INTRODUCTION

Thirty-seven frozen core sections were obtained through the JOIDES Organic Geochemistry Panel for organic geochemical study. These samples consisted of Pleistocene through Jurassic (upper Tithonian) sections from Holes 388A, 390, 390A, and 391A, 391B, 391C from the lower continental rise hills, the Blake Nose, and Blake-Bahama Basin on the outer North Atlantic continental margin. Sample site locations are shown in Figure 1.

SAMPLING AND STUDY PROCEDURES

The flat faces of the frozen half-sections were smoothed by milling upon arrival in our laboratory. Because of space limitations, black and white photographs for only those samples that have been extensively disturbed and mixed as a consequence of drilling, namely, Sections 390-3-3, 390A-6-0, 391C-9-3, and 391C-12-4 are shown in Figure 2. Each section was then described and chips of each textural type removed to determine the content of organic and carbonate carbon. These values are listed in Table 1.

The remainder of the core section was freed as thoroughly as possible of contaminating mud and characterized as shown in the flow sheet in Figure 3. The sample was freeze dried, crushed, and sieved to 150 µm. A 2-gram portion of homogenized core was analyzed for organic and carbonate carbon and for ammonium and organic nitrogen. The remaining freeze-dried material was extracted by treatment with methylene chloride for 24 hours. The methylene chloride was removed yielding an oil extract which hereafter is designated as the lipid fraction. When sufficient sample was available, elemental analyses were performed for lipid carbon, hydrogen, nitrogen, and sulfur.

The lipid fraction was chromatographed over activated silica gel using as eluants, pentane, methylene chloride, and a 50:50 mixture of methylene chloride methanol to obtain "saturate," "aromatic," and "asphaltic" fractions. The n-alkane portion of the saturate fraction was adducted in urea and analyzed by gas chromatography.

Carbon isotopic compositions were determined by closed loop combustion on an AVCO isotope ratio mass spectrometer. Measurements were made on the lipid and its aromatic and asphaltic subfractions, where possible, and on the kerogen fraction.

RESULTS

Average numerical data generated for the samples along with sub-bottom depth and age are provided in Table 2. Characterization of the lipid fractions is given in Table 3. The n-alkane distributions for those cases where sufficient sample was available are given in Figures 4 through 9.

Hole 388A

The Miocene hemipelagic clays from the lower continental rise hills from Hole 388A ranged in sub-bottom depth from 246 to 332 meters. The sedimentary section is relatively uniform in lithology. Organic content averages 0.97 ± 0.1 per cent; carbonate content, less than 0.13 per cent; and organic nitrogen content averages 757 ± 48 ppm for all but the 298-meter section which is 2554 ppm.

The amount of the total sediment as lipid ranges from 250 ppm for the bottom section then decreases to 50 ppm at the top. The ratio of lipid to total organic carbon is high, namely 2.9 per cent, at the bottom of the unit and is low and constant at 0.7 ± 0.1 per cent, for the remainder of the section. The bottom section contained 45.8 per cent of the total lipid as saturates; whereas, the remaining section is highly asphaltic, i.e., 48.4 to 64.5 per cent of the total lipid.
The carbon isotopic composition of the lipid is -27.8 $\delta^{13}C_{\text{PDB}}$ in the upper Albian sediments to a nearly constant value of -26.7 ±0.3 $\delta^{13}C_{\text{PDB}}$ in the Maastrichtian to Eocene sediments. The isotopic composition of the kerogen varies from -22.9 to -26.9 $\delta^{13}C_{\text{PDB}}$ without apparent trends with depth. Because of the small amount of lipid, we are unable to characterize it further.

**Holes 391, 391A, 391B, and 391C**

Holes 391, 391A, 391B, and 391C lie in an abyssal area in the Blake-Bahama Basin in the western North Atlantic Ocean. Five lithologic units were recognized from which samples were obtained from each.

Unit 5 consisted of Jurassic (upper Tithonian) claystone and limestone at a depth of 1393 and 1355 meters, respectively. The lower claystone is rich in organic matter with 0.81 per cent organic carbon whereas the limestone section is poor in organic matter with only 0.1 per cent organic carbon. The opposite is true for the proportion of organic carbon present as lipid. The limestone contains 2.15 per cent as opposed to 0.09 per cent in the claystone. The samples in the order of decreasing depth, contained 2.6 per cent and 11.1 per cent carbonate carbon or 21.9 per cent and 92.8 per cent as calcium carbonate, 124 and 42 ppm ammonium nitrogen, 176 and 58 ppm organic nitrogen, and 10 to 20 ppm lipid carbon.

The carbon isotopic composition of these lipid fractions are -27.3 and -26.5 $\delta^{13}C_{\text{PDB}}$ and -22.0 and -20.7 $\delta^{13}C_{\text{PDB}}$ for the kerogens.

Unit 4 consisted of upper Tithonian-lower Berriasian to upper Berriasian-lower Valanginian limestones. Five cores were obtained from between 1279 to 1022 meters depth. The bottom of the interval contains very little organic matter; the organic carbon content is 0.2 per cent which increases to 1.1 per cent at the top of the unit. The sample contained 9.7 per cent to 6.0 per cent carbonate carbon or from 80.6 per cent to 49.8 per cent as calcium carbonate, 40 to 97 ppm ammonium nitrogen, and 51 to 303 ppm organic nitrogen with the organic nitrogen increasing to the top of the section. The lipid carbon increases from a low of 10 ppm at the bottom of the unit to 790 ppm at 1185 meters, then decreases to 50 ppm at the top of the unit. The ratio of lipid to total organic carbon is low at the bottom of the unit, namely 0.62 per cent, increases to a high of 23.5 per cent at 1234 meters, then decreases steadily to 0.41 per cent at the top of the interval.

The isotopic composition of the lipid fractions is steadily lighter from the bottom to the top of the unit from -25.6 to -29.1 $\delta^{13}C_{\text{PDB}}$. The kerogen isotopic composition varies from a low of -21.6 to a high of 26.3 $\delta^{13}C_{\text{PDB}}$ with no trends with depth. Enough lipid was obtained from core-sections 391C-35-4, 391C-30-2, and 391C-26-3 for its full characterization. As shown in Table 3, all three sections were similar in their lipid properties. The saturate, aromatic, and asphaltic fractions averaged 2.8, 84.7, and 12.5 per cent respectively. The isotopic composition of the aromatic and asphaltic fractions averaged -26.8 ±0.6 $\delta^{13}C_{\text{PDB}}$. The atomic ration of hydrogen to carbon is 1.64 ±0.02.
### Table 1
Lithology of Core-Sections From DSDP Leg 44
Description and Carbonate and Organic Carbon Content of Different Textural Types

<table>
<thead>
<tr>
<th>Hole-Core-Section</th>
<th>Lithological Description</th>
<th>Textural Unit</th>
<th>Carbonate as Carbon (wt. %)</th>
<th>CaCO$_3$ (wt. %)</th>
<th>Organic Carbon (wt. %)</th>
</tr>
</thead>
<tbody>
<tr>
<td>388A-5-0</td>
<td>Clay, dark greenish gray (5Gy4/1, 5G 4/1), faintly mottled, stiff (unit no. 1); contains hard sideritic (?) concentrations (burrow fillings ?), concentrated in indistinct layers (units 2 and 3), some with soft microcrystalline sideritic (?) material; pyrite is common; composition 65%-90% clay minerals, 10%-35% quartz, up to 10% mica, and traces of feldspar</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.13</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>&lt;0.1</td>
<td>0.83</td>
<td>0.53</td>
</tr>
<tr>
<td></td>
<td></td>
<td>3</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>0.99</td>
</tr>
<tr>
<td>388A-7-3</td>
<td>Clay, dark greenish gray (5G4/1), faintly mottled, stiff, with rare siderite (?) concretions, composition 85%-95% clay minerals, 5%-15% quartz, traces of heavy minerals (rutile), chlorite, carbonate rhombs, chert, benthic foraminifers, and calcareous nannofossils</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.02</td>
</tr>
<tr>
<td>388A-7-4</td>
<td>Clay, dark greenish gray (5GY4/1), faintly mottled, stiff with rare siderite (?) concretions, composition 85%-95% clay minerals, 5%-15% quartz, traces of heavy minerals (rutile), chlorite, carbonate rhombs, chert, benthic foraminifers, and calcareous nannofossils</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.08</td>
</tr>
<tr>
<td>388A-11-0</td>
<td>Clay, grayish olive (10Y4/2) (unit no. 1) soft, plastic with dusky red (5R3/4) siderite on black (N1) pyrite (?) concretions (unit no. 2) scattered throughout sediment; sediment is badly disturbed in first 13 cm (unit a) and partially disturbed in the remaining 24 cm (units b and c) with mixture of soft grayish orange (20YR7/4) (unit no. 3) and firm fragments of drilling slurry; composition 80%-90% clay minerals, 5%-10% quartz, and traces of feldspar, mica, heavy minerals, chlorite, zeolites, carbonate rhombs, and calcareous nannofossils</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.00</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>0.58</td>
<td>4.83</td>
<td>1.42</td>
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<td></td>
<td>3</td>
<td>7.54</td>
<td>62.83</td>
<td>0.48</td>
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<td>390-3-3</td>
<td>Badly disturbed marly nannofossil ooze, pale blue-green (5BG7/2) (unit no. 1); considerable mixing due to drilling of a marly nannofossil ooze, light reddish brown (5YR6/4) (unit no. 2) with wisps of light brown (5YR5/6) (unit no. 3) admixed</td>
<td>1</td>
<td>7.36</td>
<td>61.33</td>
<td>0.25</td>
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<td></td>
<td></td>
<td>2</td>
<td>6.95</td>
<td>57.92</td>
<td>0.32</td>
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<td></td>
<td>3</td>
<td>6.84</td>
<td>57.00</td>
<td>0.37</td>
</tr>
<tr>
<td>390A-6-0</td>
<td>Nannofossil ooze, light yellowish brown (10YR6/4) (unit no. 1) with very pale brown (10YR8/3 to 10YR8/4) (unit no. 2) banding due to drilling; composition 70% calcareous nannofossils, 10% planktonic foraminifers, 3% benthic foraminifers, 5% clay minerals, 5% zeolites, 5% unidentified minerals</td>
<td>1</td>
<td>7.92</td>
<td>66.00</td>
<td>0.23</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>8.51</td>
<td>70.92</td>
<td>0.29</td>
</tr>
<tr>
<td>390A-10-0</td>
<td>Marly nannofossil ooze (20YR8/2) (unit no. 1) with 1 mm- to 8 mm-size dark olive-gray (5Y3/2) to dark gray (N3) clay filled burrows (unit no. 2)</td>
<td>1</td>
<td>7.90</td>
<td>65.83</td>
<td>0.26</td>
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<tr>
<td></td>
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<td>2</td>
<td>7.24</td>
<td>60.33</td>
<td>0.29</td>
</tr>
<tr>
<td>390A-12-0</td>
<td>Nannofossil ooze, light bluish gray (5B7/1), homogeneous; composition 55%-75% calcareous nannofossils, 15% unspecified carbonate, 25% clay minerals, 5% quartz, 2% zeolites, traces of mica, volcanic glass, heavy minerals, benthic foraminifers, radiolarians, sponge spicules</td>
<td>1</td>
<td>8.45</td>
<td>70.41</td>
<td>0.23</td>
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<td>Hole-Core-Section</td>
<td>Lithological Description</td>
<td>Textural Unit</td>
<td>Carbonate as Carbon (wt. %)</td>
<td>CaCO₃ (wt. %)</td>
<td>Organic Carbon (wt. %)</td>
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<td>------------------</td>
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<tr>
<td>391-1-2</td>
<td>Silty hemipelagic clay, soft, olive-gray (SY4/1) (unit no. 1) with small olive-gray (SY3/2) concretions (unit no. 2); clay unit from 3 to 6 cm from top of core, soft pale yellowish brown (10YR6/2) (unit no. 3)</td>
<td>1</td>
<td>0.51</td>
<td>4.25</td>
<td>1.30</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>1.24</td>
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<td></td>
<td></td>
<td>3</td>
<td>5.71</td>
<td>47.58</td>
<td>0.84</td>
</tr>
<tr>
<td>391A-3-4</td>
<td>Calcareous “silt”, yellowish gray; (SY8/1), homogeneous, stiff</td>
<td>1</td>
<td>11.34</td>
<td>94.50</td>
<td>0.15</td>
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<td>391A-4-3</td>
<td>Chalk, light bluish gray (SY7/1, SY7/2) (unit no. 1) with light gray (SY7/2) (unit no. 2) mortling layers, homogeneous</td>
<td>1</td>
<td>10.56</td>
<td>88.00</td>
<td>0.29</td>
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<tr>
<td></td>
<td></td>
<td>2</td>
<td>10.84</td>
<td>90.33</td>
<td>0.23</td>
</tr>
<tr>
<td>391A-5-0</td>
<td>Marly chalk pale olive (10Y6/2) (unit no. 1) with 1 mm to 10 mm size dark greenish gray (S4/1, S4/1) and olive-gray (S4/1) clasts (unit no. 2); clasts are mostly green siliceous clay; composition (chalk) 50% unspecified carbonate, 15% calcareous nannofossils, 35% clay minerals, traces quartz, feldspar, mica, dolomite, siliceous microfossils</td>
<td>1</td>
<td>9.34</td>
<td>77.83</td>
<td>0.55</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>2.51</td>
<td>20.92</td>
<td>1.48</td>
</tr>
<tr>
<td>391A-6-3</td>
<td>Marly chalk, yellowish gray (SY7/2) (unit no. 1) with small (3 mm) greenish gray (SY4/1) and olive-gray (SY4/1) clasts (unit no. 2) clasts; composition 40% calcareous nannofossils, 35% unspecified carbonates, 19% clay minerals, 3% sponge spicules, traces of quartz and radiolarians</td>
<td>1</td>
<td>10.54</td>
<td>87.83</td>
<td>0.30</td>
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<tr>
<td></td>
<td></td>
<td>2</td>
<td>6.44</td>
<td>53.67</td>
<td>0.90</td>
</tr>
<tr>
<td>391A-9-0</td>
<td>Siliceous muddy chalk and marly chalk, pale olive (10Y6/2) (unit no. 1) with 1 to 20-mm size greenish gray (SY3/2) (unit no. 2) clay clasts; most clasts are elongated parallel to bedding; composition 15% calcareous nannofossils, 20% unspecified carbonate, 30% radiolarians, 35% clay minerals, 3% sponge spicules, traces of quartz and radiolarians</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.85</td>
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<td></td>
<td></td>
<td>2</td>
<td>8.60</td>
<td>71.67</td>
<td>0.46</td>
</tr>
<tr>
<td>391A-10-0</td>
<td>Marly chalk, pale olive (10Y6/2) (unit no. 1) with 1 mm to 7 mm size greenish gray (S6/1) and olive-gray (S3/1) clay clasts (unit no. 2); most clasts are elongated parallel to bedding; composition 45% calcareous nannofossils, 20% unspecified carbonate, 30% clay minerals, 3% siliceous microfossils, traces foraminifers and sponge spicules</td>
<td>1</td>
<td>8.43</td>
<td>70.25</td>
<td>0.48</td>
</tr>
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<td>2</td>
<td>0.97</td>
<td>8.08</td>
<td>1.40</td>
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<tr>
<td>391A-11-0</td>
<td>Marly chalk (unit no. 1), greenish gray (SYY6/1) to pale olive-gray (10Y6/2) with 1 mm- to 15 mm-sized greenish gray and olive-gray (SY3/1) clay clasts (unit no. 2); most clasts are elongated parallel to bedding; composition 30% calcareous nannofossils, 25% unspecified carbonate, 25% clay minerals, 6% foraminifers, 25% clay minerals, 10% sponge spicules, 3% radiolarians, traces of quartz, pyrite, zeolites</td>
<td>1</td>
<td>8.64</td>
<td>72.00</td>
<td>0.43</td>
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<td></td>
<td></td>
<td>2</td>
<td>1.50</td>
<td>12.50</td>
<td>1.46</td>
</tr>
<tr>
<td>391A-12-0</td>
<td>Marly chalk (unit no. 1), pale olive (10Y6/2) with 1 mm- to 13 mm-size moderate olive-brown (SY4/4), olive-gray (SY3/2), pinkish gray (SY8/1), and light bluish gray (SB7/1) clay clasts (unit no. 2)</td>
<td>1</td>
<td>9.92</td>
<td>82.67</td>
<td>0.29</td>
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<td></td>
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<td>2</td>
<td>0.90</td>
<td>7.50</td>
<td>1.16</td>
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<tr>
<td>391A-13-0</td>
<td>Mudstone, olive-gray (SY4/2), homogeneous</td>
<td>1</td>
<td>1.43</td>
<td>11.92</td>
<td>2.71</td>
</tr>
<tr>
<td>391A-20-2</td>
<td>Claystone (unit no. 1) banded dark olive (SY3/2) and dark greenish gray (S4/1), fissile; well-defined burrows (unit nos. 2 and 3); composition 95% clay minerals, 1% sponge spicules, traces of quartz, mica, pyrite, unspecified carbonate, radiolarians</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.87</td>
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<td></td>
<td></td>
<td>2</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.56</td>
</tr>
<tr>
<td></td>
<td></td>
<td>3</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.31</td>
</tr>
<tr>
<td>Hole-Core-Section</td>
<td>Lithological Description</td>
<td>Textural Carbon (wt.%)</td>
<td>Carbonate as Carbon (wt. %)</td>
<td>Organic Carbon (wt. %)</td>
<td></td>
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<td>-------------------</td>
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<tr>
<td>391A-21-4</td>
<td>A 13-cm section of silty claystone (unit no. 1), faintly laminated from olive-black (5Y2/1), dark greenish gray (5G4/1), and olive-gray (5Y4/1), contorted due to drilling. Underlain by a 15-cm section of claystone (unit no. 2), grayish olive (10Y4/2), homogeneous, contorted due to drilling; composition 65% clay minerals, 15% mica, 10% altered minerals, 5% quartz, 5% feldspar, traces of heavy minerals.</td>
<td>1 &lt;0.1 &lt;0.83 0.54</td>
<td>2 &lt;0.1 &lt;0.83 0.82</td>
<td></td>
<td></td>
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<tr>
<td>391B-1-1</td>
<td>A 13-cm section of clay (unit no. 1) pale yellowish brown (10YR6/2), soft micaceous, composition 80% clay minerals, 6% unspecified carbonate, 5% quartz, 3% mica, 2% calcareous nannofossils, traces of sponge spicules underlain by a 19-cm section of clay (unit no. 2), dark grayish brown (5YR3/2), soft; slightly contorted due to drilling; composition same as unit 1</td>
<td>1 6.37 53.08 0.79</td>
<td>2 &lt;0.1 &lt;0.83 1.24</td>
<td></td>
<td></td>
</tr>
<tr>
<td>391C-2-0</td>
<td>Chalk (unit no. 1), very light gray (N8) with 1 mm- to 8 mm-sized dark greenish gray (5G4/1) and greenish gray (5G6/1) clasts (unit no. 2).</td>
<td>1 8.50 70.83 0.46</td>
<td>2 5.57 21.42 1.42</td>
<td></td>
<td></td>
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<tr>
<td>391C-6-3, 0-30 cm</td>
<td>Variegated claystone with silt stringers, reddish brown (2.5 YR4/4) (unit no. 1) greenish gray (5G6/1), dark gray (N3), black (N2, N1) (unit no. 2), silty stringers (unit no. 3) are white (N9); burrows (0.5-2.0 mm) occur as light colored silt lenses and stringers or very dark gray (N3) to grayish black (N2) stringers, have random orientation; contains 80% clay mineral, 15% quartz, 4% mica, 1% heavy minerals, trace pyrite.</td>
<td>1 &lt;0.1 &lt;0.83 0.68</td>
<td>2 &lt;0.1 &lt;0.83 1.07</td>
<td>3 &lt;0.1 &lt;0.83 0.34</td>
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<tr>
<td>391C-6-3, 90-115 cm</td>
<td>Claystone, olive-gray (5Y4/1) (unit no. 1) and dark gray (5Y2/1) (unit no. 2), banding with very light gray (N8) silt stringers, (unit no. 3); contains ill-defined burrows (0.5-2 cm) as white stringer and black tubes (chondrites); composition 15% quartz, 4% mica, 1% heavy minerals, trace pyrite.</td>
<td>1 &lt;0.1 &lt;0.83 1.47</td>
<td>2 &lt;0.1 &lt;0.83 0.55</td>
<td>3 &lt;0.1 &lt;0.83 0.43</td>
<td></td>
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<tr>
<td>391C-7-2</td>
<td>Silty claystone, black (2.5Y, N2.5) (unit no. 1) with light gray (5Y7/2) stringer (unit nos. 2 and 3); composition 90% clay minerals, 4% mica, 3% quartz, 2% pyrite.</td>
<td>1 &lt;0.1 &lt;0.83 0.97</td>
<td>2 &lt;0.1 &lt;0.83 1.17</td>
<td>3 &lt;0.1 &lt;0.83 2.28</td>
<td></td>
</tr>
<tr>
<td>391C-8-2</td>
<td>A 9-cm section of claystone (unit no. 1) very dark gray (2.5YR N3) to black (2.5YR, N2.5), fissile, with very thin stringers of silt (unit no. 3); followed by a 15-cm section of siltstone (unit no. 2), olive-gray (5Y5/2) and white (5Y8/1).</td>
<td>1 &lt;0.1 &lt;0.83 1.32</td>
<td>2 &lt;0.1 &lt;0.83 0.60</td>
<td>3 &lt;0.1 &lt;0.83 0.44</td>
<td></td>
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<tr>
<td>391C-9-3</td>
<td>A 21-cm section of silty claystone, black (2.5YR, N2.5) (unit no. 1) slightly fissile, with hard olive-gray (5Y5/2) (unit no. 2) silty beds; badly contorted as a result of drilling; underlain by a 15-cm section; same as unit 1, but less contorted</td>
<td>1 &lt;0.1 &lt;0.83 1.79</td>
<td>2 &lt;0.1 &lt;0.83 2.30</td>
<td></td>
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<tr>
<td>391C-10-3</td>
<td>Claystone, very dark gray (7.5 YR, N3) (unit no. 1) to greenish olive (10Y4/2) (unit no. 2) homogeneous fissile.</td>
<td>1 2.27 18.92 3.16</td>
<td>2 1.76 14.67 0.91</td>
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### TABLE 1 – Continued

<table>
<thead>
<tr>
<th>Hole-Core-Section</th>
<th>Lithological Description</th>
<th>Textural Unit</th>
<th>Carbon (wt. %)</th>
<th>CaCO₃ (wt. %)</th>
<th>Organic Carbon (wt. %)</th>
</tr>
</thead>
<tbody>
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<td>391C-11-2</td>
<td>Silty calcareous claystone, olive gray (5Y3/2) (unit no. 1), light olive (5Y5/2), greenish gray (5Y6/1), and olive-black (5Y2/1) (unit no. 2) laminated, composition 84% clay minerals, 10% calcarcous nanofossils, 4% mica, 1% quartz, 1% heavy minerals</td>
<td>1</td>
<td>7.20</td>
<td>60.00</td>
<td>0.51</td>
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<td>2</td>
<td>8.69</td>
<td>72.42</td>
<td>0.50</td>
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<tr>
<td>391C-12-4</td>
<td>Claystone, grayish black (N2) (unit no. 1), with moderate brown (5YR4/4) (unit no. 2), olive-gray (5Y3/2) (unit no. 3), and light gray (N7) (unit no. 4) mortling; badly disturbed by drilling</td>
<td>1</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
<td>1.45</td>
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<tr>
<td></td>
<td></td>
<td>2</td>
<td>&lt;0.1</td>
<td>&lt;0.83</td>
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</tr>
<tr>
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<td>3</td>
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<td>44.17</td>
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<td></td>
<td>4</td>
<td>10.80</td>
<td>90.00</td>
<td>0.46</td>
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<tr>
<td>391C-16-1</td>
<td>Limestone with shale partings and sandy siltstone beds, gray (5Y5/1) to olive gray (5Y5/2) to black (5Y2.5/1); burrowing bioturbation, cross laminations; composition 73% clay minerals, 15% calcarcous nanofossils, 5% unspecified carbonate, 2% quartz, 2% mica, 2% pyrite, 1% feldspar, trace heavy minerals</td>
<td>1</td>
<td>2.62</td>
<td>21.83</td>
<td>0.97</td>
</tr>
<tr>
<td>391C-26-3</td>
<td>A 3-cm section of limestone, light gray (N7), laminated (unit no. 1). Underlain by a 2-cm section of limestone, olive-gray (5Y5/2) (unit no. 2) with shale laminae; some burrowing olive black (5Y2/1); then a 3-cm section of unit no. 1.</td>
<td>1</td>
<td>7.45</td>
<td>62.08</td>
<td>0.62</td>
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<tr>
<td></td>
<td></td>
<td>2</td>
<td>5.68</td>
<td>47.33</td>
<td>0.89</td>
</tr>
<tr>
<td></td>
<td></td>
<td>3</td>
<td>2.83</td>
<td>23.75</td>
<td>0.56</td>
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<tr>
<td>391C-30-2</td>
<td>A 5-cm section of laminated clayey limestone (unit no. 1), light gray (5Y6/1) to light olive-gray (5Y5/2) with dark (1 mm) laminae of olive-gray (5Y3/2) claystone; underlain by a 1-cm section of claystone (unit no. 2), olive-gray (5Y3/2), fissile, 10 mm band</td>
<td>1</td>
<td>7.67</td>
<td>63.92</td>
<td>0.90</td>
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<tr>
<td></td>
<td></td>
<td>2</td>
<td>1.84</td>
<td>15.33</td>
<td>1.22</td>
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<tr>
<td>391C-35-4</td>
<td>Clayey limestone (unit no. 1), light bluish gray (5B7/1) to light gray (N7); with 1 mm to 10 mm olive-gray (5Y3/2) clay clasts (unit no. 2); composition 50% clay minerals, 20%-30% unspecified carbonate, 20%-30% calcarcous nanofossils, trace quartz, feldspar, mica, pyrite</td>
<td>1</td>
<td>7.67</td>
<td>63.92</td>
<td>0.35</td>
</tr>
<tr>
<td></td>
<td></td>
<td>2</td>
<td>4.22</td>
<td>35.17</td>
<td>2.40</td>
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<tr>
<td>391C-40-2</td>
<td>Limestone (unit no. 1), white (7.5YR, N8) with dark greenish (SG4/1) very irregular clay stringers (unit no. 2); calcispheres (calcified radiolarians [?])</td>
<td>1</td>
<td>10.06</td>
<td>83.83</td>
<td>0.07</td>
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<tr>
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<td></td>
<td>2</td>
<td>4.70</td>
<td>39.17</td>
<td>0.44</td>
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<tr>
<td>391C-48-2</td>
<td>Limestone, pale yellowish brown (10YR6/2), mottled to pale brown (5YR5/2), homogeneous, hard</td>
<td>1</td>
<td>T/S</td>
<td>T/S</td>
<td>T/S</td>
</tr>
<tr>
<td>391C-52-2</td>
<td>A 9-cm section of claystone, dark reddish brown (5YR3/2) (unit no. 1); underlain by a 3-cm section of claystone (unit no. 2) greenish gray (5G6/1) to dark greenish gray (SG4/1) banded</td>
<td>1</td>
<td>2.87</td>
<td>23.92</td>
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<td></td>
<td>2</td>
<td>2.78</td>
<td>23.17</td>
<td>0.77</td>
</tr>
</tbody>
</table>

*a* As noted in the sampling and study procedures section, textural units refer to physically different units noted on the milled half section of the core. These units are noted in the Lithological Descriptions. For homogeneous cores, chips taken for organic and carbonate carbon analyses were taken from the center of the core.
The carbon isotopic composition values for the lipid fractions of the 959 to 727 meter interval are similar, i.e., -28.5 ± 0.4 13C 0, with the exception of the 897.6-meter interval which is one del unit lighter than this average. The upper three core sections are also similar, is, -26.8 ± 0.3 13C 0. Carbon isotopic composition values of the kerogen are typically heavier than the lipid and range randomly from -21.9 to -25.9 13C 0. The algebraic difference between lipid and kerogen is large and negative for the entire unit suggesting a normal sequence of diagenesis.

Enough lipid was available from four of the nine core sections for a complete lipid analysis. The sediments toward the bottom of the unit contained slightly more saturates and a larger asphaltic fraction than the upper part. As shown in Table 3, the carbon isotopic composition of the aromatic and asphaltic fractions of the lower portion of this unit is lighter than in the upper portion.

Elemental analysis for carbon, hydrogen, nitrogen, and sulfur was completed on the lipid fraction from Sections 391-10-3 and 391-6-3. Lipid from Section 10-3 contained more elemental carbon and less oxygen than Section 6-3. However, the ratio of hydrogen to carbon lipids is within experimental error, that is, 1.64 ±0.001.

The 1-alkane distribution curves for the four core sections studies are shown in Figure 6. The general shapes of the various curves are the same. As shown in Table 3, the OEP values for these sections vary randomly with slightly higher values toward the bottom of the unit.

Unit 2 consisted of Miocene intraclastic chalk from 642 to 151 meters. Eleven samples were obtained from this unit from Holes 391A and 391C. The lithologies ranged from predominantly chalk to mudstone and claystone. The organic carbon content for eight of the eleven cores fell at or below 0.5 per cent. The carbonate content is relatively high averaging 8.9 ±2.1 per cent carbonate carbon or 74 ±15 per cent as calcium carbonate for all but the deepest lying sample, 391A-20-2, which contained less than 0.1 per cent carbonate carbon. The ammonium nitrogen content of the kerogen for the Miocene interval of Hole 391A decreases from the bottom to top of the unit from a high of 220 ppm to a low of 79 ppm. The organic nitrogen content for Hole 391A generally decreases from the oldest to the youngest sediments within this unit: from a high of 847 ppm to a low of 25 ppm. The ammonium and organic nitrogen content of Section 391C-2-0 is more like those values from unit 4 within Hole 391C than those of unit 3 from Hole 391A.

The concentration of lipid carbon relative to the dry weight of sample averages 120 ppm and ranges from a low of 20 ppm to a high of 450 ppm. The portion of organic carbon as lipid carbon averages 2.1 per cent with a range of 0.57 to 4.08 per cent. No trends in either the lipid carbon to total sediment or total organic carbon are observed with depth.

The average value of the saturate, aromatic, and asphaltic fractions in weight percent and where possible the corresponding carbon isotopic composition values in δC 13 units for these sediments are 14, 27(-24.9),
J. G. ERDMAN, K. S. SCHORNO

59 (-23.6). Only slight variations from these averages were observed throughout the unit (Table 3).

The elemental carbon content of the total lipid fraction increases from 74.9 per cent to 77.2 per cent with age or depth. Similarly, the ratio of hydrogen to carbon increases from 1.68 to 1.82 with age and depth, whereas the carbon content and H/C ratio of the lipid of Section 391A-13-0 are lower, namely 73.2 and 1.74, respectively, than the previous sections. The sulfur content of all four samples, however, steadily increases with increased depth.

As shown in Figure 7, the n-alkane distributions for Sections 391A-4-3, 391A-6-3, 391A-9-0, 391A-13-0, and 391C-2-0 are similar, with major peaks at n-C_{27}, n-C_{29}, n-C_{31} with n-C_{39} predominating. The corresponding OEP values for these samples are high, averaging 2.3

Both the shape of these curves and the OEP values indicate that petroleum genesis has not been extensive. The n-alkane distributions for Sections 391A-5-0 and 391A-11-0 are given in Figure 8. Section 391A-5-0 contains a bimodal distribution with a relatively smooth maxima at n-C_{27} and a typical recent sediment odd-predominance maxima at n-C_{31}. The OEP values for Sections 391A-5-0 and 391A-11-0 both suggest more prolific genesis of petroleum.

Unit 1 consists of Pleistocene gray hemipelagic clay. Two samples were obtained at sub-bottom depths of 1.5 and 0.1 meters from Holes 391 and 391B, respectively. Both sediments were poor in organic matter and contained less than 0.6 per cent organic carbon even though chips taken for determination of texture indicated that they were rich in organic matter.
### Table 3

Chemistry of Core-Sections From DSDP Leg 44 Extract Fractions and Their Elemental and Isotopic Composition

<table>
<thead>
<tr>
<th>Hole-Core Section</th>
<th>Lipid/Total Organic C (wt. %)</th>
<th>Elemental Composition</th>
<th>Total Lipid (wt. %)</th>
<th>Lipid Fractions</th>
<th>Aromatic (wt. %)</th>
<th>Asphaltic (wt. %)</th>
<th>OEP</th>
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<td>388A-5-0</td>
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<td>388C-7-3</td>
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<td>74.9</td>
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<td>75.7</td>
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<td>1.9</td>
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<td>0.4</td>
<td>3.2</td>
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<td>391C-13-0</td>
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<td>73.2</td>
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<td>5.8</td>
<td>9.5</td>
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<td>76.0</td>
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- $^6$C/DB: Carbon isotopic composition
- H/C: Hydrogen to carbon ratio
- N/C: Nitrogen to carbon ratio
- O/C: Oxygen to carbon ratio
- S/C: Sulfur to carbon ratio

Figure 4. Plots of normalized N-alkane composition versus carbon number for the upper and middle Miocene Core-Sections 388A-5-0, 388A-7-3, 388A-7-4, and 388A-11-0.

Figure 5. Plots of normalized N-alkane composition versus carbon number for the lower Cretaceous Core-Sections 391C-26-3, 391C-30-2, and 391C-35-4.
Figure 6. Plots of normalized N-alkane composition versus carbon number for the lower Cretaceous Core-Sections 391C-6-3, 391C-10-3, and 391C-12-4.

Figure 7. Plots of normalized N-alkane composition versus carbon number for the Miocene Core-Sections 391A-4-3, 391A-6-3, 391A-9-0, 391A-13-0, and 391C-2-0.

Figure 8. Plots of normalized N-alkane composition versus carbon number for the Miocene Core-Sections 391A-5-0 and 391A-11-0.

Figure 9. Plots of normalized N-alkane composition versus carbon number for the Pleistocene Core-Section 391-1-2.

(Table 1). The deeper sediment contained 2.28 per cent carbonate carbon (19% as calcium carbonate), 50 ppm ammonium nitrogen, 150 ppm organic nitrogen, a large fraction of both lipid carbon to total sediment (640 ppm) and to total organic carbon (12 per cent). In contrast, the shallower sediment from this unit contained more carbonate carbon (5.52 per cent) or 46 per cent as calcium carbonate, much less lipid carbon to total sediment (80 ppm) and approximately one-third the amount of lipid carbon to total organic carbon.

The carbon isotopic compositions of both the lipid and kerogen fractions of the deeper lying sediment were lighter than the shallower section. The algebraic differences between the lipid and kerogen, however, were within experimental error, -4.65 ±0.05.

Sufficient lipid was present for a total analysis in only the deeper lying sediment. The carbon content of the lipid was 52.1 per cent which is extremely low, and the nitrogen content was 1.9 per cent, which is high. This suggests that the sediment contains a considerable amount of organic matter that has undergone very little diagenetic change.

As shown in Figure 9, the n-alkane distribution for Section 391-1-2 is bimodal with a large odd predominance in the range of n-C_{23} to n-C_{33}. The curve is relatively smooth in the lower carbon number range.

**DISCUSSION**

The 37 core sections from the lower continental rise hills, Blake Nose, and the Blake-Bahama Basin range in
sub-bottom depth from near the water-sediment interface to approximately 1393 meters, and in age from Pleistocene to Jurassic. Assuming that these sediments were not previously more deeply buried, their temperatures have never been high, and probably did not exceed 45°-50°C. Their ages, however, vary by a factor of nearly two hundred. The inorganic rock matrix varies from almost pure silica-clay minerals to nearly pure calcium carbonate. In general, the organic matter reflects a photosynthetic origin with a marine or open water to nearshore environment. Thus the samples provide the opportunity to observe organic diagenesis at low temperature as a function of time and catalytic activity of the inorganic matrix.

Average odd-even predominance, OEP, is an indicator of the extent to which n-alkanes derived from the living source material are diluted by n-alkanes formed by abiogenic chemical processes following burial and cessation of life processes. Stated another way, the n-alkanes of biogenic origin, usually with an odd predominance are being diluted with those of abiogenic origin, which have a smooth carbon number distribution, that is, no odd or even predominence. Thus for source material with n-alkanes of odd carbon number predominance, the OEP value will decline as organic diagenesis proceeds toward a limiting value of 1.0. Because n-alkanes are ubiquitous in liquid petroleum, OEP can be used as a rough indicator of the amount of petroleum formed, particularly when an adjustment is made for the n-alkane concentration in the oil on the basis of the composition of the lipid fraction of the rock. Barring migration which usually is not extensive in shallowly buried rocks, the OEP value together with the lipid to total organic carbon ratio becomes an indicator of two interrelated parameters: the capacity of the source organic matter to produce liquid petroleum and the extent to which its genesis has proceeded under the given temperature-time regime (Erdman, 1975a, 1975b, Erdman et al., 1976).

Referring to Table 3, the OEP values are quite low overall and trend to lower values with increasing age. The lower OEP values in conjunction with the relatively high total organic content and the high proportion of saturates and aromatics in the lipid fraction of the older samples lead us to conclude that liquid petroleum genesis is well advanced in some of the older sediment samples. Accordingly, we conclude that (1) petroleum genesis in quantity can proceed at temperatures no higher than 45°-50°C provided there is a long interval of time and the source material is of the composition and is deposited under favorable conditions for high ultimate yield of petroleum, and (2) petroleum genesis takes place in lime sediments as readily as in clastic sediments.

To provide an accumulation with economic potential or one with a dangerous potential for blowout during uncased drilling, there must be sufficient compaction to induce migration and accumulation in adjacent coarse-grained rocks. Such compaction with subsequent migration, however, has not taken place in the sediments penetrated in these holes.

CONCLUSIONS

The 37 frozen core samples from the lower continental rise hills, the Blake Nose, and the Blake-Bahama Basin are exceptionally rich in organic material; 22 contained more than 0.5 per cent organic carbon. The organic matter was derived mainly from marine organisms. The ages of the samples ranged from Pleistocene to Jurassic. The suite of samples, therefore, provided an excellent opportunity to determine the rate of organic diagenesis and particularly the genesis of petroleum at low temperature.

A well-defined trend in organic diagenesis was observed with increasing age. In the older samples, genesis of liquid petroleum is well advanced. This suggests that for the formation of significant petroleum accumulations, deep burial is required more to provide the compaction necessary for migration rather than the elevated temperature for genesis.

ACKNOWLEDGMENT

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REFERENCES


