

6. SITE 523¹

Shipboard Scientific Party²

HOLE 523

Date occupied: 1230, 15 May 1980

Date departed: 1015, 20 May 1980

Time on site: 118 hr.

Position: 28°33.131'S; 2°15.078'W

Water depth (sea level; corrected m; echo-sounding): 4562

Water depth (rig floor; corrected m; echo-sounding): 4572

Bottom felt (m, drill pipe): 4572

Penetration (m): 193.5

Number of cores: 51

Total length of cored section (m): 182.5

Total core recovery (m): 149.2

Core recovery (%): 81.9

Oldest sediment cored:

Depth sub-bottom (m): 190.5

Nature: Nannofossil ooze

Age: middle Eocene

Principal results:

1. Obtained hydraulic piston cores from a pelagic sedimentary sequence Quaternary to middle Eocene in age and extended the calibration of the Tertiary biostratigraphic zones to Magnetostratigraphic Chron C-20-R.

2. Recognized a significant paleoceanographic event during the late middle Eocene, suggesting ocean cooling and/or that the formation of significant amounts of Antarctic ice might have started about 43 Ma (magnetostratigraphic age) or 39 Ma (radiometric age).

3. Confirmed an increase in abundance in the benthic foraminifer *Nuttalides umbonifera* when the site subsided to a given level, suggesting that species abundance might be related to the subsidence history at each site.

4. Found significant discrepancies between magnetostratigraphic and radiometric ages. With magnetostratigraphic ages based on

the assumption that the Cenozoic started 66.5 Ma in parentheses, the magnetostratigraphic and radiometric ages are respectively 41 (39) Ma and 37 Ma for the middle/upper Eocene boundary and 46 (43.57) Ma and 41 Ma for the top of NP15.

BACKGROUND AND OBJECTIVES

One of the objectives of Leg 73 drilling was to investigate the middle and late Eocene paleoceanography of the Mid-Atlantic Ridge in the Southeast Atlantic. As mentioned in the Site 522 chapter, drilling at two previous sites was done to sample upper Eocene and Oligocene sediments deposited on or near ridge crest. It was hoped that these sediments would lie above the foraminifer lysocline so they would provide materials that permitted the terminal Eocene crisis to be investigated. The primary objective of drilling at Site 523 was to sample early middle Eocene sediments to permit the investigation of the middle Eocene event. The terminal Eocene crisis was a sudden drop of ocean temperature of 4 to 5°C (Douglas and Savin, 1975; Shackleton and Kennett, 1975). The middle Eocene event, on the contrary, was a short episode when ocean temperatures were warmer (by 2–4°C) than before or immediately after (Shackleton and Kennett, 1975). We hoped to obtain samples of a continuous sequence of calcareous sediments of upper and middle Eocene age to investigate this paleoclimatic event.

Two drill sites, SA IV-6 and SA IV-7, were proposed to make sure that we would succeed in acquiring calcareous sediments that would permit us to study the crisis. The crust at SA IV-7, on Anomaly 21, was certainly older than both the middle and terminal Eocene events. The terminal Eocene sediments at SA IV-6, on Anomaly 20, may have been deposited while the site was at a ridge crest position. Dissolution at this site would be at a minimum for this event, but the ocean crust might be too young to give a sedimentary record of the earlier event. If ship time was adequate, the plans called for drilling at both sites. If, on the other hand, ship time was limited, we were to consider whether other objectives had higher priority and whether one of the two sites would fulfill the objective satisfactorily. During the cruise it became clear that only two more sites could be adequately investigated in the time remaining. For reasons given in the Site 524 chapter, we decided to drill only one of the two middle Eocene sites. Of the two we chose SA IV-7. First, the crust at this site was certain to be old enough to record the middle Eocene event. Second, previous drilling indicated that the calcite compensation depth (CCD) did not vary very much during the Eocene; thus, the degree of dissolution in sediments deposited on crust

¹ Hsü, K. J., LaBrecque, J. L., et al., *Init. Repts. DSDP*, 73: Washington (U.S. Govt. Printing Office).

² Ken J. Hsü (Co-Chief Scientist), Geologisches Institut, Eidgenössische Technische Hochschule, CH-8092 Zürich, Switzerland; John L. LaBrecque (Co-Chief Scientist), Lamont-Doherty Geological Observatory, Columbia University, Palisades, New York; Max F. Carman, Jr., Department of Geology, University of Houston, Houston, Texas; Andrew M. Gombos, Jr., Exxon Production Research Company, Houston, Texas; Anne-Marie Karpoff, Institut de Géologie, 67084 Strasbourg Cedex, France; Judith A. McKenzie, Geologisches Institut, Eidgenössische Technische Hochschule, CH-8092 Zürich, Switzerland; Stephen F. Percival, U.S. Geological Survey, Menlo Park, California; Nikolai P. Petersen, Institut für Geophysik, Universität München, D-8000 München, Federal Republic of Germany; Kenneth A. Pisciotta, Deep Sea Drilling Project, Scripps Institution of Oceanography, La Jolla, California (present address: Sohio Petroleum Company, 100 Pine Street, San Francisco, California); Richard Z. Poore, U.S. Geological Survey, Menlo Park, California (present address: U.S. Geological Survey, Reston, Virginia); Edward Schreiber, Department of Earth and Environmental Sciences, Queens College (CUNY), Flushing, New York; Lisa Tauxe, Department of Geological Sciences, Lamont-Doherty Geological Observatory, Columbia University, Palisades, New York (present address: Scripps Institution of Oceanography, La Jolla, California); Peter Tucker, Department of Geophysics, University of Edinburgh, Edinburgh, United Kingdom; Helmut J. Weissert, Department of Geological Sciences, University of Southern California, Los Angeles, California (present address: Eidgenössische Technische Hochschule, CH-8092 Zürich, Switzerland); and Ramil Wright, Department of Geology, Florida State University, Tallahassee, Florida (present address: Exxon Production Research Company, Houston, Texas).

differing in age by only 3 m.y. would not be much different. Third, as the older site, Site IV-7 would provide the longer record for the correlation of bio- and magnetostratigraphy and for the correlation of the latter with seafloor lineations. It was deemed especially important to acquire samples of late middle Eocene sediments to evaluate a part of the biostratigraphic-magnetostratigraphic correlation that was revealed by previous DSDP holes to be in error. Finally, drilling at Site IV-7 might provide evidence as to whether the seafloor spreading rate had deviated from linearity during the Eocene and Paleocene in the South Atlantic.

Another consideration in our planning was that Site IV-6 was very close to Site IV-7. If for some unforeseen reason we failed completely to achieve our primary objective at Site IV-7, we could steam back to Site IV-6 without losing a great deal of time.

The original location for Site IV-7 was chosen on the basis of continuous seismic profiling by the University of Texas Marine Science Institute (UTMSI). However, it was realized as early as the last meeting of the Ocean Paleoenvironment (OP) Panel that the sediment thickness there was minimal and might consist of only a thin veneer of Pliocene-Quaternary ooze. However, John Labrecque was able to find alternate sites on Anomalies 20 and 21 on an *Atlantis II* profile (Cruise 67, Leg 5) where thicker drape sediments were present. We designated these locations Sites IV-6 alternate and IV-7 alternate, and during the cruise we obtained permission through telegraphic messages from JOIDES to drill at the alternate sites. Site IV-7 alternate was finally drilled as Site 523.

The seismic profiling (Fig. 1) shows a local plateau apparently covered by a uniform blanket of sediments about 200 m thick. By the time we were ready to drill we

had begun to hope that we could drill the entire section above the basement with the hydraulic piston corer (HPC). Furthermore, we were not sure we could wash through such durable layers as the Miocene clay without attaching the dummy drill bit to the HPC, and this drill bit could not be attached after a remote rotary-to-HPC conversion. Therefore, we attached the HPC assembly to the drill string. We decided that if we met obstructions that prevented us from attaining (or nearly attaining) our goal, we would trip the drill string for rotary drilling and coring.

To recapitulate, our primary objective was to use the HPC to obtain a continuous sequence of upper and middle Eocene sediments. We wanted to sample sediments that had been deposited at a nearly uniform rate, were relatively free of contamination by resedimentation, and contained sufficient calcareous fossils to study the biostratigraphy, magnetostratigraphy, and paleoceanography of the mid-Eocene event. In addition, we wanted (1) to study the symmetry (or lack of it) in South Atlantic sedimentation, (2) to gather additional data concerning the terminal Eocene crisis and middle Miocene CCD crisis, (3) to sample diatom floras, if available, to provide data on the gradients of surface currents, (4) to date the seafloor lineations here (on Anomaly 21) and to clarify their temporal relation to biostratigraphic zonation, and (5) to obtain a clear and detailed history of South Atlantic seafloor spreading.

Because of poor weather and numerous difficulties with equipment, we did not satisfy all of our objectives completely. Nevertheless, the results were satisfactory. The shipboard staff therefore preferred to invest the remainder of ship time in drilling at the Cape Basin site, Site II-6, rather than in resampling at this location or in drilling at Site IV-6 alternate.

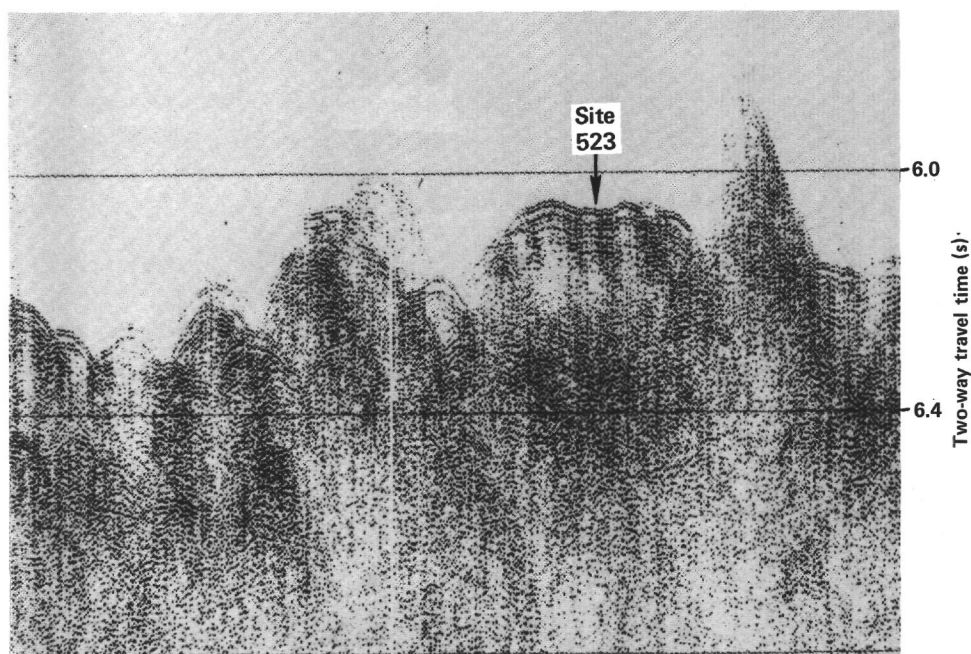


Figure 1. *Glomar Challenger* seismic profile over Site 523.

OPERATIONS

Site Approach

Site 523 was chosen from a profile of *Atlantis II* (Cruise 67, Leg 5, at 0600Z), which showed a particularly attractive pelagic drape over Magnetic Anomaly 21 basement. In order to locate the desired section, the *Glomar Challenger* steamed to a point 28°36'S and 2°21'W and took a course of 58° to follow the *Atlantis II* track (Fig. 2). The seismic equipment included 10- and 5-in.³ air guns. We lacked proper satellite fixes, which made navigation difficult, but with the excellence of navigation we had come to expect we passed directly over the profile. A beacon was dropped at 0959Z on 15 May 1980. A short survey was conducted after the beacon was dropped to make sure that relief in the area was minimal.

Coring Summary

The *Glomar Challenger* arrived on site and was in automatic positioning at 1230Z on 15 May. However, the weather, with northwesterly winds and a southwesterly swell, was unfavorable for positioning. The ship was able to hold position, but she rolled 6 to 10° at times. Toward evening, the wind shifted to the southwest, reducing pitch and roll. At 1854Z the drill crew started to make up the bottom hole assembly and prepared to run the pipes down. At 0615Z on 16 May, Hole 523 was spudded in. The HPC was started at a water depth of 4572 m (Table 1). The first core came up at 0645, but only traces of Quaternary ooze were attached to the core catcher. At 0800, the HPC bit stuck. When it was finally pulled free and recovered, it was found that the alumi-

num ring used to mark the orientation of the HPC had jammed in the seal sleeve. Hydraulic piston coring continued despite bad weather, but recovery was poor and the cores were disturbed. The recovery problem was resolved at 1830Z, when tool pusher Bill Lee discovered that the shear pins were malfunctioning and took remedial action. However, the core disturbance continued, apparently because the heavy swell (which was up to 12 ft. high) continued to come in from the southwest. These conditions persisted throughout the day of 17 May. Nevertheless, the vessel was able to hold position and to keep roll and pitch within safe operating limits.

At 0500Z on 17 May, the HPC bit encountered a hard layer at 56.4 m sub-bottom, presumably the middle Oligocene *Braarudosphaera* chalk that was expected at about this depth. The bit was able to drill through the layer, and the HPC could proceed after 3 m of interruption. Poor recovery continued; at 0940Z only a trace of ooze was recovered from Core 18 (which penetrated from 63.4 to 65.4 m sub-bottom). Some difficulty was experienced in washing from 63.4 to 65.4 sub-bottom, suggesting that a durable layer was present. Recovery in Core 19 was good, with a clump of *Braarudosphaera* chalk, mixed with pipe rust and other drilling debris, on top of Section 1. The chalk layer, which can probably be correlated with the third *Braarudosphaera* chalk bed at Site 522, was apparently the reason for the poor recovery in Core 18.

During the day the weather became calmer. Recovery percentages and core quality both improved with the weather. After Core 22 was taken the HPC samples be-

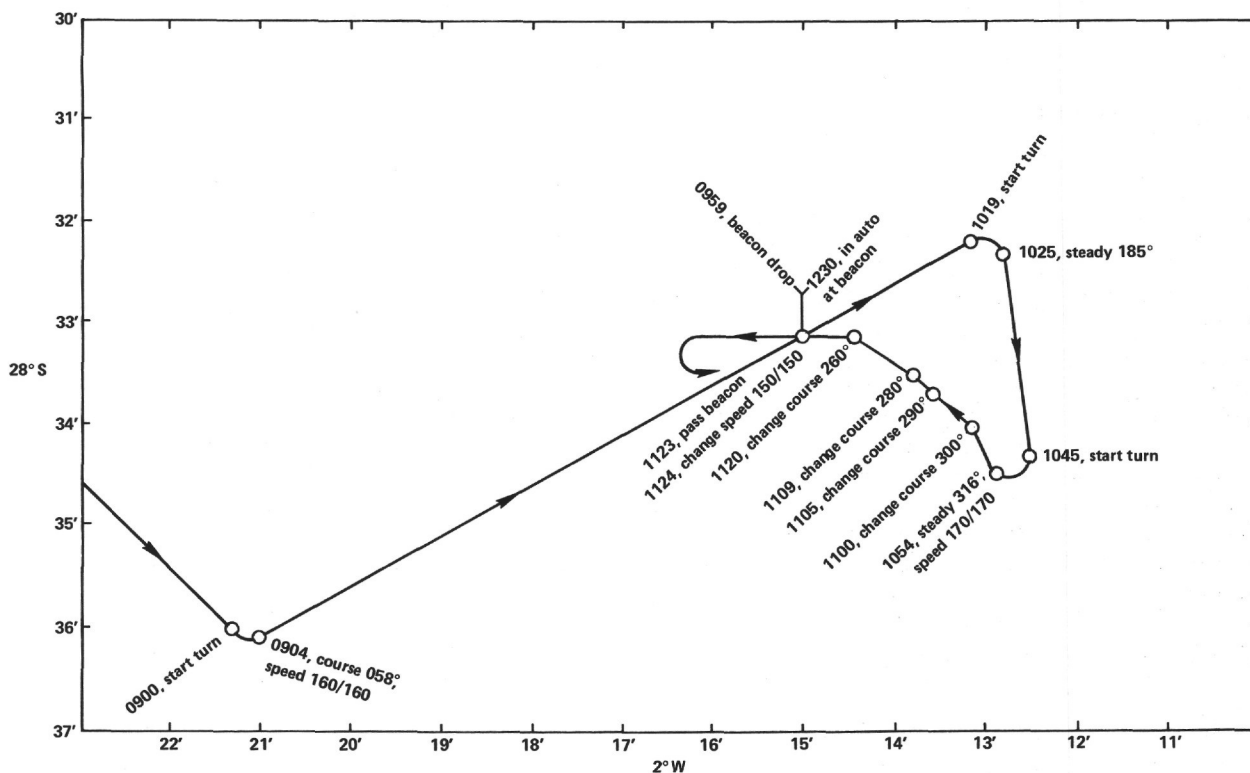


Figure 2. *Glomar Challenger* approach track, Site 523.

Table 1. Coring summary, Hole 523.

Core	Date (May 1980)	Time (hr.)	Depth from drill floor (m)	Depth below seafloor (m)	Length cored (m)	Length recovered (m)	Recovery (%)
1	16	0645	4572.0-4576.0	0-4.0	4.0	Tr	1
2	16	0844	4576.0-4580.4	4.0-8.4	4.4	2.3	52
3	16	1025	4580.4-4584.8	8.4-12.8	4.4	3.3	75
4	16	1152	4584.8-4589.2	12.8-17.2	4.4	4.1	93
5	16	1320	4589.2-4593.6	17.2-21.6	4.4	0.4	09
6	16	1440	4593.6-4598.0	21.6-26.0	4.4	4.1	93
7	16	1600	4598.0-4602.4	26.0-30.4	4.4	4.3	97
8	16	1720	4602.4-4606.0	30.4-34.0	3.6	3.7	102
9	16	1840	4606.0-4606.0	34.0-34.0	0.0	tr	—
10	16	2020	4606.0-4610.4	34.0-38.4	4.4	4.3	97
11	16	2140	4610.4-4613.9	38.4-41.9	3.5	3.4	97
12	16	2310	4613.9-4617.4	41.9-45.4	3.5	3.4	97
13	17	0035	4617.4-4621.4	45.4-49.4	4.0	4.6	115
14	17	0158	4621.4-4625.4	49.4-53.4	4.0	3.5	87
15	17	0325	4625.4-4626.4	53.4-54.4	1.0	tr	—
16	17	0500	4626.4-4628.4	54.5-56.4	2.0	2.1	105
—	17	—	4628.4-4631.4	56.4-49.4	Drill through hard layer		
17	17	0810	4631.4-4635.4	49.4-63.4	4.0	2.7	67
18	17	0940	4635.4-4637.4	63.4-65.4	2.0	tr	—
19	17	1115	4637.4-4741.4	65.4-69.4	4.0	3.5	87
20	17	1238	4741.4-4644.9	69.4-72.9	3.5	3.3	94
21	17	1400	4644.9-4649.3	72.9-77.3	4.4	4.1	93
22	17	1530	4649.3-4652.8	77.3-80.8	3.5	3.0	85
23	17	1730	4652.8-4656.8	80.8-84.8	4.0	3.7	92
24	17	1852	4656.8-4660.8	84.8-88.8	4.0	3.7	92
25	17	2030	4660.8-4664.8	88.8-92.8	4.0	4.1	100
26	17	2205	4664.8-4666.8	92.8-94.8	2.0	2.1	100
27	17	2350	4666.8-4670.8	94.8-98.8	4.0	3.0	75
28	18	0130	4670.8-4674.8	98.8-102.8	4.0	4.6	110
29	18	0310	4674.8-4678.8	102.8-106.8	4.0	tr	0
30	18	0440	4678.8-4682.8	106.8-110.8	4.0	3.1	77
31	18	0600	4682.8-4686.8	110.8-114.8	4.0	3.4	85
32	18	0730	4686.8-4689.8	114.8-117.8	3.0	2.5	83
33	18	0900	4689.8-4693.8	117.8-121.8	4.0	2.6	65
34	18	1030	4693.8-4696.8	121.8-124.8	3.0	1.9	63
35	18	1140	4696.8-4700.3	124.8-128.3	3.5	3.2	91
36	18	1320	4700.3-4702.8	128.3-130.8	2.5	1.6	64
37	18	1450	4702.8-4707.2	130.8-135.2	4.4	3.9	88
38	18	1802	4707.2-4711.2	135.2-139.2	4.0	3.2	85
39	18	1934	4711.2-4715.2	139.2-143.2	4.0	3.4	85
40	18	2104	4715.2-4718.7	143.2-146.7	3.5	2.8	85
41	18	2240	4718.7-4723.2	146.7-151.2	4.5	4.2	93
42	19	0055	4723.2-4727.2	151.2-155.2	4.0	4.4	110
43	19	0225	4727.2-4730.7	155.2-159.7	4.5	4.3	95
44	19	0350	4730.7-4735.2	159.7-162.2	3.5	3.3	94
45	19	0457	4735.2-4739.7	162.2-166.7	4.5	4.1	91
46	19	0623	4739.7-4743.2	166.7-170.2	3.5	2.1	60
47	19	0805	4743.2-4747.6	170.2-174.6	4.4	4.1	93
—	19	—	4747.6-4752.6	—	Wash—broken pipe		
48	19	1424	4752.6-4754.1	180.6-182.1	1.5	2.3	153
49	19	1536	4754.1-4758.5	182.1-186.5	4.4	4.3	97
50	19	1705	4758.5-4762.5	186.5-190.5	4.0	3.5	87
51	19	—	4762.5-4765.5	190.5-193.5	0	0	0
Total					182.5	149.2	81

Note: Dash denotes not applicable or unknown.

came suitable again for precision magnetostratigraphy. However, when Core 29 came up (at 0130Z on 18 May), the barrel was empty again, probably because the core catcher failed to prevent the cored sediment from falling out of the barrel. Throughout the drilling we had found crushed core liners, and several remedies were attempted without success. When Core 29 came up empty we changed to the flapper type of core catcher, which provided better sediment retention by minimizing the circulation of water through the cracked liners. Otherwise, hydraulic piston coring continued steadily, with good or average recovery, despite the presence of slightly lithified nodules in the oozes. At 1030Z on 19 May Core 48 was supposed to arrive on deck, but no sediments were recovered because the HPC barrel had broken near the top.

The staff met to make plans. We decided to try to drill down to the basement at this site before proceeding to drill the next site in Cape Basin. Meanwhile, an attempt to retrieve the broken barrel failed. However, the

tool sent down was smeared with ooze when it was brought up, an indication that the broken core barrel had fallen out of the drill collar. An attempt to wash 5 m down, past the broken barrel, was successful. At 1240Z, another HPC barrel was pumped down. The barrel was retrieved at 1424Z with good recovery. It was thought that the problem had been solved.

Two more cores came up, the last (Core 50) at 1705Z. Then we began to have further difficulties. First the HPC seemed to have malfunctioned. Then the sand line sent down to retrieve the core barrel came back with the end broken and the drill string had to be brought back to the rig floor. Meanwhile it was decided to wash ahead a little to "feel" the basement. However, after the string was raised a few meters, the downhole HPC barrel gave an indication of being properly seated. We had one last chance to sample the sediment above the basement. Since both seismic and magnetic data suggested that the basement was only 5 to 20 m down, we decided to wash ahead for 3 m before firing the last shot. After washing for 1.5 m the drill string seemed to encounter a hard object. It was probably the broken core barrel, because we were able to bypass the object and fire off the shear pins of the HPC (we thought). However, there was no indication that a full stroke had been achieved. At 2259Z on 19 May, having exhausted all possibilities, we terminated the drilling. The crew set back the Bowen subassembly to pull out of the hole. The drill string cleared the mudline at 2300Z, and the drill bit was on deck at 0850Z. The crew had difficulty retrieving the core barrel, which had been broken and was empty. A smear of ooze that was scraped off the core barrel and belonged to NP15 was the last core. At 1015Z on 20 May, after all the gear was secured on the rig floor, the vessel departed for Site 524.

LITHOLOGY

The dominant lithologies at Site 523 are nannofossil ooze, marly nannofossil ooze, and red clay. They range in age from middle Eocene to Quaternary. We divided the section into three lithologic units on the basis of calcium carbonate content, fossil composition, and color. The characteristics of these units are summarized in the composite sedimentary column shown in Figure 3.

Unit 1 (Cores 1-7, 0-30.2 m sub-bottom) consists of very pale brown, light yellowish brown, yellowish brown, and brownish yellow foraminifer-nannofossil ooze, nannofossil ooze, and marly nannofossil ooze. The calcium carbonate content of these pelagic sediments varies between 72 and 98%. Bioturbation features are recognized throughout the unit. Burrowing structures are most distinct in areas where alternating dark and light ooze provided sufficient color contrast to manifest the sediment mixing (e.g., Core 3).

Dark brown to dark reddish brown clay and reddish brown nannofossil clay characterize Unit 2 (Cores 8-20; 30.2-72.9 m sub-bottom). The shallowest occurrence of the dark brown clay in Core 8 defines the top of the unit. This clay, which has a carbonate content near 0%, forms the main lithology in Cores 8 through 12. The lower part of Unit 2 (Cores 13-20) consists of alternat-

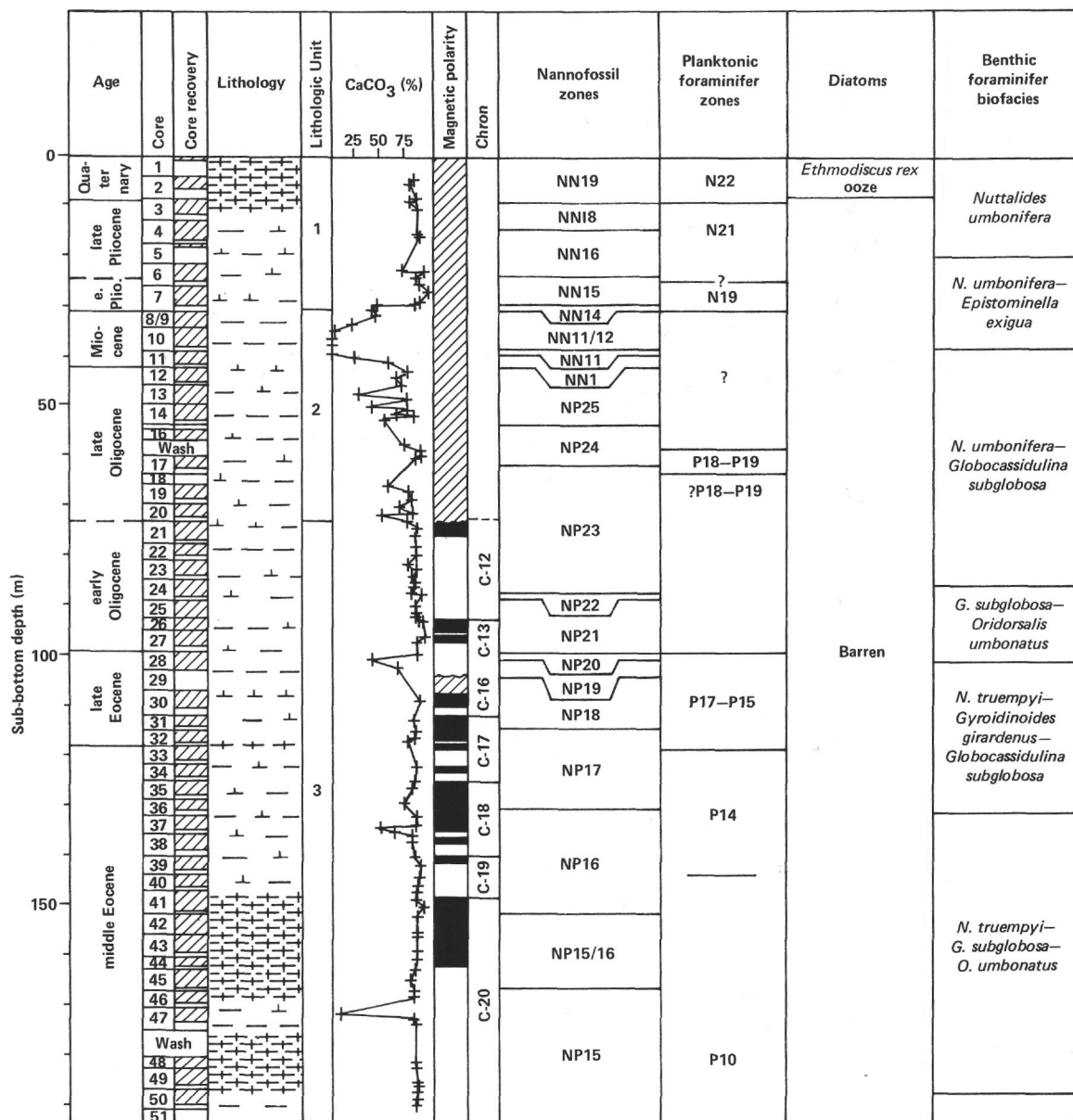


Figure 3. Stratigraphic summary, Site 523. Lithology and magnetic chronology are defined in Hsü, LaBrecque, et al. (this vol.). Solid blocks denote intervals of normal polarity; striped blocks, intervals of unknown polarity.

ing very pale brown and light yellowish brown marly nannofossil ooze and minor brown nannofossil clay. The calcium carbonate content varies between 60 and 85% for the marly nannofossil ooze and 20 and 40% for the nannofossil clay. The sediments of Unit 2 show faint burrowing; bioturbation is more visible in the areas where color changes occur. A few centimeters of white *Braarudosphaera* ooze were found in Core 19.

The main lithology of Unit 3 (Cores 21–50, 72.9–190.5 m sub-bottom) is a very pale brown and light yellowish brown nannofossil ooze and a yellowish brown marly nannofossil ooze. The calcium carbonate content varies between 76 and 96%. Several layers of brown and dark reddish brown clay up to 80 cm thick are intercalated with the nannofossil ooze sequence between 100 and 145 m sub-bottom (Cores 28, 32, 37, and 40). The

carbonate content of these layers ranges from 41 to 68%.

The percentage of foraminifers increases slightly in the lower part of Unit 3, where foraminifer-nannofossil ooze alternates with nannofossil ooze. Distinct layers of foraminifer ooze up to several centimeters thick occur in Cores 41 to 50. These layers often pass upward to nannofossil ooze that is virtually free of foraminifers. An interval rich in clay, pyroxenes, and feldspars in Core 47 is interpreted as an altered volcanoclastic layer.

BIOSTRATIGRAPHY

Summary

A middle Eocene to Quaternary sequence was cored by HPC at Site 523 (Figs. 3 and 4). The 50 cores yield

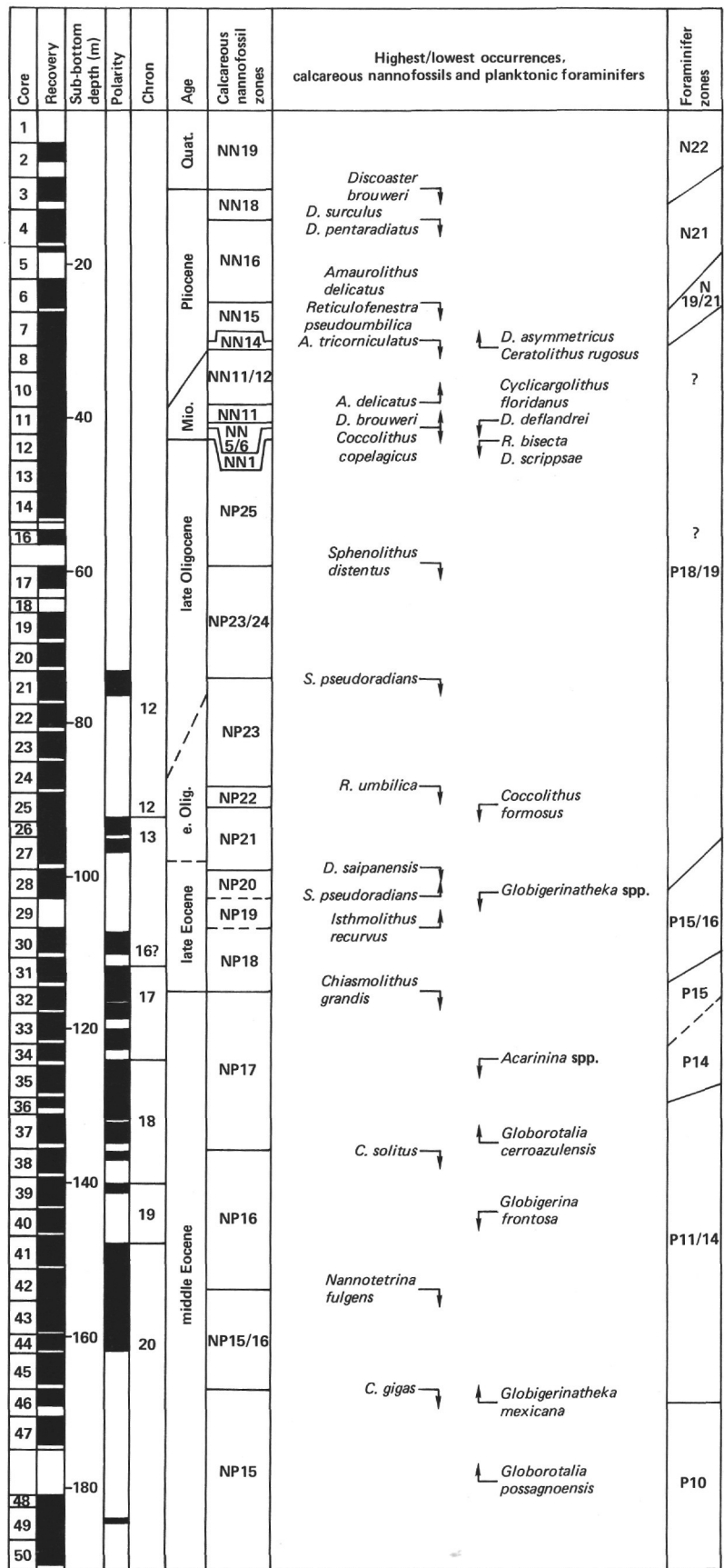


Figure 4. Biostratigraphic summary of significant calcareous and planktonic microfossils, Hole 523. Magnetic chronology as defined in Hsü, LaBrecque, et al. (this vol.). Polarity as in Fig. 3.

sediments that vary from easily zoned oozes with many well preserved microfossils to unzonable red clays with few microfossils.

The well preserved Quaternary section (Cores 1 and 2) is distinguished from the underlying well preserved upper Pliocene by the first appearance datum (FAD) of *Globorotalia truncatulinoides* in Sample 523-2,CC and the last appearance datum (LAD) of *Discoaster brouweri* in Core 3. Increasing dissolution and downhole contamination preclude planktonic foraminiferal recognition of the lower Pliocene. The top of the lower Pliocene is believed to occur in Core 6, where *Reticulofenestra pseudumbilica* is present but *Amaurolithus tricorniculatus* is not.

Biostratigraphic control is poor for the Miocene section at Site 523. Neither of the epoch's boundaries can be located with any certainty. The upper Miocene is recognized by the absence of *Ceratolithus* spp. and the presence of *Amaurolithus* spp. in Core 8. Precise zonation of the remainder of the Miocene is equivocal because of the dissolution of the planktonic foraminifers.

The first certain Oligocene sediment was encountered in Core 12, which was recognized as belonging to NP25 (late Oligocene). Although nannofossil zonation is clear down to the upper/lower Oligocene boundary (between Samples 523-21,CC and 523-23,CC), the foraminiferal fauna in this same sequence is subject to strong dissolution except in Sample 523-17,CC, which is zoned as P18/19. The planktonic foraminiferal fauna between Samples 523-23,CC and 523-26,CC is little better preserved, but it also appears to belong to Zones P18/19. The nannofossils in this sequence yield progressively older assemblages until Zone NP20 (late Eocene) is encountered. This zone, which is recognized by the presence of *D. saipanensis* and *D. barbadiensis*, occurs in Sample 523-28,CC. The first downhole occurrence of *Globigerina theka*, a solution-resistant genus, in the foraminiferal assemblage in Sample 523-28,CC provides added confirmation of the late Eocene data.

The late/middle Eocene boundary is tentatively placed between Samples 523-31,CC and 523-32,CC on the basis of the occurrence of *Chiasmolithus grandis* at the top of Core 32 and the occurrence of *G. semiinvoluta* in Core 31. Strong dissolution has reduced the diversity of the planktonic foraminifers in this interval.

Although the hole terminated in middle Eocene sediments not far above basement, the presence of reworked lower Eocene planktonic and benthic foraminifers suggests the presence of a lower Eocene section just above the basement.

Calcareous Nannoplankton

Samples 523-1,CC to 523-3-1, 102–103 cm are Quaternary in age and zoned as NN19 because of the presence of *Pseudoemiliania lacunosa* Gartner.

Samples 523-3-2, 25–26 cm to 523-3,CC are late Pliocene in age and zoned as NN18 because of the occurrence of *Discoaster brouweri* Tan Sin Hok. The interval from Sample 523-4-2, 100–101 cm to 523-6-2, 71–72 cm is assigned to the late Pliocene and Zone NN16 because

of the presence of *D. surculus* Martini and Bramlette. The occurrence of *Reticulofenestra pseudumbilica* Gartner in the interval from Sample 523-6-3, 71–72 cm to 523-7-3, 39–40 cm means it belongs to the early Pliocene and Zone NN15. Sample 523-7,CC belongs to the early Pliocene and Zone NN14 because of the co-occurrence of *D. asymmetricus* Gartner and *Amaurolithus tricorniculatus* (Gartner).

The interval from Sample 523-8-1, 51–52 cm to Sample 523-10-3, 50–51 cm is assigned to the late Miocene and Zones NN11/12 because of the presence of *Amaurolithus* spp. and the absence of *Ceratolithus* spp. Samples 523-10,CC to 523-11-1, 145–146 cm are late Miocene in age and belong to Zone NN11 according to the occurrence of *D. quinqueramus* Gartner. Sediments that belong to the early to middle Miocene and Zones NN5/6 occur in Sample 523-11-2, 60–61 cm, as indicated by the co-occurrence of *Cyclicargolithus floridanus* (Hay and Roth), *D. brouweri* Tan Sin Hok, and *D. deflandrei* Bramlette and Riedel. Early Miocene NN1 sediments are found from Sample 523-11-3, 5–6 cm to 523-11,CC according to the occurrence of *Coccolithus eopelagicus* Bramlette and Riedel.

Sediments from Sample 523-12-1, 120–121 cm to 523-14,CC are assigned to the late Oligocene and Zone NP25 because of the presence of *R. bisecta* (Hay, Mohler, and Wade) and *Dictyococites scrippsae* Bukry and Percival. The core catcher samples from Cores 15 to 16 are placed in the late Oligocene and Zone NP24 because of the presence of *Sphenolithus predistentus* Bramlette and Wilcoxon and *S. distentus* Bramlette and Wilcoxon. Samples 523-17,CC to 523-24-2, 94–95 cm belong to Zone NP23 because of the presence of *S. pseudoradians* Bramlette and Wilcoxon. The *Braarudosphaera* ooze occurs in Sample 523-19-1, 52–53 cm. Early Oligocene sediments belonging to Zone NP22 occur from Sample 523-24-3, 52–53 cm to 523-25-1, 82–83 cm because of the presence of *R. umbilica* (Levin). *Cyclococcolithina formosa* (Kamptner) marks the early Oligocene and Zone NP21 for the interval from Sample 523-25-2, 82–83 cm to 523-27,CC.

The presence of *D. saipanensis* Bramlette and Riedel indicates that late Eocene NP20 sediment is found from Sample 523-28-1, 72–73 cm to 523-28,CC. The core catcher sample from Core 29 belongs to Zone NP19 according to the absence of *S. pseudoradians* Bramlette and Wilcoxon. The interval from Sample 523-30-1, 120–121 cm to 523-31,CC is assigned to the late Eocene and Zone NP18 because of the absence of *Isthmolithus recurvus* Deflandre. Samples 523-32-1, 57–58 cm to 523-37,CC are assigned to the middle Eocene and Zone NP17 because of the presence of *Chiasmolithus grandis* (Bramlette and Riedel). Samples 523-38-1, 89–90 cm to 523-46-1, 60–61 cm are middle Eocene and belong to Zone NP16 because of the occurrence of *C. solitus* (Bramlette and Sullivan). Samples 523-42-3, 54–55 cm to 523-45,CC are undifferentiated Zones NP15/16 because of the occurrence of *Nannotetrina fulgens* (Stradner). *C. gigas* (Bramlette and Sullivan), a Zone NP15 marker species, is found from Sample 523-46-1, 70–71 cm to Sample 523-50,CC.

Planktonic Foraminifers

Quaternary through middle Eocene planktonic foraminifers were recovered in Hole 523 (Fig. 3). In general, the middle Eocene, late Pliocene, and Quaternary assemblages are relatively diverse and moderately well preserved, whereas the late Eocene, Oligocene, Miocene, and early Pliocene assemblages are sparse and poorly preserved because of intense dissolution.

The Quaternary assemblages in Samples 523-1, CC and 523-2, CC contain abundant specimens of *Globorotalia crassaformis*, *G. inflata*, *G. truncatulinoides*, and *Globigerina bulloides*. Samples 523-3, CC through 523-6, CC yield similar assemblages but lack *Globorotalia truncatulinoides*. Samples 523-3, CC through 523-6, CC are thus assigned to the Pliocene, and the Pliocene/Quaternary boundary is between Samples 523-3, CC and 523-2, CC.

Sample 523-7, CC contains sparse Pliocene microfossils, but the core is highly disturbed and the foraminifers may be present because of downhole cavings. Samples 523-8, CC through 523-10, CC contain fragments of planktonic foraminifers and a few moderately well preserved Pliocene and Quaternary taxa (such as *Globorotalia tosaensis* and *G. inflata*). Intact specimens of planktonic foraminifers in Samples 523-8, CC through 523-10, CC are clearly downhole contaminants.

Strongly dissolved assemblages continue in Samples 523-11, CC through 523-16, CC. Only a few specimens of such long ranging early Miocene to Eocene taxa as *Catapsydrax dissimilis* and *Globorotaloides suteri* were found in this interval. The calcareous nannofossils suggest that the Oligocene/Miocene boundary occurs between Samples 523-11, CC and 523-12, CC.

A well preserved Oligocene (Zone P18/P19) assemblage occurs in Sample 523-17, CC, but strongly dissolved assemblages return again down to Sample 523-23, CC. Pliocene and Quaternary contaminants occur sporadically throughout this interval. Samples 523-24, CC through 523-26, CC contain meager assemblages that are referred to Zones P18/P19 because of the associated Oligocene nannofossils. Sample 523-27, CC is barren, and *Globigerinatheka* spp. occur in Sample 523-28, CC. Calcareous nannofossils place the Eocene/Oligocene boundary between Samples 523-27, CC and 523-28, CC.

Samples 523-28, CC through 523-32, CC consist almost exclusively of specimens and fragments of *Globigerinatheka*. These low diversity assemblages are the product of intense dissolution. The preservation of the foraminifers improves in Samples 523-33, CC to 523-35, CC, and representatives of *Acarinina* and *Subbotina* begin to appear in the fauna. The middle/late Eocene boundary is placed between Sample 523-31, CC and 523-32, CC, but the position of this boundary may be primarily a function of dissolution intensity. Samples 523-36, CC and 523-37, CC are almost barren of planktonic foraminifers.

Planktonic foraminifers are common and moderately well preserved in Samples 523-38, CC through 523-50, CC. *A. bullbrooki*, *Globigerina frontosa*, *Globigeri-*

natheka senni, other *Globigerinatheka* spp., and *Pseudohastigerina micra* are common to abundant in most samples. Specimens of *Globigerinatheka* are common to abundant in Samples 523-38, CC through 523-45, CC, whereas they occur sparsely in Samples 523-46, CC through 523-50, CC.

Besides being made more difficult by repeated instances of downhole contamination, the interpretation of the middle Eocene fauna in Hole 523 is complicated by reworking. For example, the early Eocene taxa *Morozovella marginodentata* and *M. subbotinae* occur in Sample 523-41, CC. In addition, several key zonal markers, such as *Orbulinoides beckmanni* and *M. aragonensis*, are absent or poorly represented, although they are described as being present in Site 19 on the west side of the ridge.

Diatoms

At Site 523, *Ethmodiscus rex* ooze is abundant in Sample 523-1, CC and common in Sample 523-2, CC. In both of these cores, clay particles are a major constituent of the sediment. All other core samples are barren of diatoms.

Benthic Foraminifers

The benthic foraminifers in Hole 523 are common, reasonably diverse, and generally well preserved. The predominant faunal elements of the upper Miocene to Quaternary sequence are *Nuttalides umbonifera*, *Globocassidulina subglobosa*, *Oridorsalis umbonatus*, and *Epistominella exigua*. In the upper part of the sequence *N. umbonifera* comprises from 40 to 65% of the assemblage. In the lower part of the sequence, this species comprises less than 20% of the fauna; it is a significant contributor to the fauna but does not dominate, as it does in the upper sequence. The proportional contribution of *E. exigua* increases significantly in this lower sequence. The transition between the upper and lower assemblages occurs at 23 m sub-bottom (Section 523-6-2). It can be inferred, by interpolation from the sedimentation rate curve (Fig. 5), that this transition occurred 3.3 Ma. By analogy with recent abyssal foraminiferal distributions (Schnitker, 1980), this mid-Pliocene transition would seem to mark an influx of Antarctic Bottom Water (AABW) into the eastern South Atlantic. It was at approximately this time that the AABW was being generated by increased glacial activity in Antarctica (Watkins and Kennett, 1972). Lying just above this interval in Hole 523 is alternating light and dark calcareous ooze that may be the product of glacially induced fluctuations in dissolution. Evidence of glaciation also occurs in the North Atlantic at this time (Berggren, 1972).

Most of the Miocene section in Hole 523 is greatly condensed as a result of dissolution. The only well preserved benthic foraminifer fauna occurs in the uppermost Miocene, as described above, and the lowest Miocene. The early Miocene fauna, as well as the fauna of the late Miocene and part of the early Oligocene, was dominated by *N. umbonifera* and *G. subglobosa*, with subordinate but significant amounts of *O. umbonatus*

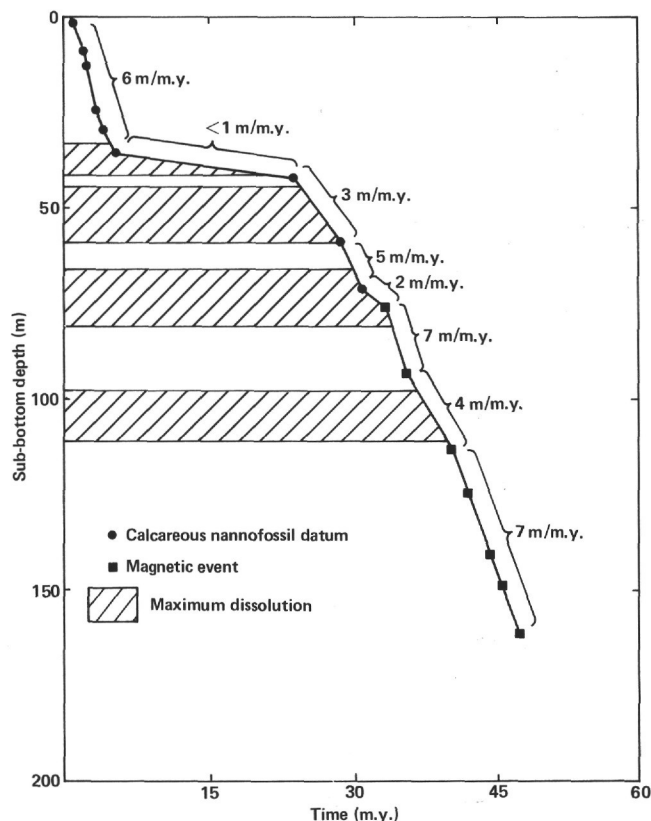


Figure 5. Sediment rates (not corrected for compaction), Hole 523.

and *Gyroidinoides girardenus*. This fauna persists down-hole to the interval between Cores 22 and 27 (79–97 m sub-bottom), where the proportion of *N. umbonifera* declines from an average value of 15% to a value of 3%. The remainder of the early Oligocene fauna that occurs below this interval is characterized by *O. umbonatus* and *Globocassidulina subglobosa*, with subordinate amounts of *Nonion havaneuse*, *Gyroidinoides girardenus*, and *Astrononion pusillum*. By interpolation from the sedimentation rate curve (Fig. 5), this downhole decline in *Nuttalides umbonifera* occurs between 34.6 and 37.1 Ma.

It was during this part of the latest Eocene to earliest Oligocene that Antarctic ice began to build up and the Antarctic Bottom Water began to form (Kennett et al., 1972). If the abundance of *N. umbonifera* was as indicative of the presence of AABW in the Oligocene as it is today, the early Oligocene increase in specimens of this species in Hole 523 may be a consequence of bottom water formation at this time. Site 523 lay at a depth of 3450 to 3550 m at this time (Fig. 6).

The upper Eocene and part of the middle Eocene sequence in Hole 523 (down to Core 37) is characterized by *N. truempyi*, with accessory amounts of *O. umbonatus*, *G. girardenus*, and *Globocassidulina subglobosa*. A significant shift in the subordinate fauna occurs in Core 37 (134 m sub-bottom; interpolated age of 44.0 m.y.), where *Gyroidinoides girardenus* declines and the dominant assemblage is *N. truempyi*–*Globocassidulina subglobosa*–*O. umbonatus*.

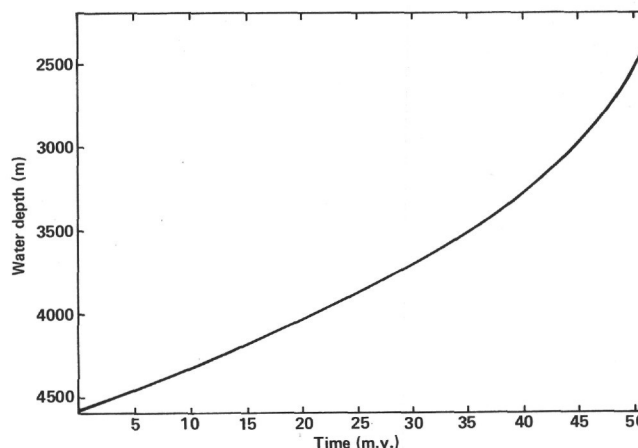


Figure 6. Bathymetry of Site 523 based on subsidence curves which are corrected for sediment load (Berger and von Rad, 1972).

Dissolution

At least three parameters provide indices of dissolution at Site 523: (1) the ratio of benthic to planktonic foraminifers (the proportion of benthic foraminifer tests); (2) the ratio of fragments to whole tests in the planktonic foraminifers; and (3) calcium carbonate content. Figure 7 illustrates the first two parameters. The curves behave sympathetically, with the percentage of fragments a more sensitive measure of dissolution during episodes of low to moderate dissolution and the percentage of benthic foraminifers a more sensitive indicator during episodes of moderate to intense dissolution. The fluctuations in calcium carbonate content are illustrated in Figure 3. These indices reveal that Site 523 lay below the CCD during the earliest Pliocene and late Miocene and at or near the CCD twice in the late Oligocene and during the latest Eocene.

The Quaternary and upper Pliocene section is characterized by fluctuating levels of dissolution (e.g., 10–15 m sub-bottom) as a consequence of glacial epochs. The sediments in this interval are alternately light and dark as a result of these dissolution fluctuations. The samples in this sequence are spaced too far apart to permit the distinction of individual epochs.

The intervals in intense dissolution in the middle and upper Miocene are merged in Hole 523 in a condensed interval characterized by very low carbonate content (Fig. 3) and almost no foraminifers. There are two intervals of strong dissolution in the Oligocene section. The uppermost occurs between 42 and 58 m sub-bottom in the uppermost Oligocene section (Nannofossil Zone NP25). The lower interval lies between 65 and 80 m sub-bottom, just above the upper/lower Oligocene boundary (in Nannofossil Zone NP23/24, just above(?) Magnetic Chron C-12).

A broad dissolution peak lies in the upper Eocene between 97 and 110 m sub-bottom. Below this peak the dissolution indices decline gradually, and the preservation of calcareous microfossils is quite good below 140 m sub-bottom (in the middle Eocene sequence). There is a sharp spike at 133 to 134 m sub-bottom. It is not clear if

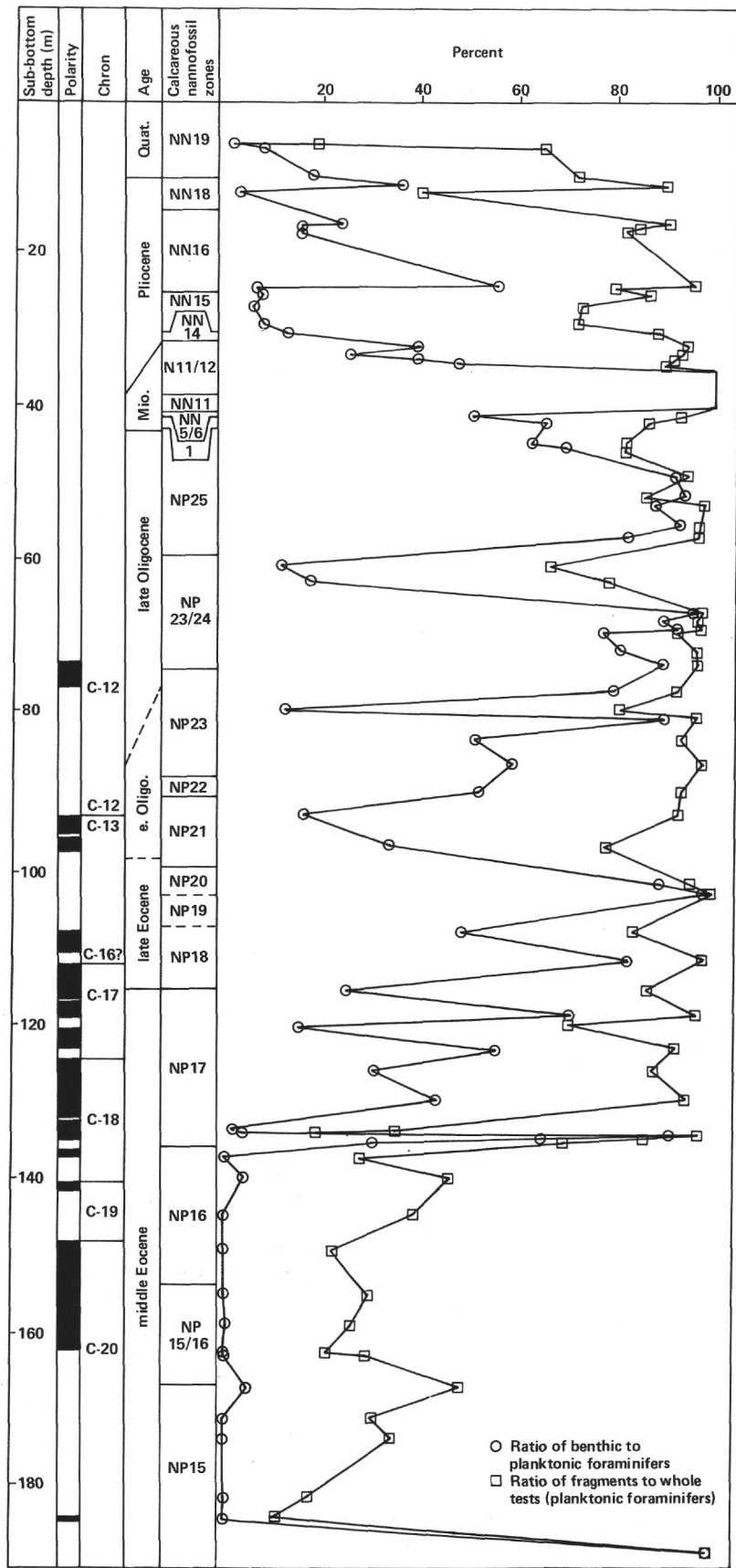


Figure 7. Dissolution indices, Hole 523. See text for discussion. Polarity as in Fig. 3.

this very marked and short-lived change in dissolution indices is a product of middle Eocene ocean chemistry or post-depositional diagenesis.

There is a curious clay layer in the middle Eocene part of Hole 523. It occurs at 134 m sub-bottom (interpolated age of 44.0 m.y.) and exhibits low carbonate content (47%) and a high dissolution index from the foraminiferal data. The curious feature of this interval is the remarkably low dissolution that occurs immediately above and below the clay layer. This sharp increase in dissolution over a very short interval (interpolated duration of 0.4 m.y.) is difficult to explain. It may represent an extremely rapid and short-lived rise in CCD during the mid-Eocene, or it may represent diagenetic effects.

SEDIMENTATION RATES

Sedimentation rates in Hole 523 were calculated from calcareous nannofossil datums in the Miocene to Quaternary portion of the hole and from magnetic polarity reversal events in the mid-Eocene to Oligocene portion (Fig. 5). The chronology of the paleomagnetic datums is

essentially that of LaBrecque et al. (1977) as modified by Mankinen and Dalrymple (1979) using the new decay and abundance constants recommended by the International Union of Geological Sciences Subcommittee on Geochronology.

The early Pliocene to Quaternary rates are generally similar to those of the mid-Eocene to Oligocene, whereas sedimentation rates in the Miocene, especially in the upper Miocene, are very low (<1 m/m.y.) because of strong dissolution during this interval. Episodes of mild dissolution in the late Oligocene and late Eocene are reflected in the reduced sedimentation rates at these times.

PALEOMAGNETISM

Figure 8 summarizes the magnetic polarity stratigraphy at Site 523. All samples from Cores 19 through 50 (~60 m sub-bottom to the bottom of the hole) were demagnetized in an alternating field of 175 Oe. Paleomagnetic results from the overlying 60 m (Cores 1-18) are not included in this discussion because poor core recovery did not permit a continuous and detailed magnetic stratigraphy to be developed. Discrete polarity

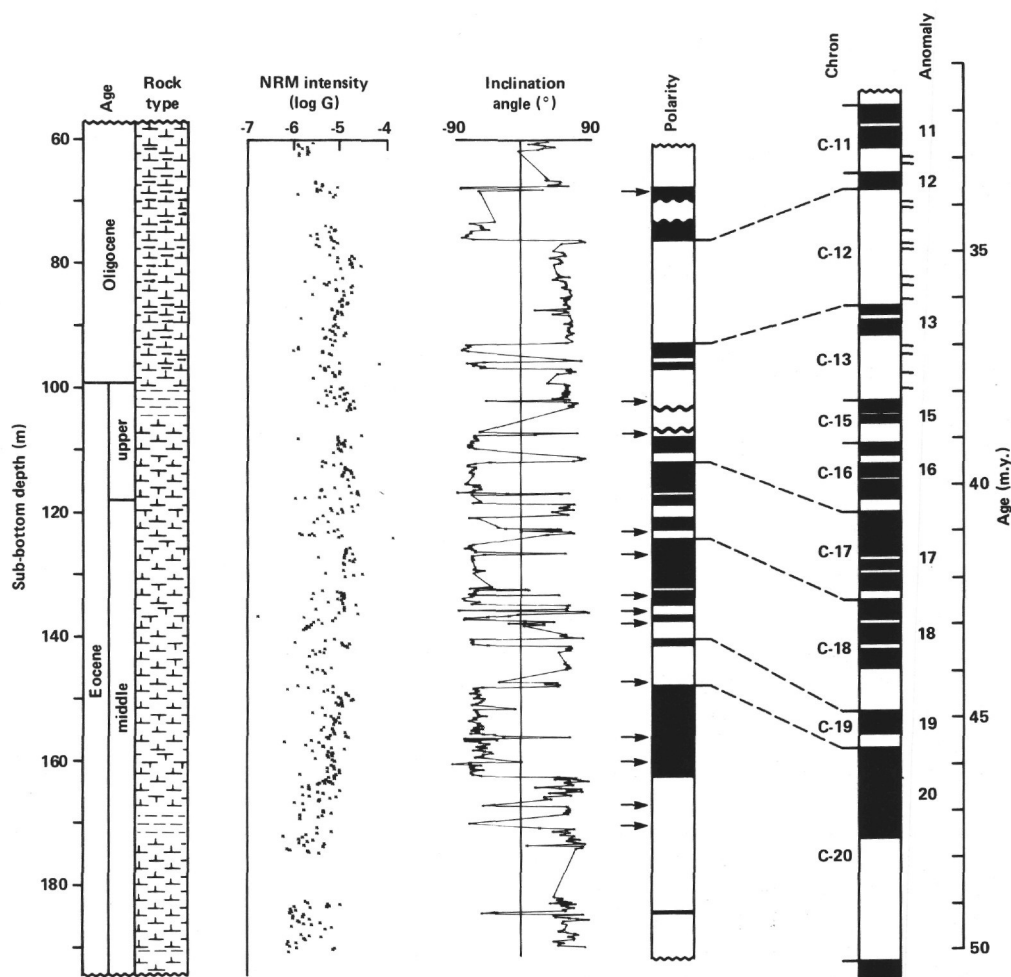


Figure 8. Summary of magnetic polarity stratigraphy at Site 523. Polarity as in Fig. 3. Lithology is defined in Hsü, LaBrecque, et al. (this vol.). Arrows identify the positions of polarity units that are defined by single samples and are therefore tentative (see Tauxe et al., this vol.).

measurements for this upper interval are given in the core description forms.

Chron C-11 through C-20 are represented by the Oligocene-middle Eocene sediments recovered at Site 523. Although the site was located on Marine Magnetic Anomaly 21, equipment failures precluded the recovery of sediment all the way to basement (Chron C-21). From about 80 m sub-bottom downward, the principal trend in the sedimentary sequence is a decrease in NRM intensity and an increase in sedimentation rate.

The paleomagnetic results from Site 523 extend the Paleogene magnetostratigraphy well into the middle Eocene. When combined with the results from Sites 522 and 524, these data provide one of the better Paleogene paleomagnetic sequences published to date.

PHYSICAL PROPERTIES

Core recovery was poor in the upper part of the section because of fairly heavy seas. Water cores, mush, and vertical bedding characterize these cores. As the sea calmed recovery improved. Sediments were more indurated at this site than at the other sites drilled on this leg; the induration is reflected in the higher values of bulk density and lower values of water content. The induration is not reflected in the acoustic velocities measured throughout the column, however; these values remain comparable to those measured at the previously drilled sites. The physical property data are summarized in Figure 9.

Several measurements of thermal conductivity were carried out on the sediments recovered from the lower part of the hole. The results are given in Table 2. The average and range are consistent with the values observed for the previous sites. There appears to be little variation in thermal conductivity with increasing depth or age.

INORGANIC GEOCHEMISTRY

The results of the analysis of the interstitial water chemistry of three samples from Site 523 are summarized in Table 3. No significant downhole trends are evident in this short section of sediments.

CORRELATION OF GEOPHYSICAL DATA TO DRILLING RESULTS

The site originally proposed for Site IV-7 was inappropriate because the seismic data for the site indicated a high probability that nearly all of the sediment cover was missing. The site eventually drilled is midway through Anomaly 21 at a water depth of 4573 m. The age of the site was predicted to be 50 m.y., middle Eocene.

The seismic data were gathered with 5- and 10-in.³ air guns with a band pass of 80 to 200 cps. As can be seen from Figure 1, the site is in a region of pelagic drape with a minimum of regional relief. The site was chosen for this reason because of our experience at previous sites, Site 522 in particular. The seismic data indicate 0.25 s or approximately 190 m of sediment. A more exact estimate is difficult to make because the reflections

off the acoustic basement are diffuse, probably as a result of the increasing frequency of turbidite deposits observed near the bottom of the drilled section. The turbidites may have come from the adjacent minor peak during the early history of the basement. A reflector at 0.08 s may correspond to the layer of *Braarudosphaera* ooze.

Although the hole did not reach the basement, extrapolated sedimentation rates indicate that the basement is 200 m sub-bottom.

SUMMARY AND CONCLUSIONS

The primary objective of drilling at Site 523 was to obtain a precisely dated calcareous Oligocene and Eocene sequence for studies in paleoceanography. On the whole, we achieved this aim, but the preservation of the foraminifers was poor in some of the upper Oligocene and upper Eocene sediments. The correlation of the biostratigraphy and magnetostratigraphy is good for 60 m of the section ranging from middle Oligocene to middle Eocene in age.

Lithostratigraphy

The draped sediments at this site are entirely pelagic. The sequence is divided into three units: (1) Pliocene-Quaternary oozes, (2) Miocene red clays, and (3) Eocene-Oligocene oozes. The first unit is equivalent to the A(lbatross), B(lake), and C(hallenger) formations, the second is equivalent to the D(iscovery) formation, and the lowest is equivalent to the E(ndeavor), F(ram), G(azzelle), and G(rampus) formations at Sites 17 to 20, which were drilled during Leg 3 (see Maxwell et al., 1970). The threefold division is a manifestation of the effect of the Miocene CCD rise.

Correlation of Biostratigraphy and Magnetostratigraphy with Seafloor Lineations

Recovery was less than perfect in the upper third of the cored section, and core disturbance rendered the quality of the paleomagnetic data fair to poor. Furthermore, much of the Neogene section had undergone extensive dissolution. Therefore, we did not compare the magnetostratigraphy of the upper 60 m of sediments to the pattern of the seafloor lineations. Magnetostratigraphic Chrons C-11 to C-20 were recognized in the lower sequence. The correlation of the magnetostratigraphy to seafloor anomaly patterns is excellent from Chron C-17 to C-20 (Tauxe et al. and Poore et al., this vol.). The datum levels (highest occurrences, HOs, and lowest occurrences LOs) of the following key species (species useful for zonation) have been calibrated magnetostratigraphically: HO *Sphenolithus distentus*, HO *S. pseudoradians*, HO *Reticulofenestra umbilica*, HO *Coccolithus formosus*, HO *Discoaster saipanensis*, LO *S. pseudoradians*, LO *Isthmolithus recurvus*, HO *Chiasmolithus grandis*, HO *C. solitus*, HO *Nannotetrina fulgens*, HO *C. gigas*, HO *Globigerinatheka* spp., HO *Acarinina* spp., LO *Globorotalia cerroazulensis*, HO *Globigerina frontosa*, LO *Globorotalia possagnoensis*, and LO *Globigerinatheka mexicana*. Interested readers should refer to

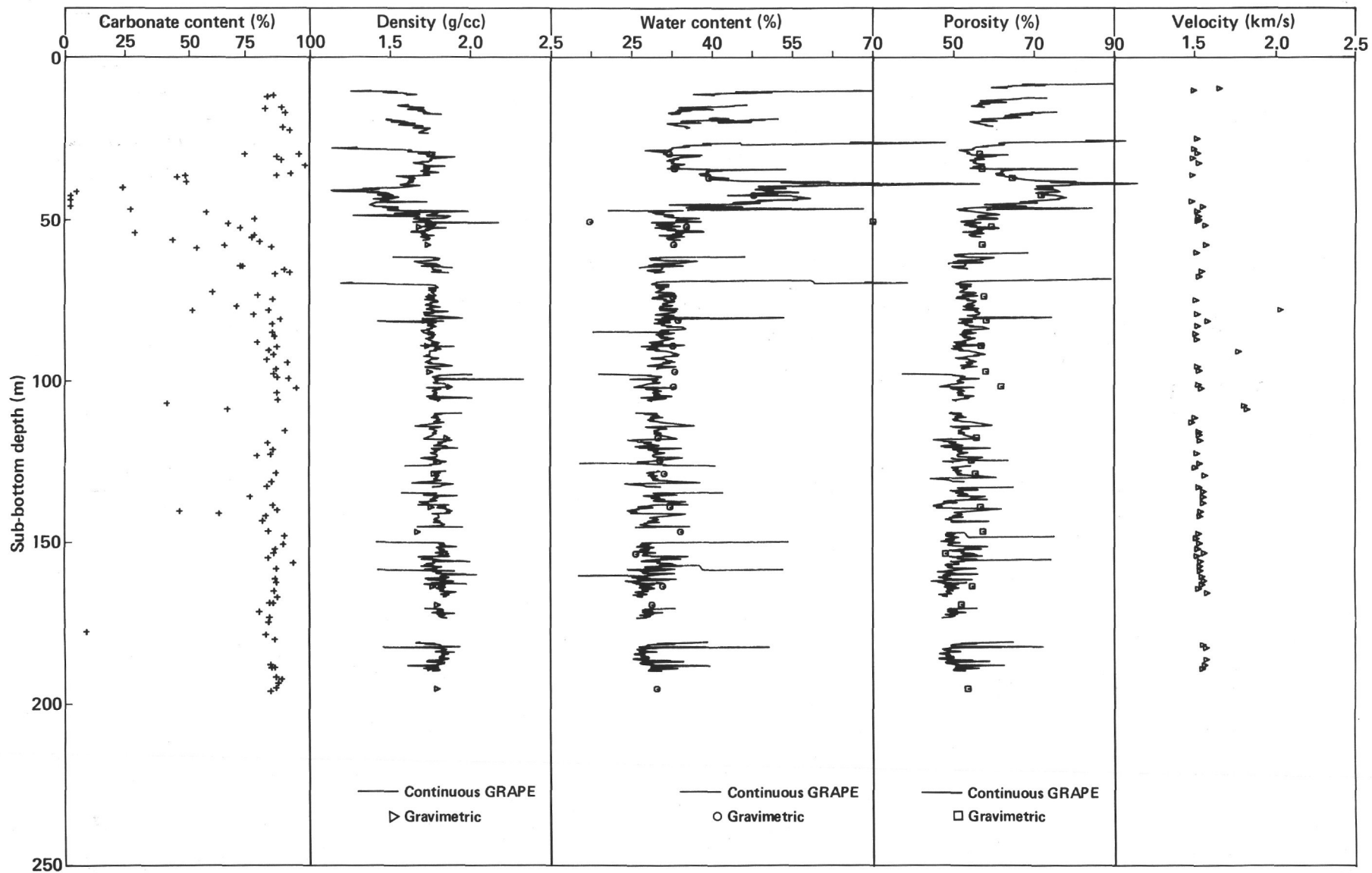


Figure 9. Summary of physical properties, Site 523. Velocity data are for perpendicular beds.

Table 2. Thermal conductivity measurements at Site 523.

Core-Section	Thermal conductivity ^a (W/m °C)
42-2	1.93
45-2	1.30
47-2	1.47
50-2	1.63

^a Average value: 1.58; standard deviation: 0.63.

Table 3. Summary of shipboard geochemical data.

Core-Section (intervals in cm)	Sub-bottom depth (m)	pH	Alkalinity (meq/l)	Salinity (‰)	Calcium (mmol/l)	Magnesium (mmol/l)	Chlorinity (‰)
8-1, 144-150	30.4-34.0	7.338	2.577	38.0	10.66	49.91	19.37
21-2, 144-150	72.9-77.3	7.236	2.494	35.5	10.51	53.73	19.61
30-2, 143-150	106.8-110.8	7.282	2.294	36.8	10.80	51.45	19.20

the summary of the biostratigraphy of this site and of this cruise (Poore et al., this vol.), as well as to range charts, for further information.

The magnetostratigraphy of the deep-sea Paleogene sequence is being worked out for the first time. The study of the Contessa sections near Gubbio, Italy (Lowrie et al., in press) affords the only comparison. We note a number of discrepancies. For example, the HO of *C. grandis* was found near the top of Chron C-18 at Contessa, but it is well up in Chron C-17 at Site 523. We suspect that the upper range has been truncated at Con-

tessa; the dissolution of nannofossils there is considerable (Perch-Nielsen, pers. comm., 1982). For that reason we believe the nannofossil datum levels at Site 523 to be more reliable.

The planktonic foraminifers at Site 523 have been subjected to various degrees of dissolution, and the dissolution causes difficulties when one tries to correlate the nannofossil foraminifer zones. For example, the middle/late Eocene boundary is usually placed at either the top of the range of the foraminifer *Truncorotaloides rohri* (P14/P15 boundary) or the top of the range of the nannofossil *C. grandis* (NP17/NP18 boundary; see Hardenbol and Berggren, 1978). At this site, the P14/P15 foraminifer zonal boundary was defined by the HO of *Acarinina* spp.; the last common occurrence of this species is at the bottom of Chron C-17, considerably lower than the HO of *C. grandis*, which is near the top of Chron C-17. Rare occurrences of the foraminifer species in overlying cores suggest that the species LAD might indeed be higher and that the position of the HO in the biostratigraphic summary of this site report may be due to the truncation of the range by dissolution.

Calcite Dissolution

The Neogene record of dissolution is obscured by the slow sedimentation rate (see Fig. 10). The seafloor at Site 523 subsided to greater depths during the Miocene than at the other Leg 73 sites, so the sediments of the whole epoch belong to the pleistolytic or hololytic red clay facies. One specimen barren of nannofossils was recognized from NN11/12, indicating the rise of the

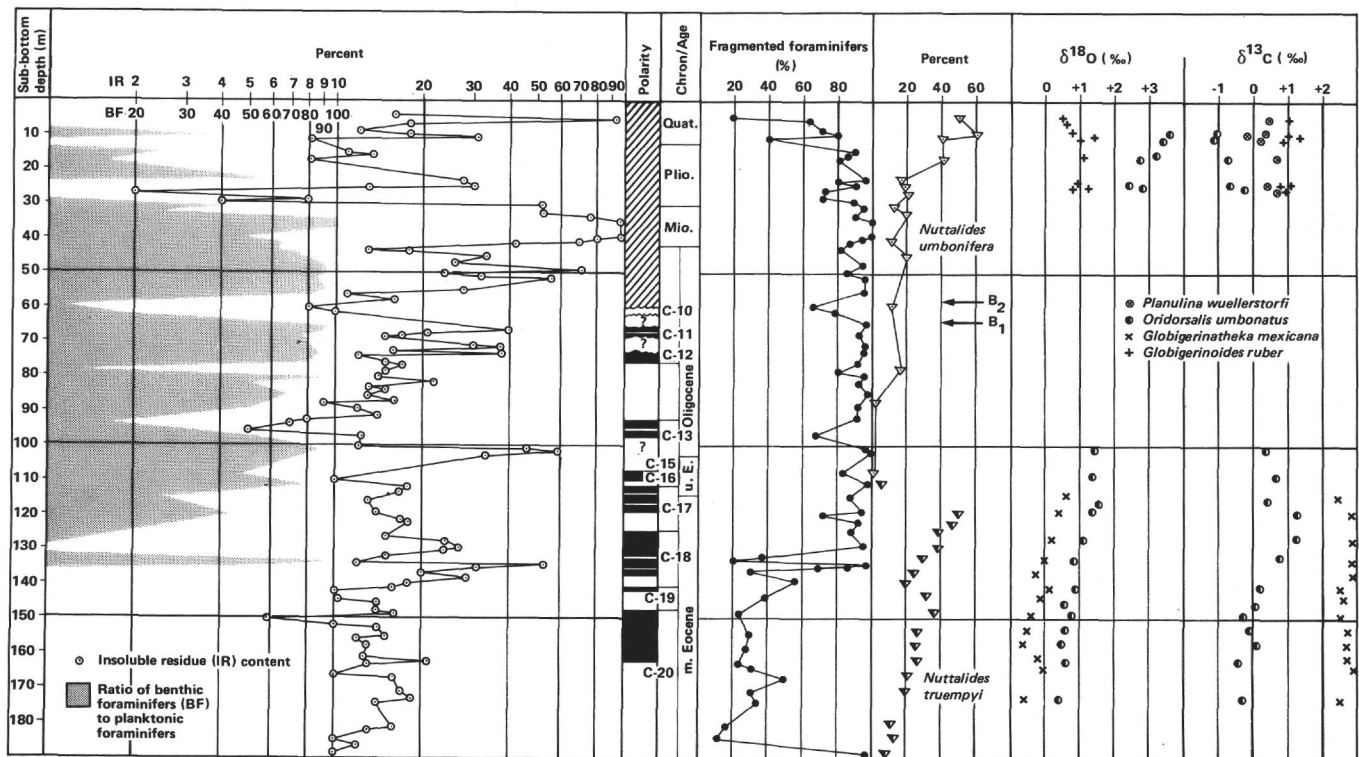


Figure 10. Paleocyanographic data, Site 523. Polarity as in Fig. 3. Arrows B₁ and B₂ indicate the stratigraphic levels of two *Braarudosphaera* chalk layers.

CCD above 4200 m during that time. Nannofossil Zones NN2 to 4 and NN7 to 10 were not identified. Either the marker fossils for those zones are more susceptible to dissolution or dissolution was more intense during those times.

Insoluble residue (IR) content indicates that episodes of dissolution took place during the early Oligocene shortly before the *Braarudosphaera* blooms (Chron C-11 and C-12). A more significant event coincided approximately in timing with the Eocene/Oligocene boundary (Chron C-13-R). This boundary event produced marls with 60% IR content, 90% benthic foraminifer (BF) content, and almost 100% fragmented foraminifer tests (FFT). The equivalent late Eocene facies on the other side of the Mid-Atlantic Ridge (Site 19) has been named Gazelle ooze, and it has an average IR content of only 32%, with a maximum close to 50% (Hsü and Andrews in Maxwell, Von Herzen, Andrews, et al., 1970, p. 449; Pimm, 1970, p. 502). Thus, during the Eocene, as in later times, dissolution seems to have been more intense on the east than the west side of the Mid-Atlantic Ridge.

Another sharp dissolution event during the late middle Eocene is evident from the sediment at 135 m sub-bottom, a horizon that has been dated as belonging to Chron C-18-N2 and lies near the top of NP16, within the *Truncotaloides rohri* zone. Insoluble residue content increased from about 10 to 50%, benthic foraminifer content from less than 10 to more than 90%, and foraminifer test fragments from about 50 to more than 90%. As Figure 10 shows, this event coincides in timing with the beginning of shifts in oxygen and carbon isotopes; it also coincides with the beginning of an increase in the proportion of *Nuttalides truempyi* in the benthic assemblage. The water depth of the site at that time was about 3050 m, and this dissolution event marked a brief but significant excursion of the lysocline to a water depth of less than 3000 m. A corresponding dissolution event has been found at Site 19 at 99.6 m sub-bottom (Sample 19-7-4, 3-4 cm), the base of the Gazelle ooze. The mesolytic sediment there lies within Nannofossil Zone NP16 and the *T. rohri* Foraminifer Zone, and it has an IR content of about 50% (Pimm, 1970, p. 502).

The calcite dissolution has greatly affected the sedimentation rate. The Pliocene-Quaternary sediments were deposited at an average rate of 6 m/m.y. (1.2 m/m.y. for IR). The Miocene sediments are less than 12 m thick and were deposited at a rate of about 0.65 m/m.y. (0.5 m/m.y. for IR). The Oligocene sediments were deposited at a rate of 4.5 m/m.y. (1.0 m/m.y. for IR). The Eocene rate was about 8 m/m.y. (1.2 m/m.y. for IR). As at other Leg 73 sites, the rate at which insoluble residue accumulated was reduced during the Miocene.

Paleoceanography

A significant change in oceanic environments took place during the late Eocene some 45 Ma (Chron C-18-N2). The benthic population began to change; toward the end of middle Eocene it was dominated by *Nuttalides truempyi*. Both the benthic and planktonic foraminifers show oxygen-isotope shifts of about 1 ‰ during the late middle Eocene, indicating a rapid cooling of ocean temperatures and/or the first significant accumulation of Antarctic ice. The temperatures remained little changed during the late Eocene and probably began to decrease again toward the beginning of the Oligocene (as shown by the evidence from Site 522). An increase in *N. umbonifera* took place 3 to 4 m.y. later, in the early Oligocene, when the site sank to a water depth of 3500 m, but the increase still took place 3 to 4 m.y. earlier than at Site 522. At each site, the abundance of this species is clearly related to the history of subsidence. The trend of the oxygen-isotope shift during the Eocene is in general agreement with that noted by Shackleton and Kennett (1975) for the Southern oceans and by Borsma and Shackleton (1977) for the South Atlantic. The shifts in the carbon isotopes of the benthic foraminifers were parallel to the oxygen shifts in the isotopes, and both trends were positive during the late middle Eocene.

We found no clear correlation between the dissolution events and paleoclimatic temperatures. Although the beginning of the late middle Eocene oxygen shift was marked by a dissolution event, the CCD change seems to have reversed direction soon after the rise, so that much of the upper middle and upper Eocene sediments are oligolytic (<20% IR). The early Oligocene cooling was also marked by a very brief dissolution event, followed by a dramatic depression of CCD. Oligocene dissolution events took place before and after the *Braarudosphaera* blooms. Apparently the seafloor at Site 523 was positioned near the Oligocene lysocline, so that slight changes in levels produced rapid facies changes.

Two *Braarudosphaera* chalk layers have been identified at Site 523. Both appear during Chron C-10, and they have ages of about 30.5 and 31.5 m.y., respectively. They can be correlated with the upper two *Braarudosphaera* chalk layers at Site 522. The blooms took place at times when the CCD was deep.

During the Miocene, the seafloor remained below the lysocline, and at times it was below the CCD, so only pleistolytic and hololytic sediments accumulated. The lysocline rapidly became depressed during the early Pliocene, so the sediments changed from pleistolytic to mesolytic to oligolytic within 1 m.y. or so. This remarkable depression in CCD, which is well known, is discussed in Hsü et al. (this vol.).

Basement Age and Seafloor Spreading Rate

The basement was not reached by drilling because of mechanical failures. We penetrated a thick section of sediments belonging to Chron C-20-R but did not reach the oldest sediment at the site, which lies on Seafloor Anomaly 21. However, the drilling depth lies very close to the seismically predicted basement depth, and the drill hole may have bottomed a few meters above the basement.

The ages of several magnetostratigraphically calibrated datums can be compared with radiometric dates to yield information on the question of seafloor spread-

ing rate. LaBrecque et al. (1977) noted that the Eocene magnetostratigraphic ages predicted on the basis of a linear rate of seafloor spreading during the Cenozoic are consistently older than the corresponding radiometric ages. They suspected significant deviations from a linear seafloor spreading rate. If we use the HO of *Chiasmolithus grandis* near the top of Chron C-17 as the middle/late Eocene boundary, the magnetostratigraphic age would be about 41 m.y., appreciably younger than that indicated by Bolli, Ryan, et al. (1978). However, the age is still significantly older than the radiometric age given by Odin and Curry (1981) for the NP17/NP18 boundary (37 ± 1.5 m.y.). In addition, magnetostratigraphy indicates that the top of NP15, although not accurately dated, should not be higher than the top of Chron C-20-N; in other words, it should be older than 46 m.y. (Tauxe et al., this vol.). Radiometry, however, indicates that the datum should be younger than 41 m.y. If the minimum probable age of 63.5 m.y. is used for the beginning of the Cenozoic, the magnetostratigraphic ages for the tops of NP17 and NP15 would be 39 and 43.5 m.y., respectively. The discrepancies are reduced to 2 and 2.5 m.y., respectively, but they still lie outside the quoted experimental error of radiometric dating. We cannot rule out the possibility that seafloor spreading did deviate slightly from linearity during the middle Eocene. More probably the radiometric dates by Odin and Curry are too young (see Hsü et al., this vol.).

REFERENCES

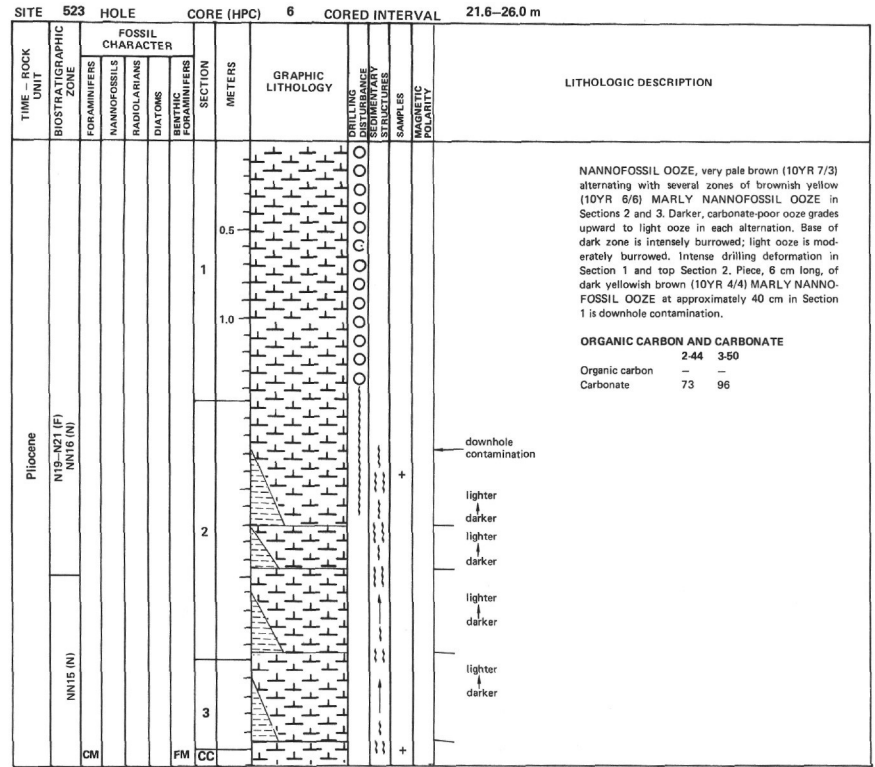
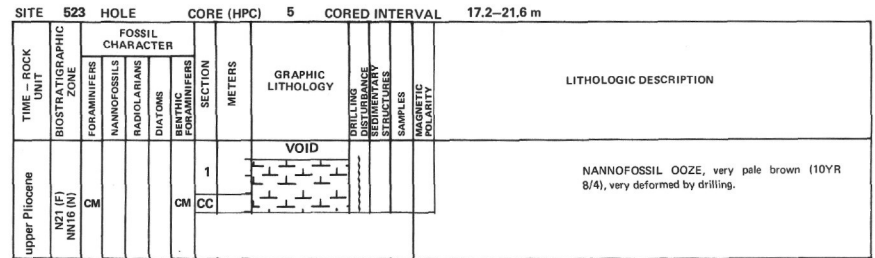
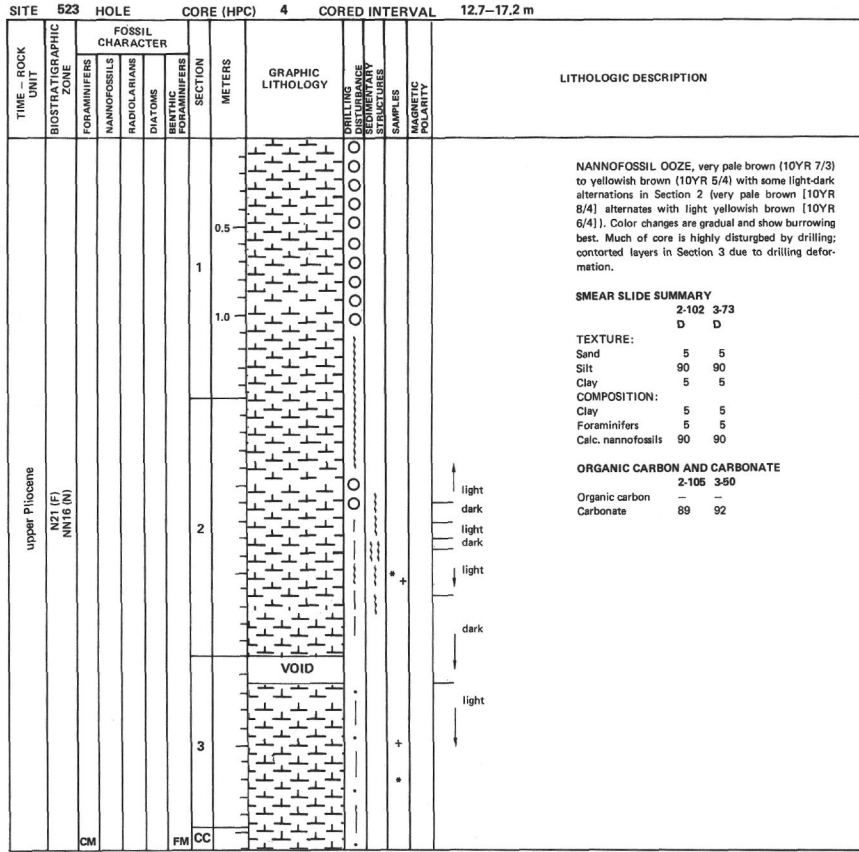
- Berger, W. H., and von Rad, U., 1972. Cretaceous and Cenozoic sediments from the Atlantic Ocean. In Hayes, D. E., Pimm, A. C., et al., *Init. Repts. DSDP*, 14: Washington (U.S. Govt. Printing Office), 787-954.
- Berggren, W. A., 1972. Late Pliocene-Pleistocene glaciation. In Laughton, A. S., Berggren, W. A., et al., *Init. Repts. DSDP*, 12: Washington (U.S. Govt. Printing Office), 953-964.
- Boersma, A., and Shackleton, N., 1977. Tertiary oxygen and carbon isotope stratigraphy, Site 357 (mid latitude South Atlantic). In Supko, P., Perch-Nielsen, K., et al., *Init. Repts. DSDP*, 39: Washington (U.S. Govt. Printing Office), 911-924.
- Bolli, H. M., Ryan, W. B. F., et al., 1978. *Init. Repts. DSDP*, 40: Washington (U.S. Govt. Printing Office).
- Douglas, R. C., and Savin, S. M., 1975. Oxygen and carbon analyses of Tertiary and Cretaceous microfossils from Shatsky Rise and other sites in the North Pacific Ocean. In Larson, R. L., Moberly, R., et al., *Init. Repts. DSDP*, 32: Washington (U.S. Govt. Printing Office), 509-520.
- Hardenbol, J., and Berggren, W. A., 1978. A new Paleogene numerical time-scale. *Stud. Geol. Tulsa Okla.*, 6:205-212.
- Kennett, J. P., Burns, R. E., Andrews, J. E., Churkin, M., Davies, T. A., et al., 1972. Australian-Antarctic continental drift: paleocirculation changes and Oligocene deep-sea erosion. *Nat. Phys. Sci.*, 239(91):51-55.
- LaBrecque, J. L., Kent, D. V., and Cande, S. C., 1977. Revised magnetic polarity time scale for Late Cretaceous and Cenozoic time. *Geology*, 5:330-335.
- Lowrie, W., Alvarez, W., Napoleone, G., Perch-Nielsen, K., Premoli Silva, I., et al., in press. Paleogene magnetic stratigraphy in Umbrian pelagic carbonate rocks: the Contessa sections, Gubbio. *Geol. Soc. Am. Bull.*
- Mankinen, E. A., and Dalrymple, G. B., 1979. Revised geomagnetic polarity time scale for interval 0-5 m.y.B.P. *J. Geophys. Res.*, 84: 615-626.
- Maxwell, A. E., Von Herzen, R. P., et al., 1970. *Init. Repts. DSDP*, 3: Washington (U.S. Govt. Printing Office).
- Maxwell, A. E., Von Herzen, R. P., Andrews, J. E., Boyce, R. E., Milow, E. D., and Shipboard Scientific Party, 1970. Summary and conclusions. In Maxwell, A. E., Von Herzen, R. P., et al., *Init. Repts. DSDP*, 3: Washington (U.S. Govt. Printing Office), 441-471.
- Odin, G. S., and Curry, D., 1981. L'échelle numérique des temps paléogènes en 1981. *C. R. Hebd. Seances Acad. Sci. Ser. B*, 293: 1003-1006.
- Pimm, A. C., 1970. Carbon carbonate results, Leg 3. In Maxwell, A. E., Von Herzen, R. P., et al., *Init. Repts. DSDP*, 3: Washington (U.S. Govt. Printing Office), 495-508.
- Schnitker, D., 1980. Quaternary deep-sea benthic foraminifers and bottom water masses. *Ann. Rev. Earth Planet. Sci.*, 8:343-370.
- Shackleton, N. J., and Kennett, J. P., 1975. Paleotemperature history of the Cenozoic and the initiation of Antarctic glaciation: oxygen and carbon isotope analyses in DSDP Sites 277, 279 and 281. In Kennett, J. P., Houtz, R. E., et al., *Init. Repts. DSDP*, 29: Washington (U.S. Govt. Printing Office), 743-755.
- Watkins, N. D., and Kennett, J. P., 1972. Regional sedimentary disconformities and upper Cenozoic changes in bottom water velocities between Australia and Antarctica. *Antarct. Res. Ser.*, 19:273-293.

Date of Initial Receipt: August 12, 1982

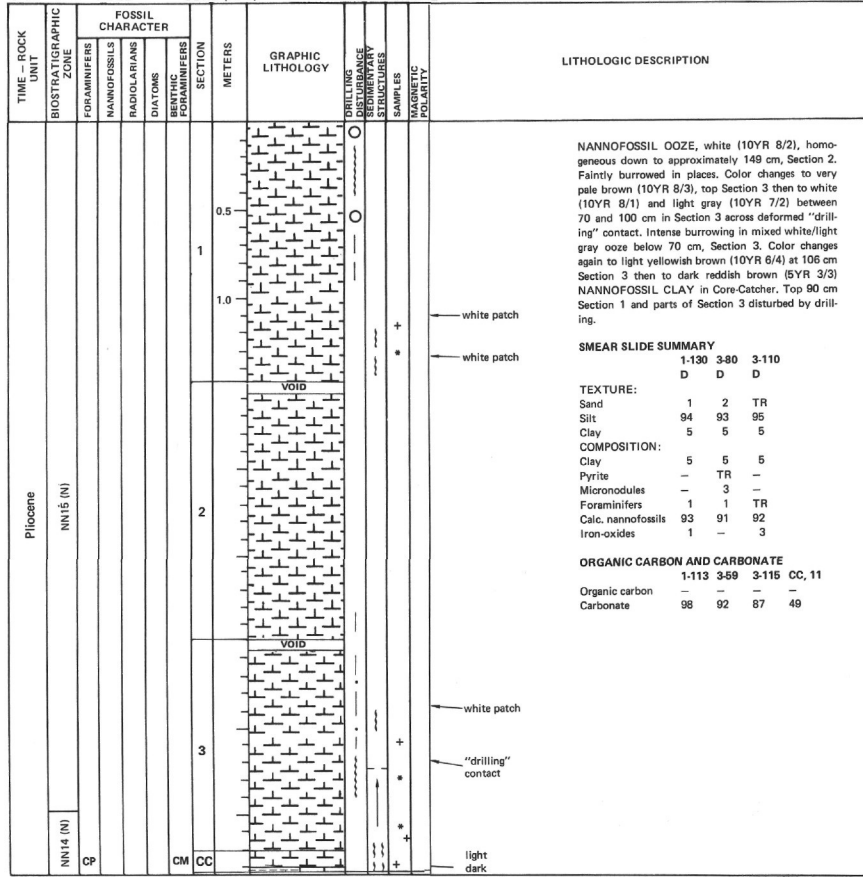
SITE 523 HOLE CORE (HPC) 1 CORED INTERVAL 0.0-4.0 m									
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS				
Quaternary	N22 (F) NN19 (N)	AG	AP	CP	CC				Trace of FORAMINIFER NANNOFOSSIL OOZE in Core-Catcher.

SITE 523 HOLE CORE (HPC) 2 CORED INTERVAL 4.0-8.4 m									
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS				
Quaternary	N22 (F) NN19 (N)	AG	AP	FG	CC				<p>FORAMINIFER-NANNOFOSSIL OOZE, very pale brown (10YR 7/4-7/6-8/3) alternating with darker, light yellowish brown (10YR 6/4) intervals in Section 2. Burrows common throughout but best seen along color contacts. Top Section 1 disturbed by drilling.</p> <p>SMEAR SLIDE SUMMARY 1-126 D</p> <p>TEXTURE: Sand 5 Silt 85 Clay 10</p> <p>COMPOSITION: Clay 5 Carbonate unspec. 5 Foraminifers 5 Calc. nannofossils 85 Diatoms TR Iron-oxides TR</p> <p>ORGANIC CARBON AND CARBONATE 1-146 2-36 Organic carbon - Carbonate 85 83</p>

SITE 523 HOLE CORE (HPC) 3 CORED INTERVAL 8.4-12.8 m									
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS				
Quaternary	N21 (F) NN18 (N)	CM	CM	CC					<p>FORAMINIFER NANNOFOSSIL OOZE, very pale brown (10YR 7/3-8/3) alternating with (darker) light yellowish brown (10YR 6/4) MARLY NANNOFOSSIL OOZE on scale of 5-80 cm. Alternations usually consist of darker intensely burrowed marly ooze grading upward into lighter less distinctly burrowed ooze with burrowed but relatively sharp upper contact with overlying marly ooze. Top 60 cm of Section 1 are soupy, but contain some zones of brown (10YR 4/3) MARLY NANNOFOSSIL OOZE.</p> <p>SMEAR SLIDE SUMMARY 1-98 1-112 D D</p> <p>TEXTURE: Sand 10 10 Silt 85 80 Clay 5 10</p> <p>COMPOSITION: Clay 5 5 Carbonate unspec. - 5 Foraminifers 10 5 Calc. nannofossils 80 80 Iron-oxides - 5</p> <p>ORGANIC CARBON AND CARBONATE 1-76 1-110 2-96 Organic carbon - - Carbonate 88 82 90</p>



SITE 523 HOLE CORE (HPC) 7 CORED INTERVAL 26.0-30.4 m



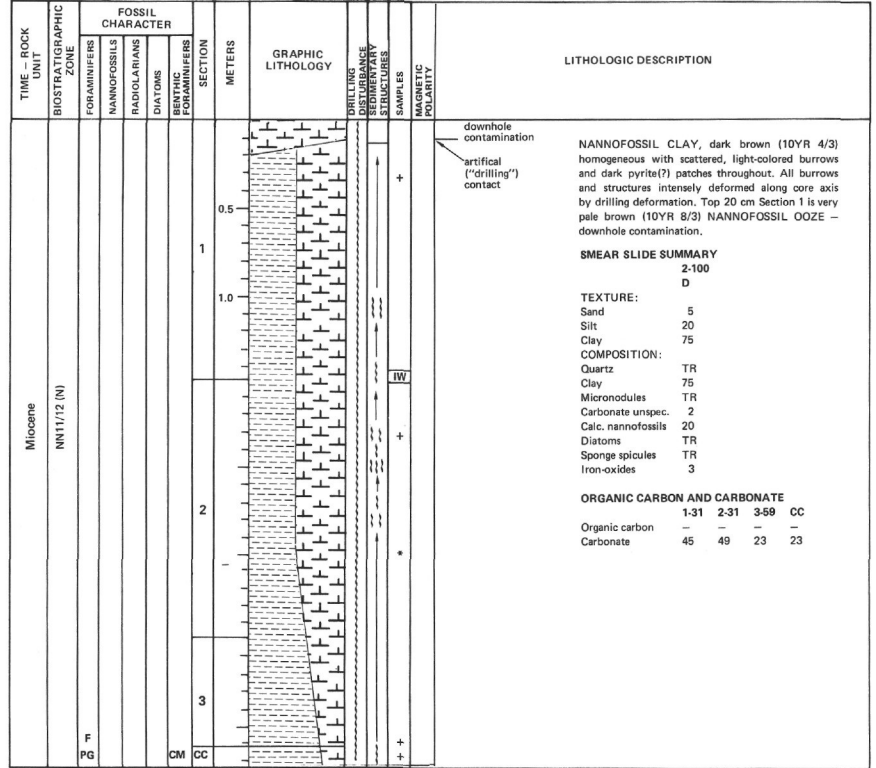
SMEAR SLIDE SUMMARY

	1-130	3-80	3-110
D	D	D	D
TEXTURE:			
Sand	1	2	TR
Silt	94	93	95
Clay	5	5	5
COMPOSITION:			
Clay	5	5	5
Pyrite	-	TR	-
Micronodules	-	3	-
Foraminifers	1	1	TR
Calc. nannofossils	93	91	92
Iron-oxides	1	-	3

ORGANIC CARBON AND CARBONATE

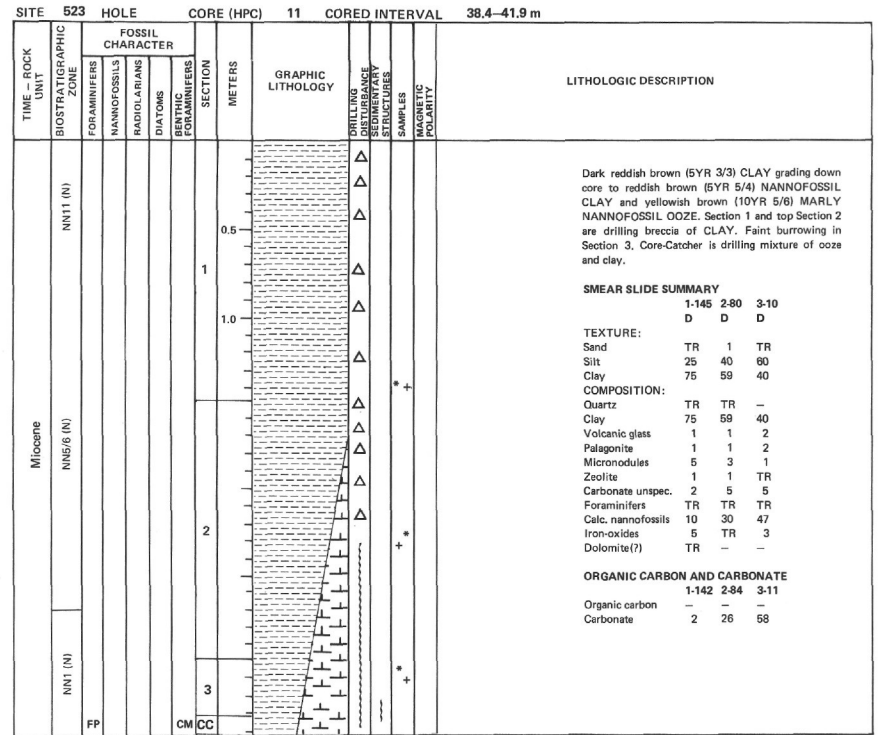
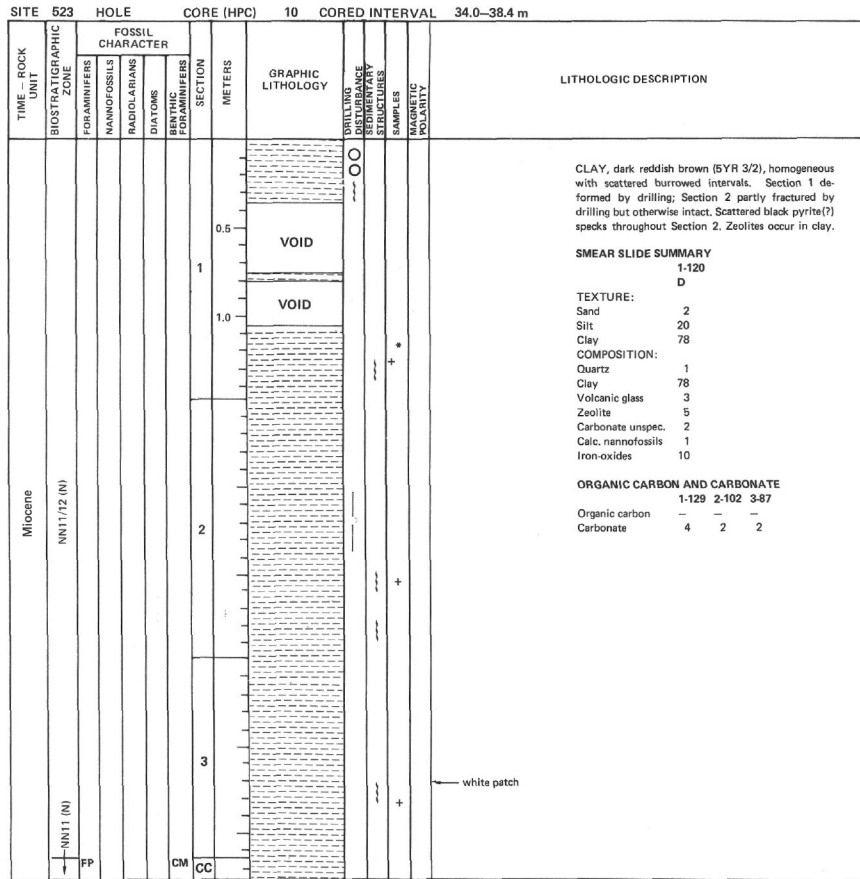
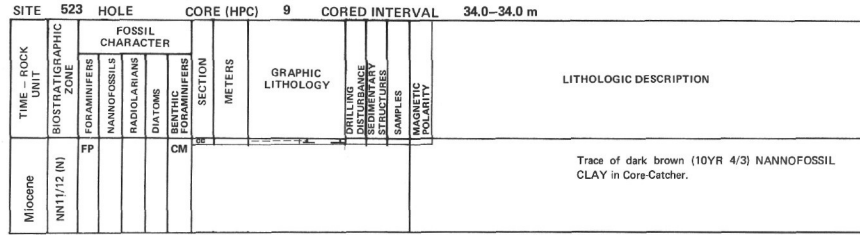
	1-113	3-59	3-115	CC, 11
Organic carbon	-	-	-	-
Carbonate	98	92	87	49

SITE 523 HOLE CORE (HPC) 8 CORED INTERVAL 30.4-34.0 m

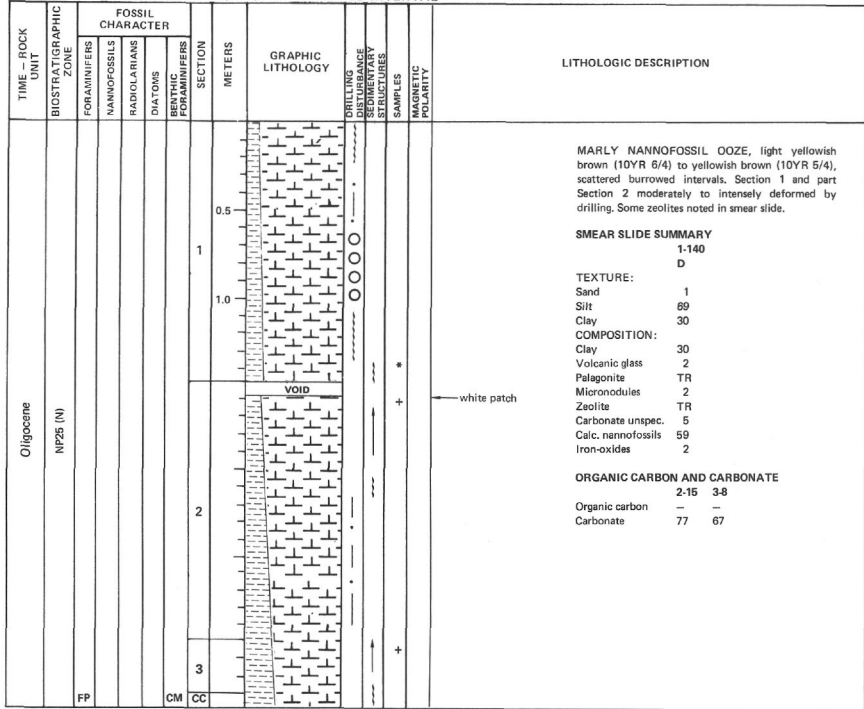


SMEAR SLIDE SUMMARY

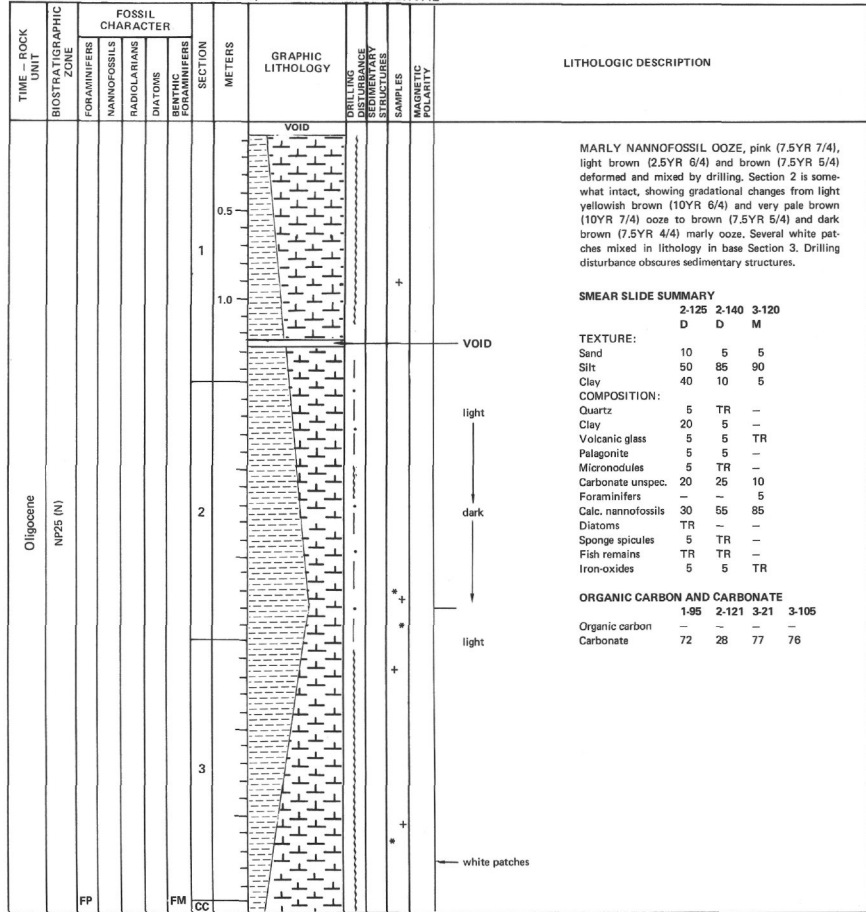
	2-100			
D	D			
TEXTURE:				
Sand	5			
Silt	20			
Clay	75			
COMPOSITION:				
Quartz	TR			
Clay	75			
Micronodules	TR			
Carbonate unspc.	2			
Calc. nannofossils	20			
Diatoms	TR			
Sponge spicules	TR			
Iron-oxides	3			
ORGANIC CARBON AND CARBONATE				
	1-31	2-31	3-59	CC
Organic carbon	-	-	-	-
Carbonate	45	49	23	23

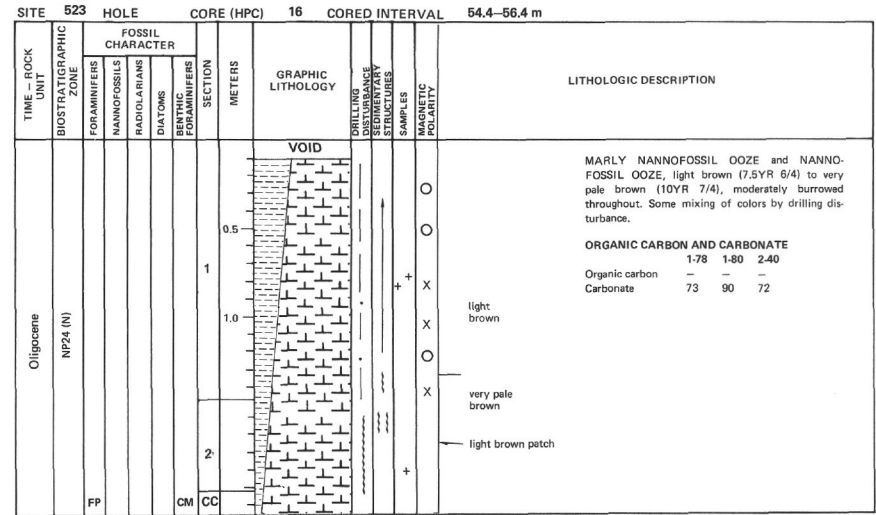
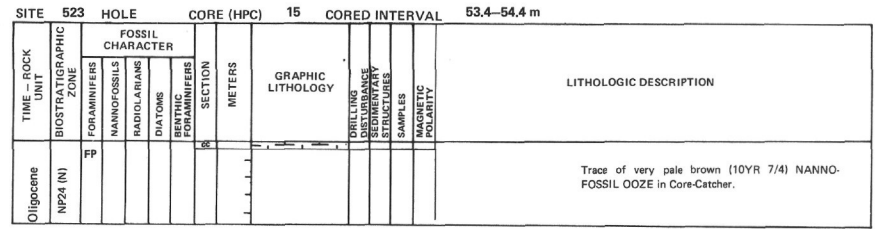
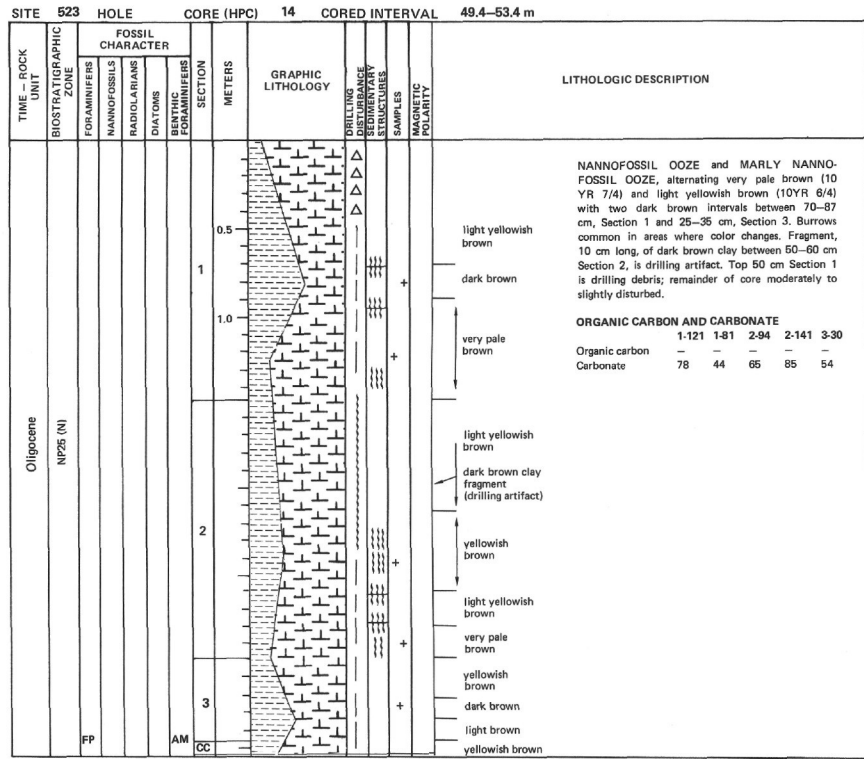


SITE 523 HOLE CORE (HPC) 12 CORED INTERVAL 41.9-45.4 m



SITE 523 HOLE CORE (HPC) 13 CORED INTERVAL 45.4-49.4 m





SITE 523 HOLE CORE (HPC) 17 CORED INTERVAL 59.4-63.4 m

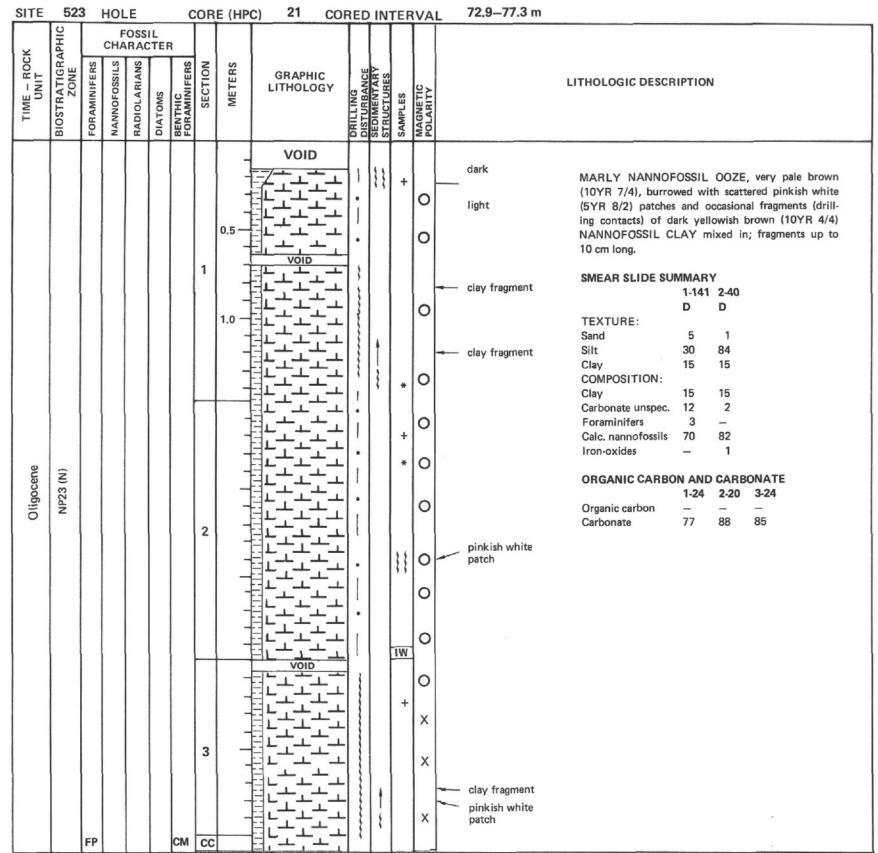
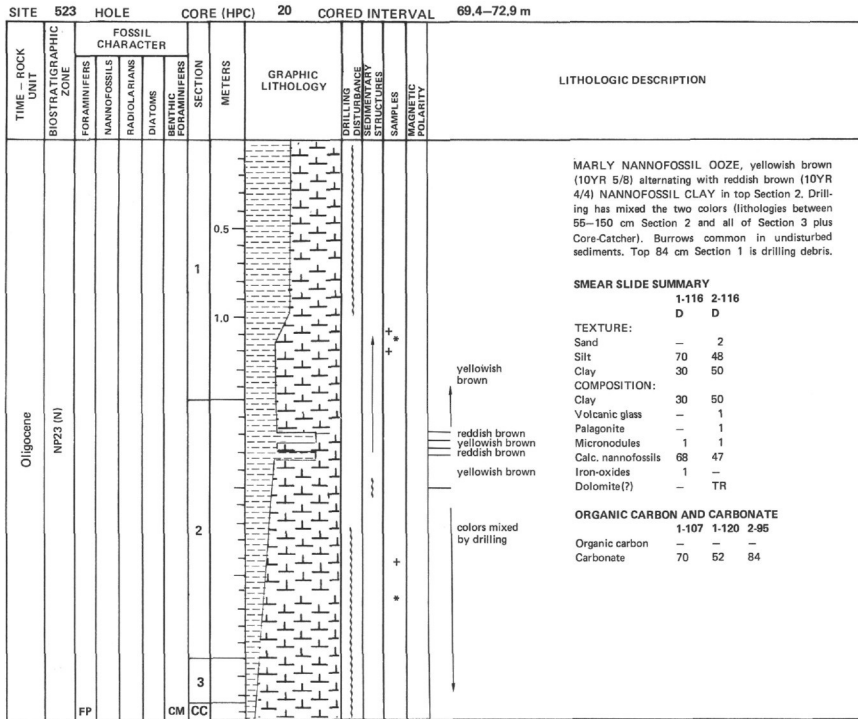
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
lower Oligocene	NP22 (N)					1	VOID					pinkish white patch	
												pinkish white patch	
						1.0						SMEAR SLIDE SUMMARY	
												1-71 1-140	
												D M	
												TEXTURE:	
												Sand - TR	
												Silt 100 90	
												Clay - 10	
												COMPOSITION:	
												Quartz TR -	
												Clay - 5	
												Volcanic glass TR 3	
												Palagonite - 2	
												Micronodules - TR	
												Carbonate unspec. 20 20	
												Foraminifers - 15	
												Calc. nannofossils 80 55	
												Sponge spicules - TR	
												Fish remains - TR	
												ORGANIC CARBON AND CARBONATE	
												1-80 1-124	
												Organic carbon - -	
												Carbonate 92 86	
												pink patches	

SITE 523 HOLE CORE (HPC) 18 CORED INTERVAL 63.4-65.4 m

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
Oligocene	NP23 (N)											Trace of NANNOFOSSIL OOZE in Core-Catcher.	

SITE 523 HOLE CORE (HPC) 19 CORED INTERVAL 65.4-69.4 m

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
Oligocene	NP23 (N)					1	VOID					drilling mixture of rust and ooze	
												Braarudosphaera ooze	
						1.0						yellow brown	
												red brown	
												yellow brown	
												red brown	
												yellow brown	
												dark yellow brown	
												very pale brown	
												grading to	
												light yellowish brown	
												very pale brown	
												yellow brown	
												very pale brown	
												SMEAR SLIDE SUMMARY	
												1-85 1-80	
												M M	
												TEXTURE:	
												Sand - -	
												Silt 100 70	
												Clay - 30	
												COMPOSITION:	
												Clay - 30	
												Volcanic glass - TR	
												Micronodules - TR	
												Carbonate unspec. 90 -	
												Calc. nannofossils 10 69	
												Iron-oxides - 1	
												ORGANIC CARBON AND CARBONATE	
												1-86 2-27 3-7	
												Organic carbon - -	
												Carbonate 60 79 85	



SITE 523 HOLE CORE (HPC) 22 CORED INTERVAL 77.3-80.8 m

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDEMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
		FORAMINIFERS											
		CM	CC										
Oligocene	NP23 (N)					0.5					X	NANNOFOSSIL OOZE, very pale brown (10YR 7/4-8/4), homogeneous, faintly burrowed with scattered white patches in Section 2; white patches have grayish haloes. Slight color change from lighter to darker ooze in Section 1 as noted. Top 50 cm Section 1 is drilling debris. ORGANIC CARBON AND CARBONATE 1-142 2-121 Organic carbon — — Carbonate 85 86	
						1			X	lighter			
						1.0			X	darker			
									X				
									X				
									X				
									X				
									X				
									X				
									X				
									X				
									X				
									X	white patches			
									X				
									X				

SITE 523 HOLE CORE (HPC) 23 CORED INTERVAL 80.8-84.8 m

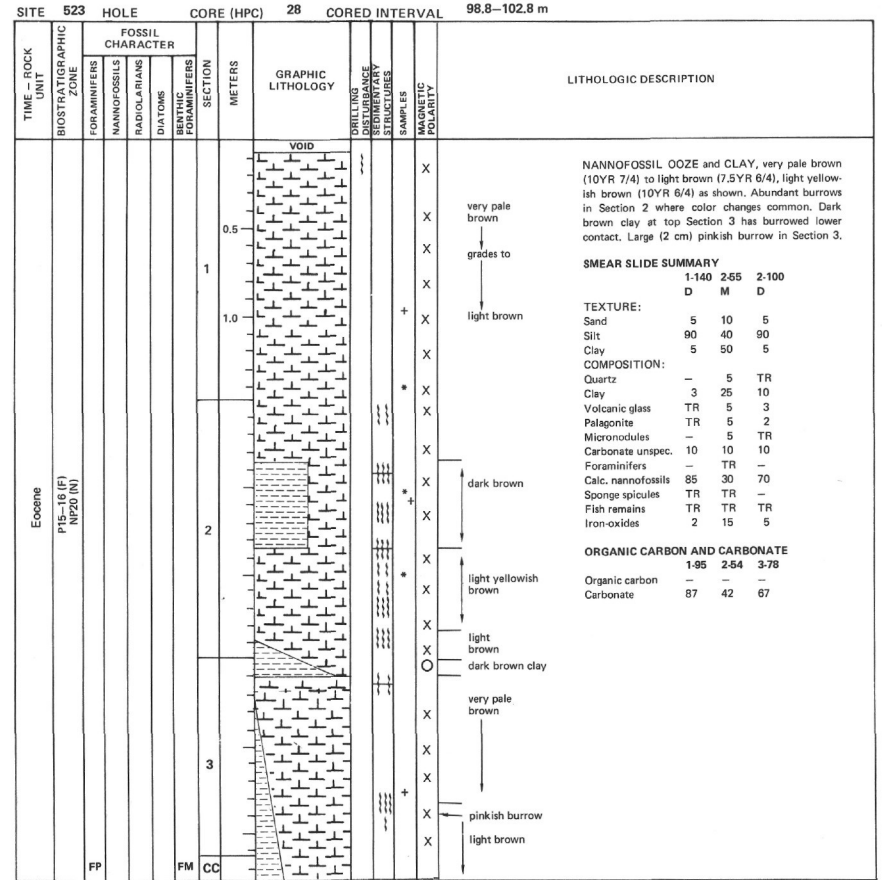
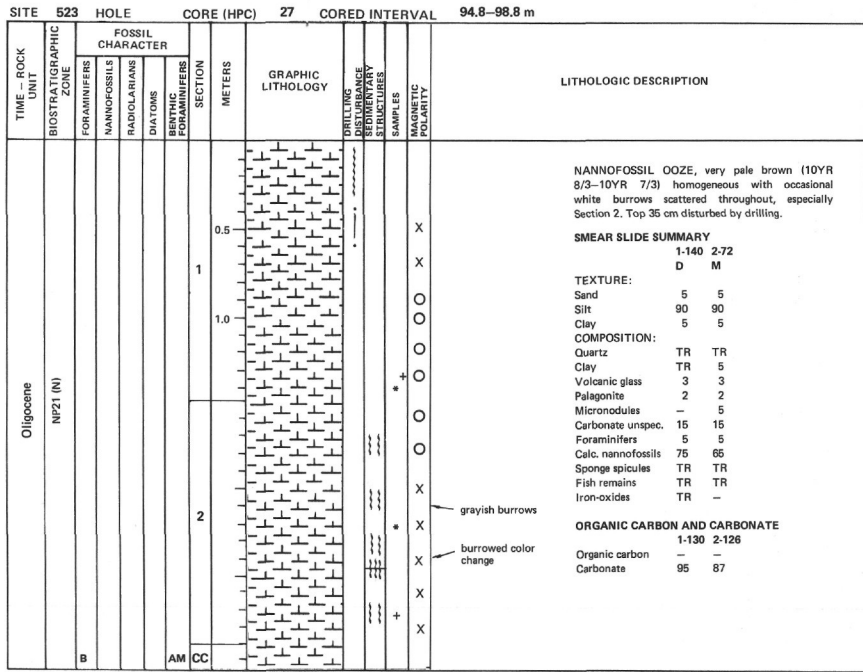
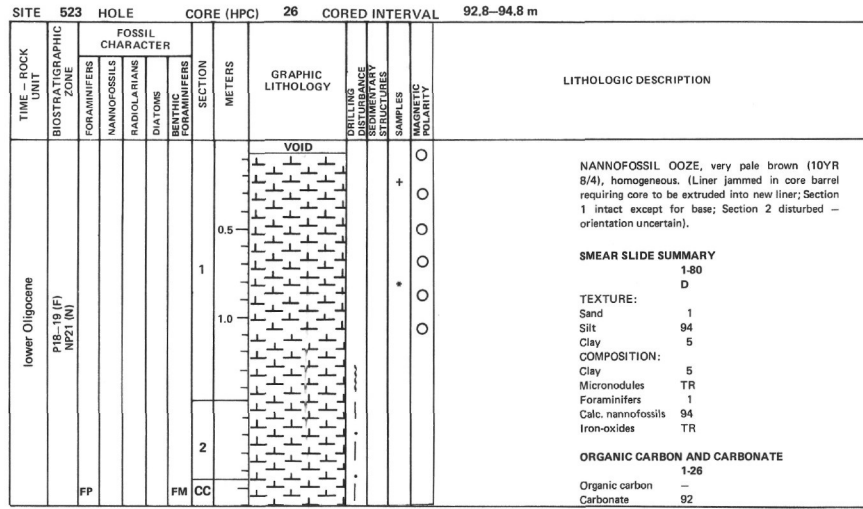
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDEMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
		FORAMINIFERS											
		CM	P	CC									
lower Oligocene	NP23 (N)						VOID				X	NANNOFOSSIL OOZE, very pale brown (10YR 7/4) with one marly, light yellowish brown (10YR 8/4) interval in Section 1. Burrows common between 84-105 cm, Section 1 but faint or sparse throughout rest of core. Pinkish white patches in base Section 2 and in Section 3. Minimal drilling deformation. SMEAR SLIDE SUMMARY 1-100 2-144 D M TEXTURE: Sand — 1 Silt 85 90 Clay 15 9 COMPOSITION: Clay 15 9 Palagonite TR — Micronodules 1 — Carbonate unsp. 2 — Foraminifers TR TR Calc. nanofossils 82 90 Iron-oxides — 1 ORGANIC CARBON AND CARBONATE 1-98 2-98 3-52 Organic carbon 1-98 2-98 3-52 Carbonate 79 87 84	
						0.5				X	very pale brown		
						1				X	light yellowish brown		
						1.0				X	very pale brown		
										X			
										X			
										X			
										X			
										X			
										X			
										X			
										X			
										X	white patch with grayish halo		
										X	2 pinkish white patches		
										X			

SITE 523 HOLE CORE (HPC) 24 CORED INTERVAL 84.8-88.8 m

TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIAATOMS						
Oligocene	P18-18 (F) NP23 (N)	FP				0.5	[Pattern]			X pinkish patches	<p>NANNOFOSSIL OOZE, very pale brown (10YR 7/4) with scattered 1-5 cm pinkish white (5YR 8/2) patches mixed or haloed with grayish ooze. Pink color possibly from rhodochrosite (not obvious in smear slides). Abundant diffuse, yellowish brown (10YR 6/4) spots (3-5 mm across) scattered throughout ooze may be bioturbation features. Minimal core disturbance.</p> <p>SMEAR SLIDE SUMMARY</p> <p>1-100 2-88</p> <p>D D</p> <p>TEXTURE:</p> <p>Sand 3 4</p> <p>Silt 87 80</p> <p>Clay 10 16</p> <p>COMPOSITION:</p> <p>Clay 8 11</p> <p>Carbonate unsp. 10 5</p> <p>Foraminifers 2 4</p> <p>Calc. nannofossils 80 80</p> <p>ORGANIC CARBON AND CARBONATE</p> <p>1-89 2-89 3-32</p> <p>Organic carbon - - -</p> <p>Carbonate 86 93 91</p>
						1.0				X pinkish patch	
	VOID										
Oligocene	P18-19 (F) NP21 (N)	FP				2	[Pattern]			X pinkish patch with gray halo	<p>X pinkish and grayish patches</p> <p>X pinkish patches</p>
						VOID					
	VOID										
Oligocene	NP22 (N)	FP				3	[Pattern]			X white patches	
						VOID					

SITE 523 HOLE CORE (HPC) 25 CORED INTERVAL 88.8-92.8 m

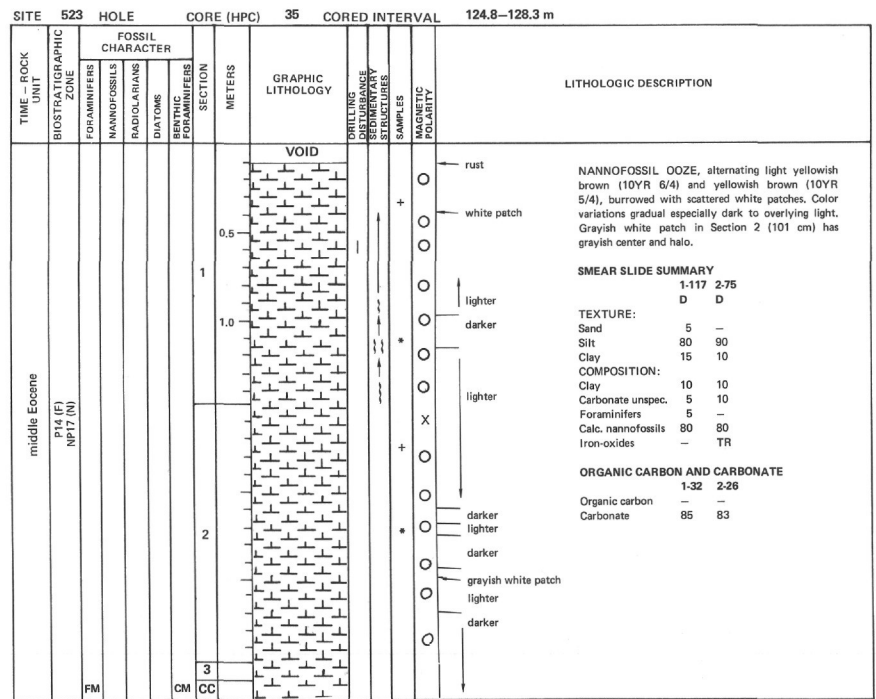
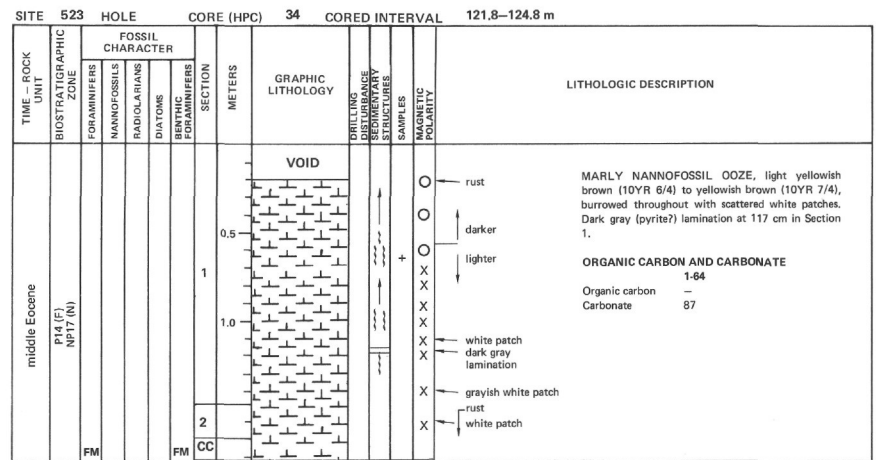
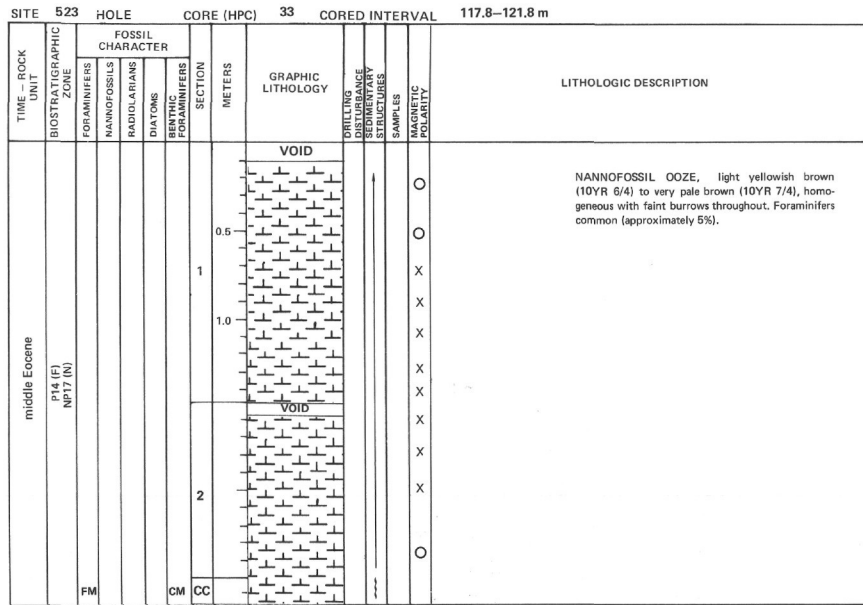
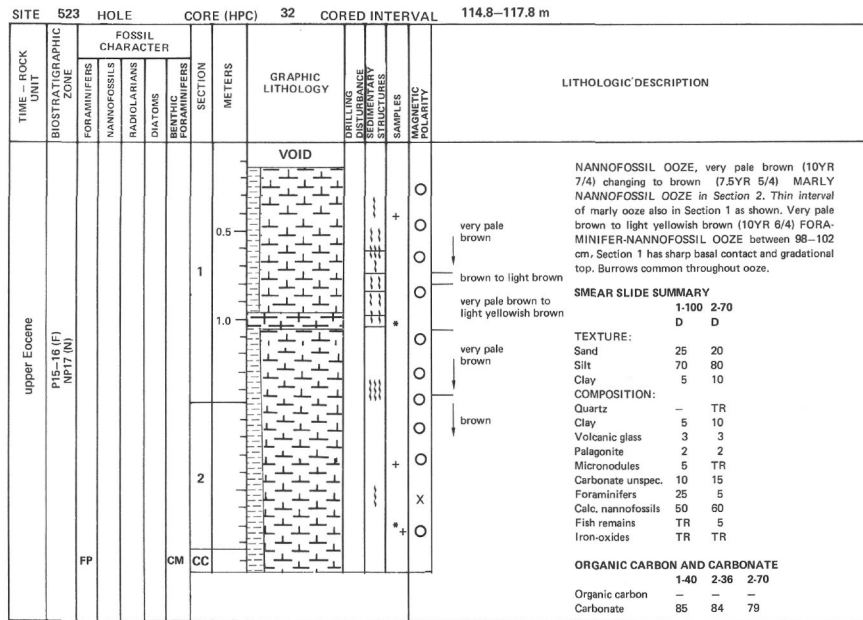
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIAATOMS						
Oligocene	P18-19 (F) NP21 (N)	FP				0.5	[Pattern]			X	<p>NANNOFOSSIL OOZE, very pale brown (10YR 7/4), abundant yellowish brown (10YR 6/4) spots suggest bioturbation. Few scattered pinkish white patches noted in Sections 2 and 3. Core very similar to Core 24. Minimal drilling disturbance.</p> <p>SMEAR SLIDE SUMMARY</p> <p>1-80</p> <p>D</p> <p>TEXTURE:</p> <p>Sand -</p> <p>Silt 90</p> <p>Clay 10</p> <p>COMPOSITION:</p> <p>Clay 10</p> <p>Foraminifers TR</p> <p>Calc. nannofossils 90</p> <p>Iron-oxides TR</p> <p>Dolomite(?)/rhodochrosite(?) TR</p> <p>ORGANIC CARBON AND CARBONATE</p> <p>1-130 2-130 3-92</p> <p>Organic carbon - - -</p> <p>Carbonate 87 86 87</p>
						1.0				X	
	VOID										
Oligocene	P18-19 (F) NP21 (N)	FP				2	[Pattern]			X	<p>X pinkish white patch</p>
						VOID					
	VOID										
Oligocene	NP22 (N)	FP				3	[Pattern]			X	<p>X white patches</p>
						VOID					

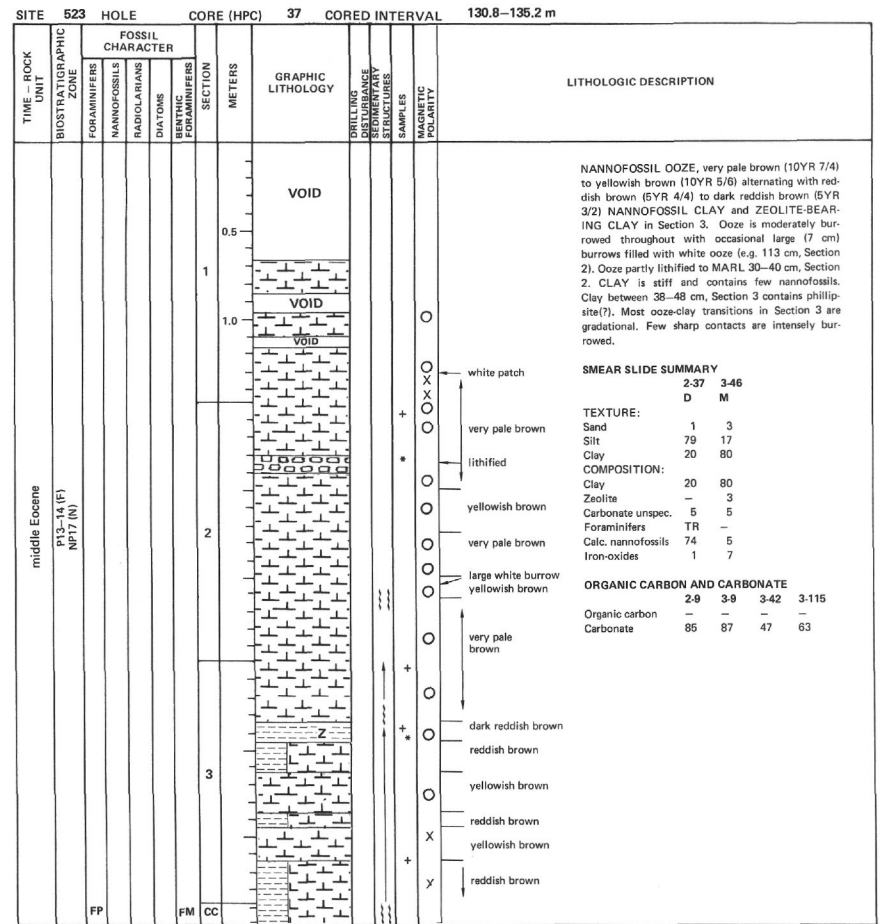
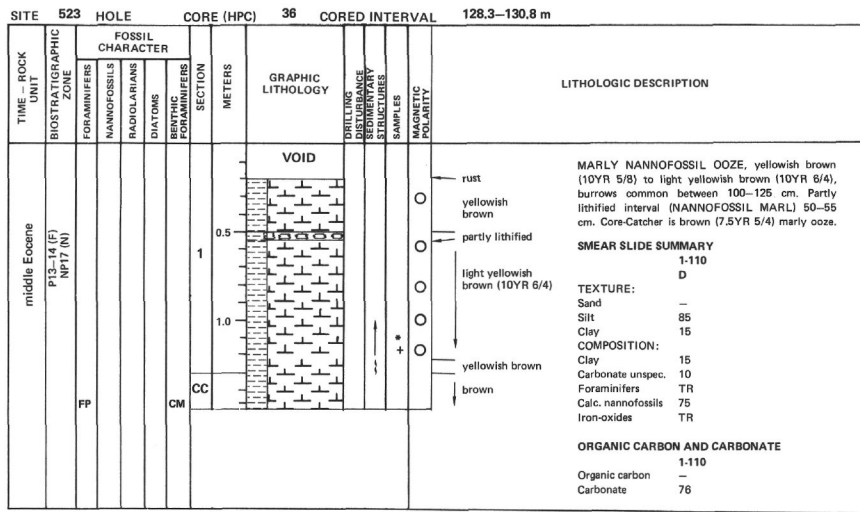


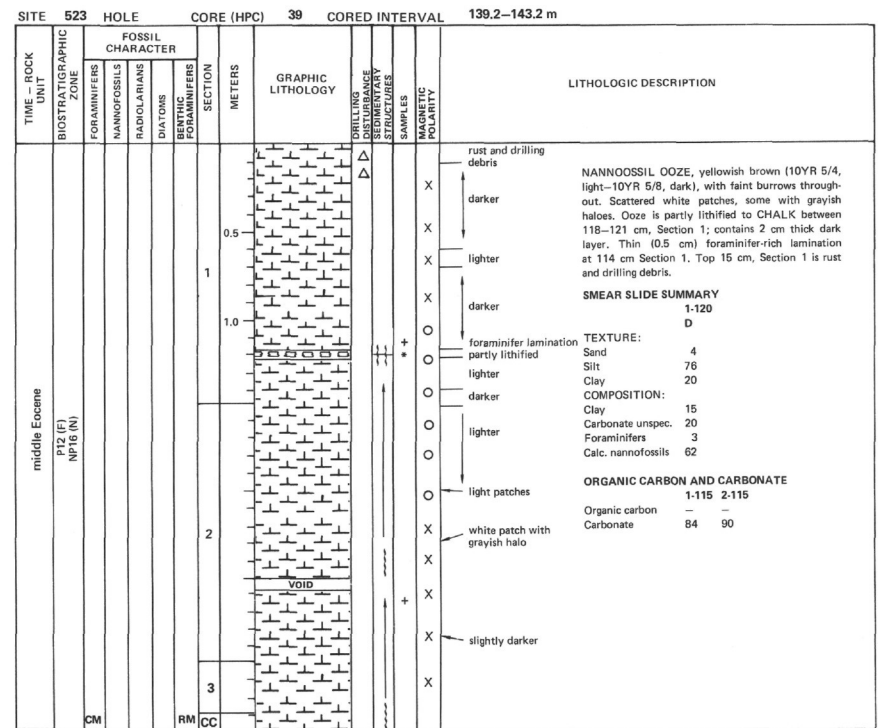
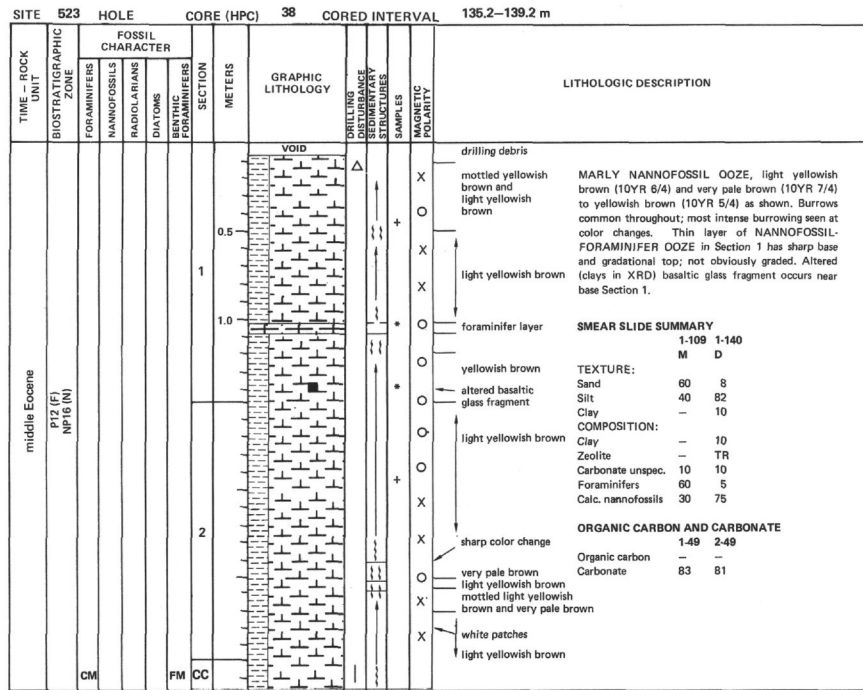
SITE 523 HOLE CORE (HPC) 29 CORED INTERVAL 102.8-106.8 m													
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
Eocene	P15-16 (F) NP18 (N)	FP											Trace of NANNOFOSSIL OOZE in Core-Catcher.

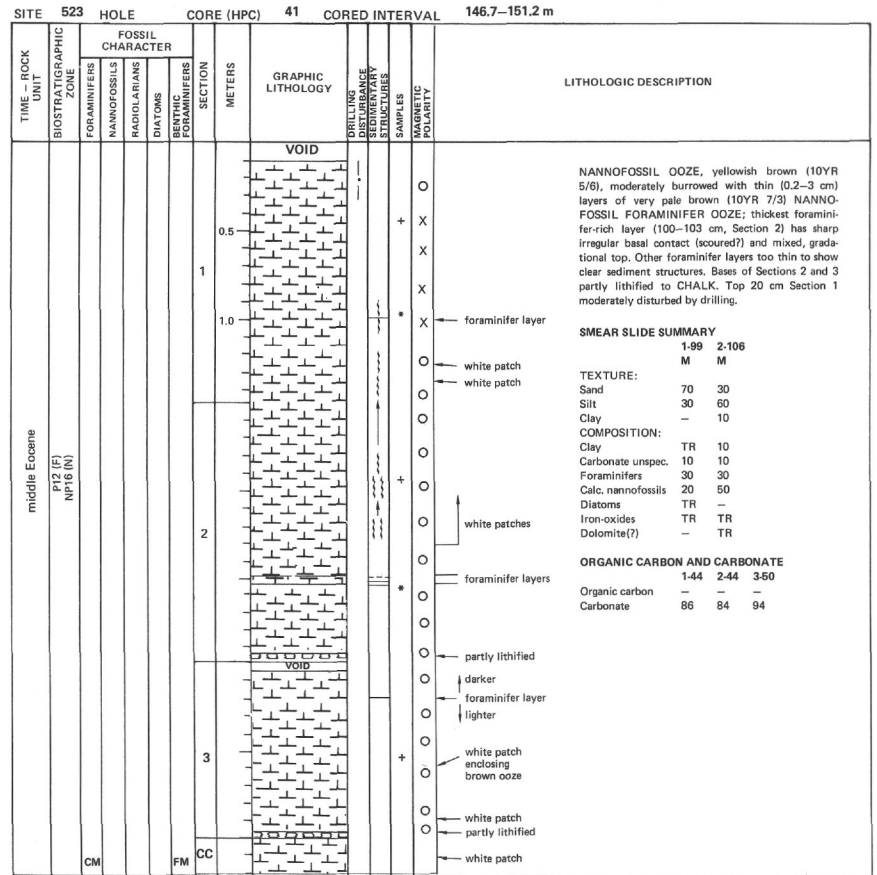
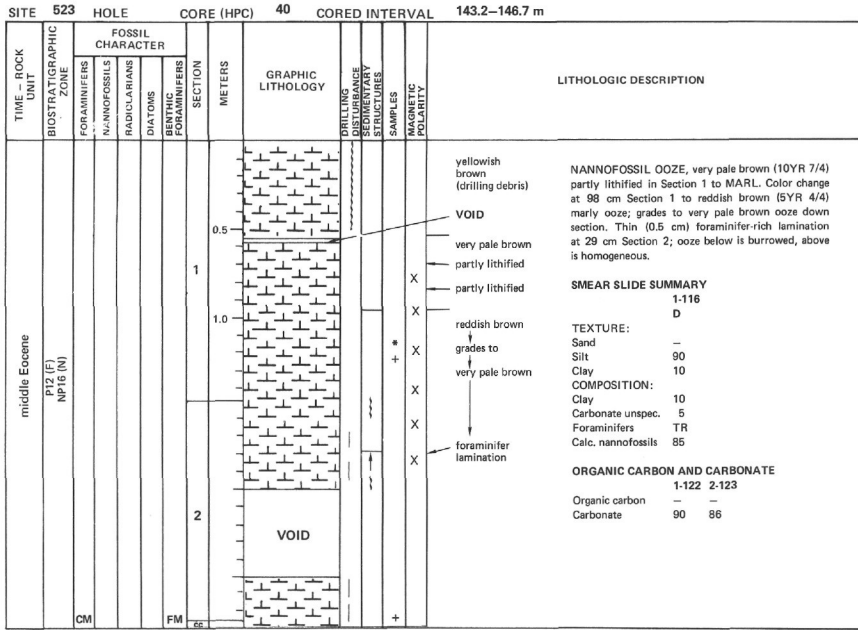
SITE 523 HOLE CORE (HPC) 30 CORED INTERVAL 106.8-110.8 m													
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
upper Eocene	P15-16 (F) NP18 (N)	FP											<p>NANNOFOSSIL OOZE, light yellowish brown (10YR 6/4), burrowed throughout, with single large pink (7.5YR 8/4) burrow between 120-130 cm, Section 2. Thin layer of drilling rust at 55 cm Section 1 intruded in core.</p> <p>ORGANIC CARBON AND CARBONATE Organic carbon - 2-100 Carbonate 90</p>

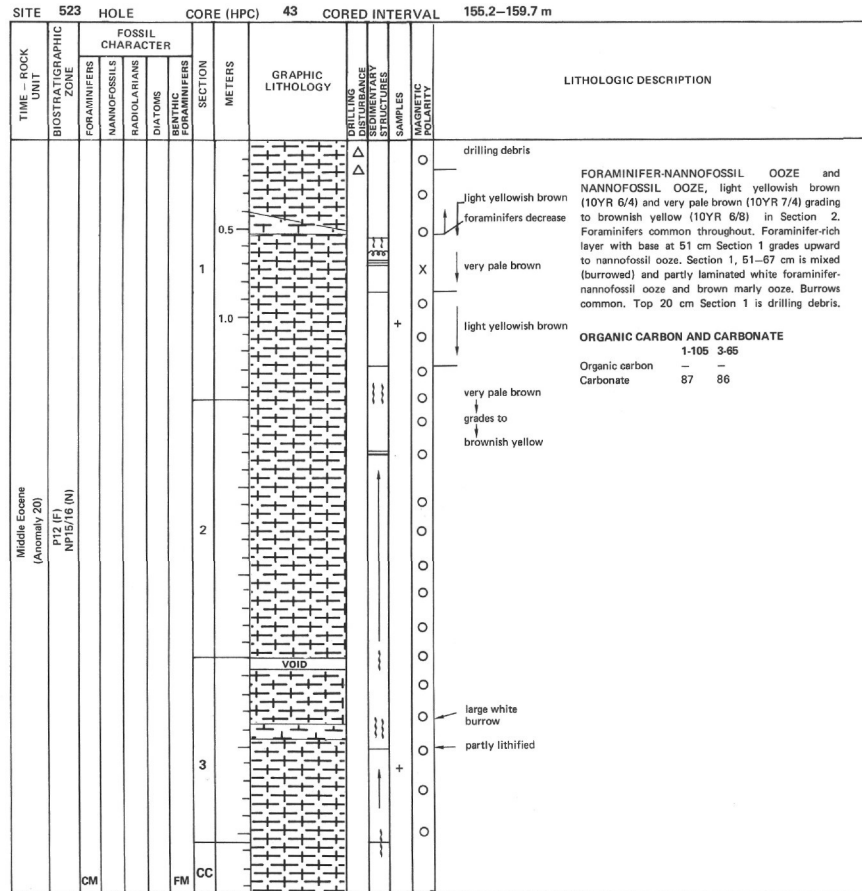
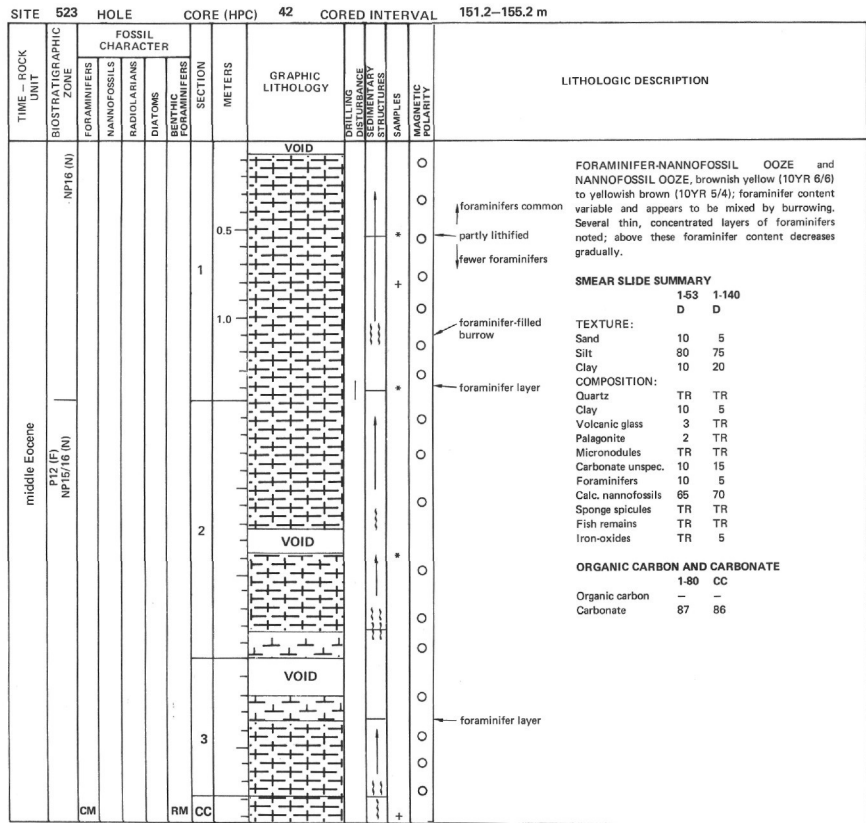
SITE 523 HOLE CORE (HPC) 31 CORED INTERVAL 110.8-114.8 m													
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLES	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS								
upper Eocene	P15-16 (F) NP18 (N)	FP											<p>NANNOFOSSIL OOZE, very pale brown (10YR 7/4) to light yellowish brown (10YR 6/4), burrowed throughout. Minimal core disturbance.</p> <p>SMEAR SLIDE SUMMARY 1-140 D</p> <p>TEXTURE: Sand 5 Silt 90 Clay 5</p> <p>COMPOSITION: Quartz TR Clay 5 Volcanic glass 3 Palagonite 2 Carbonate unsp. 15 Calc. nannofossils 75 Fish remains TR Iron-oxides TR</p> <p>ORGANIC CARBON AND CARBONATE Organic carbon - 2-80 Carbonate 83</p>

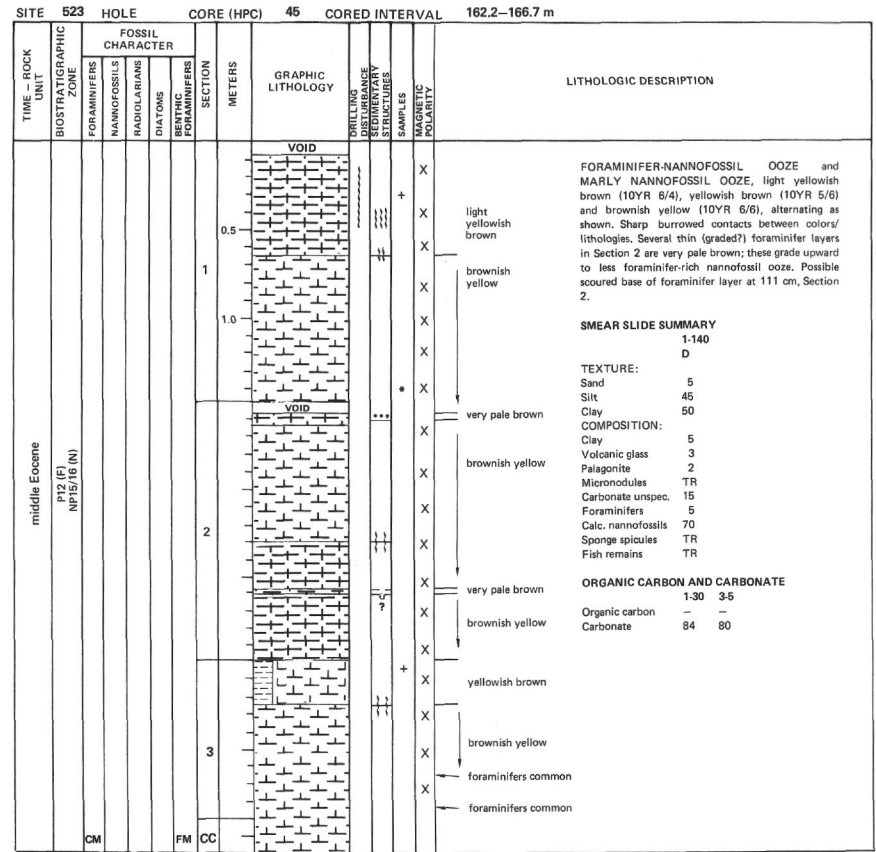
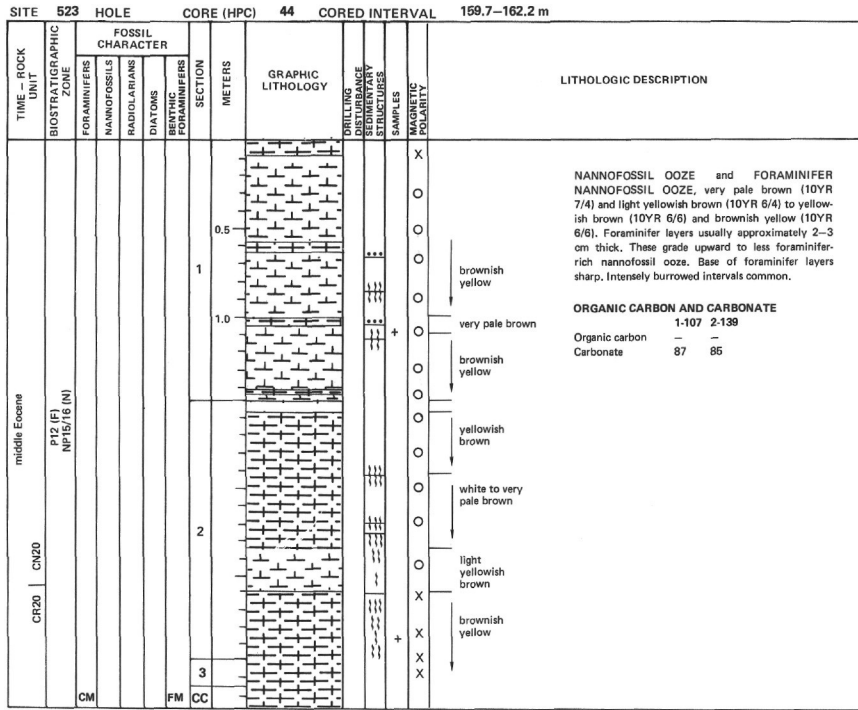


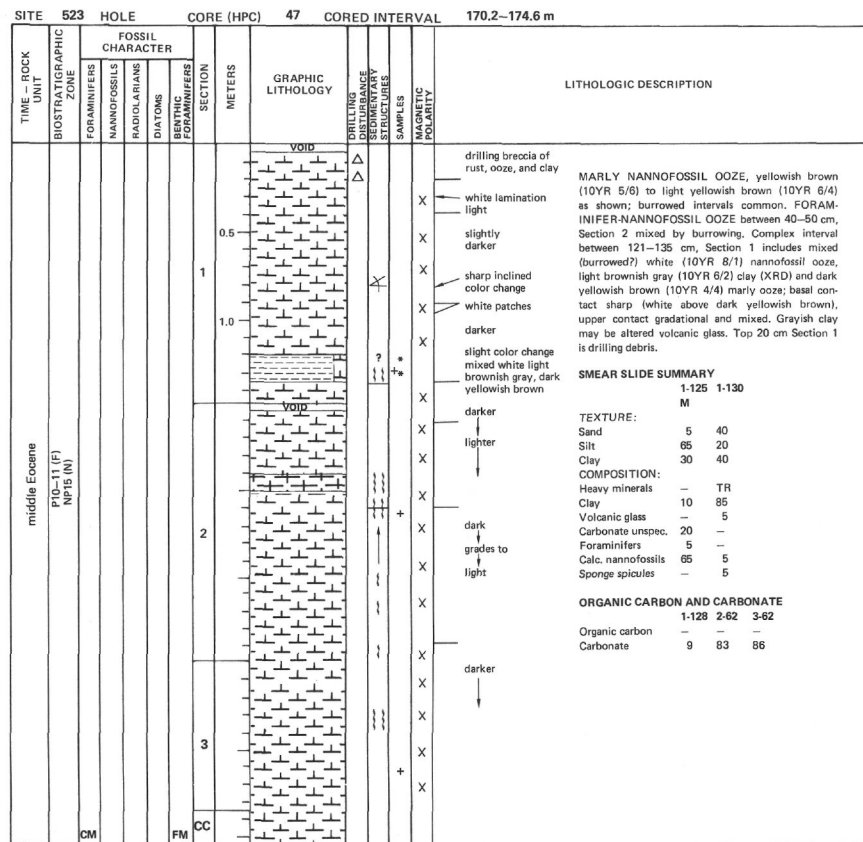
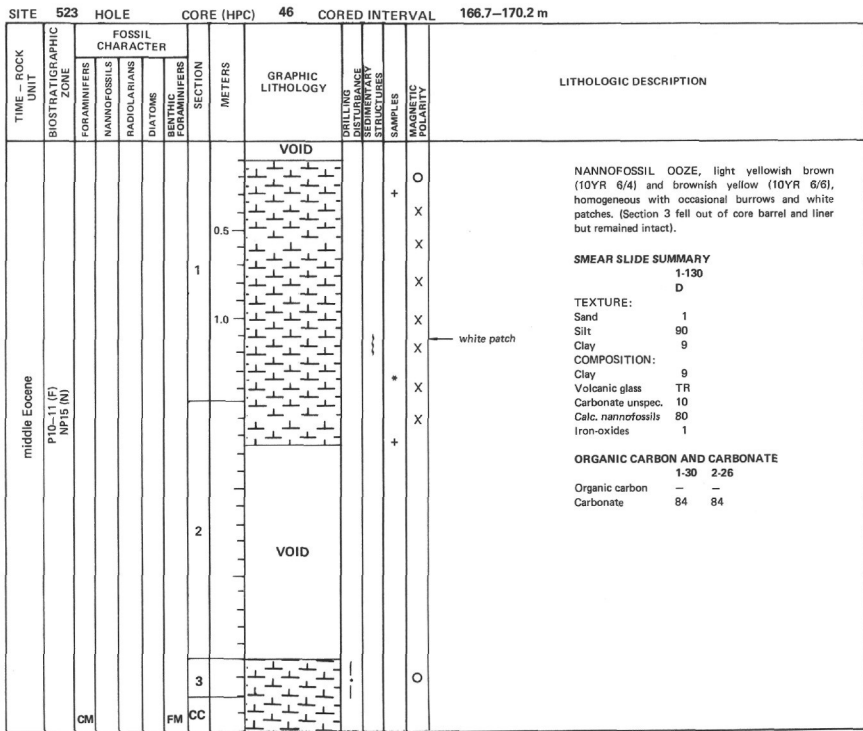












SITE 523		HOLE CORE (HPC) 48		CORED INTERVAL 180.6-182.1 m							
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION	METERS	GRAPHIC LITHOLOGY	DISURBANCE	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION	
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS							DIATOMS
middle Eocene	P10-11 (F) NP15 (N)	CM				1			+	X	<p>NANNOFOSSIL OOZE, light yellowish brown (10 YR 6/4), slightly burrowed, homogeneous.</p> <p>SMEAR SLIDE SUMMARY 1-116 D</p> <p>TEXTURE: Sand 3 Silt 87 Clay 10</p> <p>COMPOSITION: Clay 10 Carbonate unspec. 5 Foraminifers 3 Calc. nannofossils 82 Iron-oxides TR</p> <p>ORGANIC CARBON AND CARBONATE Organic carbon 1.95 2.35 Carbonate 84 86</p>

SITE 523		HOLE CORE (HPC) 49		CORED INTERVAL 182.1-186.5 m													
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER			SECTION	METERS	GRAPHIC LITHOLOGY	DISURBANCE	MAGNETIC POLARITY	LITHOLOGIC DESCRIPTION							
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS							DIATOMS						
middle Eocene	P10-11 (F) NP15 (N)	CM				2				<p>NANNOFOSSIL OOZE, light yellowish brown (10YR 6/4) to very pale brown (10YR 7/4-7/3), partly lithified to MARL in places. Thin layer of FORAMINIFER-NANNOFOSSIL OOZE in Section 3 has sharp basal contact and gradational top. Burrows common in Section 2. Contorted laminations and microfaults at 100-110 cm, Section 2.</p> <p>SMEAR SLIDE SUMMAR 3-83 D</p> <p>TEXTURE: Sand 2 Silt 85 Clay 13</p> <p>COMPOSITION: Clay 10 Palagonite TR Carbonate unspec. 15 Foraminifers 2 Calc. nannofossils 73</p> <p>ORGANIC CARBON AND CARBONATE Organic carbon 1.32 3.30 3.87 Carbonate 85 87 87</p>							
											RM	CC	3				X

SITE 523 HOLE		CORE (HPC) 50		CORED INTERVAL 186.5-190.5 m										
TIME - ROCK UNIT	BIOSTRATIGRAPHIC ZONE	FOSSIL CHARACTER				SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	SEDIMENTARY STRUCTURES	SAMPLER	METERS	POLARITY	LITHOLOGIC DESCRIPTION
		FORAMINIFERS	NANNOFOSSILS	RADIOLARIANS	DIATOMS									
middle Eocene	P10-11 (F) NP15 (N)						0.5	VOID						yellowish brown very pale brown light olive brown VOID
							1.0							very pale brown yellowish brown gray halo around burrow very pale brown
							2.0							light yellowish brown yellowish brown very pale brown
						3.0	VOID						light yellowish brown yellowish brown white patch light yellowish brown mixed very pale brown and gray	
	CM													
		RM												
		CC												

NANNOFOSSIL OOZE and MARLY NANNOFOSSIL OOZE, yellowish brown (10YR 5/6), light yellowish brown (10YR 6/4) to very pale brown (10YR 7/3); several intervals lithified to MARL in Section 1 - one piece at approximately 25 cm appears nodular and has dark, clay-rich horizontal layer in it. Thin FORAMINIFER-NANNOFOSSIL OOZE with sharp basal contact occurs in Section 2. Base Section 3 and Core-Catcher is mixed very pale brown (10YR 7/3) and gray (pyrite?) ooze.

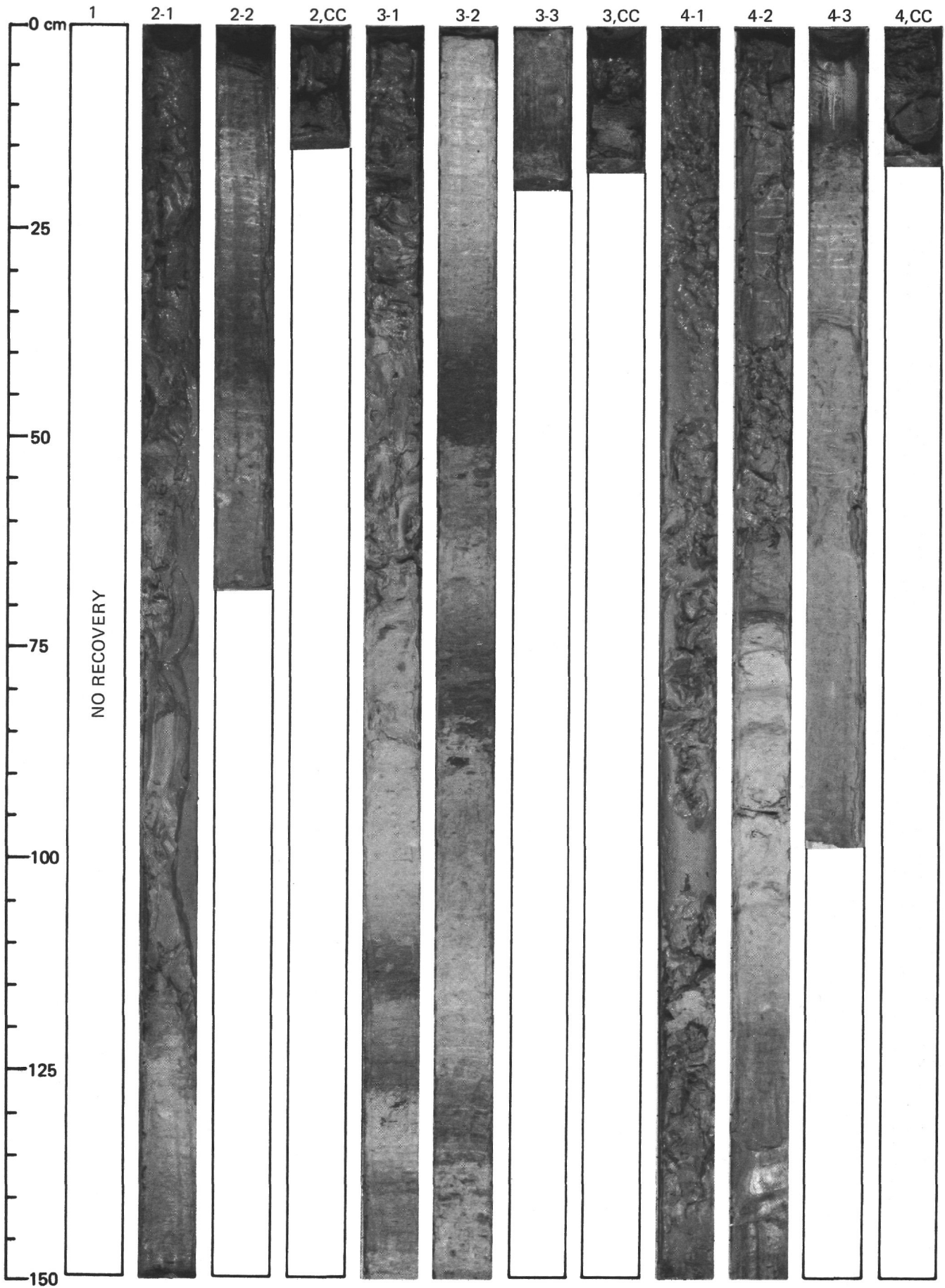
SMEAR SLIDE SUMMARY

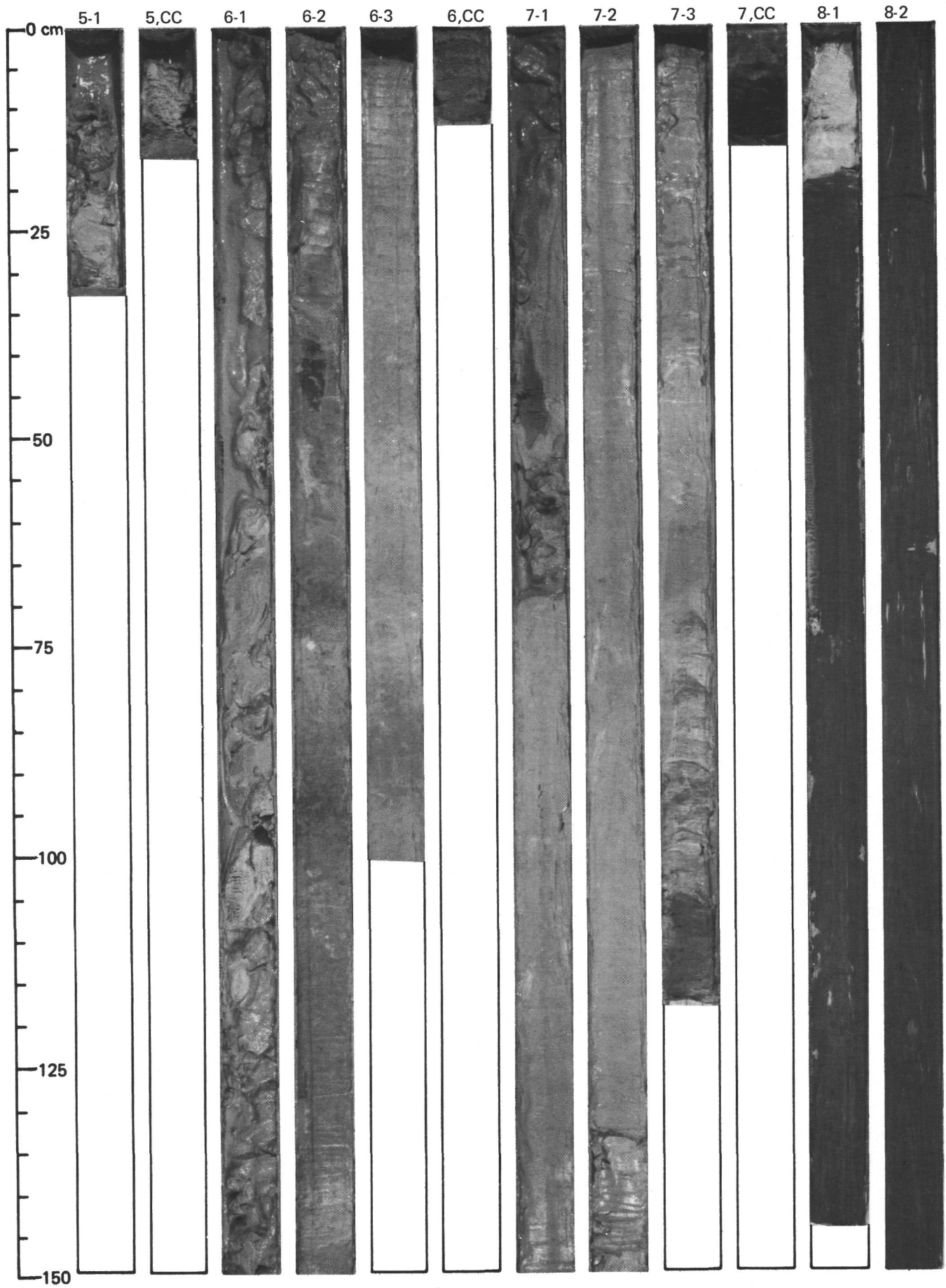
	3-4	3-29
TEXTURE:	D	M
Sand	5	5
Silt	85	85
Clay	10	10
COMPOSITION:		
Clay	10	10
Palagonite	TR	-
Carbonate unsp. spec.	25	25
Foraminifers	5	5
Calc. nannofossils	60	60

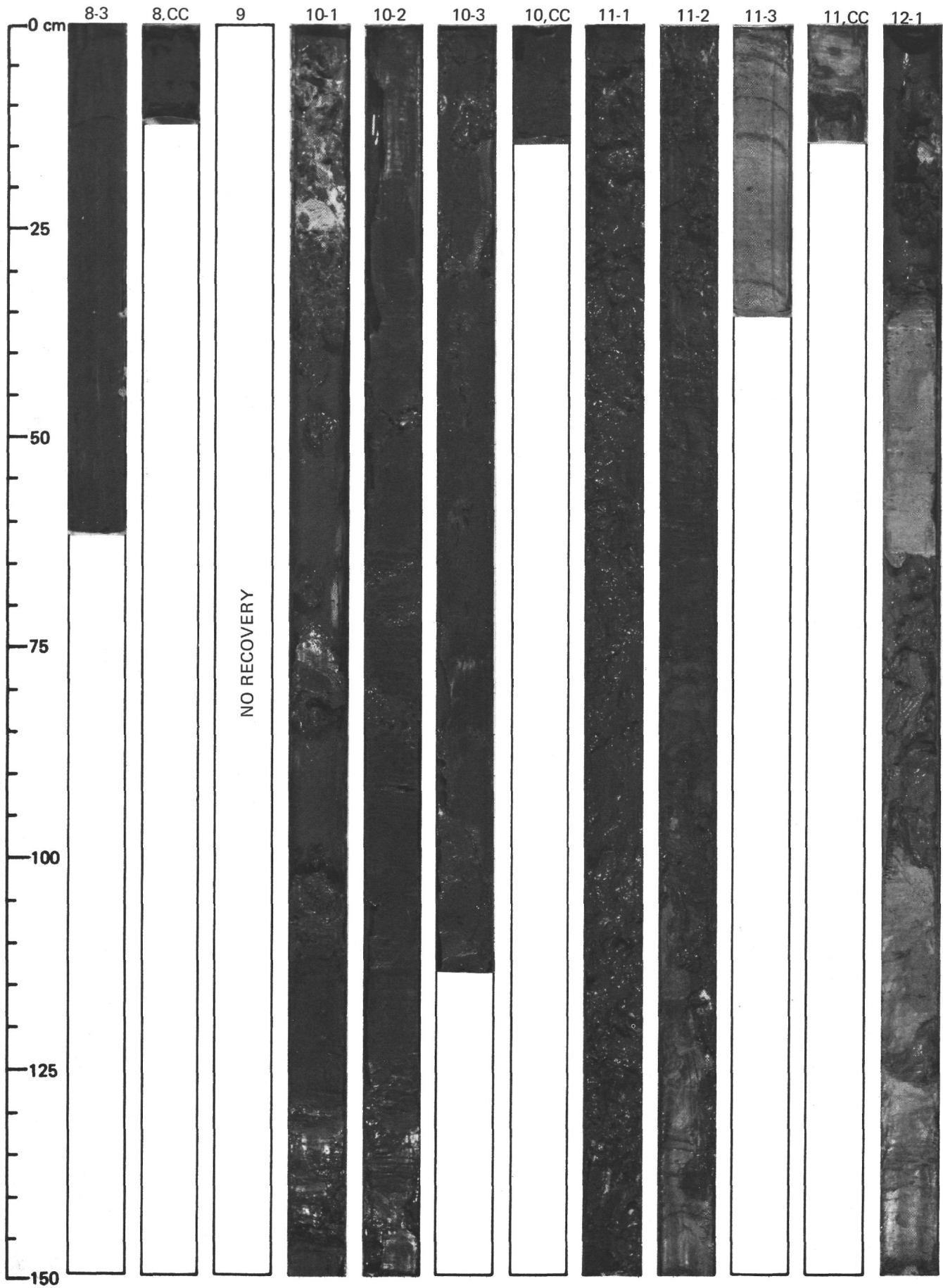
ORGANIC CARBON AND CARBONATE

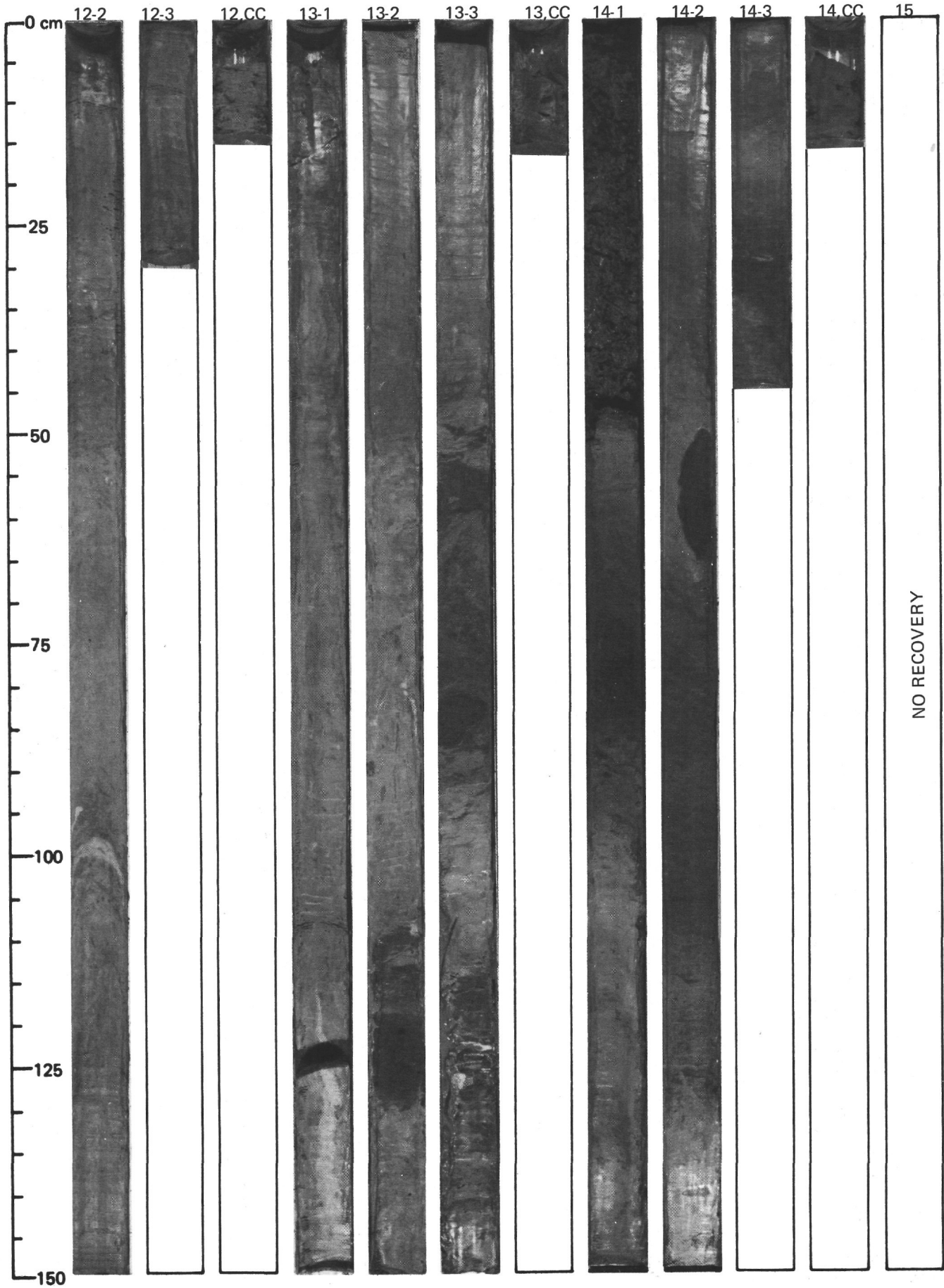
	1-79	2-79	3-29
Organic carbon	-	-	-
Carbonate	88	87	84

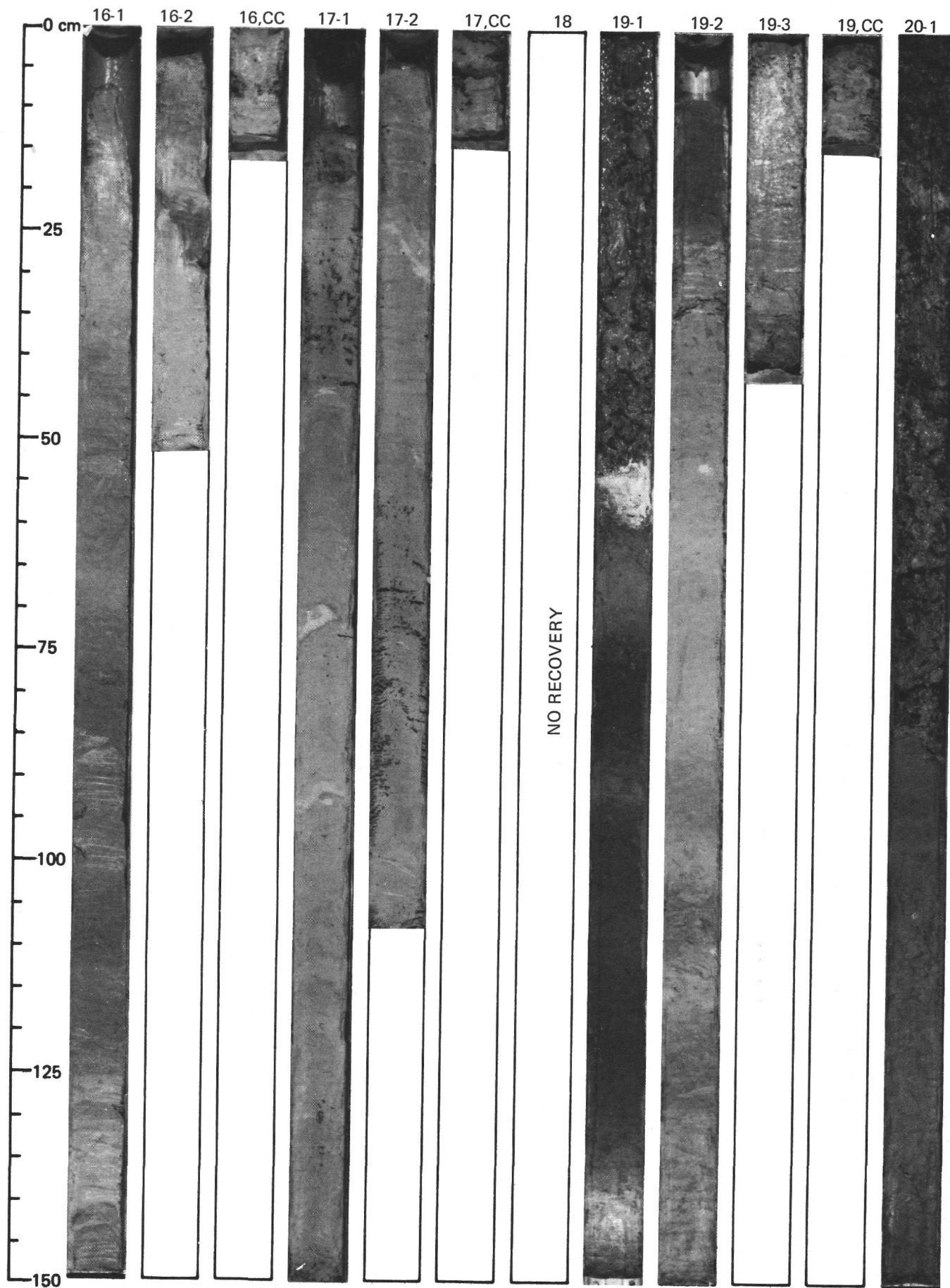
NOTE: Core 51, 193.5-197.8 m: No recovery.

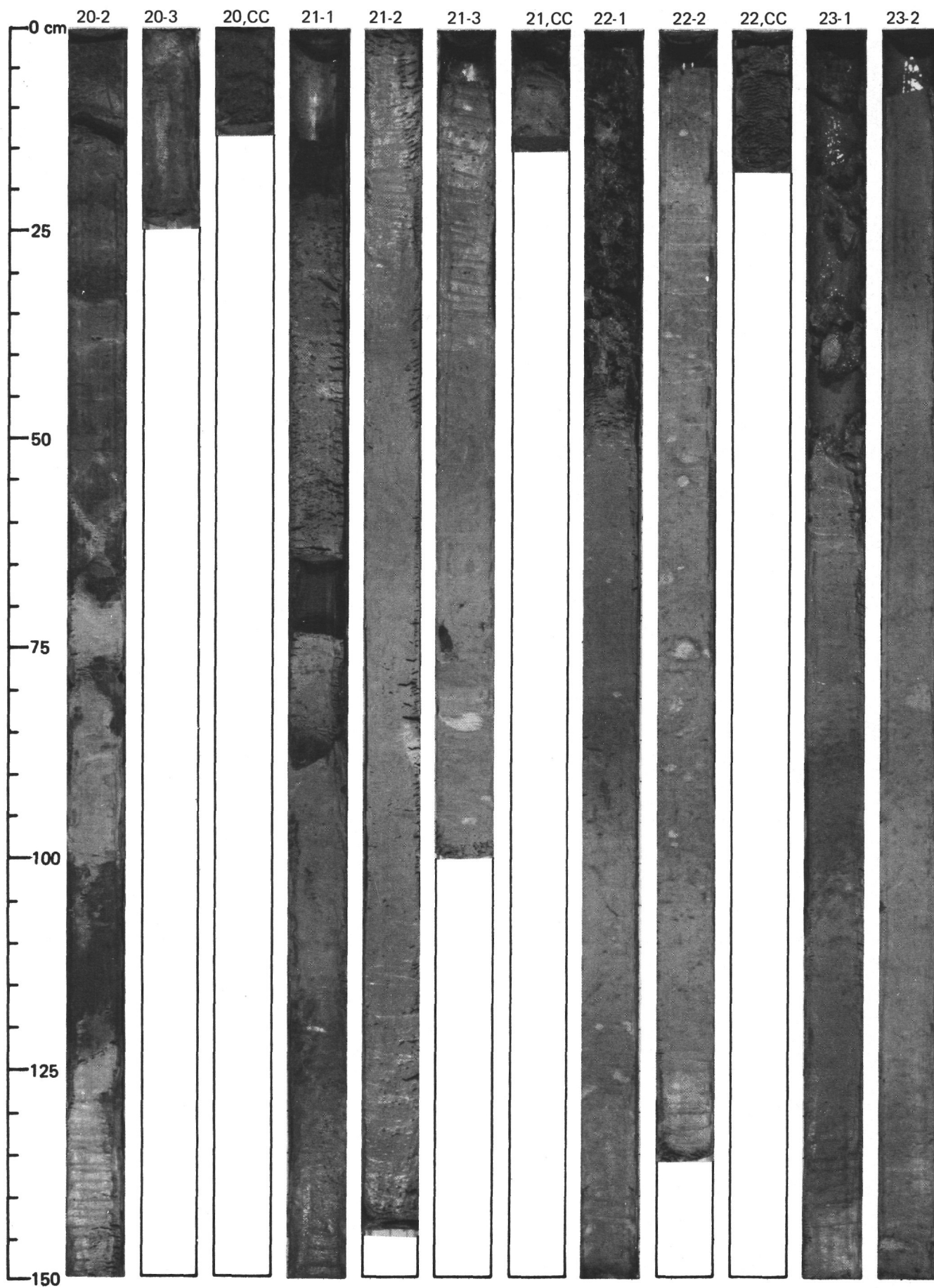


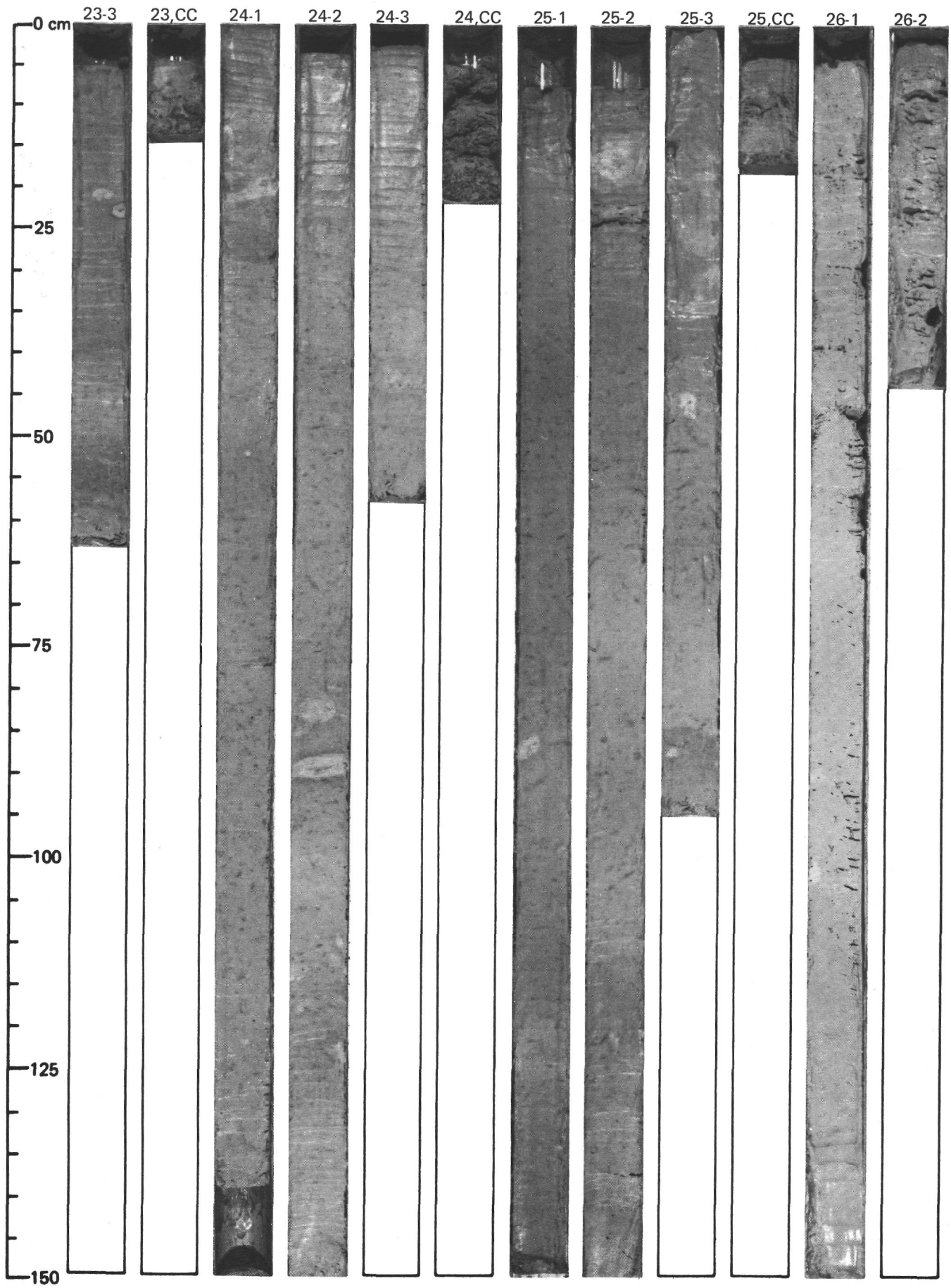


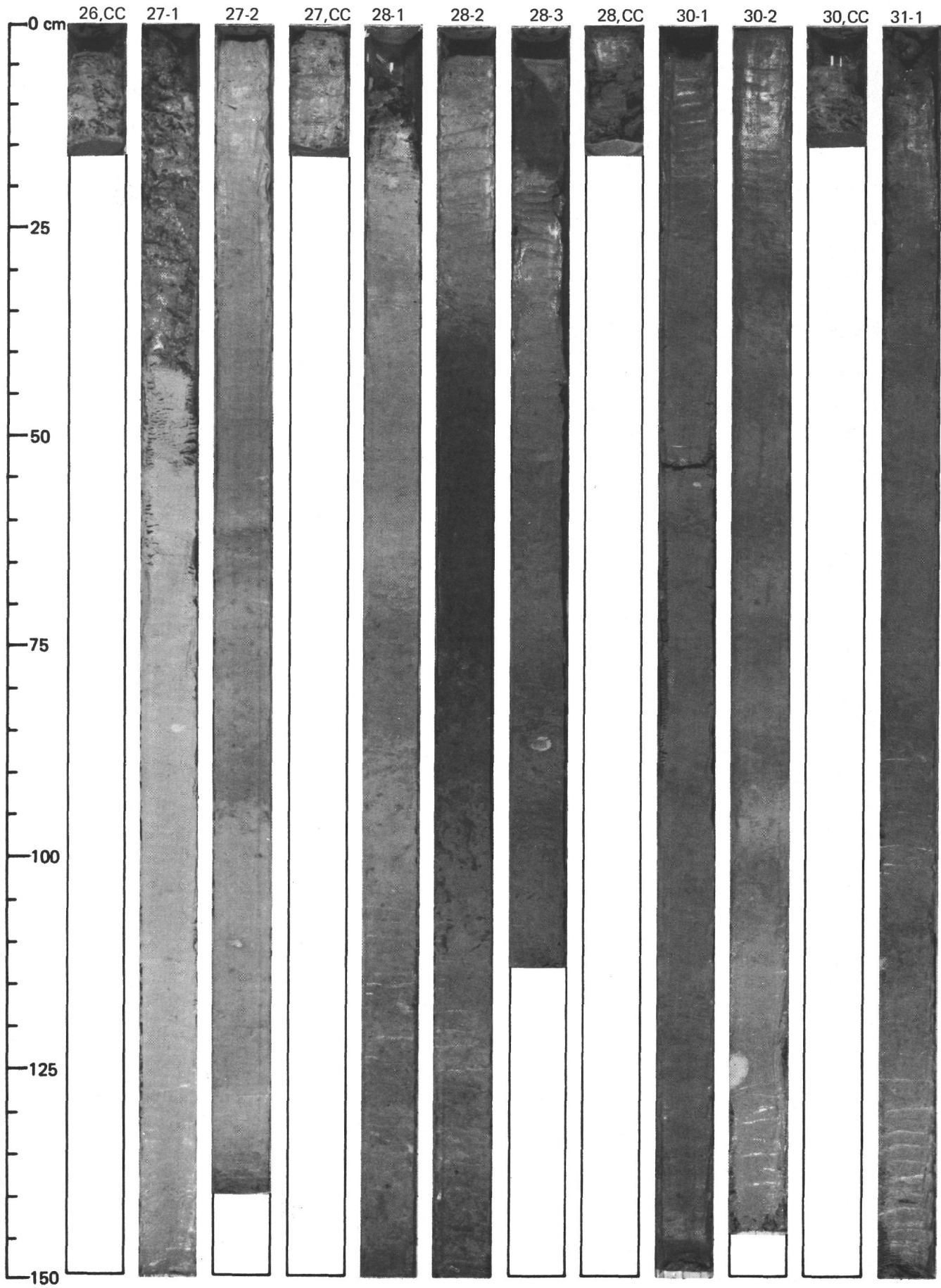


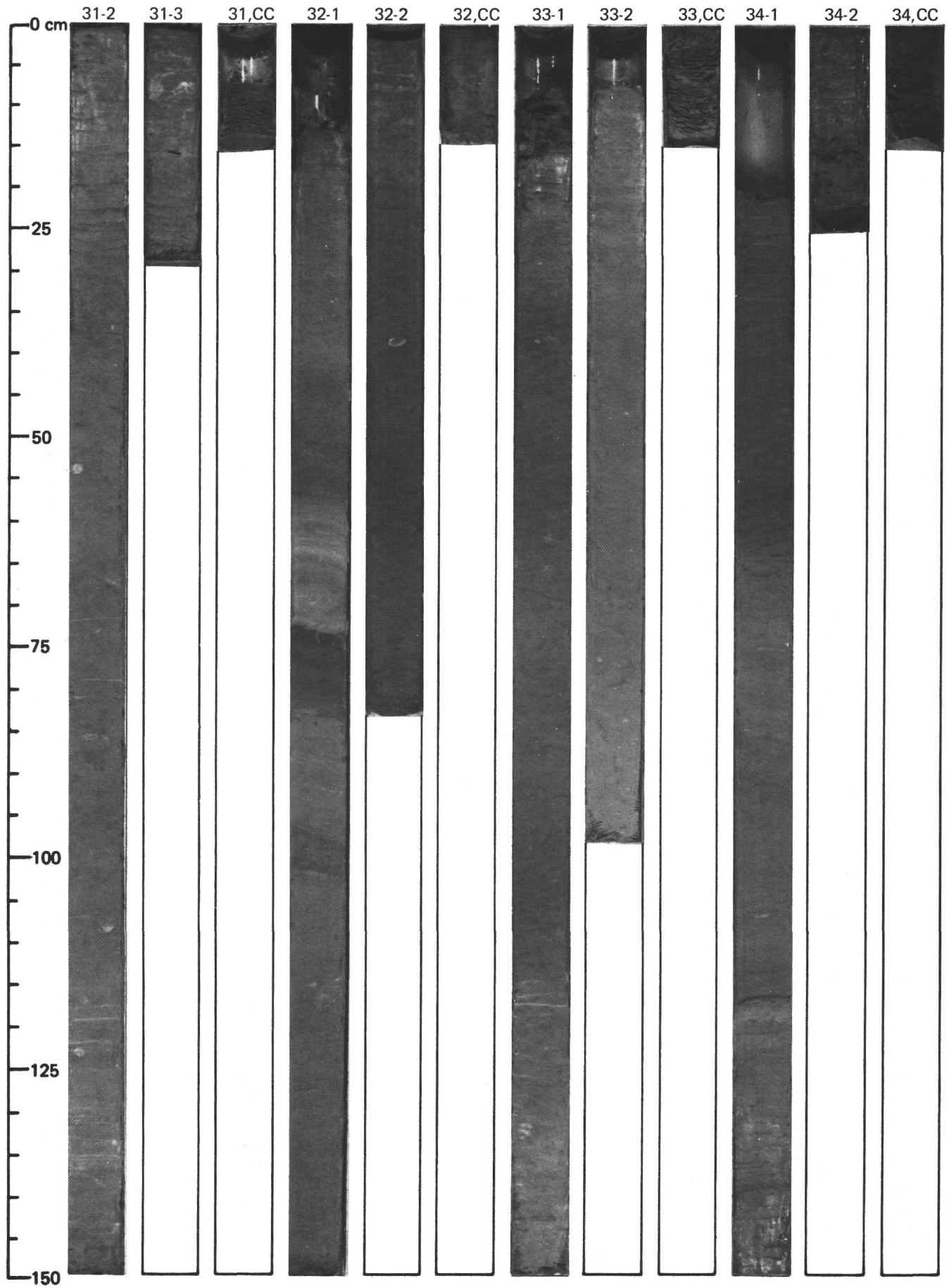


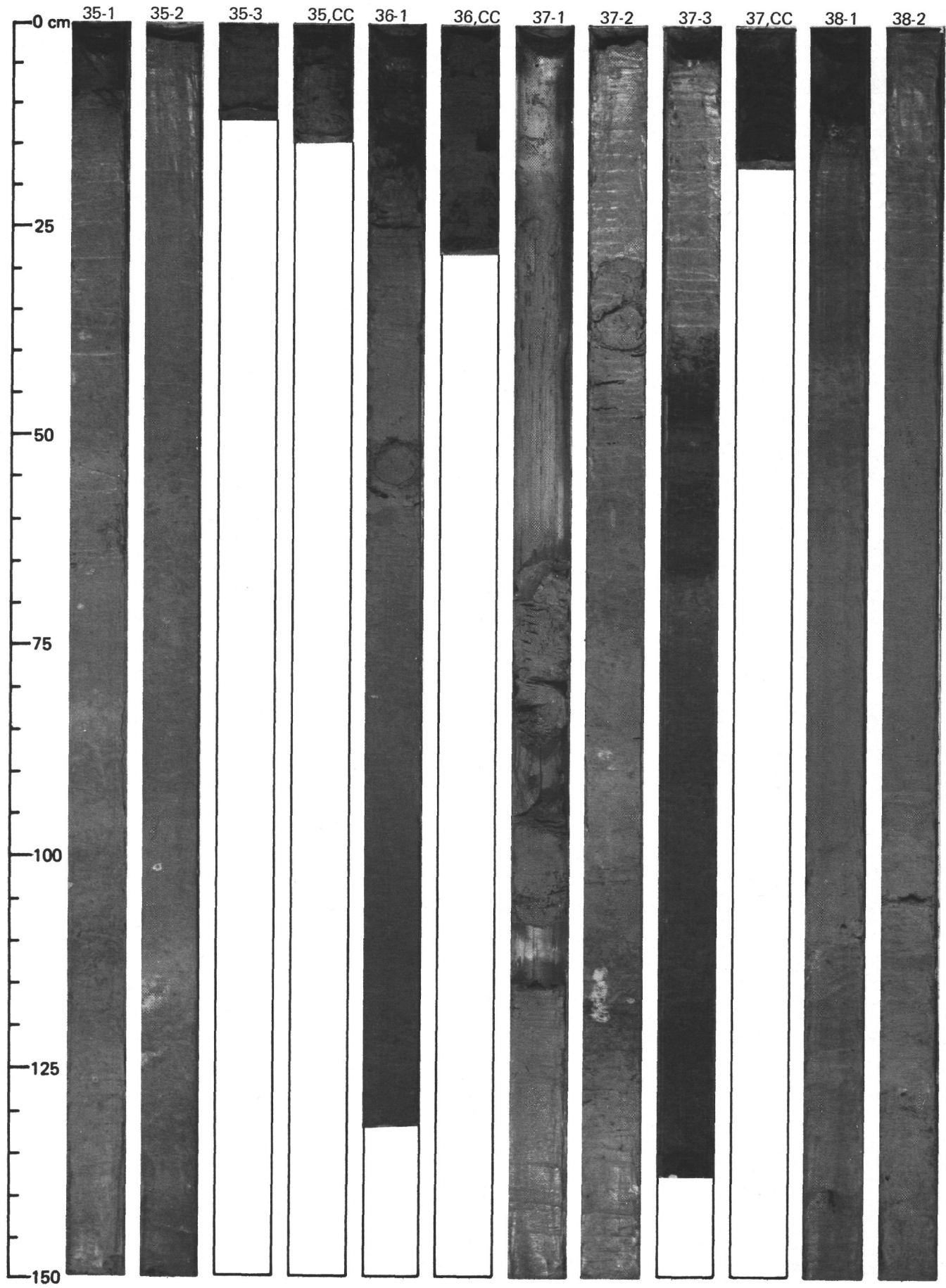












SITE 523 (HOLE 523)

