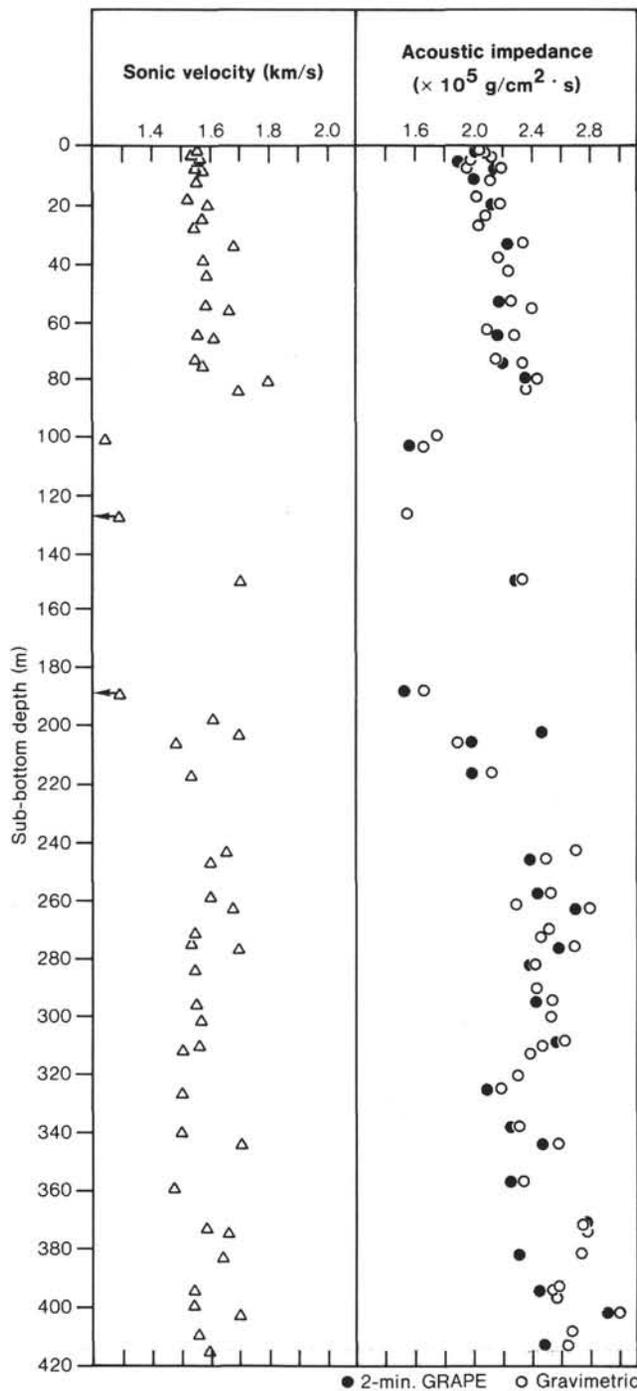


Figure 8. Acoustic characteristics of sediments at Site 568. (Data point with arrow indicates information off the scale.)

mud line to 1.65 to 1.70 km/s at 420 m. A gap between 85 and 198 m with a few widely distributed data points reflects the area in which degassing within the sediment caused excessive attenuation.

Shear and compressive strengths are plotted in Figure 9. These data are incomplete, but those presented reflect the low strengths induced by the aforementioned degassing. A rapid increase in strength is observed between 200 and 230 m sub-bottom, below which the sediment becomes indurated, precluding further measurements.

Thermal conductivities were performed using both the full and half-space probes, depending on the degree of consolidation of sediment. The data obtained are not corrected to *in situ* conditions and are shown in Table 2.

Discussion

The character of physical properties at this site closely resembles that of Site 496 physical properties determined during Leg 67. The Pleistocene/Pliocene boundary is correlated with the upper interval of rapid change in physical properties (230–240 m), as observed here and at Site 496. The lower interval (360–375 m), showing a similar trend, is not directly related to such a feature, however, it does lie near the early/middle Miocene boundary.

The presence of external stress in the system at this site may be appreciated in excess pore pressures monitored by the *in situ* pore-water samples, sigmoidal veins, and fractures. The effect of this stress in the sediment here is linked to the dewatering capabilities of the system and the foliated nature of the sediment and mudstone. *In situ* pore pressure responses to the downhole interstitial water sampler indicated pressures exceeding hydrostatic pressure. These high pressures may be indicative of slow drainage conditions that do not allow dissipation of excess pore fluid (see Taylor and Bryant, this volume).

Physical properties traits linked to hydrates existent at this site are difficult to assess. This difficulty is expected because the hydrate is most likely present in dispersed form in the majority of the section. Chemical analysis of gas molecular composition suggests that dispersed hydrate layers exist between 190 and 345 m and are dispersed with some solid hydrate between 391 and 410 m sub-bottom. Lithostratigraphic observations show interspersed layering of highly fractured units within most cohesive blocks of mudstone, and this may represent the distinct style limiting hydrate distribution. The physical properties of sediment recovered along with the hydrate in Core 43, however, are similar to those of surrounding facies.

GEOPHYSICS

Seismic Reflection Records

Site 568 is located about midway between seismic records GUA-13 and GUA-18; GUA-18 was made using a maxipulse sound source to increase penetration over that obtained in GUA-13. At the position of Site 568 the two lines are plotted about 1 km apart. The multiple satellite fixes recorded during drilling place Site 568 941 m north and 524 m west of Site 496.

Both records GUA-13 and GUA-18 display essentially the same geologic features, and only 13 is shown here (Fig. 10). The GUA-13 display shows a sequence of reflections immediately below the seafloor, representing a hemipelagic unit that laps onto a dipping section of higher amplitude, which represents a sediment lobe. Sonic logging gives a similar result. Using the corrected PDR (precision depth recorder) depth at the site gives a

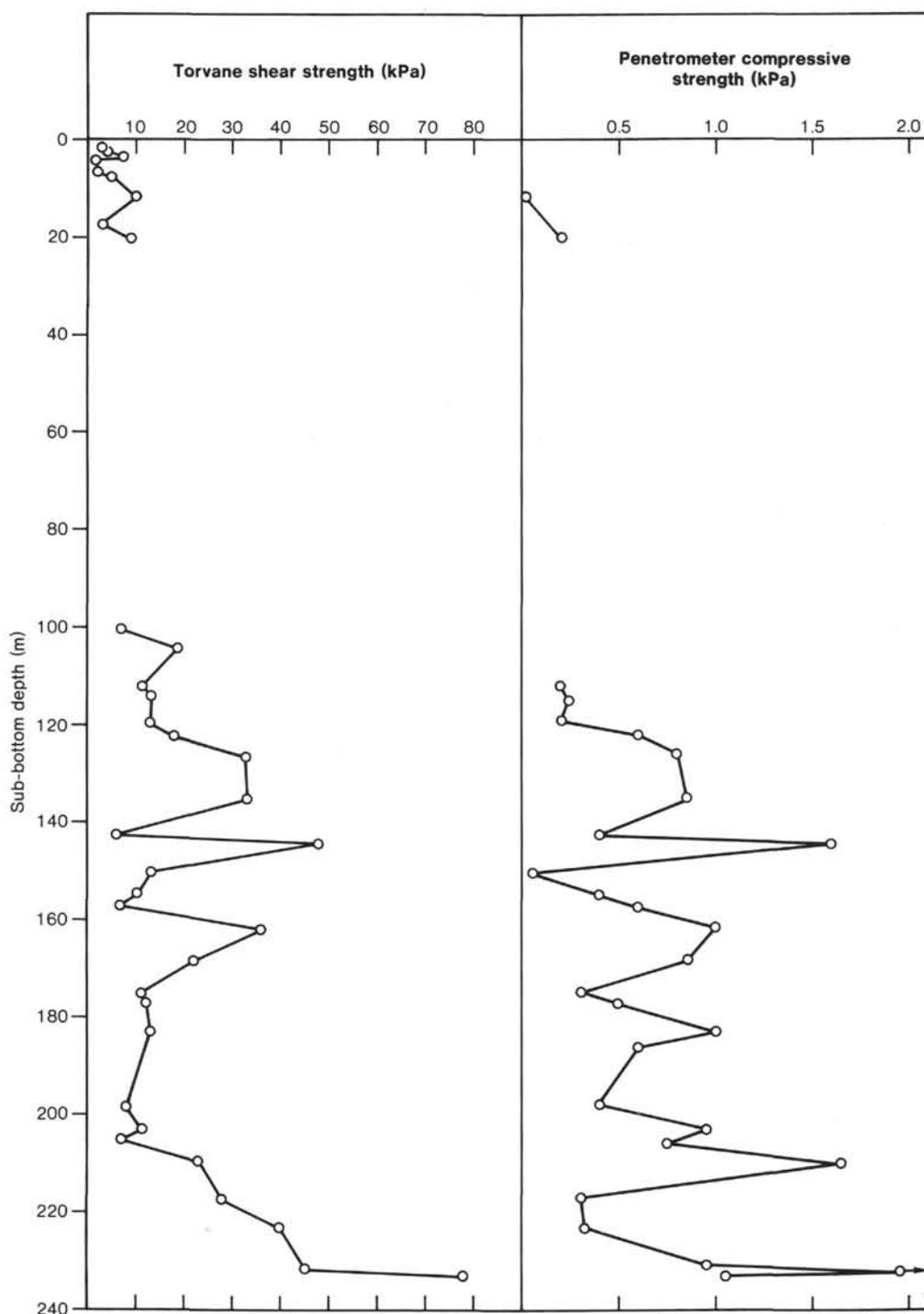


Figure 9. Downhole strength measurements from Site 568 sediments.

BSR (bottom simulating reflection) depth of 469; using the drill stem depth of bottom felt gives a BSR depth of 445. The former is more compatible with the seismic record because it is an acoustic measurement. Therefore drilling at Site 568 ended about 44 m above the base of hydrate reflection. The dipping section of the lobate unit is cut by a strong reflection subparallel to the seafloor. This latter reflection is a BSR that marks the base of the gas hydrate.

Calculations concerning the depth of the BSR were first made using a value from semblance diagrams of 1.8 to 1.85 km/s. This velocity gives a depth to the BSR of about 500 m. During the drilling at Site 568 it became apparent that the semblance velocities for the upper part of the section were high, and by 280 m a good statistical population could be established that showed velocities of 1.5 to 1.6 km/s. Using these values and projecting to a value of 1.8 km/s at the BSR gave a BSR depth of 460 m,

Table 2. Thermal conductivities for Hole 568 sediments.

Section (core-section)	Sub-bottom depth (m)	Thermal conductivity (mcal/cm°C · s)
13-5	114.40	1.626
16-6	144.80	1.557
18-5	162.88	1.874
33-3	304.67	2.955
42-6	397.67	4.691

which agreed, as is discussed later, with the BSR depth determined from a measured temperature gradient.

In Situ Temperature and Pressure

The *in situ* water temperature probe, fitted for the first time with a pressure measuring system, was run six

times in Hole 568 (Table 3; Fig. 11). Temperature measurements from the first four lowerings of the probe define a changing gradient with depth that is generally high for a convergent margin and higher than values derived from the drilling at Site 496. The fifth temperature measurement was probably made with incomplete penetration, because the sampler contained seawater. The sixth reading is similarly suspect because of seawater in the sampler and an implied steep temperature gradient. The last two runs were made in stiff sediment as seen in the cores, which may explain the suspect readings.

The temperature gradient defined by the first four runs gives a base of hydrate depth corresponding to the BSR depth of 460 m (see Organic Geochemistry section).

Interpretation of the pressure measurements is reported by von Huene in a chapter on the evidence of elevated pore pressure (this volume).

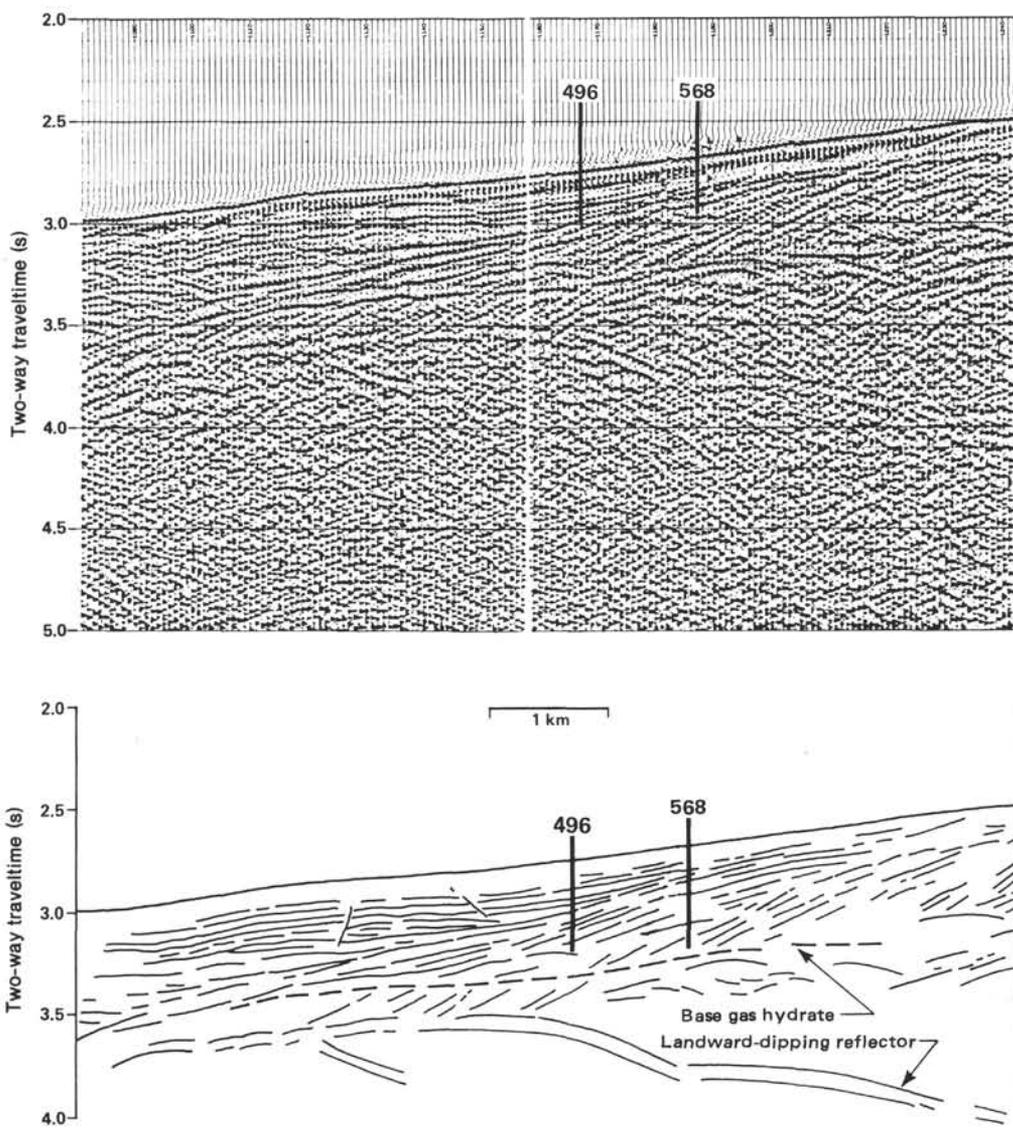


Figure 10. Part of seismic record GUA-13 across the midslope; an interpretive line drawing shows the locations of Sites 496 and 568 and illustrates the structure of the slope deposits and a base of gas hydrate BSR.

Table 3. *In situ* temperature–pressure measurements.

Station	Depth sub-bottom (m)	Bit height (m)	Min. temp. (°C)	Probe temp. (°C)	Mud-line temp. (°C)	Bit pressure (psi)	Probe pressure (psi)	Mud-line pressure (psi)	Comments
1	89.4	10.0	3.6	6.9	2.9	3172	3278	2858	
2	137.2	3.0	2.6	10.1	2.9	3263	3400	(avg. 2843) 2736	
3	186.2	10.5	10.5	13.2 (13.8?)	None	3290	3560	(avg. 2771) 2869	May have pulled out once.
4	282.3	6.0	11.4	16.6	None	3480	3450?	(avg. 2849) 2806	Pressure is suspect.
5	359.9	10.5	2.9	8.7	2.8	3563	3680 (after pore water sampling)	(avg. 2771) 2928 (avg. 2879)	Water sample and temperature are suspect.
6	417.7	6.0	5.2	18.0	None	—	—	—	Water sample and temperature are suspect.

Note: — indicates no data.

Downhole Logging

Two logging runs provided caliper, natural gamma, sonic velocity, formation density, and neutron porosity logs from 85.5 to 410 m (Fig. 12). A cursory study of the caliper, sonic, and density traces was made from field prints. The sonic log reflects the base of the Pliocene unconformity and fixes its time intercept at the contact between the hemipelagic and lobate sediment sequences in the seismic reflection record. Above the contact the sonic trace is without character and has a constant velocity of 1.5 km/s but at the contact a series of thin higher-velocity beds are seen that also deflect the density and caliper logs; below the contact the trace has low amplitude oscillations around a fairly constant 1.6 km/s value. At 365 m, the trace becomes noisy, shows an average velocity of 1.8 km/s, and corresponds to a very rugose part of the hole; no apparent change in lithology is noted for this interval. The section below 365 m has abundant fractures and begins to show local steep dips that may influence the rugosity of the hole.

The formation density trace indicates a density increase at the Pliocene/Pleistocene boundary and again at the Miocene/Pliocene boundary. It reflects well the massive mudstone section between 180 and 305 m; below 305 m the trace appears noisy, corresponding to the rugosity of the hole.

There are no obvious effects from gas hydrate in the logs when the boundaries of dispersed and recovered hydrate are compared with the sonic log and the events that can be attributed to a lithologic change are recognized. There is no increase in velocity at hydrate boundaries, as expected from geophysical studies elsewhere (Kvenvolden and Barnard, 1983). It seems that the hydrate is dispersed in quantities that are insufficient to be detected by the logging instruments but are sufficient to produce a BSR.

PALEOMAGNETISM

At least one oriented sample was taken from each core section that was not clearly disturbed by drilling. Results of stepwise alternating field demagnetization on selected samples are plotted in Figure 13. Several of the

samples show significant changes in intensity and direction in low fields, less than 50 Oe, which indicates the presence of low-stability secondary components. It was found that many samples from this hole changed their directions by 50° or more when left for 12 hr. in the ambient magnetic field in the van, which was about 0.4 Oe. Paleomagnetic measurements were made on samples from the first 22 cores. The results appear in Figure 14. Most of the samples were later demagnetized in fields of 150 Oe. The results, even after demagnetization, did not show any systematic stratigraphic pattern and bore a close resemblance to random distribution. It was suspected that drilling disturbance was responsible for at least part of this behavior, although low magnetic stability and the conditions under which the sediments were deposited also play a role.

GEOCHEMISTRY

A major objective at Site 568 was to study the geochemistry related to the occurrence of gas hydrates in oceanic sediments. This site was located near Site 496 of Leg 67 (von Huene, Aubouin et al., 1980) in order to argument the geological and geochemical information obtained previously. Although no gas hydrates were observed directly during drilling at Site 496, salinity and chlorinity measured on pore waters from sediments there were anomalously low, suggesting the presence of gas hydrates (Hesse and Harrison, 1981). In addition, refined interpretations of geophysical records show that a bottom simulating reflector (BSR), indicative of the base of the zone of gas hydrates, is present at the site (von Huene et al., 1982); the BSR had not been recognized at the same time Site 496 was drilled. Thus both chemical and geophysical information indicated that gas hydrates should be present in this part of the landward slope of the Middle America Trench, and Site 568 was drilled to provide detailed data from which to interpret the occurrence and origin of gas hydrates in oceanic sediments offshore Guatemala.

Gas Analyses

Gases were obtained from gas pockets that developed as sediment separated in the core liner because of gas ex-

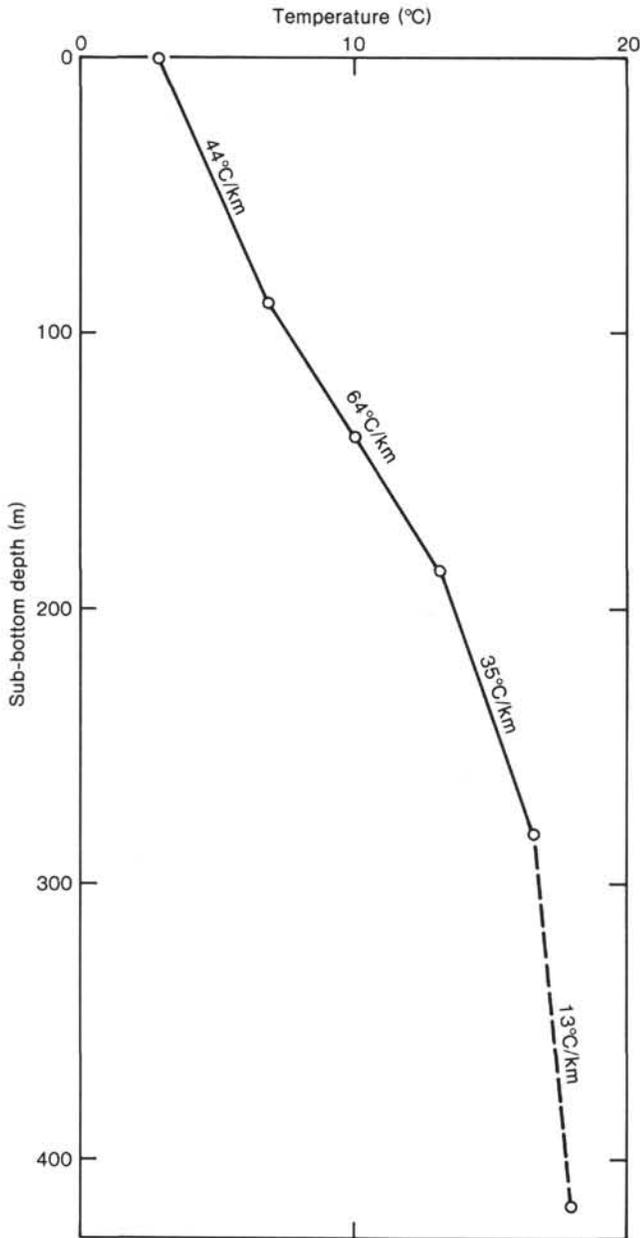


Figure 11. Summary of temperature measurements made with the *in situ* probe. Geothermal gradient between points of measurement is also shown.

pansion. Gas pockets were sampled by means of vacuumainers, and the collected gases were analyzed by gas chromatography, providing a measure of the volumetric composition of gases within the core liner.

Hydrocarbon gases methane (C_1), ethane (C_2), propane (C_3), isobutane ($i-C_4$), normal butane ($n-C_4$), neopentane ($neo-C_5$), isopentane ($i-C_5$), and normal pentane ($n-C_5$), are present in most samples from Site 568, with C_1 being the most abundant gas found (Table 4). Ratios of C_1/C_2 generally decrease with depth (Fig. 15). The sudden decrease in C_1/C_2 ratios in the interval between 220 and 280 m is because of greatly diminished concentrations of C_1 ; C_2 concentrations remain about the same, as in samples above and below this interval.

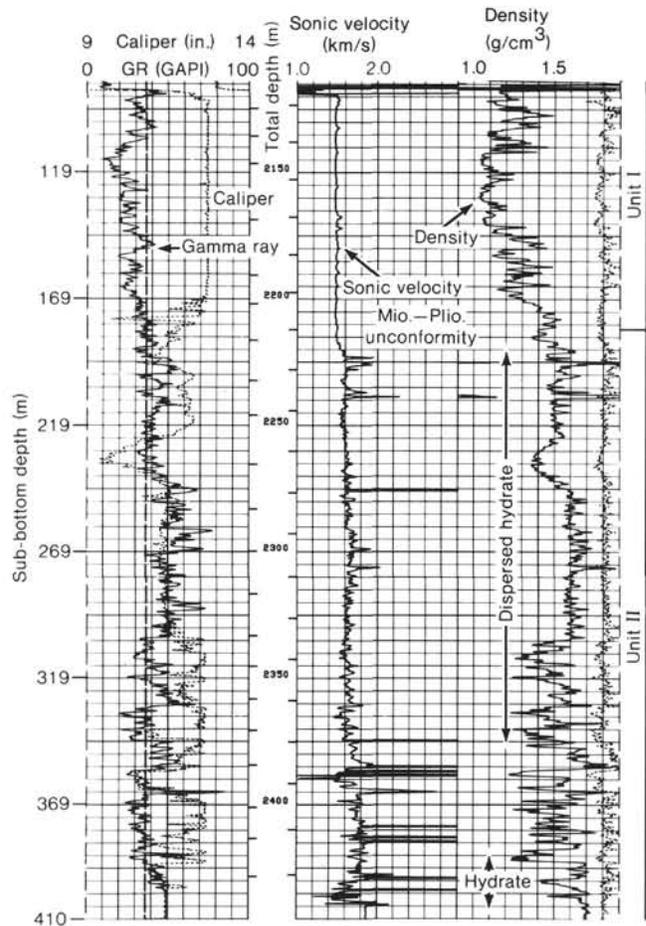


Figure 12. Annotated logs from Site 568.

Over all, the C_1/C_2 ratios decrease with depth from 20 to 417 m by almost three orders of magnitude.

Gas Hydrates

The relative distributions of hydrocarbon gases in the sediments at Site 568 suggests that gas hydrates are present in the intervals from 190 to 345 and from 391 to 410 m sub-bottom. In these intervals C_1 and C_2 dominate. The larger hydrocarbons are present but in very low concentrations. In contrast, the concentrations of these larger hydrocarbons are greater elsewhere in the core (Table 4). The interpretation of the presence of hydrate is based on consideration of the molecular sizes of cages within the gas hydrate structure (Hand et al., 1974). Results from Site 565 suggested that both structure I and II gas hydrates are present, because gases larger than $i-C_4$ apparently are excluded. If only structure I gas hydrate were present, gases larger than C_2 would be excluded from the structure. At Site 568 the presence of structure I gas hydrates is suggested in the intervals where C_1 and C_2 dominate and the larger hydrocarbon gases are at low concentrations. No solid gas hydrates were observed, however, in the interval from 190 to 345 m. Rather the fine-grained sediment generally was disrupted because of the presence of gas; the gas hydrate likely had been dispersed in the sediment and had de-

Table 4. Distribution of hydrocarbon gases in Hole 568.

Section (core-section)	Sub-bottom depth (m)	C ₁ (%)	C ₂ (ppm)	C ₃ (ppm)	i-C ₄ (ppm)	n-C ₄ (ppm)	neo-C ₅ (ppm)	i-C ₅ (ppm)	n-C ₅ (ppm)
3-5	18	65.0	2.2	1.50	0.25	0.08	—	0.23	0.44
4-4	28	69.0	2.6	2.00	0.46	0.14	—	0.33	0.09
5-5	39	72.0	3.0	3.10	1.10	0.13	—	0.42	0.09
6-4	47	80.0	6.5	4.00	1.30	0.22	0.01	0.71	0.05
7-3	56	75.0	4.3	3.60	1.20	0.20	—	1.40	0.14
8-5	68	71.0	6.3	4.70	1.60	0.38	0.20	0.88	0.07
9-4	75	82.0	6.7	5.60	2.20	0.42	0.10	1.90	0.12
10-4	85	82.0	14.0	6.70	3.00	0.60	0.10	3.20	0.22
12-4	104	75.0	10.0	4.30	1.80	0.37	—	2.70	0.11
13-4	114	69.0	11.0	3.80	1.60	0.34	—	3.30	0.04
14-6	126	95.0	41.0	3.80	2.00	0.35	0.01	2.40	0.14
15-6	135	68.0	41.0	1.70	1.10	0.13	—	0.64	0.09
16-4	145	90.0	54.0	3.50	2.00	0.42	—	0.50	0.13
17-4	153	76.0	49.0	1.70	1.10	0.19	—	1.30	0.17
18-3	160	74.0	56.0	3.30	2.30	0.39	0.09	4.20	0.36
19-7	176	69.0	65.0	1.70	1.30	0.24	0.09	1.80	0.29
20-6	184	50.0	36.0	0.87	0.73	0.12	0.10	0.51	0.08
22-5	203	78.0	130.0	0.79	0.95	0.06	0.10	0.21	0.23
23-2	208	70.0	94.0	0.62	0.65	0.06	0.10	0.04	0.19
24-4	220	72.0	110.0	0.58	0.85	0.03	0.20	0.26	0.04
25-6	232	65.0	95.0	0.61	0.50	0.04	0.10	0.06	0.13
27-3	247	6.2	130.0	0.49	0.79	0.05	0.10	0.13	0.11
28-4	259	3.8	110.0	0.61	0.59	0.04	0.20	0.25	0.06
29-5	270	68.0	110.0	0.56	0.62	—	0.30	0.04	0.08
30-6	281	67.0	150.0	0.52	0.77	0.04	0.24	0.27	0.07
32-6	300	61.0	100.0	0.55	0.19	0.04	0.13	0.04	0.09
33-6	310	71.0	120.0	0.58	0.48	0.05	0.20	0.05	0.20
34-6	319	67.0	140.0	0.71	0.63	0.05	0.20	0.16	0.20
35-4	327	76.0	170.0	0.75	0.75	0.05	0.35	0.25	0.13
36-6	339	68.0	150.0	0.67	0.82	0.05	0.35	0.23	0.09
37-4	346	59.0	116.0	0.67	0.94	0.06	0.45	0.46	0.04
38-6	358	78.0	230.0	0.77	1.20	0.05	0.78	0.48	0.03
39-3	363	70.0	440.0	0.67	0.78	0.05	0.15	0.08	0.04
40-3	374	75.0	560.0	0.97	1.60	0.09	0.53	0.08	0.03
41-5	386	72.0	540.0	1.30	1.50	0.17	0.45	0.32	0.04
42-6	398	76.0	800.0	0.73	1.50	—	0.18	0.14	0.10
43-4	404	65.0	670.0	0.52	0.99	0.03	0.12	—	0.13
44-4	414	78.0	1000.0	0.53	1.60	—	0.25	0.03	0.07

Note: — indicates not detected.

composed during the drilling process. In the interval between 391 and 410 m solid gas hydrate was observed, and the distribution of hydrocarbon gases recovered from the core also indicated the presence of gas hydrates.

Solid pieces of gas hydrates were recovered from fractures in a tuffaceous mudstone in the interval from 403.2 to 404.7 m (Fig. 16). A sample of this gas hydrate was placed in a pressure device. The sample decomposed, producing about 30 volumes of gas to 1 volume of water. This water had a salinity of about 5.5‰. A second sample, which had been decomposing in the laboratory for about 15 min., was also placed in the pressure vessel; the volumetric ratio of gas to water for this sample was about 7. These ratios clearly demonstrate that solid gas hydrate had been found. If these samples had been out-gassing water in the form of ice, the maximum possible ratios would have been about 3. The distributional pattern of hydrocarbon gases resulting from the decomposition of the gas hydrate within the pressure device is dominated by C₁ and C₂.

The sub-bottom interval from 410 to 417 m (total depth of the hole) does not contain gas hydrates as indicated by the distribution of gases. This distribution is typical of those intervals observed earlier where gas hy-

drates are presumed to be absent. In summary, the distribution of hydrocarbon gases at Site 568 are interpreted as follows:

Sub-bottom depth (m)	Gas occurrence
0-190	Biogenic gas followed by gas resulting from early diagenesis
190-345	Gas hydrate dispersed in fine-grained sediment
345-391	Gas in sediments in a nonhydrated state
391-410	Dispersed and solid gas hydrate
410-417	Gas in a nonhydrated state

Pressure Core Barrel

The pressure core barrel (PCB), described in detail by Kvenvolden et al. (1983), was deployed three times at Site 568. Two PCBs recovered pressurized cores from which gas samples were taken. Results are summarized in Table 5, and the details of the experiment are given in Kvenvolden and McDonald (this volume). Aliquots of gas samples from PCB-3 were analyzed by gas chroma-

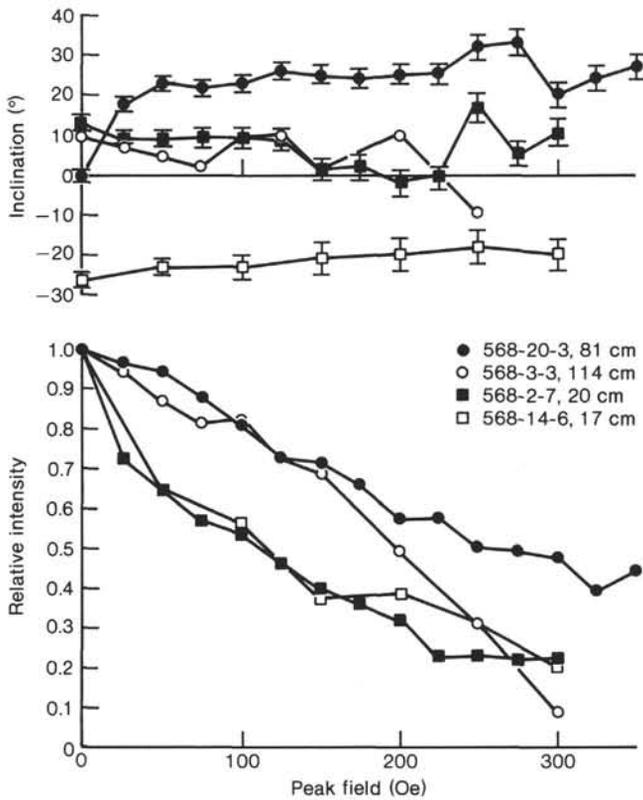


Figure 13. Alternating field demagnetization of selected samples from Hole 568.

tography, and the remaining gas was retained for shore-based isotopic and molecular analyses.

Interstitial Water Chemistry

The inorganic parameters measured were calcium, magnesium, chlorinity, salinity, alkalinity, and pH. The results of these measurements are shown in Figure 17. The results obtained from squeezed sediment samples essentially duplicate the results obtained at Site 496 of Leg 67 (Hesse and Harrison, 1981). This duplication is expected because of the close proximity of Site 496 and Site 568.

Summary

1. Gas hydrates occur in the sediments at Site 568. The molecular distribution of hydrocarbon gases, the low values of chlorinity and salinity, and the disrupted nature of the sediments indicate that gas hydrates are dispersed in sediment at sub-bottom depths between 190 and 345 m.

2. Solid gas hydrate was observed in a fractured tuffaceous mudstone at a depth of about 403 m. When this hydrate decomposed, between 7 and 30 volumes of mainly C₁ were released per volume of water, and this water had a salinity of about 5.5‰, indicating the freshwater nature of the water of the gas hydrate structure. The gas mixture released during decomposition of the gas hydrate contained mainly C₁ and C₂. The presence of dispersed gas hydrate between 190 and 345 m and of solid gas hydrates in fractures between 391 and 410 m suggests that gas hydrates are stratified at Site 568.

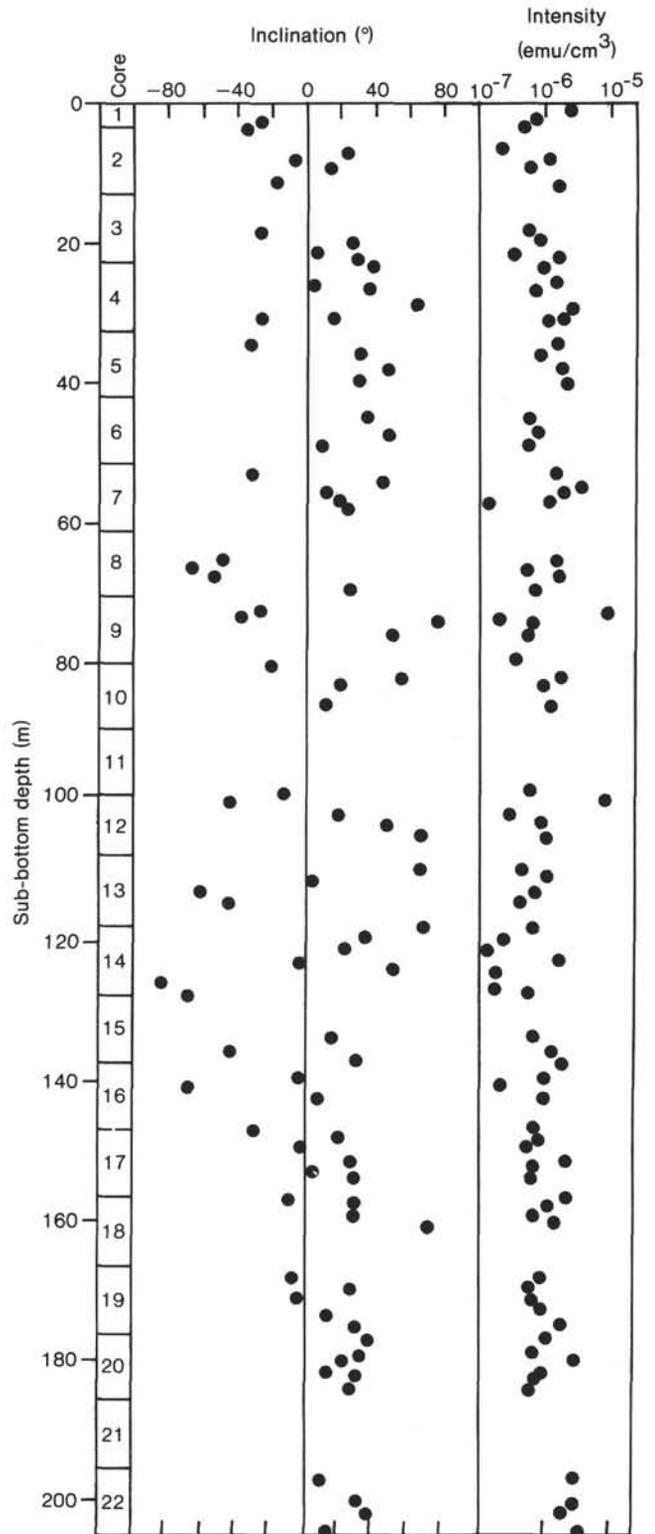


Figure 14. Stratigraphic plot of NRM (natural remanent magnetization) data for Hole 568.

SUMMARY AND CONCLUSIONS

Site 568 is on the upper part of the Middle America Trench slope, in 2031 m of water, 4000 m above and about 47 km landward from the Trench axis. It is lo-

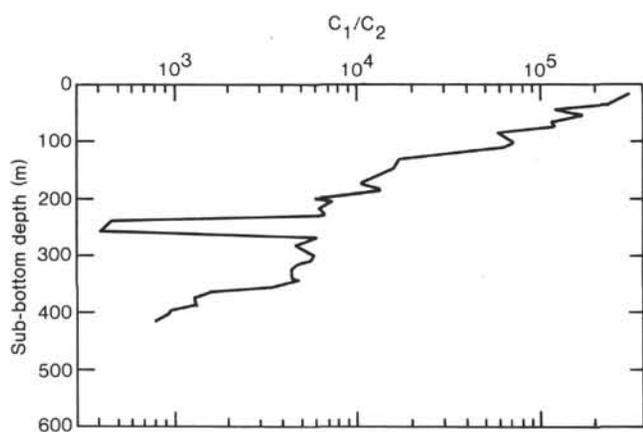


Figure 15. Ratios of methane/ethane (C_1/C_2) with depth at Site 568.

cated about 1 km upslope from Leg 67 Site 496, which had been abandoned after a gas composition was encountered, which suggested possible thermogenic components and perhaps gas hydrate. The major objective at Site 568 was a detailed monitoring of the gas through the whole section to study the formation of gas hydrate. Seismic record GUA-13 was reprocessed to bring out the shallow structure and the base of the gas hydrate seismic reflector. Within the Safety Panel recommendations, Site 568 could be drilled to 100 m above the reflector or to the base of gas hydrate as calculated from temperature data; during the drilling, permission was obtained to penetrate at least to the equivalent depth of Site 496.

Gas hydrates were recognized at Site 568 both from gas composition and as recovered pieces of gas hydrate. The molecular distribution of hydrocarbon gases and the low values of chlorinity and salinity indicate that gas hydrates are dispersed in sediment and coexist as a hydrated and nonhydrated sequence in the 418 m drilled. The sequence encountered is as follows: 0–190 m, biogenic gas followed by gas resulting from early diagenesis; 190–345 m, gas hydrate dispersed in fine-grained sediment; 345–391 m, gas in sediments in a nonhydrated state; 391–410 m, dispersed and solid gas hydrate; 410–417 m, gas in a nonhydrated state. The sediment sequence containing the gas in hydrated and nonhydrated form is a mudstone of very low permeability and porosity: fracturing probably provides the principal porosity and permeability. In the main hydrate section (190–345 m), no hydrate was observed visually despite a careful inspection of each core as soon as it came on deck nor was there an unusual amount of gas in the cores. The hydrate was recognized by a gas composition containing mainly C_1 and C_2 . Larger hydrocarbon molecules were present only in low abundance, because they are excluded from the gas-hydrate structure. The hydrate observed visually occurred at about 403 m. It was a white icelike substance filling fractures in a tuffaceous mudstone (50% glass shards). When the hydrate decomposed, between 7 and 30 volumes of mainly C_1 were released per volume of water. The water was fresh, indicating the fresh-water nature of the gas hydrate structure. The quantity of gas relative to water released on decomposition, the composition of the released water and the pore



Figure 16. Gas hydrate (white icelike material) embedded in sediment from a sub-bottom depth of about 403 m at Site 568 (Sample 568-43-4, 40–60 cm).

water, and the gas composition demonstrate that gas hydrate had been recovered.

Downhole logging shows that the gas hydrate occurs in concentrations so low that neither velocity nor densi-

Table 5. Results from pressure core barrel, Site 568.

PCB no.	Core	Sub-bottom depth (m)	Core recovered		Initial pressure (psig)	Pressure before venting (psig)	Sediment water (%)	Gas released (l)	Gas/water vol. ratio
			(m)	(l)					
1	11	89.4-98.8	1.05	14	3000	3000	50	n.d.	n.d.
2	21	185.7-195.4	1.61	22	2900	3100	52	7.4	0.7
3	31	282.3-292.0	1.16 ^a	15	2000	2300	n.d.	n.d.	n.d.

Note: n.d. = not determined.

^a Unpressurized core.

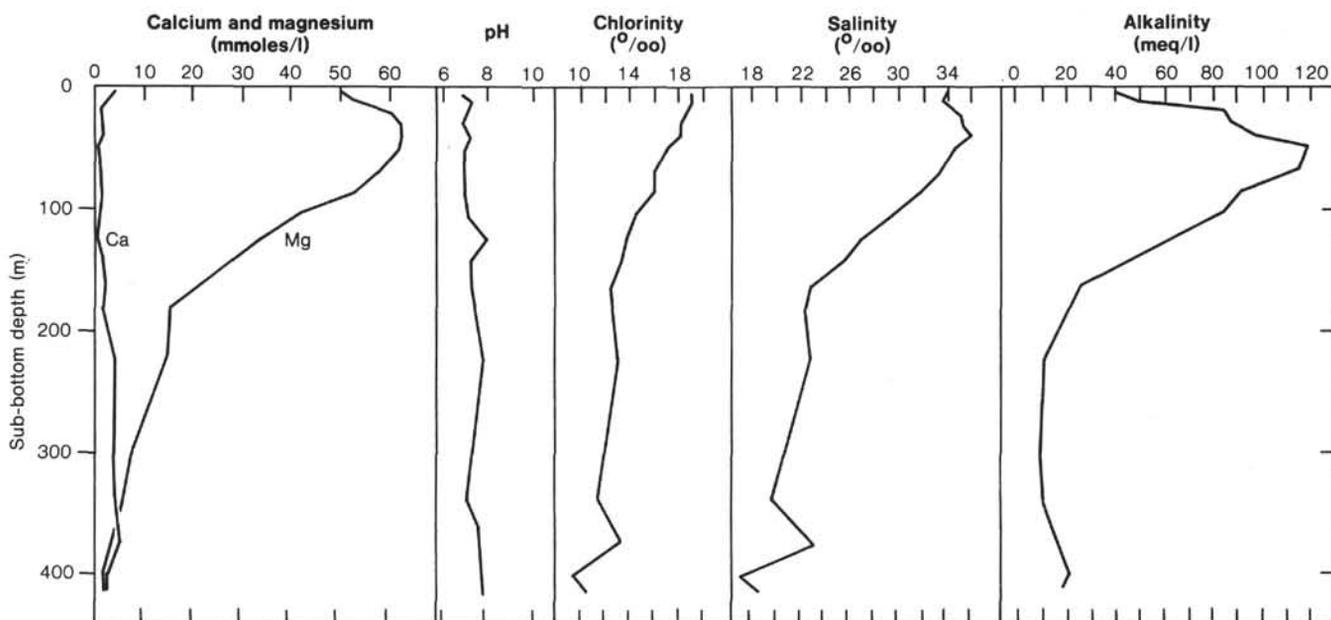


Figure 17. Chemistry of interstitial water at Site 568.

ty are affected by the gas hydrate any more than they are affected by very small changes in lithology. The lithology is very uniform, showing excursion of less than 100 m/s in velocity and 0.2 g/cm³, in density. Beds of limestone only 10 cm thick in the core are clearly seen in the logs. The hydrate certainly does not occur in sheets of solid material or it would have been seen in the logs and recovered in cores.

The lithology is much like that from Site 496. An upper unit (0-182 m) mainly comprises Recent to late Pleistocene massive olive gray mudstone. This mud is homogeneous, poorly reflective, and shows an almost straight sonic velocity trace at 1500 m/s. The mud is a hemipelagic sequence that unconformably drapes a lobate prograding sequence, as seen in seismic reflection records. The lobate sequence comprises a second lithologic unit (182-418 m) of mudstone that is mottled and bioturbated, although the Pliocene portion appears to belong to the hemipelagic cover. The lower mudstone is Miocene and becomes veined and fractured toward its base with reworked Eocene material. The lower cores show occasional inclined bedding that is near vertical in the last cores. This unit appears to be a Miocene prograding depositional lobe that has been covered by hemipelagic Pliocene and Pleistocene sediment. The lobe is positioned on the northwest bank of San José Canyon,

which suggests an overbank feature. Several such lobes can be seen in seismic records as part of the slope sediment covering the acoustic basement surface.

The analysis of benthic foraminifers assemblages indicates subsidence throughout the Miocene, and both uplift and subsidence in the Pliocene and Pleistocene. This history is more varied than that indicated by paleocologic analysis at Site 496 where the same Miocene trend is seen. It seems that the vertical component of motion during Neogene tectonism is more than a simple uplift or subsidence; the preliminary results will be refined and corrected for the effects of sea level change.

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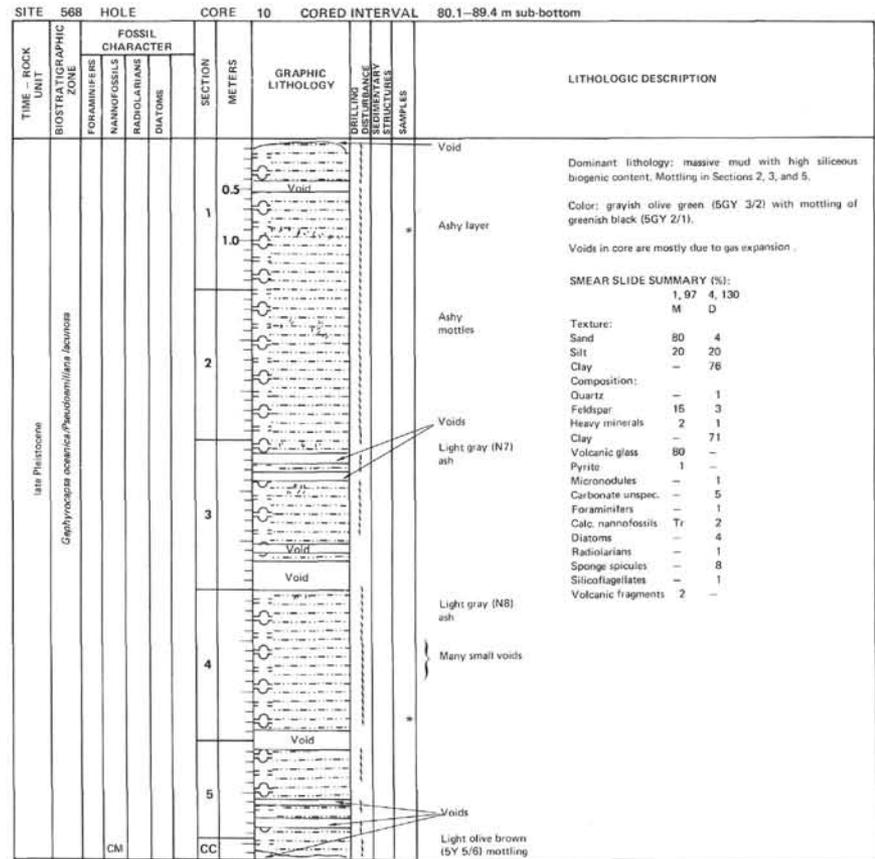
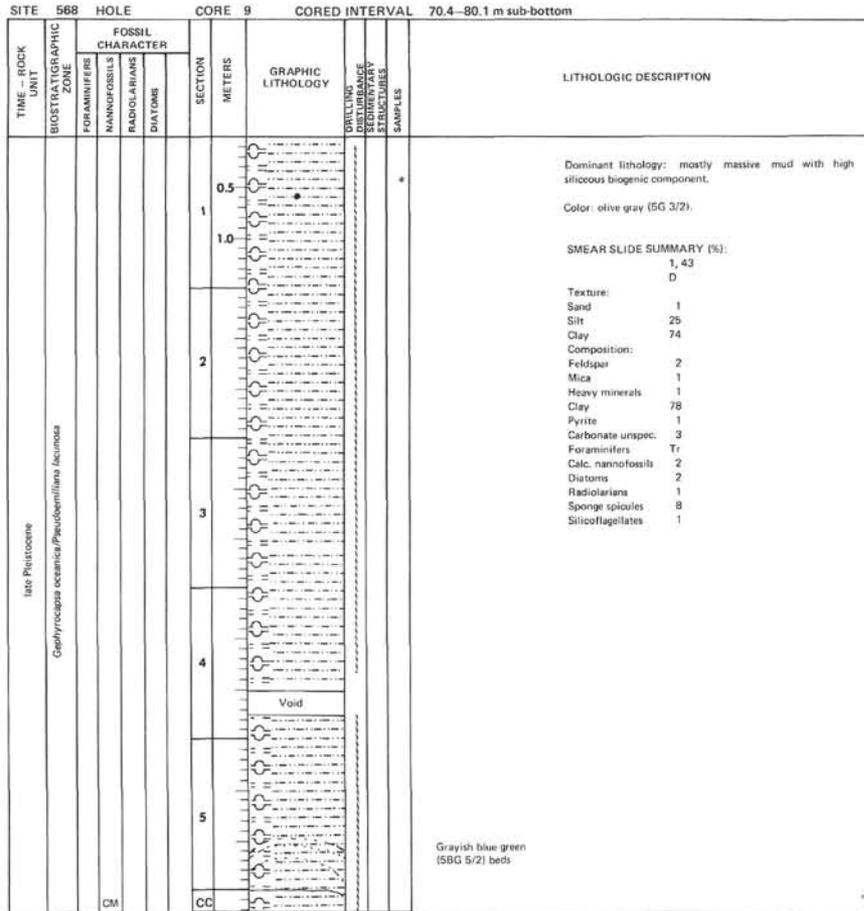
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SITE 568 HOLE CORE 3 CORED INTERVAL 13.0-22.7 m sub-bottom		FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE OBSERVATIONS	LITHOLOGIC DESCRIPTION
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FORAMINIFERS	DIATOMS				
				0.5	Void		Dominant lithology: massive, structureless mud. H ₂ S odor throughout core. Gas expansion cracks. Some greenish black (5G 2/1) mottling. Color: olive gray (5Y 3/2). SMEAR SLIDE SUMMARY (%): 1, 120 4, 123 6, 93 D D M Texture: Sand 2 1 15 Silt 20 15 20 Clay 78 84 65 Composition: Quartz 4 1 2 Feldspar 2 1 - Mica Tr - - Heavy minerals Tr Tr - Clay 73 80 60 Volcanic glass - - 25 Glauconite 1 2 1 Pyrite - Tr - Carbonate unsp. 2 2 - Calc. nanofossils 5 - - Diatoms 4 3 1 Radiolarians 2 1 - Sponge spicules 1 1 1 Fish remains 1 1 -
				1.0			
				2			
				3			
				4		5G 2/1 mottle	
				5	Void		
				6	Void Light gray (N7) ash 10Y 4/2 mottle		
				7	Void		
				CC			

SITE 568 HOLE CORE 4 CORED INTERVAL 22.7-32.3 m sub-bottom		FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE OBSERVATIONS	LITHOLOGIC DESCRIPTION
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FORAMINIFERS	DIATOMS				
				0.5			Dominant lithology: mud with high siliceous biogenic component. Mostly massive, structureless. Some mottling throughout. H ₂ S odor, gas expansion cracks throughout core. Color: mostly olive gray (5Y 3/2). SMEAR SLIDE SUMMARY (%): 3, 145 CC D M Texture: Sand 5 10 Silt 15 20 Clay 80 70 Composition: Quartz 5 5 Mica Tr - Heavy minerals 3 1 Clay 68 - Volcanic glass 10 20 Glauconite 2 2 Foraminifera 1 - Calc. nanofossils - 1 Diatoms 5 - Radiolarians 3 - Sponge spicules 2 - Plant debris 1 1
				1.0			
				2	Void		
				3			
				4		Black (N1) mottle	
				5	Void		
				6	Void Greenish black (5GY 2/1) mottle and pale greenish yellow (10Y 8/2) ash		
				CC	Voids Ashy layer		

SITE 568 HOLE		CORE 7		CORED INTERVAL 51.4-61.6 m sub-bottom																																																																														
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION																																																																												
		FORAMINIFERS	NANNOFOSSILS																																																																															
		RADIOLARIANS	DIATOMS																																																																															
Late Pleistocene	<i>Gephyrocapsa oceanica/Pseudonella lacunosa</i>			0.5		<p>Dominant lithology: mud. Mostly massive, although some lamination occurs throughout the core. Gas expansion cracks.</p> <p>Color: mostly olive gray (5Y 3/2).</p> <p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <thead> <tr> <th></th> <th>1, 138</th> <th>3, 107</th> <th>5, 47</th> </tr> <tr> <th></th> <th>M</th> <th>M</th> <th>D</th> </tr> </thead> <tbody> <tr> <td>Texture:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Sand</td> <td>25</td> <td>70</td> <td>2</td> </tr> <tr> <td>Silt</td> <td>35</td> <td>30</td> <td>10</td> </tr> <tr> <td>Clay</td> <td>40</td> <td>-</td> <td>88</td> </tr> <tr> <td>Composition:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Quartz</td> <td>20</td> <td>4</td> <td>2</td> </tr> <tr> <td>Feldspar</td> <td>5</td> <td>1</td> <td>-</td> </tr> <tr> <td>Heavy minerals</td> <td>-</td> <td>-</td> <td>1</td> </tr> <tr> <td>Clay</td> <td>35</td> <td>-</td> <td>60</td> </tr> <tr> <td>Volcanic glass</td> <td>30</td> <td>90</td> <td>-</td> </tr> <tr> <td>Glauconite</td> <td>2</td> <td>-</td> <td>1</td> </tr> <tr> <td>Carbonate unsp. spec.</td> <td>-</td> <td>-</td> <td>2</td> </tr> <tr> <td>Calc. nannofossils</td> <td>-</td> <td>-</td> <td>18</td> </tr> <tr> <td>Diatoms</td> <td>-</td> <td>-</td> <td>Tr</td> </tr> <tr> <td>Radiolarians</td> <td>-</td> <td>-</td> <td>Tr</td> </tr> <tr> <td>Silicoflagellates</td> <td>-</td> <td>-</td> <td>1</td> </tr> <tr> <td>Fish remains</td> <td>-</td> <td>-</td> <td>1</td> </tr> </tbody> </table>		1, 138	3, 107	5, 47		M	M	D	Texture:				Sand	25	70	2	Silt	35	30	10	Clay	40	-	88	Composition:				Quartz	20	4	2	Feldspar	5	1	-	Heavy minerals	-	-	1	Clay	35	-	60	Volcanic glass	30	90	-	Glauconite	2	-	1	Carbonate unsp. spec.	-	-	2	Calc. nannofossils	-	-	18	Diatoms	-	-	Tr	Radiolarians	-	-	Tr	Silicoflagellates	-	-	1	Fish remains	-	-	1
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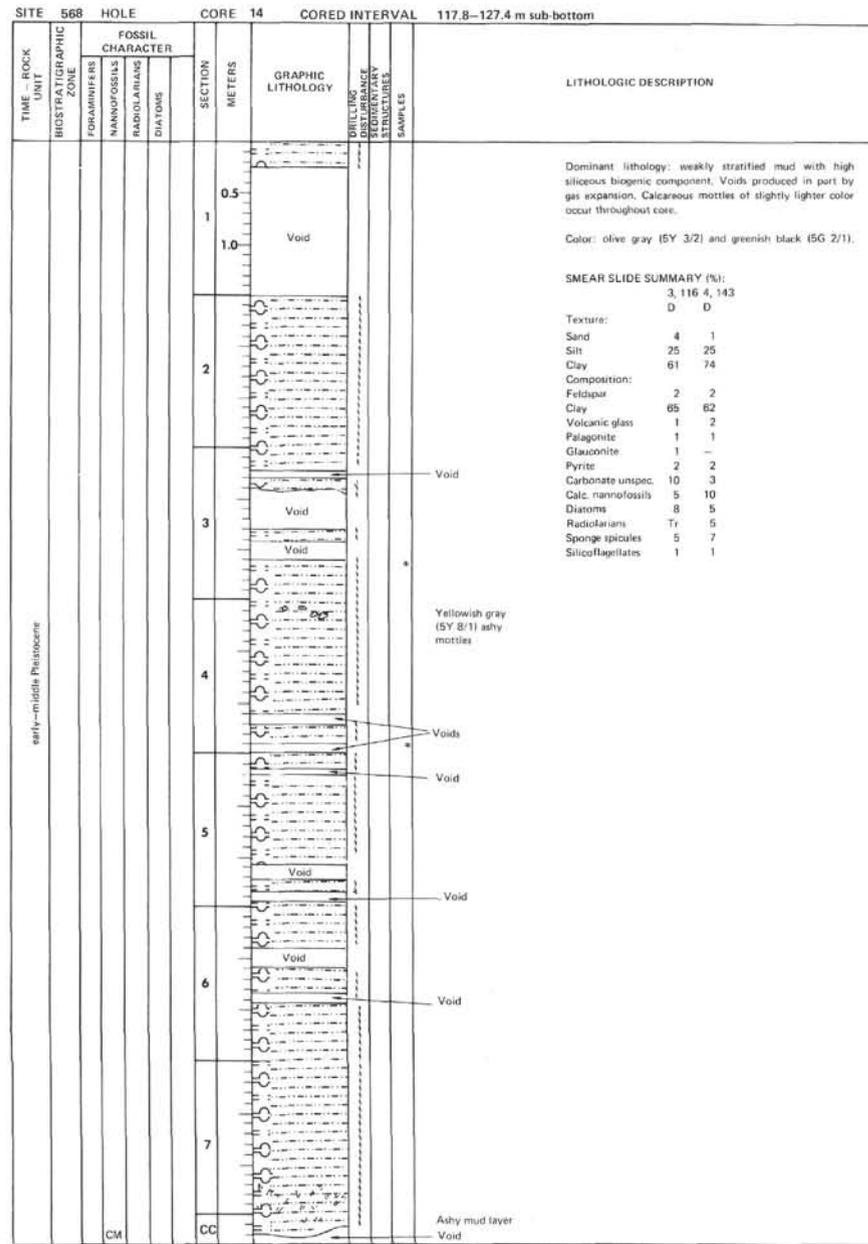
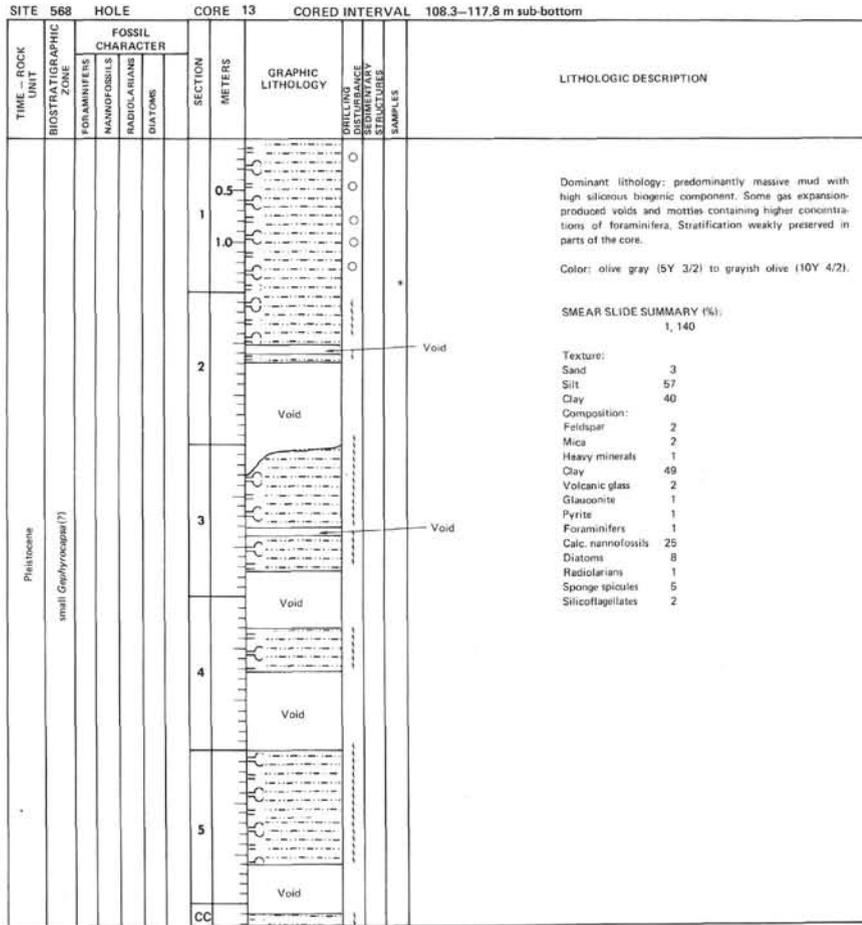


SITE 568 HOLE CORE 11 CORED INTERVAL 89.4-98.8 m sub-bottom

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION																																																					
		FORAMINIFERS	MAMMOFOSILS	RADIOLARIANS						DIATOMS																																																				
Pleistocene	small <i>Gephyrocapsa</i> ?																																																													
		CM			0.5 1.0			<p>Void</p> <p>Dominant lithology: massive mud with siliceous biogenic component.</p> <p>Color: moderate olive brown (5Y 4/4).</p> <p>This core was taken using the Pressure Core Barrel.</p> <p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <tr> <td></td> <td>1, 97</td> <td>CC</td> </tr> <tr> <td></td> <td>D</td> <td>M</td> </tr> <tr> <td>Texture:</td> <td></td> <td></td> </tr> <tr> <td>Sand</td> <td>8</td> <td>70</td> </tr> <tr> <td>Silt</td> <td>20</td> <td>25</td> </tr> <tr> <td>Clay</td> <td>72</td> <td>5</td> </tr> <tr> <td>Composition:</td> <td></td> <td></td> </tr> <tr> <td>Quartz</td> <td>2</td> <td>-</td> </tr> <tr> <td>Heavy minerals</td> <td>Tr</td> <td>-</td> </tr> <tr> <td>Clay</td> <td>56</td> <td>7</td> </tr> <tr> <td>Volcanic glass</td> <td>2</td> <td>40</td> </tr> <tr> <td>Carbonate unspec.</td> <td>5</td> <td>-</td> </tr> <tr> <td>Calc. nannofossils</td> <td>20</td> <td>-</td> </tr> <tr> <td>Diatoms</td> <td>5</td> <td>1</td> </tr> <tr> <td>Radiolarians</td> <td>Tr</td> <td>-</td> </tr> <tr> <td>Sponge spicules</td> <td>7</td> <td>-</td> </tr> <tr> <td>Silicoflagellates</td> <td>Tr</td> <td>-</td> </tr> <tr> <td>Volcanic fragments</td> <td>2</td> <td>2</td> </tr> </table> <p>White (N6) ash mottle</p>		1, 97	CC		D	M	Texture:			Sand	8	70	Silt	20	25	Clay	72	5	Composition:			Quartz	2	-	Heavy minerals	Tr	-	Clay	56	7	Volcanic glass	2	40	Carbonate unspec.	5	-	Calc. nannofossils	20	-	Diatoms	5	1	Radiolarians	Tr	-	Sponge spicules	7	-	Silicoflagellates	Tr	-	Volcanic fragments	2	2
	1, 97	CC																																																												
	D	M																																																												
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Composition:																																																														
Quartz	2	-																																																												
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Volcanic glass	2	40																																																												
Carbonate unspec.	5	-																																																												
Calc. nannofossils	20	-																																																												
Diatoms	5	1																																																												
Radiolarians	Tr	-																																																												
Sponge spicules	7	-																																																												
Silicoflagellates	Tr	-																																																												
Volcanic fragments	2	2																																																												
				CC				Void																																																						

SITE 568 HOLE CORE 12 CORED INTERVAL 98.8-108.3 m sub-bottom

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION																																																																																							
		FORAMINIFERS	MAMMOFOSILS	RADIOLARIANS						DIATOMS																																																																																						
Pleistocene	small <i>Gephyrocapsa</i> ?																																																																																															
		CM			0.5 1.0			<p>Dominant lithology: massive, structureless mud with high siliceous biogenic component. Some laminations are preserved. Gas expansion cracks. Most small voids in the core are due to gas expansion.</p> <p>Color: grayish olive green (5GY 3/2), grading through Section 2 to olive gray (5Y 3/2).</p> <p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <tr> <td></td> <td>1, 80</td> <td>2, 80</td> <td>5, 103</td> </tr> <tr> <td></td> <td>D</td> <td>D</td> <td>M</td> </tr> <tr> <td>Texture:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Sand</td> <td>5</td> <td>3</td> <td>20</td> </tr> <tr> <td>Silt</td> <td>20</td> <td>17</td> <td>75</td> </tr> <tr> <td>Clay</td> <td>75</td> <td>80</td> <td>5</td> </tr> <tr> <td>Composition:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Quartz</td> <td>5</td> <td>4</td> <td>-</td> </tr> <tr> <td>Feldspar</td> <td>2</td> <td>-</td> <td>15</td> </tr> <tr> <td>Mica</td> <td>1</td> <td>-</td> <td>-</td> </tr> <tr> <td>Heavy minerals</td> <td>1</td> <td>-</td> <td>Tr</td> </tr> <tr> <td>Clay</td> <td>42</td> <td>55</td> <td>20</td> </tr> <tr> <td>Volcanic glass</td> <td>2</td> <td>2</td> <td>60</td> </tr> <tr> <td>Glauconite</td> <td>3</td> <td>1</td> <td>-</td> </tr> <tr> <td>Pyrite</td> <td>3</td> <td>-</td> <td>-</td> </tr> <tr> <td>Carbonate unspec.</td> <td>2</td> <td>8</td> <td>2</td> </tr> <tr> <td>Foraminifers</td> <td>3</td> <td>Tr</td> <td>-</td> </tr> <tr> <td>Calc. nannofossils</td> <td>15</td> <td>3</td> <td>-</td> </tr> <tr> <td>Diatoms</td> <td>5</td> <td>4</td> <td>1</td> </tr> <tr> <td>Radiolarians</td> <td>5</td> <td>3</td> <td>Tr</td> </tr> <tr> <td>Sponge spicules</td> <td>1</td> <td>4</td> <td>-</td> </tr> <tr> <td>Silicoflagellates</td> <td>-</td> <td>3</td> <td>-</td> </tr> </table>		1, 80	2, 80	5, 103		D	D	M	Texture:				Sand	5	3	20	Silt	20	17	75	Clay	75	80	5	Composition:				Quartz	5	4	-	Feldspar	2	-	15	Mica	1	-	-	Heavy minerals	1	-	Tr	Clay	42	55	20	Volcanic glass	2	2	60	Glauconite	3	1	-	Pyrite	3	-	-	Carbonate unspec.	2	8	2	Foraminifers	3	Tr	-	Calc. nannofossils	15	3	-	Diatoms	5	4	1	Radiolarians	5	3	Tr	Sponge spicules	1	4	-	Silicoflagellates	-	3	-
	1, 80	2, 80	5, 103																																																																																													
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Mica	1	-	-																																																																																													
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Diatoms	5	4	1																																																																																													
Radiolarians	5	3	Tr																																																																																													
Sponge spicules	1	4	-																																																																																													
Silicoflagellates	-	3	-																																																																																													
				CC				Void																																																																																								
								Olive black (5Y 2/1) ashy layers																																																																																								

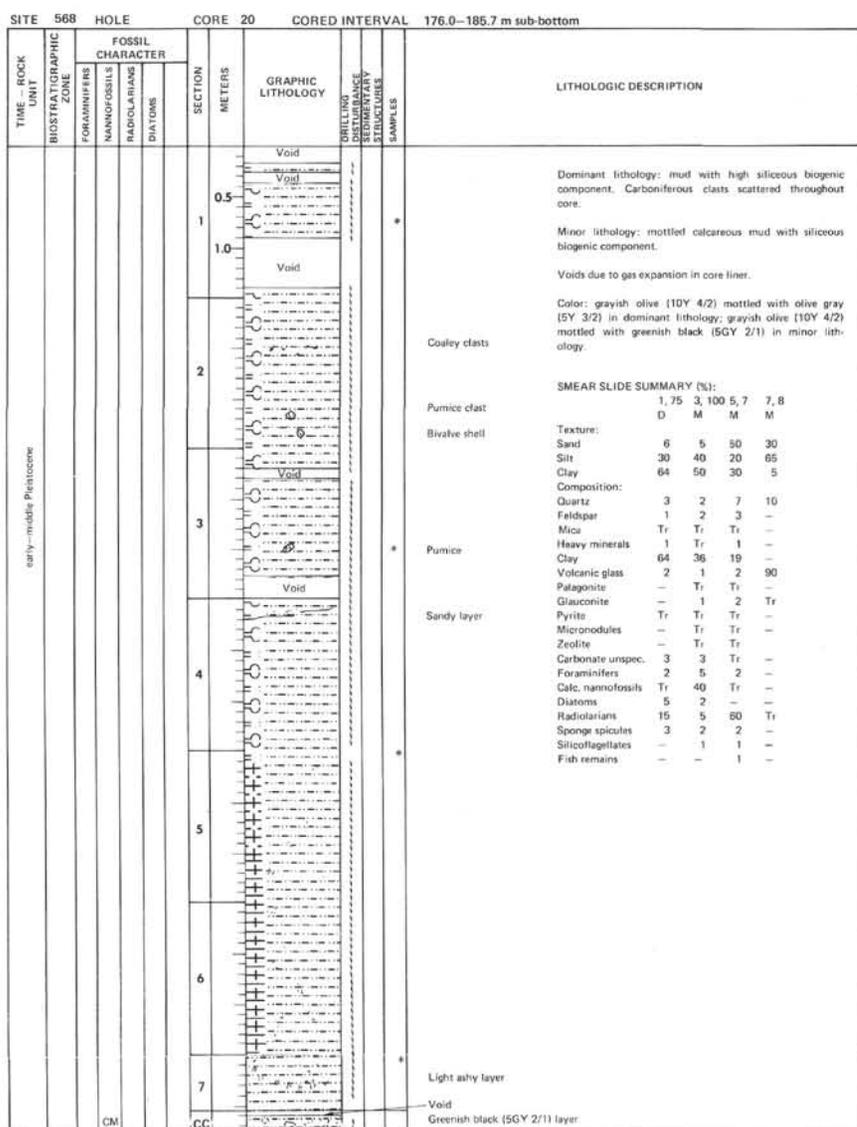
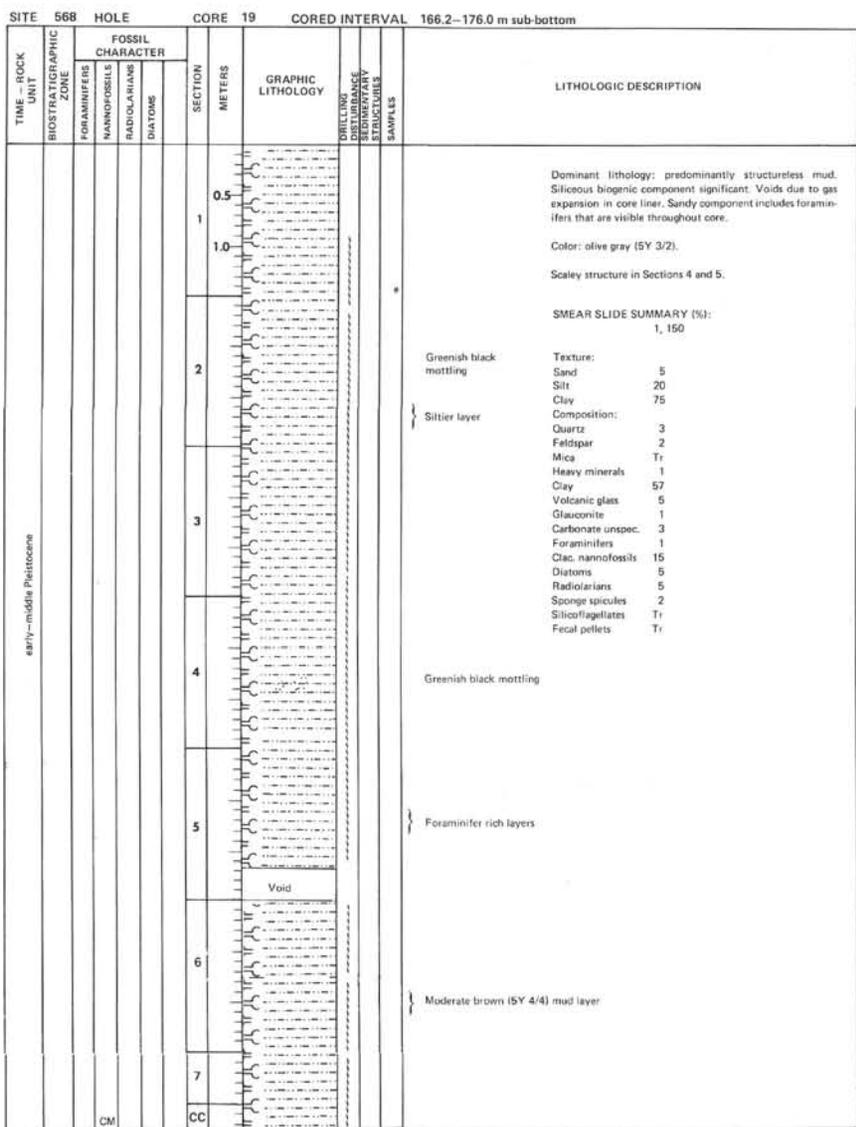


SITE 568 HOLE		CORE 15		CORED INTERVAL 127.4-137.2 m sub-bottom						
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE RECORD CORRECTION SUSPENSIVE STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION		
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS					DIATOMS	
early-middle Pleistocene	FP							Dark (N3) layer		
								0.5	Void	Dominant lithology: soupy mud with darker mud layers. Voids due in most cases to gas expansion. Color: grayish olive (10Y 4/2) to olive gray (5Y 3/2). SMEAR SLIDE SUMMARY (%): Foraminifera 1, 19 D
								1.0	Void	
								2	Void	Texture: Sand 3 Silt 10 Clay 87 Composition: Feldspar 1 Heavy minerals Tr Clay 62 Volcanic glass 1 Glauconite 1 Pyrite 2 Carbonate unspec. 3 Foraminifera Tr Calc. nanofossils 20 Diatoms 4 Radiolarians 2 Sponge spicules 4 Silicoflagellates Tr
								3	Void	
								4	Void	
5	Void									
6	Void									
								CC		

SITE 568 HOLE		CORE 16		CORED INTERVAL 137.2-146.9 m sub-bottom						
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE RECORD CORRECTION SUSPENSIVE STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION		
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS					DIATOMS	
early-middle Pleistocene	FP							Dominant lithology: predominantly massive mud with some lamination visible in lower sections. Color: olive gray (5Y 3/2), interbedding in Sections 5 and 6 with dark greenish gray (5GY 4/1). SMEAR SLIDE SUMMARY (%): 4, 18, 5, 80 M Texture: Sand 20 20 Silt 80 30 Clay - 50 Composition: Quartz 10 15 Feldspar 5 10 Mica - Tr Heavy minerals 3 Tr Clay - 44 Volcanic glass 78 1 Glauconite - 1 Pyrite 2 2 Carbonate unspec. 2 3 Foraminifera - 1 Calc. nanofossils - 5 Diatoms - 4 Radiolarians - 3 Sponge spicules - 1		
								0.5	Void	
								1.0	Void	
								2	Void	Light greenish gray (5GY 8/1) mottle
								3	Void	Ashy?
								4	Void	Pale greenish yellow (10Y 3/2) ashy mottle
5	Void	Sandier horizon								
6	Void	Sedimentary breccia								
			Sandy horizon							
			Greenish black layer							
				CC						

SITE 568 HOLE CORE 17 CORED INTERVAL 146.9-156.5 m sub-bottom		SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE			
		0.5	Void	* Pale olive (10Y 6/2) burrow Dominant lithology: predominantly massive mud with some faint lamination in most sections. Most small voids due to gas expansion. Color: olive gray (5Y 3/2). SMEAR SLIDE SUMMARY (%): 1, 20, 3, 36 CC D D M Texture: Sand 10 5 10 Silt 20 15 80 Clay 70 80 10 Composition: Quartz 10 3 3 Feldspar 5 1 - Mica 1 Tr - Heavy minerals 2 1 - Clay 64 80 10 Volcanic glass 2 - 87 Glauconite 1 - - Calc. nannofossils 2 5 - Diatoms 1 - - Radiolarians 3 3 - Sponge spicules 4 2 - Silicoflagellates - Tr - Fish remains - Tr - Plant debris - Tr -
		1.0	Void	
		2	Void	* Pumice pebbles and burrowing in this horizon
		3	Void	
		4	Void	Ashy layer
		5	Void	
		6	Void	Voids
		7	Void	Ashy mud
		CC		* Light greenish (5G 8/1) ash

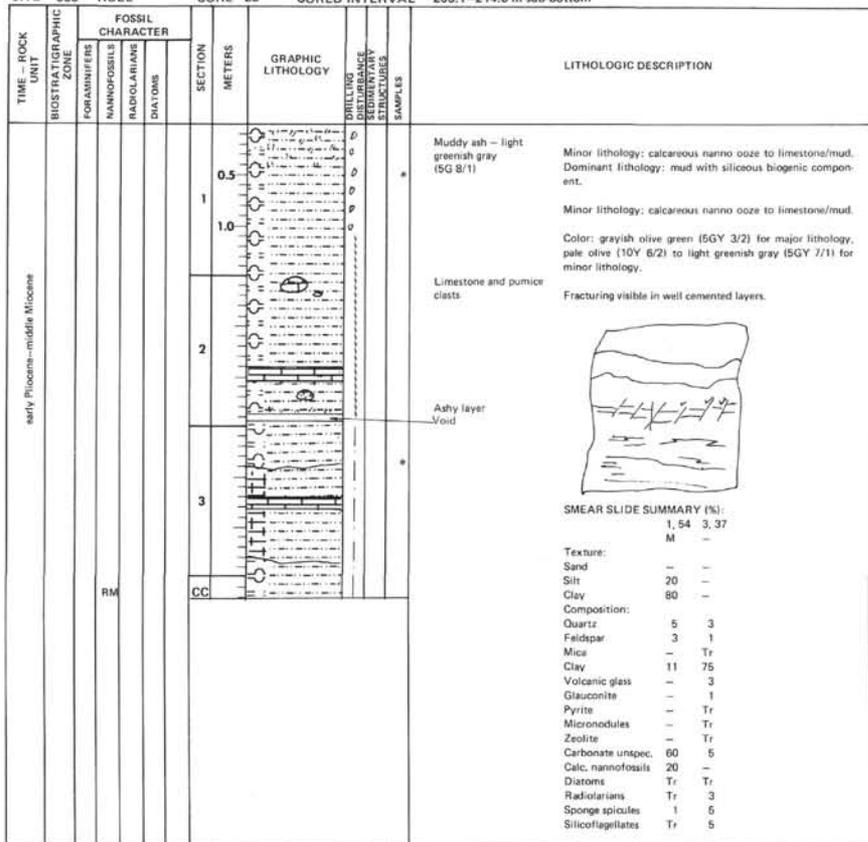
SITE 568 HOLE CORE 18 CORED INTERVAL 156.5-166.2 m sub-bottom		SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE			
		0.5	Void	* Dominant lithology: predominantly structureless and with faint lamination preserved in Section 4. Voids due to gas expansion. Color: olive gray (5Y 3/2). SMEAR SLIDE SUMMARY (%): 3, 65, 3, 100, 4, 32 D D D Texture: Sand 20 10 20 Silt 25 15 25 Clay 55 75 55 Composition: Quartz 5 3 15 Feldspar 2 2 5 Heavy minerals 1 1 2 Clay 52 63 39 Volcanic glass 1 - 2 Palagonite 2 Tr - Glauconite - 1 3 Pyrite - Tr 5 Micronodules - Tr - Zeolite - Tr - Carbonate unsp. - 3 3 Foraminifers 1 Tr - Calc. nannofossils 5 - 5 Diatoms 15 10 10 Radiolarians 10 3 10 Sponge spicules 5 9 1
		1.0	Void	
		2		Frothy texture from gas bubbling greenish black (5GY 2/1) burrow
		3		
		4		
		5		Greenish black mottling
		CC		Void Greenish black (5GY 2/1) mottling



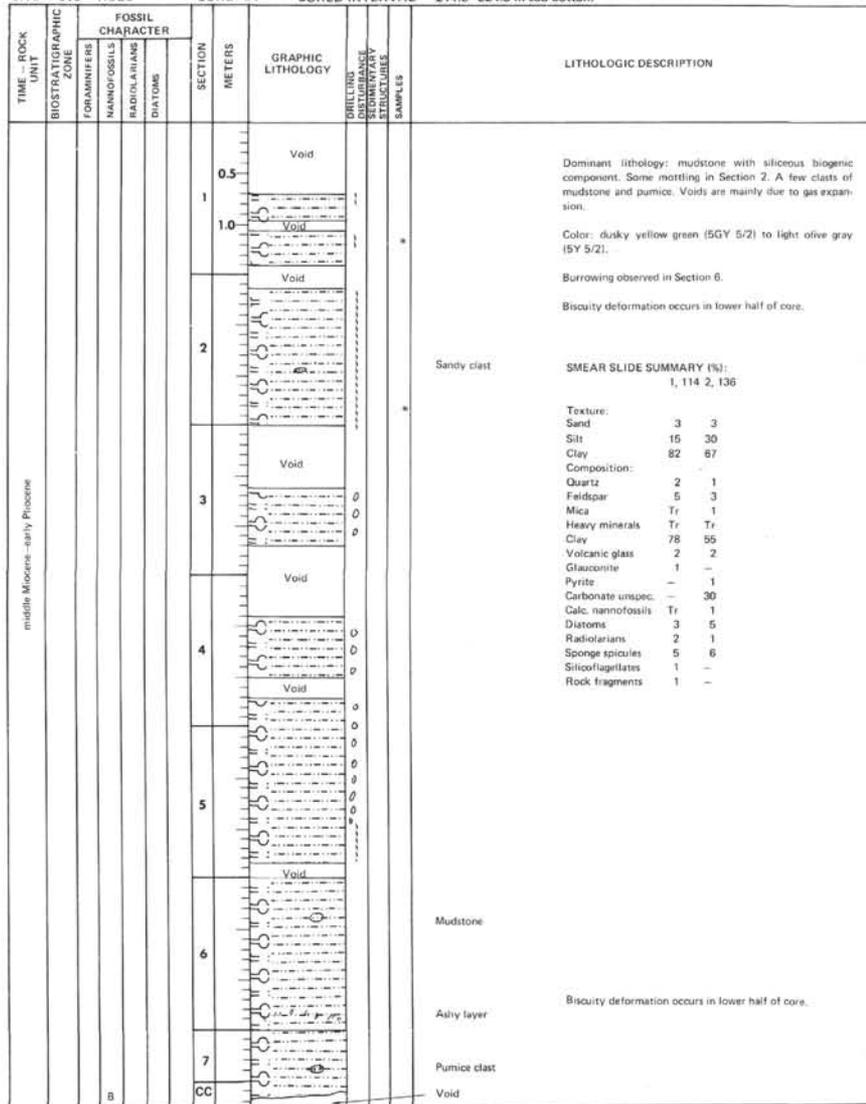
SITE 568		HOLE		CORE 21		CORED INTERVAL 185.7-195.4 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE DESTRUCTURE SAMPLES	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	RADIOLARIANS				
early-middle Pleistocene							<p>Pressure Core Barrel.</p> <p>Dominant lithology: soupy, structureless mud with high siliceous biogenic content.</p> <p>Minor lithology: from Core Catcher: calcareous mudstone containing clasts of darker color.</p> <p>Color: dominant lithology is grayish olive (10Y 4/2), minor lithology is pale olive (10Y 4/2), with darker clast of black to pale green (L.u., N1 to 10G 6/2).</p> <p>SMEAR SLIDE SUMMARY (%): 1, 126</p> <p>Texture: Sand 5 Silt 20 Clay 75</p> <p>Composition: Quartz 7 Feldspar 3 Mica Tr Heavy minerals Tr Clay 73 Glauconite 1 Pyrite Tr Micronodules Tr Zeolite Tr Carbonate unsp. 2 Foraminifers 3 Calc. nannofossils Tr Diatoms Tr Radiolarians 5 Sponge spicules 3 Silicoflagellates 3</p>

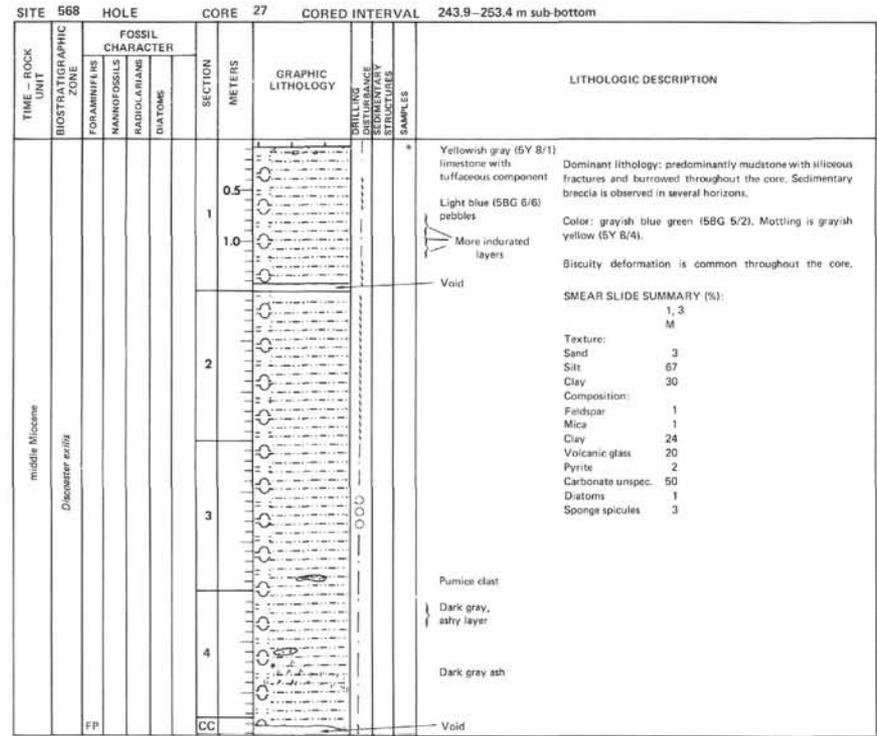
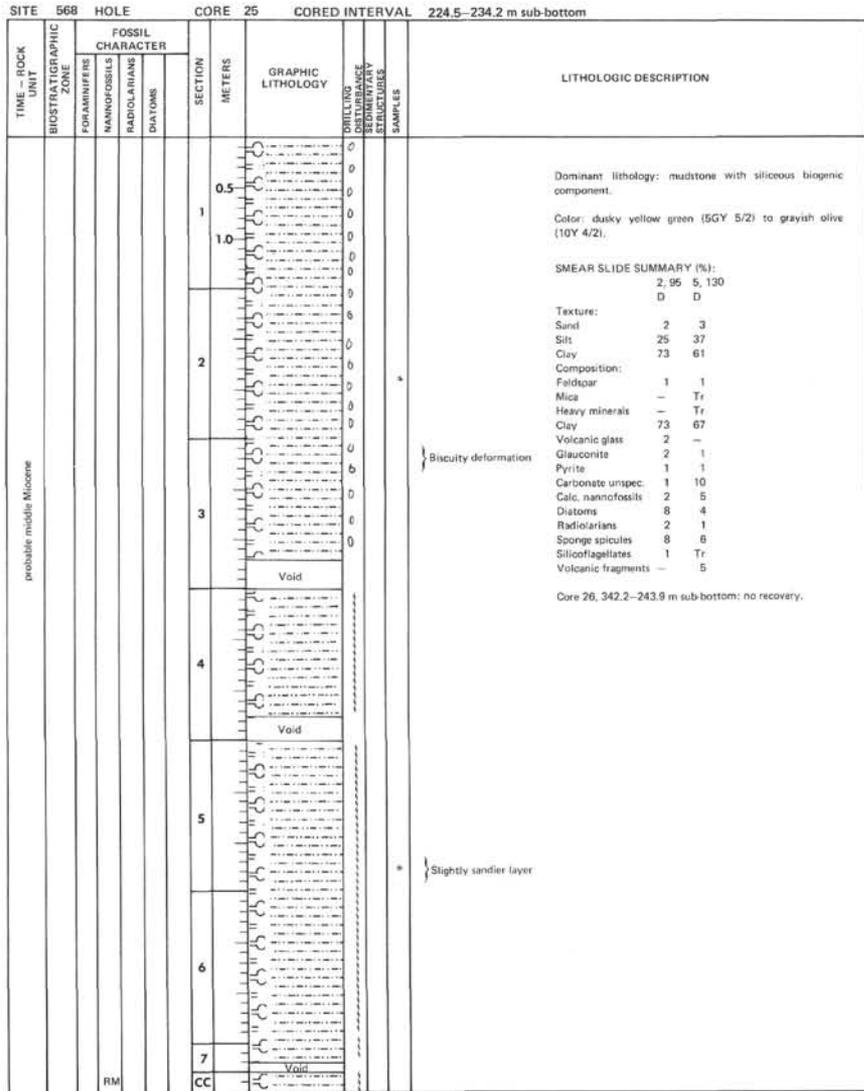
SITE 568		HOLE		CORE 22		CORED INTERVAL 195.4-205.1 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE DESTRUCTURE SAMPLES	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	RADIOLARIANS				
middle Miocene-early Pliocene							<p>Dominant lithology: predominantly massive mud with high siliceous biogenic component. Small mudstone clasts and a few sandier horizons occur.</p> <p>Minor lithology: from top of Section 1, micro-conglomeratic limestone, shot through with fractures.</p> <p>Color: major lithology is grayish olive green (5GY 3/2) Minor is light olive gray (5Y 5/2).</p> <p>Dark pebbles in both lithologies are black (N1) to moderate blue green (5BG 4/6).</p> <p>SMEAR SLIDE SUMMARY (%): 1, 130 3, 77 CC</p> <p>Texture: Sand 5 20 30 Silt 30 40 30 Clay 65 40 40</p> <p>Composition: Quartz 5 15 10 Feldspar 3 5 5 Mica - Tr 3 Heavy minerals Tr 1 Tr Clay 63 78 72 Volcanic glass 5 40 - Glauconite - 1 - Carbonate unsp. - Tr - Foraminifers 3 - Tr Calc. nannofossils - Tr Tr Diatoms 3 Tr 3 Radiolarians 5 Tr Tr Sponge spicules 9 Tr 2 Silicoflagellates 3 Tr 2 Rock fragments - - 3</p>

SITE 568 HOLE CORE 23 CORED INTERVAL 205.1-214.8 m sub-bottom



SITE 568 HOLE CORE 24 CORED INTERVAL 214.8-224.5 m sub-bottom





SITE	668	HOLE	CORE	28	CORED INTERVAL	253.4-263.1 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS			
		DIATOMS			DRILLING SEQUENCE	SEMI-DIAGRAM	SAMPLES
					1.0		
					2		SMEAR SLIDE SUMMARY (%): 1, 45 1, 47 7, 36 Texture: Sand 30 8 10 Silt 20 15 10 Clay 50 77 80 Composition: Quartz - 3 2 Feldspar 5 2 - Mica - 2 - Heavy minerals Tr - - Clay 62 65 70 Volcanic glass 25 10 2 Glauconite - 1 1 Pyrite 5 - - Carbonate unsp. - - 5 Foraminifers - - 2 Calc. nannofossils - - 10 Radiolarians 1 1 1 Sponge spicules 2 - - Fish remains - - 1
					3		
					4		Void
					5		
					6		Void
					7		
							Slightly calcareous

SITE	568	HOLE	CORE	29	CORED INTERVAL	263.1-272.6 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS			
		DIATOMS			DRILLING SEQUENCE	SEMI-DIAGRAM	SAMPLES
					1.0		
					2		Ashy layer SMEAR SLIDE SUMMARY (%): 1, 20 1, 127 3, 147 D M M Texture: Sand 5 70 - Silt 30 30 - Clay 65 - - Composition: Quartz 1 - - Feldspar 3 15 15 Mica Tr 2 1 Heavy minerals - 2 - Clay 50 - - Volcanic glass - 82 82 Glauconite Tr - - Pyrite 1 1 2 Micronodules 2 - - Carbonate unsp. 20 - - Calc. nannofossils 5 - - Diatoms 2 - - Radiolarians 1 - - Sponge spicules 2 - -
					3		Light olive gray (5Y 6/1) ash
					4		Ashy, sandy layer
					5		Ashy, sandy layer
					6		Void
					7		Disturbed ash layer
							Void

SITE 568		HOLE		CORE 32		CORED INTERVAL 292.0-301.6 m sub-bottom				
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DISTURBANCE STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION			
		FORAMINIFERS	DIATOMS							
early middle Miocene <i>Sphenolithus heteromorphus</i>				0.5			Dominant lithology: mudstone (effervesces slightly with HCl). Burrowed throughout core, fractured by ubiquitous small veination; these are often kinked, anastomosing, or both. Color: grayish olive (10Y 4/2) to dark olive gray (5GY 4/1).			
				1				Breccia		
				1.0						
				2						
				3						Yellow (5Y 6/4) burrowing
				4						Breccia
				5						Breccia
6						Breccia				
CC							Pale olive nanno ooze Ash layer			

SITE 568		HOLE		CORE 33		CORED INTERVAL 301.6-311.3 m sub-bottom				
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DISTURBANCE STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION			
		FORAMINIFERS	DIATOMS							
early middle Miocene <i>Sphenolithus heteromorphus</i>				0.5			Breccia			
				1				Dominant lithology: very fine mudstone, cut throughout by small, anastomosing veins. Burrows occur throughout the core. Some scaly fabric.		
				1.0				Minor lithology: claystone, in gradational contact with overlying mudstone, and exhibiting same sedimentary structures. Color: grayish olive (10Y 4/2) to pale olive (10Y 6/4).		
				2						
				3						
				4						Burrowed and fractured breccia
				5						Breccia
6										
CC										

SITE 568 HOLE CORE 38 CORED INTERVAL 350.2-359.9 m sub bottom		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG SYMBOLS FOR SEDIMENTARY STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION																																																																								
TIME ROCK UNIT	BIOSTRATIGRAPHIC ZONE					FOSSIL CHARACTER	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS																																																																			
early Miocene	Helicospira anguligera	0.5			<p>Dominant lithology: mudstone with ash admixture. Drilling disturbance obscures sedimentary structures.</p> <p>Color: mostly grayish olive (10Y 4/2).</p> <p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <tr> <td></td> <td>1, 67</td> <td>3, 4</td> <td>6, 97</td> </tr> <tr> <td></td> <td>D</td> <td>M</td> <td>M</td> </tr> </table> <p>Texture:</p> <table border="1"> <tr> <td>Sand</td> <td>20</td> <td>-</td> <td>5</td> </tr> <tr> <td>Silt</td> <td>5</td> <td>30</td> <td>80</td> </tr> <tr> <td>Clay</td> <td>15</td> <td>70</td> <td>15</td> </tr> </table> <p>Composition:</p> <table border="1"> <tr> <td>Quartz</td> <td>5</td> <td>10</td> <td>3</td> </tr> <tr> <td>Feldspar</td> <td>3</td> <td>-</td> <td>2</td> </tr> <tr> <td>Mica</td> <td>1</td> <td>Tr</td> <td>Tr</td> </tr> <tr> <td>Clay</td> <td>65</td> <td>45</td> <td>-</td> </tr> <tr> <td>Volcanic glass</td> <td>12</td> <td>40</td> <td>-</td> </tr> <tr> <td>Glaucinite</td> <td>-</td> <td>1</td> <td>3</td> </tr> <tr> <td>Pyrite</td> <td>-</td> <td>-</td> <td>3</td> </tr> <tr> <td>Carbonate unspec.</td> <td>-</td> <td>5</td> <td>60</td> </tr> <tr> <td>Foraminifers</td> <td>-</td> <td>-</td> <td>5</td> </tr> <tr> <td>Calc. nannofossils</td> <td>10</td> <td>Tr</td> <td>15</td> </tr> <tr> <td>Radiolarians</td> <td>3</td> <td>-</td> <td>10</td> </tr> <tr> <td>Silicoflagellates</td> <td>Tr</td> <td>-</td> <td>-</td> </tr> <tr> <td>Fish remains</td> <td>1</td> <td>-</td> <td>-</td> </tr> </table>		1, 67	3, 4	6, 97		D	M	M	Sand	20	-	5	Silt	5	30	80	Clay	15	70	15	Quartz	5	10	3	Feldspar	3	-	2	Mica	1	Tr	Tr	Clay	65	45	-	Volcanic glass	12	40	-	Glaucinite	-	1	3	Pyrite	-	-	3	Carbonate unspec.	-	5	60	Foraminifers	-	-	5	Calc. nannofossils	10	Tr	15	Radiolarians	3	-	10	Silicoflagellates	Tr	-	-	Fish remains	1	-	-
			1, 67	3, 4		6, 97																																																																							
			D	M		M																																																																							
		Sand	20	-		5																																																																							
		Silt	5	30		80																																																																							
		Clay	15	70		15																																																																							
		Quartz	5	10		3																																																																							
Feldspar	3	-	2																																																																										
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4																																																																													
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SITE 568 HOLE CORE 39 CORED INTERVAL 359.9-369.6 m sub bottom		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG SYMBOLS FOR SEDIMENTARY STRUCTURES SAMPLES	LITHOLOGIC DESCRIPTION																																																						
TIME ROCK UNIT	BIOSTRATIGRAPHIC ZONE					FOSSIL CHARACTER	FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS																																																	
early Miocene	Helicospira anguligera	0.5			<p>Dominant lithology: mudstone containing ash and mottling from bioturbation. Blue green pebbly clasts in Section 1.</p> <p>Color: grayish olive (10Y 4/2).</p> <p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <tr> <td></td> <td>M</td> <td>D</td> </tr> <tr> <td>Texture:</td> <td>20</td> <td>5</td> </tr> <tr> <td>Sand</td> <td>70</td> <td>60</td> </tr> <tr> <td>Silt</td> <td>10</td> <td>39</td> </tr> <tr> <td>Clay</td> <td>5</td> <td>2</td> </tr> <tr> <td>Quartz</td> <td>8</td> <td>2</td> </tr> <tr> <td>Feldspar</td> <td>1</td> <td>Tr</td> </tr> <tr> <td>Heavy minerals</td> <td>62</td> <td>74</td> </tr> <tr> <td>Clay</td> <td>7</td> <td>-</td> </tr> <tr> <td>Volcanic glass</td> <td>5</td> <td>1</td> </tr> <tr> <td>Pyrite</td> <td>-</td> <td>1</td> </tr> <tr> <td>Micronodules</td> <td>5</td> <td>4</td> </tr> <tr> <td>Carbonate unspec.</td> <td>-</td> <td>1</td> </tr> <tr> <td>Foraminifers</td> <td>2</td> <td>5</td> </tr> <tr> <td>Calc. nannofossils</td> <td>1</td> <td>2</td> </tr> <tr> <td>Diatoms</td> <td>2</td> <td>2</td> </tr> <tr> <td>Radiolarians</td> <td>2</td> <td>2</td> </tr> <tr> <td>Sponge spicules</td> <td>2</td> <td>2</td> </tr> </table>		M	D	Texture:	20	5	Sand	70	60	Silt	10	39	Clay	5	2	Quartz	8	2	Feldspar	1	Tr	Heavy minerals	62	74	Clay	7	-	Volcanic glass	5	1	Pyrite	-	1	Micronodules	5	4	Carbonate unspec.	-	1	Foraminifers	2	5	Calc. nannofossils	1	2	Diatoms	2	2	Radiolarians	2	2	Sponge spicules	2	2
			M	D																																																							
		Texture:	20	5																																																							
		Sand	70	60																																																							
Silt	10	39																																																									
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Heavy minerals	62	74																																																									
Clay	7	-																																																									
Volcanic glass	5	1																																																									
Pyrite	-	1																																																									
Micronodules	5	4																																																									
Carbonate unspec.	-	1																																																									
Foraminifers	2	5																																																									
Calc. nannofossils	1	2																																																									
Diatoms	2	2																																																									
Radiolarians	2	2																																																									
Sponge spicules	2	2																																																									
1		Void																																																									
1.0		Void																																																									
2		Void		Ashy layer - light olive gray (5Y 6/1)																																																							
3		Void																																																									
4		Void																																																									
RM																																																											

SITE 568		HOLE		CORE 40		CORED INTERVAL 369.6-379.3 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NAKNOFOSSILS				
early Miocene	<i>Helicophaera amygdalata</i>						Dominant lithology: mudstone with veins and fractures throughout the core. Bedding planes are not horizontal. Burrowing is common. Color: grayish olive (10Y 4/2). SMEAR SLIDE SUMMARY (%): 1, 134 4, 140 D D Texture: Sand 5 - Silt 35 20 Clay 60 80 Composition: Quartz - 5 Feldspar 1 - Mica Tr - Heavy minerals Tr 1 Clay 80 65 Volcanic glass - 10 Glauconite Tr 1 Pyrite 1 - Micronodules 1 - Carbonate unspcc. 2 10 Foraminifers 1 - Calc. nanofossils 5 - Diatoms 2 - Radiolarians 5 - Sponge spicules 2 1
				0.5	Void		
				1.0	Void		
				2			
				3			
				4			
				5			
				6			
				CC			

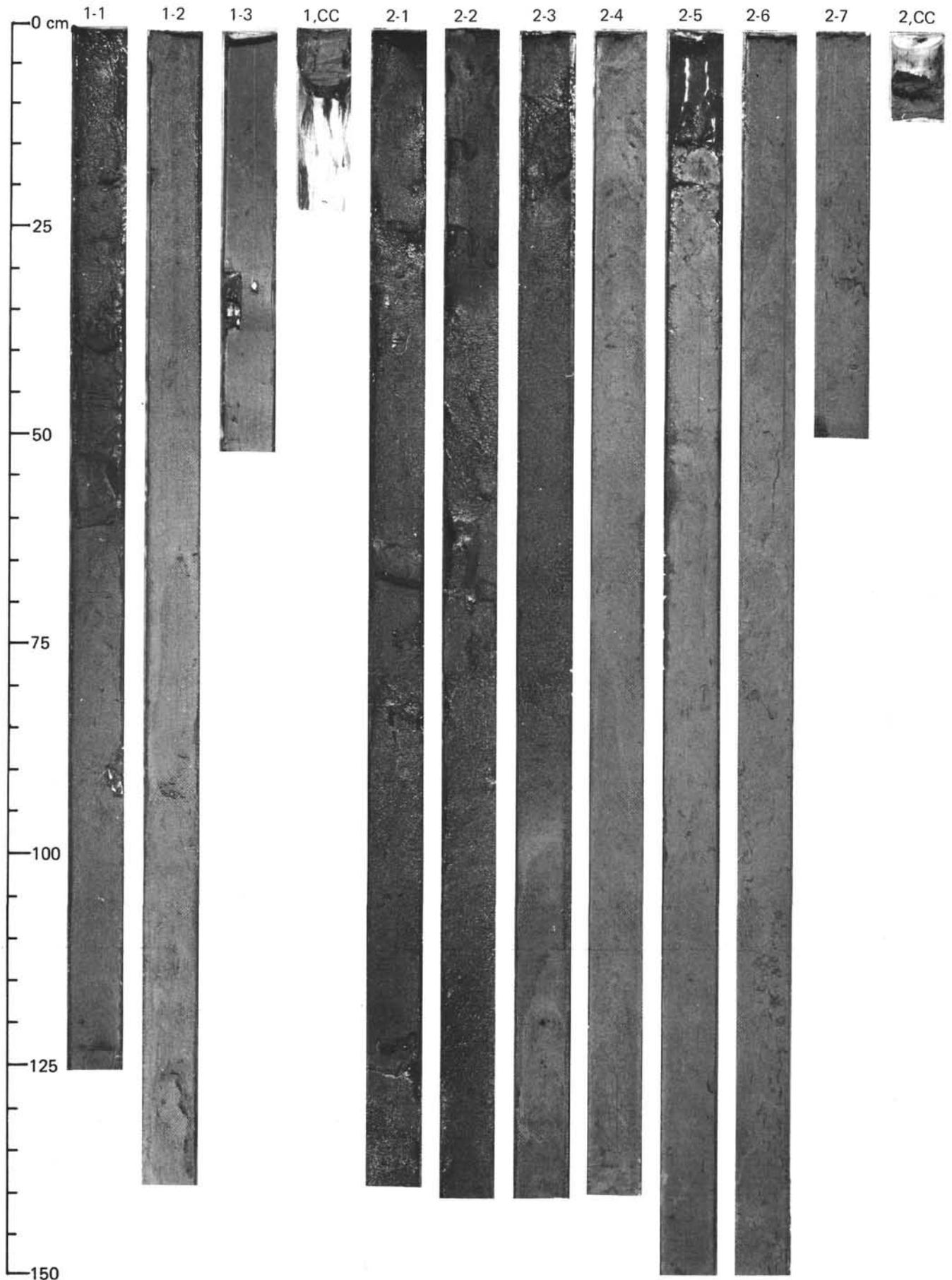
SITE 568		HOLE		CORE 41		CORED INTERVAL 379.3-389.0 m sub-bottom	
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NAKNOFOSSILS				
early Miocene	<i>Helicophaera amygdalata</i>						Dominant lithology: mudstone with veins, fractures, and burrow mottling throughout the core. Fracture angles dip about 30°. Scaley deformation. Color: grayish olive (10Y 4/2). SMEAR SLIDE SUMMARY (%): 1, 120 CC D M Texture: Sand 25 90 Silt 60 15 Clay 15 5 Composition: Quartz 2 5 Feldspar 3 10 Heavy minerals Tr 1 Clay 70 5 Volcanic glass 3 40 Glauconite Tr 5 Pyrite - 2 Carbonate unspcc. 10 10 Foraminifers 2 5 Calc. nanofossils 1 - Diatoms 2 5 Radiolarians 4 7 Sponge spicules 2 5 Silicoflagellates 1 -
				0.5	Void		
				1.0	Void		
				2			
				3			
				4			
				5			
				6			
				CC			

SITE 568 HOLE		CORE 42		CORED INTERVAL 389.0-398.7 m sub-bottom				
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILL LOG OBSERVATIONS	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS				
early Miocene	<i>Sphenolithus belemnos</i> - <i>Triquetrorhabdulus carinatus</i>				0.5			<p>Dominant lithology: mudstone with abundant burrowing. Bedding dips up to 70° to horizontal in several sections. Veinational and fracturing are common.</p> <p>Color: dusky yellowish green (10GY 3/2) with abundant clasts and mottles of light olive brown (5Y 5/6).</p>
					1.0			
					2.0			
					3.0			
					4.0			
					5.0			
					6.0			
			7.0					
			CC				70° apparent dip	
							60° apparent dip	

SITE 568 HOLE		CORE 43		CORED INTERVAL 398.7-408.9 m sub-bottom																																																																																								
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILL LOG OBSERVATIONS	LITHOLOGIC DESCRIPTION																																																																																				
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS					DIAATOMS																																																																																			
early Miocene	<i>Sphenolithus belemnos</i> - <i>Triquetrorhabdulus carinatus</i>				0.5			<p>Dominant lithology: mudstone with abundant burrows, veinational, and fracturing. Gas hydrate crystals found in Section 4, 45-83 cm.</p> <p>Color: grayish olive green (5GY 3/2) to pale olive (10Y 6/2) with mottles of pale olive (10Y 6/2) to moderate olive brown (5Y 4/4).</p>																																																																																				
					1.0																																																																																							
					2.0																																																																																							
					3.0																																																																																							
					4.0			<p>Void</p> <p>Ashy layer</p> <p>Gas hydrate crystals</p>																																																																																				
								<p>SMEAR SLIDE SUMMARY (%):</p> <table border="1"> <thead> <tr> <th></th> <th>2, 40</th> <th>4, 33</th> <th>4, 71</th> </tr> <tr> <th></th> <th>D</th> <th>D</th> <th>M</th> </tr> </thead> <tbody> <tr> <td>Texture:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Sand</td> <td>30</td> <td>5</td> <td>50</td> </tr> <tr> <td>Silt</td> <td>40</td> <td>17</td> <td>30</td> </tr> <tr> <td>Clay</td> <td>30</td> <td>78</td> <td>20</td> </tr> <tr> <td>Composition:</td> <td></td> <td></td> <td></td> </tr> <tr> <td>Quartz</td> <td>34</td> <td>10</td> <td>3</td> </tr> <tr> <td>Feldspar</td> <td>15</td> <td>2</td> <td>15</td> </tr> <tr> <td>Mica</td> <td>-</td> <td>-</td> <td>Tr</td> </tr> <tr> <td>Heavy minerals</td> <td>2</td> <td>1</td> <td>2</td> </tr> <tr> <td>Clay</td> <td>27</td> <td>75</td> <td>44</td> </tr> <tr> <td>Volcanic glass</td> <td>3</td> <td>3</td> <td>20</td> </tr> <tr> <td>Glauconite</td> <td>1</td> <td>1</td> <td>-</td> </tr> <tr> <td>Pyrite</td> <td>4</td> <td>2</td> <td>1</td> </tr> <tr> <td>Carbonate unspec.</td> <td>3</td> <td>-</td> <td>5</td> </tr> <tr> <td>Foraminifera</td> <td>2</td> <td>1</td> <td>-</td> </tr> <tr> <td>Calc. nannofossils</td> <td>3</td> <td>3</td> <td>2</td> </tr> <tr> <td>Diatoms</td> <td>-</td> <td>Tr</td> <td>3</td> </tr> <tr> <td>Radiolarians</td> <td>3</td> <td>1</td> <td>1</td> </tr> <tr> <td>Sponge spicules</td> <td>3</td> <td>1</td> <td>5</td> </tr> </tbody> </table>		2, 40	4, 33	4, 71		D	D	M	Texture:				Sand	30	5	50	Silt	40	17	30	Clay	30	78	20	Composition:				Quartz	34	10	3	Feldspar	15	2	15	Mica	-	-	Tr	Heavy minerals	2	1	2	Clay	27	75	44	Volcanic glass	3	3	20	Glauconite	1	1	-	Pyrite	4	2	1	Carbonate unspec.	3	-	5	Foraminifera	2	1	-	Calc. nannofossils	3	3	2	Diatoms	-	Tr	3	Radiolarians	3	1	1	Sponge spicules	3	1	5
	2, 40	4, 33	4, 71																																																																																									
	D	D	M																																																																																									
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Carbonate unspec.	3	-	5																																																																																									
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Calc. nannofossils	3	3	2																																																																																									
Diatoms	-	Tr	3																																																																																									
Radiolarians	3	1	1																																																																																									
Sponge spicules	3	1	5																																																																																									

SITE 568 HOLE CORE 44 CORED INTERVAL 408.9-417.7 m sub-bottom

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	STRUCTURE	DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS							
					0.5						<p>Dominant lithology: mudstone with abundant burrows, veination, and fracturing. Burrows are calcareous. Bedding plane is subvertical.</p> <p>Color: grayish olive green (5GY 3/2) with mottles of moderate olive brown (5Y 4/4).</p> <p>SMEAR SLIDE SUMMARY (%): 5, 50</p> <p>Texture: Sand: 15 Silt: 30 Clay: 55</p> <p>Composition: Quartz: 2 Feldspar: 8 Heavy minerals: Tr Clay: 69 Volcanic glass: 5 Pyrite: 2 Carbonate unsp. : 5 Calc. nannofossils: 1 Diatoms: 2 Radiolarians: 1 Sponge spicules: 5</p>
					1						
					1.0						
					2						
					3						
					4						
					5						
					CC						



1-1

1-2

1-3

1,CC

2-1

2-2

2-3

2-4

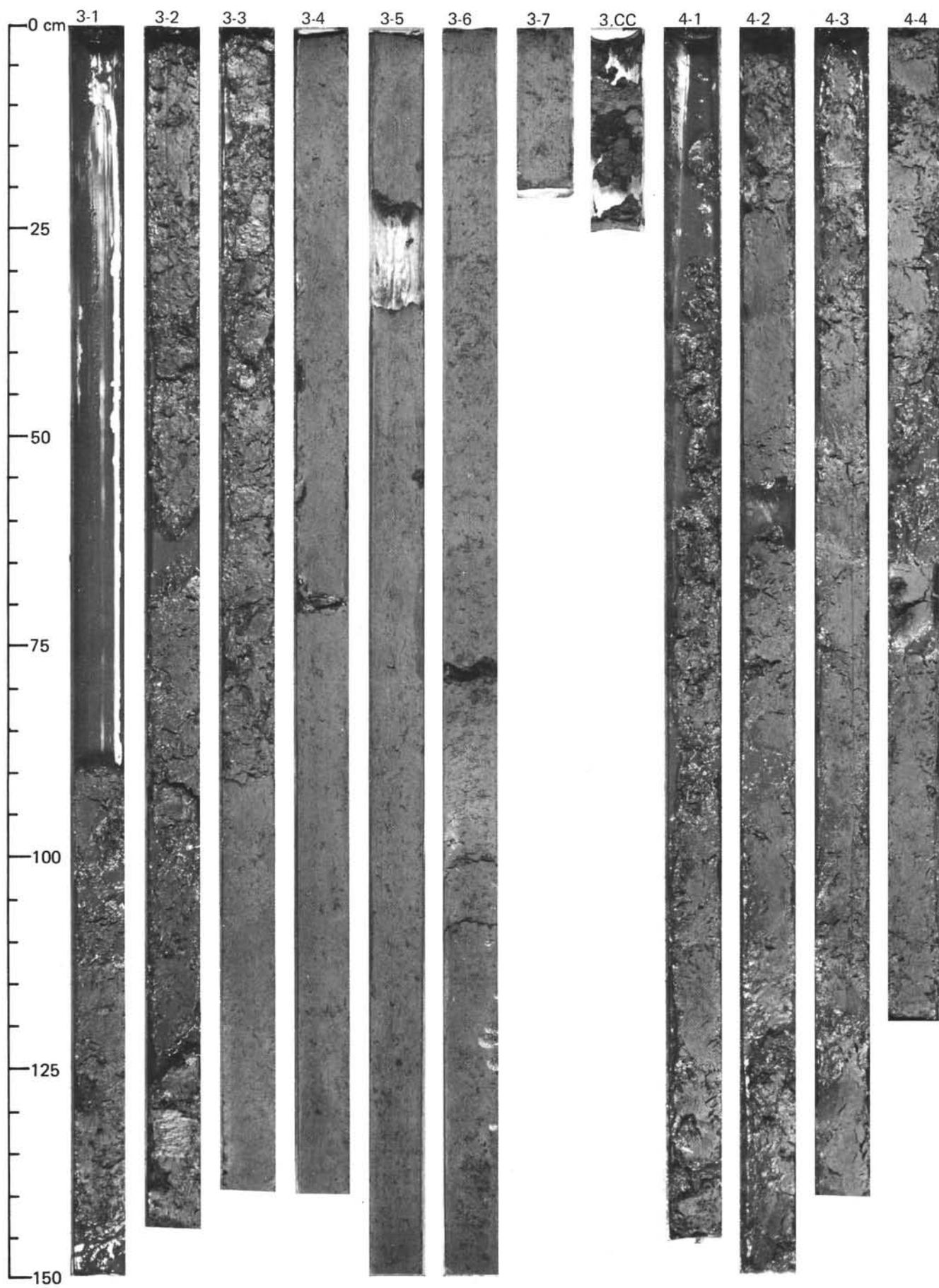
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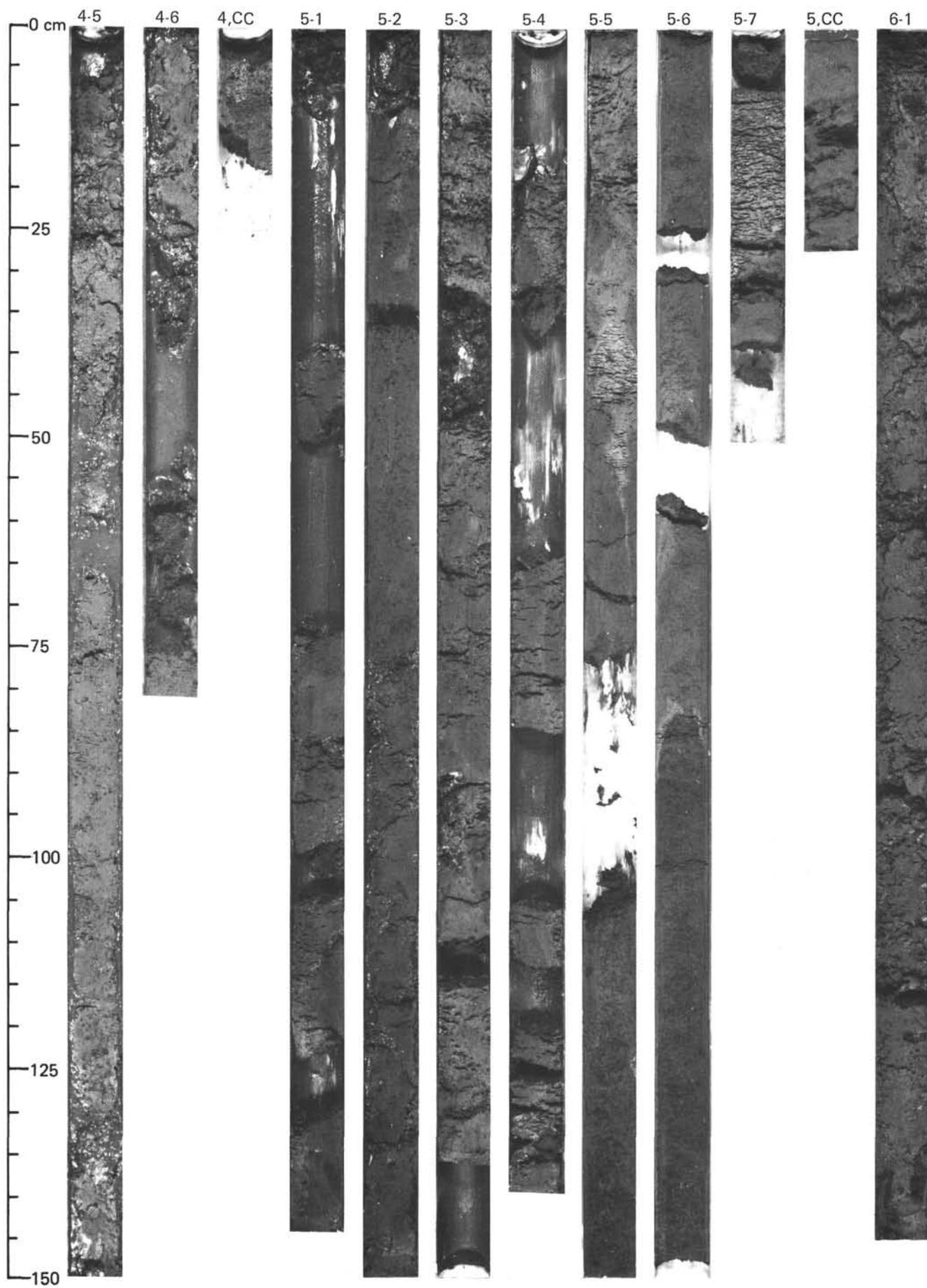
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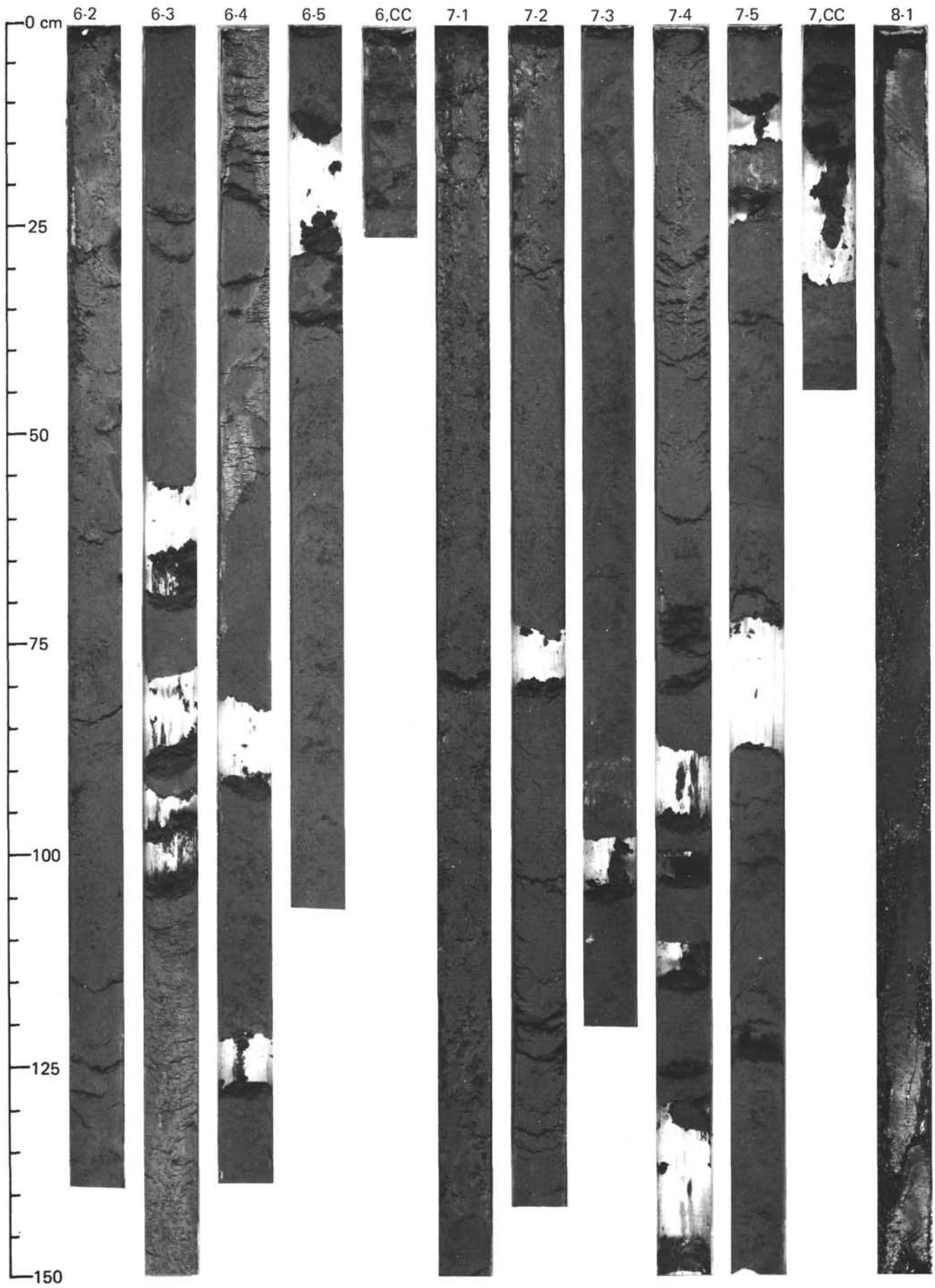
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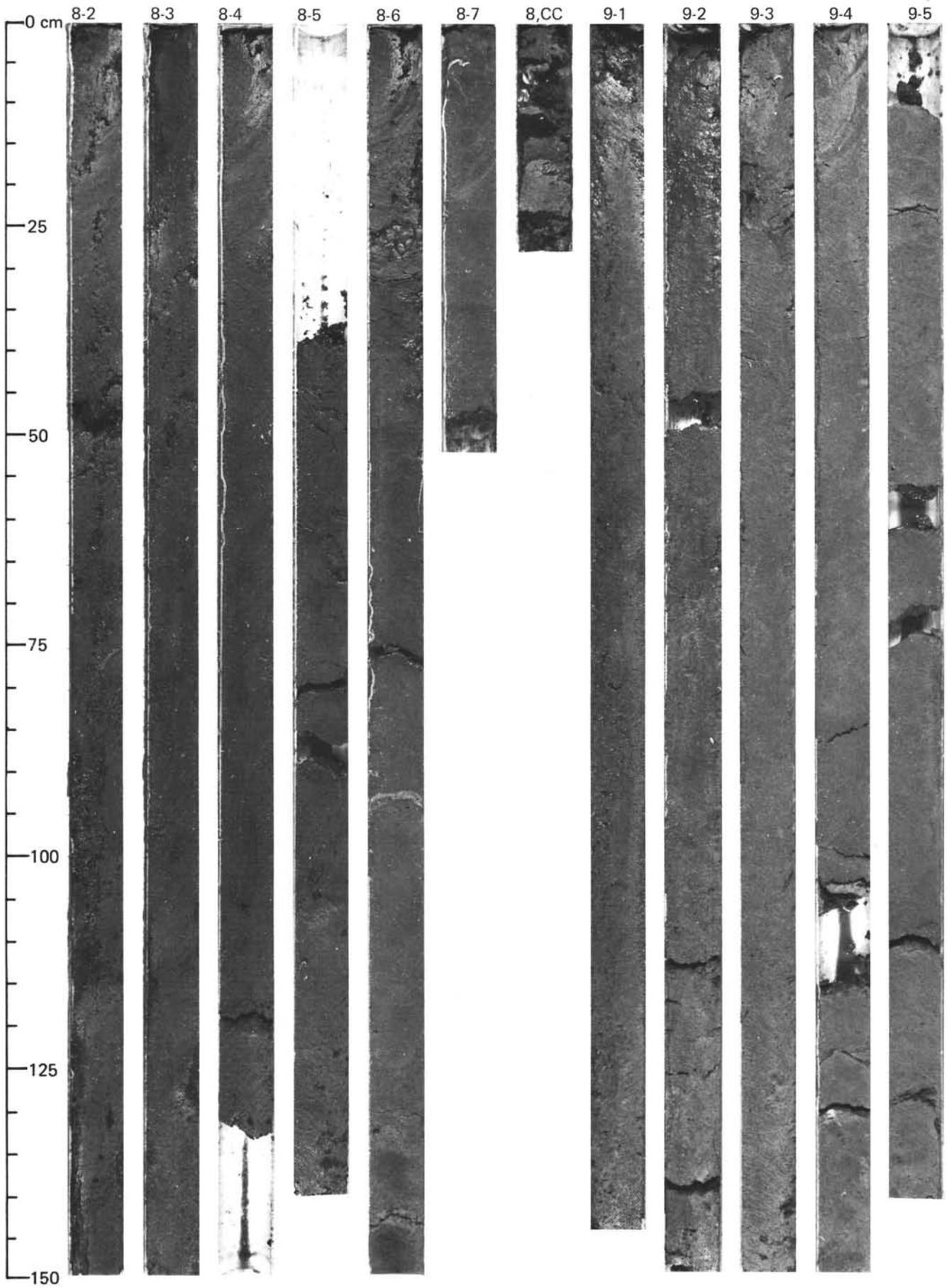
2,CC











8-2

8-3

8-4

8-5

8-6

8-7

8,CC

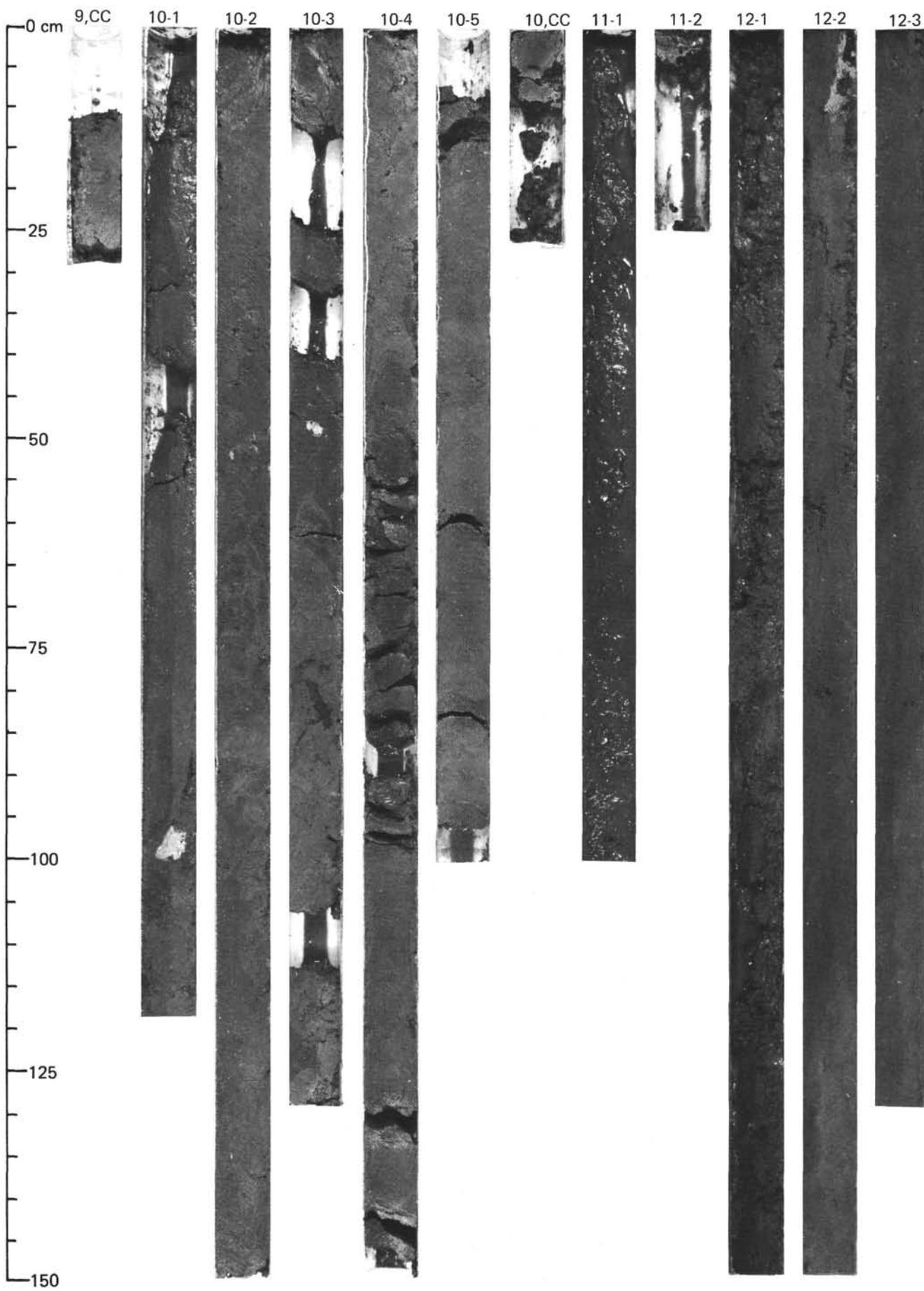
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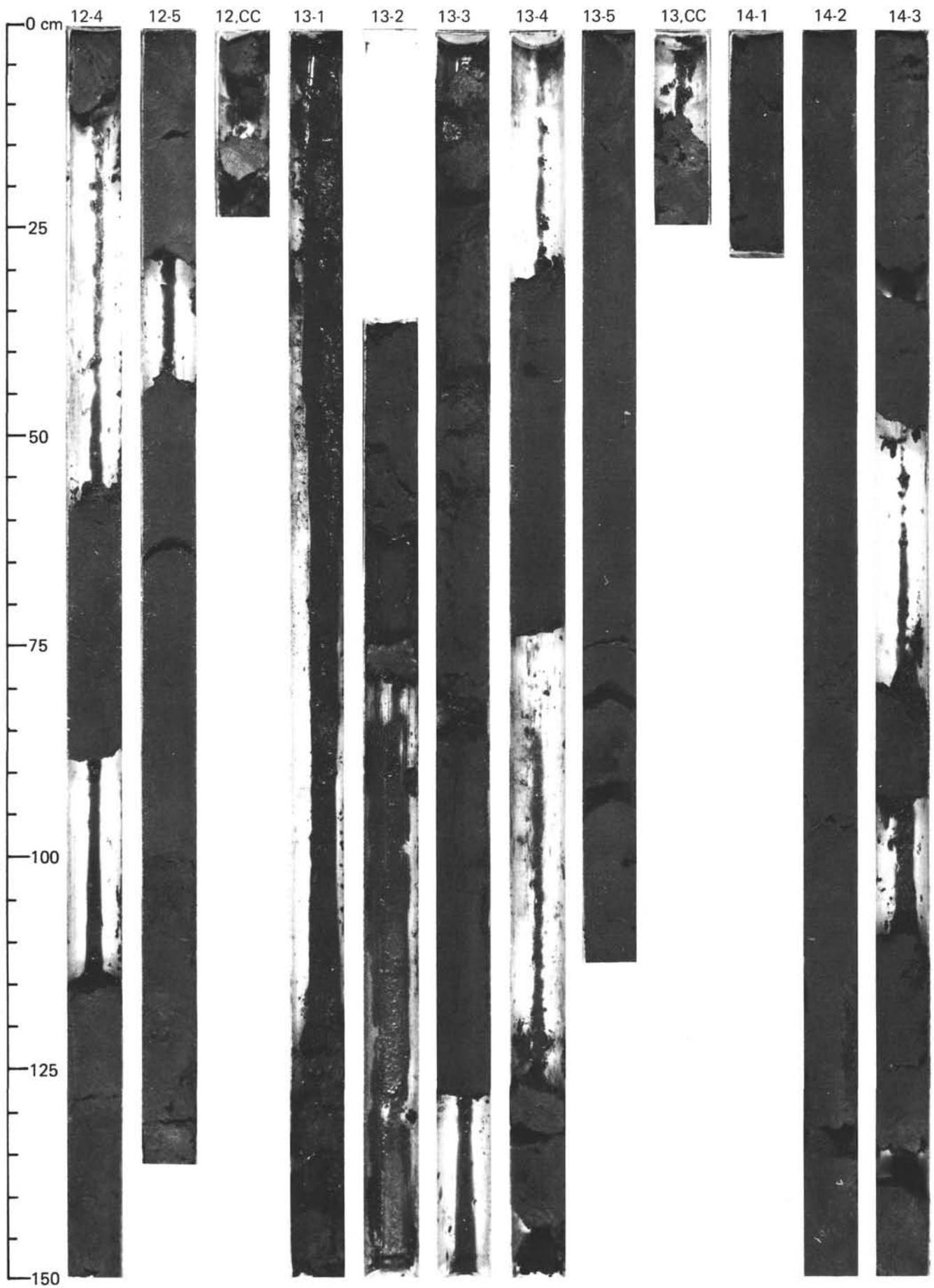
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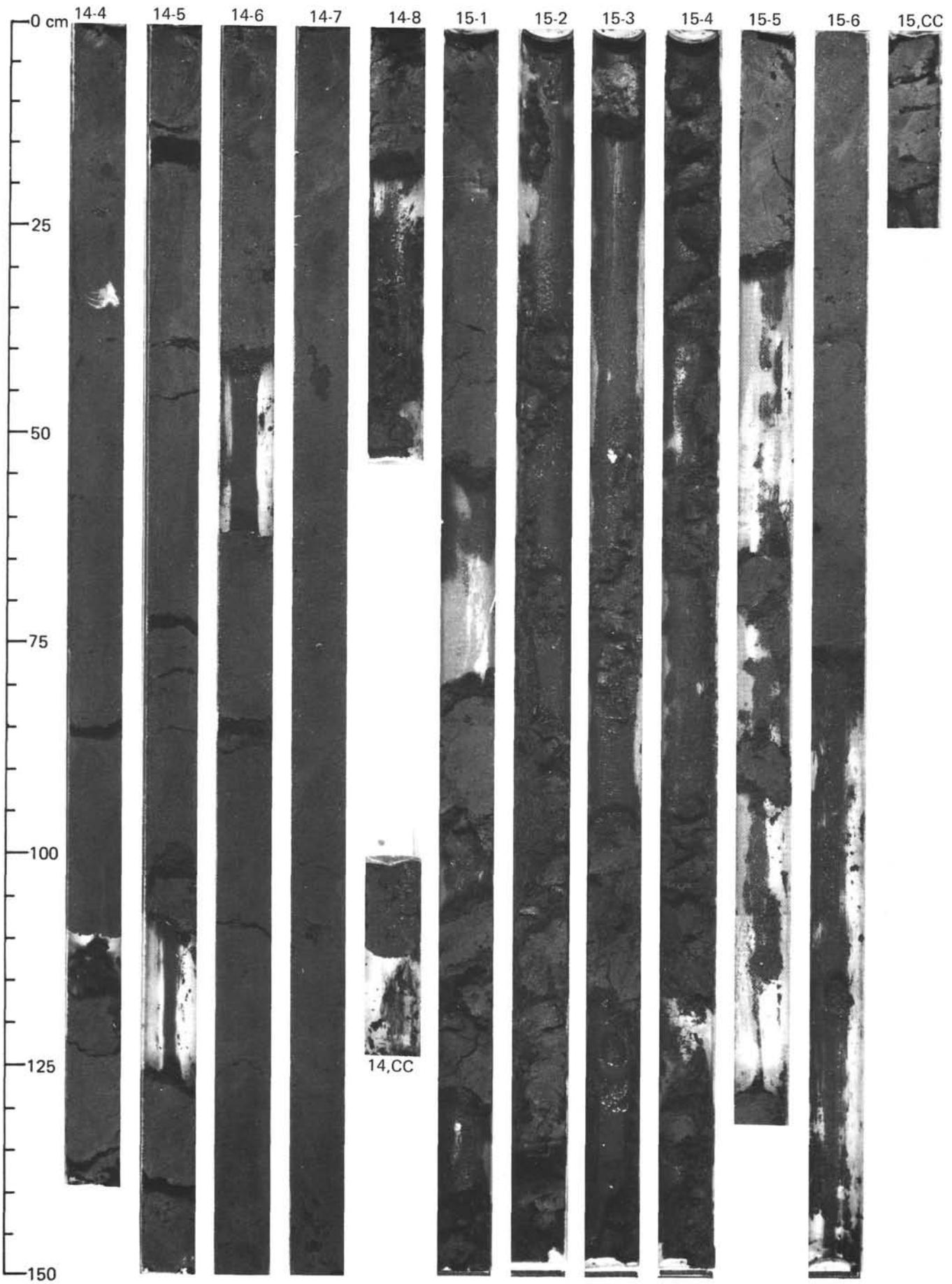
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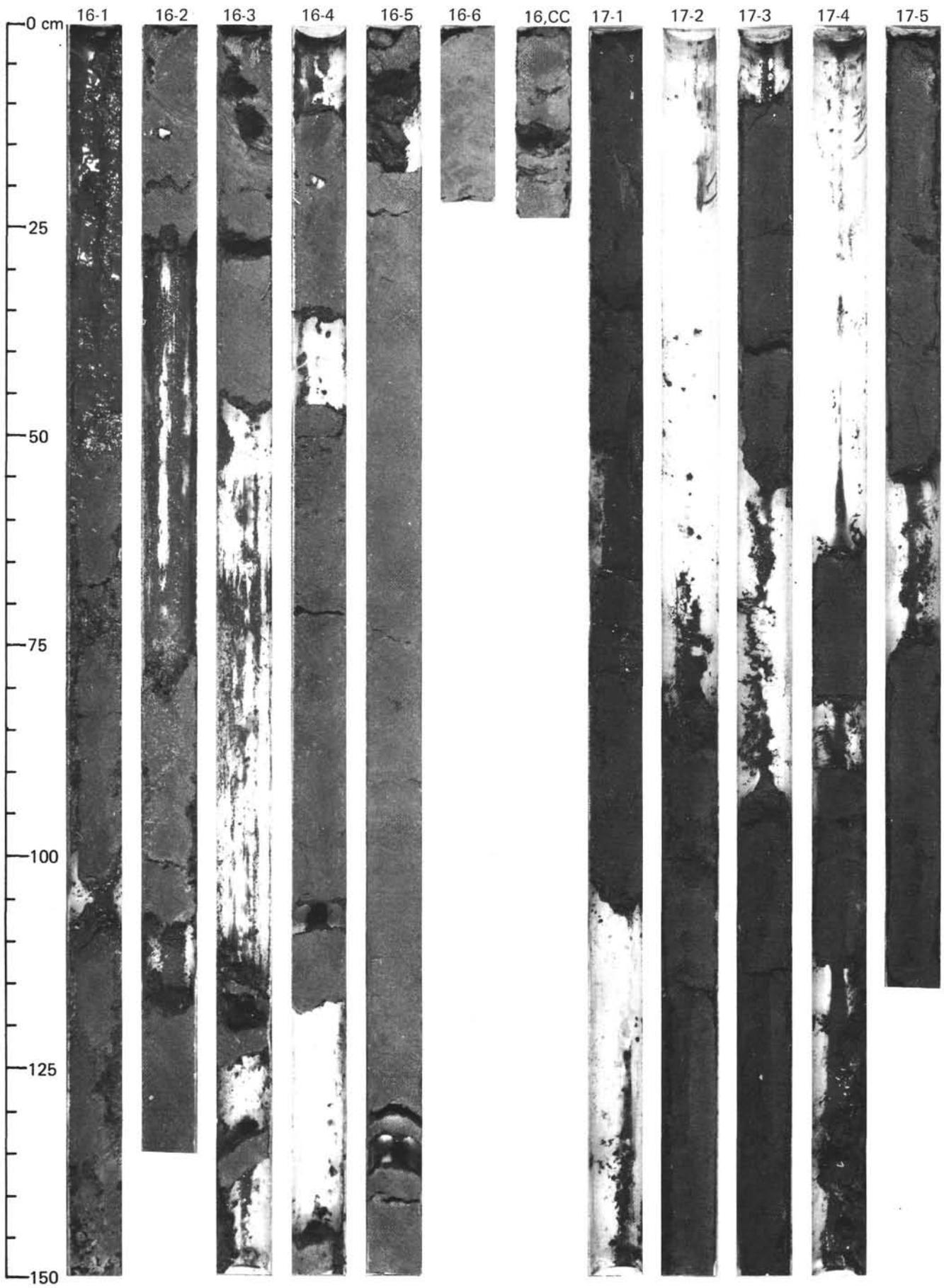
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9-5

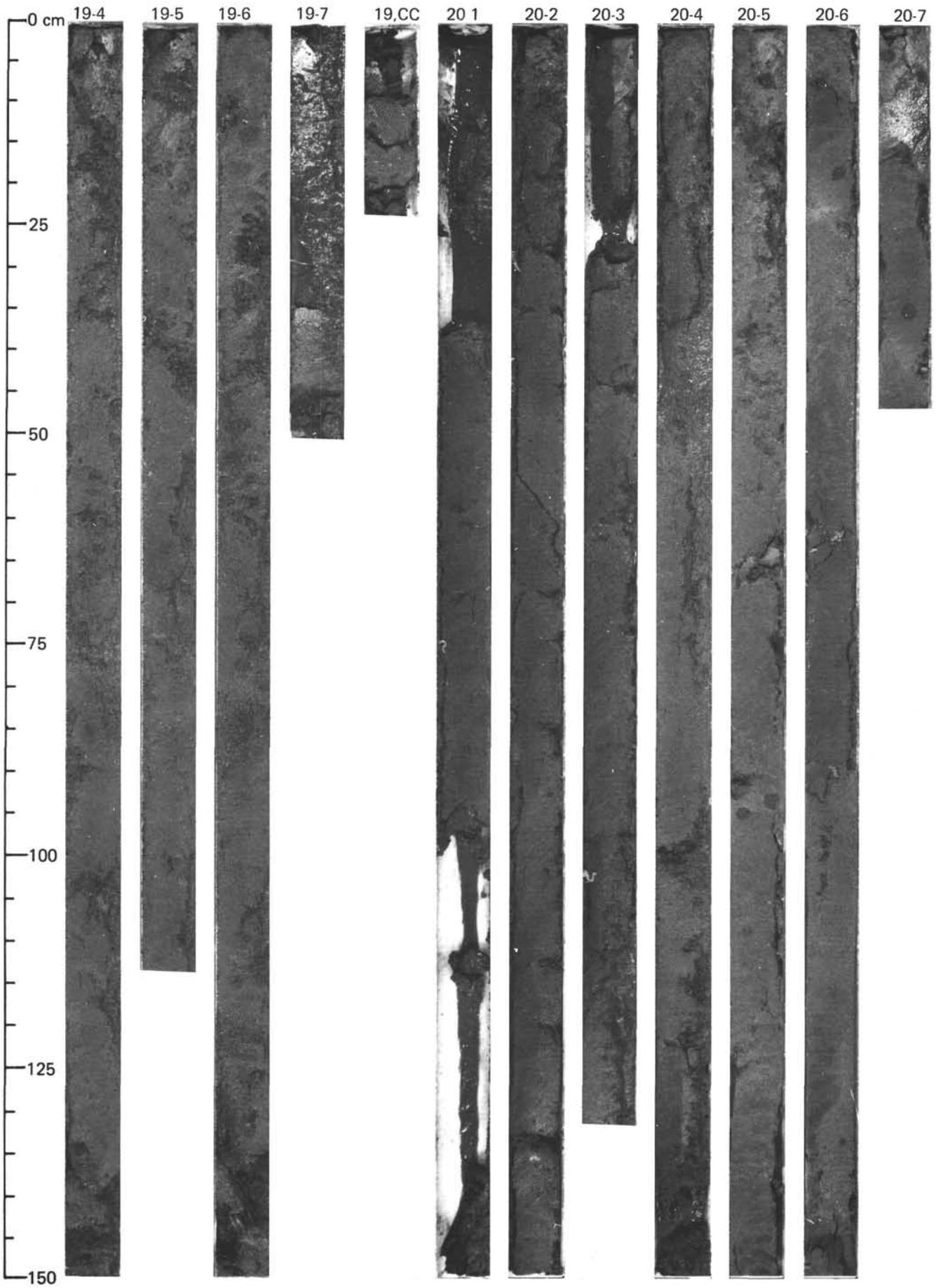












19-4

19-5

19-6

19-7

19,CC

20 1

20-2

20-3

20-4

20-5

20-6

20-7

