## The Shipboard Scientific Party1

## SITE DATA

Locality: Northern flank of Naturaliste Plateau Position: lat 33°47.69'S long 112°28.42'E

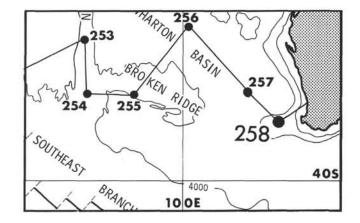
Dates Occupied: 25-29 October 1972

Water Depth: 2793 meters

Penetration: 258, 525 meters; 258A, 123.5 meters

Number of Cores: 34 Oldest Datable Sediment Cored:

Depth (subbottom): 435-444.5 meters (Core 22) Nature: Ferruginous clay Age: Mid Albian



Principal Results: Miocene to Recent nannofossil ooze (114 m) overlies a thick Cretaceous section (at least 411 m). Basement was not reached at this site. The Cretaceous section consists of 149 meters of Cenomanian to Santonian nannofossil chalk and silicified limestone above 251 meters of mid Albian to Cenomanian ferruginous detrital clay. The lowermost 11 meters of section sampled consist of unfossiliferous detrital sandstone and silty clay. There is a major disconformity between Santonian and late Miocene sediments.

# **BACKGROUND AND OBJECTIVES**

Site 258 is located in 2793 meters of water atop an apparent high on the northern flank of the Naturaliste Plateau (Figure 1). The Naturaliste Plateau is a submerged western extension of the southwest Australian continental margin. Summit depths of the plateau are near 2500 meters. Its southern edge is rather steep and trends east-west, parallel to the southern Australian margin and the Diamantina Fracture Zone (Heezen and Tharp, 1965). The northern slope of the Plateau is, by contrast, rather gentle. The Naturaliste Plateau and Broken Ridge to its west, seem to lie on a single trend parallel to the Diamantina Fracture Zone and Ob Trench but are separated by a deep-water gap about 600 km wide. Burkle et al. (1967) have reported retrieving an Upper Cretaceous (Turonian) core from the plateau near Site 258. This sample was a pale orange to white pelagic chalk. They concluded that, in the Cretaceous, limestone was laid down in water depths similar to those of today.

The stratigraphy of the Naturaliste Plateau is important for the understanding of the early history of the Indian Ocean and the breakup of Gondwanaland. In particular, the oldest sediment of marine facies retrieved from the plateau might give a maximum age for the Indian Ocean. We also hoped to sample an equivalent of the Santonian limestone retrieved from Broken Ridge (Site 255), thereby linking the history of the two structures. Further, because of the critical location of the plateau at the southwestern tip of Australia, we expected that the effects of the initiation of the Circumpolar Current in the Oligocene (?) would be reflected in the stratigraphy of the site.

The seismic profiles near the site (Figure 2) and the published profiles in Burkle et al. (1967), show about 0.85 sec DT of sediment conformably overlying a rough acoustic basement. Profiles onto and away from the site (Chapter 12, this volume, fig. 14 and 15) show that this basement return is discontinuous atop the plateau and is not necessarily correlated with the basement reflection in the Wharton Basin. An intermediate reflector at the site, at 0.3 sec DT, crops out to the north and to the east.

#### **OPERATIONS**

Site 258 was approached from the northwest. Our approach was somewhat delayed by very strong head-

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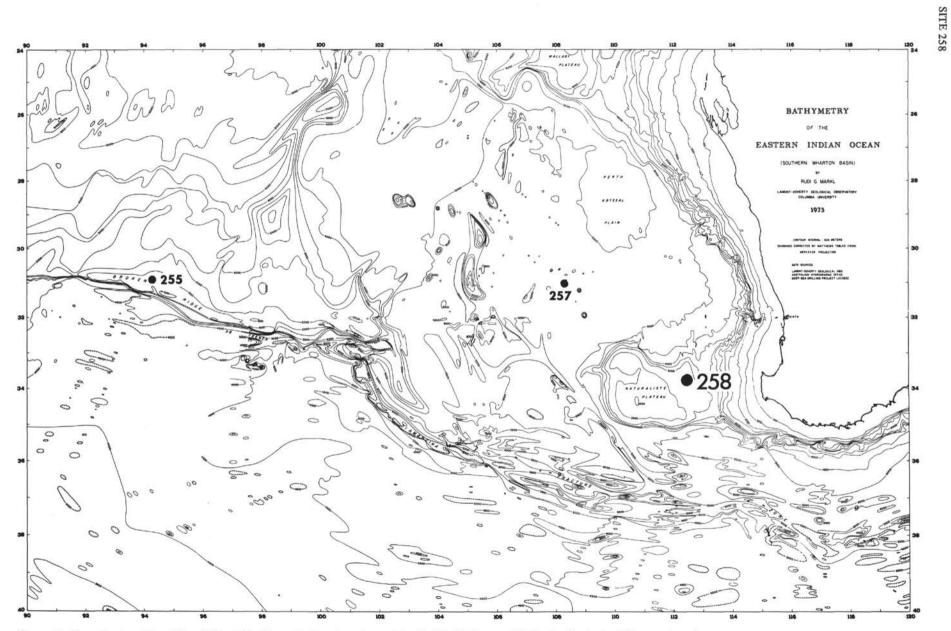


Figure 1. Base chart and locality of Site 258. (Compiled and contoured by R. Markl, Lamont-Doherty Geological Observatory.)

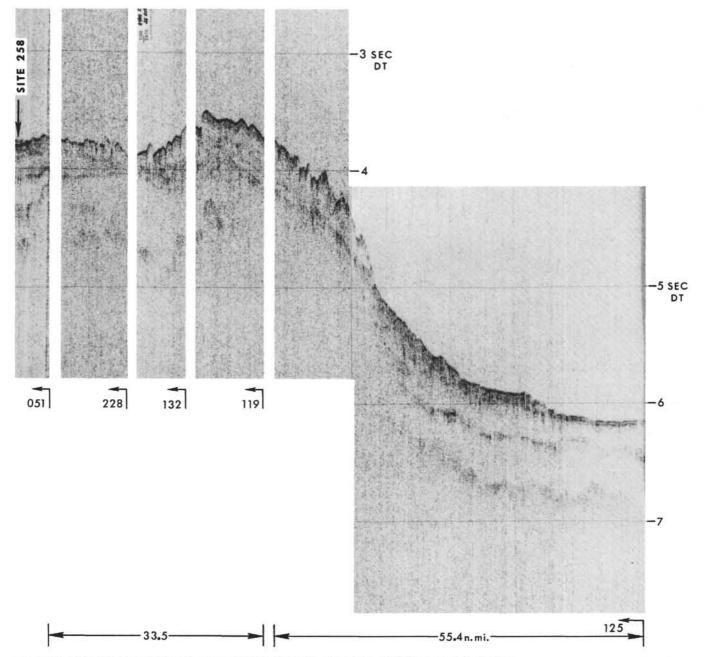


Figure 2. Seismic reflection profile approaching Site 258, taken from D/V Glomar Challenger.

winds and swells. At one time on passage from Site 257 our forward speed was reduced to 6.9 knots from a nominal 10 knots. Arriving at a point about 8.05 km (5 miles) northeast of the site, we turned onto a southwesterly course for our final approach, passing over the site in the early afternoon of 25 October. We then turned back over the site and dropped the beacon, while underway at 5 knots, at 1412, 25 October. By the time of our arrival on site, the winds had dropped to 10 mph, but confused swells from three different directions gave us considerable initial problems in positioning since the ship would not take up a stable heading. The water depth at the site was 2793 meters (corrected). Bottom was reached and the first core brought onboard at 2310, after which we cored discontinuously down to a subbottom depth of 525 meters. Table I gives vital statistics for the cores cut at Site 258. No downhole temperature measurements were attempted at this site and operations proceeded with all possible speed through 26 October. The weather remained calm and sunny, although weather forecasts and satellite photos showed a complex depression approaching from the west. By the morning of 27 October it was quite apparent that the storm was going to pass very close to us and that operations might be somewhat hampered. However, it was equally clear that time was very short

	TAE	L	E 1		
Cores	Cut	at	Site	258	

					Len	oth	
	Date		Depth from	Depth Below		Reco-	Reco-
			Drill Floor	Sea Floor	Cored	vered	
10.000	(Oct.	ALC: 10		(m)			very
Core	1972)	Time	(m)	(m)	(m)	(m)	(%)
Hole 258							
1	25	2310	2803.0-2812.5	0-9.5	9.5	9.6	101
2	26	0005	2812.5-2822.0	9.5-19.0	9.5	4.0	42
Drilled	77. T. I	(*******	2822.0-2850.5	5.050 B 5.750	0500.000	0.000	
3	26	0115	2850.5-2860.0	47.5-57.0	9.5	9.0	95
Drilled	20	0115	2860.0-2888.5	41.0 01.0	2.0	2.0	20
4	26	0225	2888.5-2898.0	85.5-95.0	9.5	0.6	6
Drilled	20	0225		63.3-93.0	9.5	0.0	0
	0.0	0040	2898.0-2926:5	100 5 100 0	0.5	4.0	
5	26	0340	2926.5-2936.0	123.5-133.0	9.5	4.2	44
Drilled			2936.0-2945.5				
6	26	0456	2945.5-2955.0	142.5-152.0	9.5	6.1	64
7	26	0600	2955.0-2964.5	152.0-161.5	9.5	3.7	39
Drilled			2964.5-2974.0	152.0-161.5	9.5	3.7	
8	26	0713	2974.0-2983.5	171.0-180.5	9.5	CC	0
9	26	0835	2983.5-2994.0	280.5-290.0	9.5	1.3	14
Drilled			2994.0-3002.5				
10	26	1104	3002.5-3009.0	199.5-206.0	6.5	2.0 <sup>a</sup>	31
Drilled			3009.0-3018.5				
11	26	1253	3018.5-3028.0	215.5-225.0	9.5	4.0 <sup>a</sup>	42
Drilled	-	0.000	3028.0-3037.5		23025		
12	26	1410	3037.5-3047.0	234.5-244.0	9.5	9.5 <sup>a</sup>	100
Drilled	20	1110	3047.0-3056.5	251.5 211.0	2.0	7.0	100
13	26	1555	3056.5-3066.0	253.5-263.0	9.5	5.0 <sup>a</sup>	53
15	26	1755	3066.0-3075.5	263.0-272.5	9.5	1.5 <sup>a</sup>	16
Drilled	20	1/55		203.0-272.5	9.5	1.5	10
	26	1040	3075.5-3085.0	202 0 201 5	0.5	8.0 <sup>a</sup>	04
15	20	1940	3085.0-3094.5	282.0-291.5	9.5	8.0-	84
Drilled	24		3094.5-3104.0		0.5	a 73	00
16	26	2135	3104.0-3113.5	301.0-310.5	9.5	7.6 <sup>a</sup>	80
Drilled			3113.5-3123.0			2	-
17	26	2325	3123.0-3132.0	320.0-329.0	9.0	7.5 <sup>a</sup>	79
Drilled			3132.0-3142.0			10.772.20	
18	27	0205	3142.0-3151.5	339.0-348.5	9.5	5.5 <sup>a</sup>	58
Drilled			3151.5-3161.0				
19	27	0425	3161.0-3170.5	358.0-367.5	9.5	CCa	0
Drilled			3170.5-3180.0				
20	27	0651	3180.0-3189.5	377.0-386.5	9.5	0.8 <sup>a</sup>	9
Drilled			3189.5-3208.5				
21	27	0900	3208.5-3218.0	405.5-415.0	9.5	4.5 <sup>a</sup>	47
Drilled	21	0700	3218.0-3237.0	100.0 110.0	2.0	4.0	
22	27	1030	3237.0-3246.5	435.0-444.5	9.5	6.7	71
Drilled	21	1050		455.0444.5	9.5	0.7	11
	07	1015	3245.5-3275.0	470 0 401 5	0.5		17
23	27	1245	3275.0-3284.5	472.0-481.5	9.5	1.6	17
Drilled			3284.5-3313.0				
24	27	1542	3313.0-3322.5	510.0-519.5	9.5	7.5	79
25	28	0700	3322.5-3328.0	519.5-525.0	5.5	5.3	96
Fotal <sup>C</sup>					230.5	115.55	50
Hole 258A							
				19-10-00-00-00-00-00-00-00-00-00-00-00-00-		2 × 2 × 2 × 1	
1 <sup>b</sup>	28	1105	2803.0-2812.5	0-9.5	9.5	1.5	16
1A	28	1155	2803.0-2812.5	0-9.5	9.5	1.7	18
Drilled			2812.5-2822.0				
2	28	1312	2822.0-2831.5	19.0-28.5	9.5	4.5	47
2 3	28	1430	2831.5-2844.0	28.5-38.0	9.5	9.5	100
4	28	1548	2844.0-2850.5	38.0-47.5	9.5	8.5	89
5	28	1653	2850.5-2860.0	47.5-57.0	9.5	9.1	96
Drilled	20	1000	2860.0-2869.5		2.0		
6	28	1848	2869.5-2879.0	66.5-76.0	9.5	9.3	98
Drilled	20	1040		00.0-70.0	9.5	2.5	90
	20	1057	2879.0-2898.0	05 0 104 5	0.5	7.0	74
7	28	1957	2898.0-2907.5	95.0-104.5	9.5	7.0	
8	28	2107	2907.5-2917.0	104.5-114.0	9.5	8.0	84
9	28	2220	2917.0-2926.5	114.0-123.5	9.5	7.9	83
Total <sup>C</sup>							

TABLE 1 – Continued

<sup>a</sup>Core was overcored to the depth of the top of the following core.

<sup>b</sup>Two cores were cut from the interval 0-9 meters subbottom. Due to confusion in the laboratory, these were designated 258A-1 and 258A-1A, rather than the correct designations 258A-1 and 258B-1. Core 258A-1 is essentially the same as 258A-1A and has not been considered in subsequent discussions.

<sup>c</sup>Totals do not include figures from Core 258-A.

and that Site 258 was giving up some surprising and extremely valuable information. With the goodwill of the drilling crews, operations were continued into extreme weather conditions. The ship performed excellently and at the time we, reluctantly, suspended operations we were operating in winds of 50-60 mph and breaking 15- 20-foot seas. In these conditions there was simply not enough power available to hold the ship on site over the hole and at the same time work pipe, besides which the motion of the ship was such that it was becoming impossible to work safely on the rig floor.

The decision to pull out of the hole was made at 1630 27 October. At first we pulled pipe to within 30 meters of the sea bed in the hope of being able to retain the hole and wait out the worst of the storm. However, by 1800 it was obvious that conditions were not going to improve in the immediate future, so we let go of the hole and held position over the beacon with the pipe suspended about 200 meters above the bottom through the night. Wind and sea showed little sign of abating by midnight but soon after dawn, conditions began to moderate to the point where we could consider drilling a second hole.

At 0700, 28 October we started down again for Hole 258A. Time was very short, so rather than attempt deeper penetration, we decided to use the available time to sample some of the gaps in the coring of the upper portion of Hole 258. Nine cores were cut. The last core of Leg 26 was brought aboard at 2220, 28 October. We then pulled pipe and got underway at 0535, 29 October. After a brief postsite survey we departed for Fremantle at 0630.

### LITHOLOGY

At Hole 258, 525 meters of sediment were drilled and discontinuously cored. Recovery totaled 115.5 meters from the 230.5 meters cored, or approximately 50%. A second hole, 258A, was drilled in order to study in detail the Tertiary-Quaternary and Cretaceous-Tertiary boundaries at the site. In addition to a mudline core, Hole 258A was cored in two continuous sequences: from 19 to 76 meters and from 95.0 to 123.5 meters below the sea floor. Five lithostratigraphic units are recognized. The youngest consists of coccolith ooze. Unit 2, consisting of interbedded chalk and siliceous limestone, can be subdivided conveniently into two subunits according to the distribution of the siliceous limestone and the petrography of the chalk. The third unit is a transitional sequence passing down into the detrital clay of the fourth stratigraphic unit. The succession is completed by a lowermost unit of glauconitic sand and silty mudstone (Table 2).

## Unit 1

Unit 1 consists of 114 meters of foram-bearing, micarb-bearing, or sponge-bearing coccolith ooze. The color of the sediments varies from light gray, greenishgray, and yellow-gray to bluish-gray. The sediments show no obvious compositional trends with stratigraphic depth, although foram-, micarb-, and spongerich varieties were recorded below 95 meters. The detrital clay component makes up approximately 2% of the ooze. Other accessories include trace amounts of glauconite, quartz, pyrite, and submicroscopic opaque and translucent ferruginous aggregates. Core 1 contains 16% aragonite, the only occurrence of this polymorph of calcium carbonate in the Leg 26 cores. This may indicate rapid accumulation of sediment leading to the preservation of the metastable form. Some gypsum is found in the  $< 2\mu$  fraction from the lower part of Unit 1.

#### Unit 2

Subunit 2a consists mostly of light gray, very light gray, or yellowish-gray micarb and/or foram-bearing chalk. The uppermost 5 meters of the unsilicified component of the subunit include a few intervals which are still sufficiently unconsolidated to be termed ooze. Numerous beds less than 5 cm thick of silicified limestone occurring throughout the unit show the same sedimentary structures as the chalk and obviously are derived from them by diagenesis. The silicified limestone and much rarer chert constitute approximately 13% of the recovered material; this may be greater than the true value because of preferential recovery of harder rock. The contact with Unit 1 was not cored in Hole 258; however, examination of the Geolograph drilling records indicated that resistant silicified strata were encountered first at 112 meters below the sea floor. This boundary was cored in Hole 258A, where the uppermost chert occurred at a depth of 114 meters. At approximately 124 meters in Hole 258 fragments of dark brown and light olive-gray chert containing Inoceramus sp. fragments were observed. Diagenesis of most of the silicified limestone is incomplete, and all effervesce at least moderately in dilute hydrochloric acid.

Within the chalk, foraminiferal content ranges from approximately 2% to 7%, and microcrystalline calcite forms 5% to 20% of the sediment. In general, the microcrystalline calcite averages less than 10%.

Subunit 2b consists of the basal 60 meters of micarb coccolith chalk and coccolith micarb chalk. Microcrystalline carbonate is much more abundant than in the

Unit/ Subunit	Hole/ Core	Depth Below Sea Floor (m)	Thickness (m)	Description
1	258/1-4 258/1-9	114	114	Light gray, greenish-gray, bluish-grey, and yellow-gray sponge-, micarb-, and foram= bearing coccolith ooze
2a	258/5-10 258A/9	114-203	89	Yellowish-gray and very light gray foram- and micarb-bearing chalk and silicified limestone
2b	258/11-13	203-263	60	Light greenish-gray and yellowish= gray coccolith micarb and micarb coccolith chalk
3	258/14, 15	263-285	22	Interstratified dark greenish= gray zeolite-rich detrital clay, olive-black ferruginous clay, light olive-green coccolith detrital clay, and light olive= green coccolith-rich micarb chalk
4	258/15-24	285-514	229	Brownish-black and olive-black ferruginous detrital clay
5	258/24, 25	514-525	11	Olive-gray and greenish-gray fine-grained glauconite detrital sandstone; dusky brown glauconitic detrital silty clay

TABLE 2 Lithologic Summary, Site 258

younger subunit 2a and ranges from 25% to 70%. Foraminifera generally are present only in trace amounts and rarely exceed 2%. Accessory components include glauconite, submicroscopic opaque and translucent ferruginous aggregates, pyrite, and dolomite rhombs, all in trace proportions. Very rare traces of quartz and mica were noted. Detrital clay makes up about 2% of the sediment. Throughout Subunit 2b bioturbation was extensive.

The uppermost Subunit 2b sediments were not recovered. However, in view of the reliability of the Geolograph drilling records, proven in positioning the top of Subunit 2a, the top of Subunit 2b can be placed confidently at 203 meters below the sea floor.

## Unit 3

Unit 3 consists of 22 meters of passage bed lithologies transitional between the biogenic ooze of Unit 2 and the older detrital clay of Unit 4. The upper part of the unit above approximately 282 meters consists of dark greenish-gray zeolite-rich detrital clay and olive-black, black, or greenish-black ferruginous detrital clay, interbedded with light olive-green foram-bearing coccolith-rich micarb chalk. Cristobalite is a major component of the sediments of Unit 3. The zeolite-rich clay contains up to 20% clinoptilolite. The prominence of the zeolite component is understandable in view of the very low sedimentation rates obtained paleontologically for this part of the Cretaceous section. In smear slides the ferruginous clay contains well over 20% pyritic and otherwise ferruginous amorphous translucent and opaque aggregates. The basal 3 meters of Unit 3 consist of light olive-green zeolite and micarb-bearing coccolith-rich and coccolith detrital clay.

### Unit 4

Unit 4 consists of 229 meters of brownish-black to olive-black ferruginous detrital clay. This clay varies compositionally between zeolite-, coccolith-, and micarb-bearing varieties according to the particular accessory component. Both the top and the bottom of the unit were observed in the recovered core at 285 and 514 meters, respectively. The upper boundary is transitional and the lower abrupt.

Throughout, Unit 4 shows very little lithologic variation. Bioturbation is ubiquitous, in the form of prolific horizontal or subhorizontal burrowing. With apparent increase in compaction, the bioturbation in older horizons becomes more difficult to recognize and is easily mistaken for lamination.

The precise nature of the ferruginous material cannot be determined at this stage. Finely divided pyrite, responsible for the unit's coloration, is widely disseminated, as is subordinate silt-sized framboidal pyrite. However, much of the ferruginous material is in the form of submicroscopic opaque and translucent aggregates. X-ray analysis is necessary for its precise identification. The total ferruginous component averages about 5%-8%. Scattered pyrite nodules, up to 3 cm in diameter, occur below 408 meters. Zeolite throughout the unit ranges from trace amounts up to 10%. It appears to be particularly concentrated within lighter olive-black bioturbated areas. Coccoliths are also widespread, but are absent in at least the basal 40 meters. On the average they form less than 10% of the detrital clays. Foraminifera are present in trace amounts, or at the very most 2%-3%. Microcrystalline calcite, in amounts varying from 1% to 10% was noted at rare intervals in this unit. Siderite occurs widely in trace

amounts beneath 323-meter depths in the hole. At 325.3 meters a 10-cm band of medium-sand grade siderite was recovered, but in all other instances this carbonate occurs only as dispersed minute specks or very thin veinlets and nodules.

## Unit 5

The basal 12 meters of the drilled sequence consists of fine-grained and minor medium-grained, very wellsorted glauconitic sandstone overlying dusky brown silty mudstone (Figure 3). The sandstone contains abundant carbonate cements and veinlets, in places calcitic, but commonly sideritic or dolomitic. Macrofossil debris is not uncommon, and the sand shows some bioturbation, including vertical burrows. Some of the very fine-grained sandstone has a chert-like degree of induration, and some of these sediments have been fractured and infilled by either the overlying sediments or by sedimentary dikes. The fine-grained sand also displays finely developed lamination and grading in units varying from a few millimeters up to 6 cm in thickness. Much of the well-sorted graded sand also displays dispersed pebble and granule grade black or brown fine-grained lithic fragments. Slight load-casting of the sand down into thin silty interbeds was observed at 517 meters. Irregular dark laminae less than 0.5 cm thick with splendent pyrite granules occur at 515.5 meters down the hole. The basal 9.5 meters of Unit 5 constst almost entirely of dusky brown ferruginous detrital clay with some laminae of very fine silty sand. More than half of the clay is kaolinite (57%), accompanied by montmorillonite (13%), potash feldspar (16%), and detrital quartz (3%). The rest of the sediment consists of hematite (11%) and a trace of magnetite. Pyrite is common, as separate nodules up to 7 mm in diameter, and as 1-mm-thick horizontal stringers up to 1.5 cm long situated in black laminae of similar thickness. The petrography of both the silty ferruginous clay and the sand appears to be the same. Much of the



Figure 3. Volcanogenic sandstone. The sediment framework consists of fine-grained porphyritic basaltic volcanic fragments, variably turbid grains of brown and green basaltic glass, and iron oxides set in an abundant calcite cement. Plain light; field width 0.8 mm.

sediments of Units 4 and 5 was apparently derived from the erosion of basaltic volcanic rocks.

## SHIPBOARD GEOCHEMICAL MEASUREMENTS

Routine analyses for salinity, pH, and alkalinity were conducted on interstitial water samples squeezed from 10 sediment samples taken at depths in Hole 258 from 17.5 meters to 481.5 meters below the sea floor. In addition, pH was measured on the uppermost two samples of unsqueezed sediment by the punch-in method before the core recoveries became too stiff for the electrodes. No geochemical sampling was performed on sediment from Hole 258A. The sampling and analytical techniques are described in the report for Site 250. The results for Site 258 are summarized in Table 3 and are illustrated in Figure 4.

## Results

The regional near-bottom salinity at 3000 meters cited by Wyrtki et al. (1971) is  $34.7 \text{ }_{0/00}$ . Salinities in the interstitial waters exceed this value and remained relatively constant, ranging from 34.9 to  $35.5 \text{ }_{0/00}$ throughout the hole down to and including the 260.0meter level. Salinity values obtained from the four lowermost samples (below 290 m) are lower than any of the values above, ranging from 27.8 (at 481.5 m maximum depth) to  $34.2 \text{ }_{0/00}$  at 345.5 meters below the sea floor.

#### pH

The coupled punch-in and flow-through pH measurements differed by 0.38-0.4 pH unit, punch-in values being consistently higher. Flow-through pH in the uppermost sample was 7.29, below the normal range for seawater of 7.8-8.2. Values decreased with depth to a minimum of 6.71 at 260 meters below the sea floor. In the next lower sample, at 290 meters, pH increased to 7.26 and remained constant at 7.14 for the lowermost three samples. Thus, the ferruginous detrital clay sequence appears to be characterized by a distinct preferred pH value.

#### Alkalinity

The uppermost two samples from the hole had a maximum alkalinity of 3.23 meq/kg. From this depth, values decreased down the hole consistently to a minimum of 0.98 meq/kg at 290 meters. The two samples below this depth had higher alkalinities, 1.47 and 1.34 meq/kg, respectively, both values being lower than any value above 290 meters. No measurement was taken on the lowermost sample.

#### PHYSICAL PROPERTIES

The physical properties measured at Site 258 were porosity, acoustic velocity, bulk density, and thermal conductivity. The methods are described in the Explanatory Notes (Chapter 2). The results are shown in the hole summary diagram.

#### Density, Porosity, and Water Content

The mean bulk density of the upper carbonate ooze layers was 1.63 g/cc. The density increased slowly in the

TABLE 3 Summary of Shipboard Geochemical Measurements, Site 258

Sample (Interval in cm)	Depth Below Sea Floor (m)	Lab Temp (°C)	pH Punch-in/ Flow-through	Alkalinity (meq/kg)	Salinity (°/••)
(Reference seawater)	-	-	8.19/8.14	2.38	35.8
2-2, 144-150	17.5	22.3	7.66/7.29	3.23	35.2
3-5, 144-150	57.0	22.3	7.67/7.27	3.23	35.2
5-2, 144-150	131.5	22.3	/7.06 <sup>a</sup>	2.93	35.5
7-3, 144-150	161.5	22.5	/6.95 <sup>a</sup>	2.66	34.9
11-3, 144-150	225.0	22.2	/6.91 <sup>a</sup>	1.66	35.2
13-2, 144-150	260.0	22.2	/6.71 <sup>a</sup>	1.76	35.2
15-5, 143-150	290.0	22.0	/7.26 <sup>a</sup>	0.98	33.6
18-2, 140-150	345.5	22.2	/7.14 <sup>a</sup>	1.47	34.2
21-1, 0-10	405.5	21.6	/7.14 <sup>a</sup>	1.34	34.0
23-2, 140-150	481.5	22.1	/7.16 <sup>a</sup>		27.8

<sup>a</sup>Too stiff to measure punch-in.

underlying sediments to 1.75 g/cc at the bottom of the hole. Porosity and water content decrease irregularly with increasing depth.

## Acoustic Velocity and Acoustic Impedance

The acoustic velocities are very irregular, following the varied lithology. The average velocity of the upper unconsolidated ooze section is 1.58 km/sec. The velocity in the underlying sediments generally ranges from 1.6 to 1.9 km/sec. Lithified chalk and silicified limestone layers near 85 meters have velocities between 1.7 and 2.1 km/sec, near 118 meters a velocity of 3.0 and near 191 meters a velocity of 3.7. An indurated glauconite sandstone at 517 meters has a velocity of 3.3 km/sec.

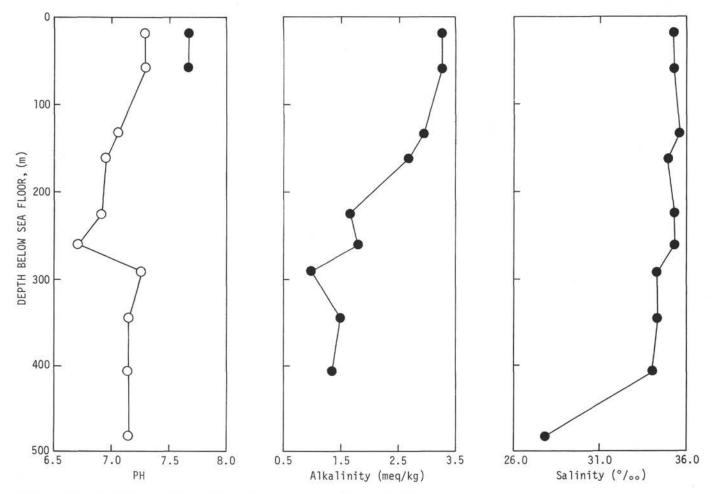


Figure 4. Graphic summary of geochemical measurements taken at Site 258.

Most of the consolidated sediments have a significant anisotropy.

The acoustic impedance increases generally with pronounced breaks for the hard indurated layers. A weak seismic reflection at about 0.09 sec DT may be associated with the thin high-velocity silicified limestone layers near 85 meters. There is an obvious decrease in acoustic impedance at about 270 meters for the strong reflector at 0.30 sec DT. The reflector at 0.55 sec DT likely arises from the contact between the ferruginous clay and underlying sandy mudstones at about 514 meters. There is a high velocity glauconite sandstone layer in the latter section at 517 meters.

#### CORRELATION OF SEISMIC REFLECTION PROFILE AND DRILLING RESULTS

An on-site sonobuoy profile was run at Site 258 (Figure 5). Four prominent reflections are seen at 0.09, 0.30, 0.55, and 0.84 sec DT subbottom. We believe that Hole 258 penetrated to the 0.55 sec DT reflection before weather won out.

Table 4 shows a correlation between depths of lithologic contrasts, impedance contrasts, and reflection

times, plus a calculation of the average velocity to reflection depth. The upper reflection corresponds to an increase in induration in Unit 1 and not to the boundary (unconformity) between Units 1 and 2 at 114 meters. The 0.3 sec DT reflection corresponds to the contact between Units 3 and 4 at 285 meters, and the 0.55 sec DT reflector is the contact with the indurated sands at 514 meters. Using an average velocity of 1.9 km/sec, the implied basement depth is 798 meters. There may be a very weak impedance contrast at or near 285 meters. It seems that at this site reflections are composed of returns from sequences of high-velocity stringers, such as silicified limestones and sands, rather than massive lithologic changes. With limited recovery of these hard stringers, it is very hard to deduce changes in impedance contrast.

## PALEONTOLOGY

#### **Biostratigraphic Summary**

In the two holes at Site 258, 114 meters of upper Miocene-Recent overlying 411 meters of middle Albian to Santonian sediments were penetrated. The Tertiary

258 33°48'S 112°29'E 2793 meters

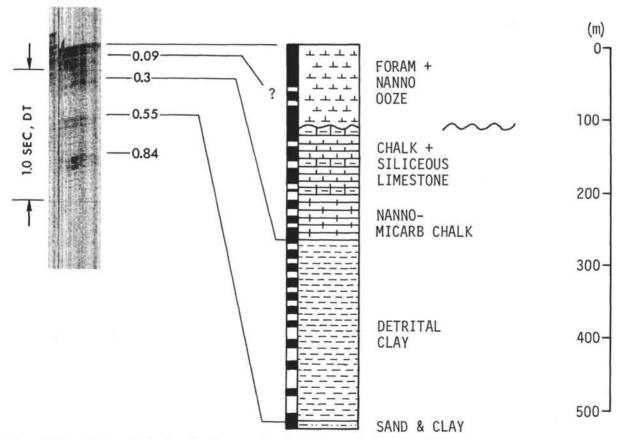


Figure 5. Correlation of seismic reflection record and drilling results from Site 258.

NATURALISTE PLATEAU

1	TABLE 4			
Seismic Reflectors and	Lithologic	Changes	at Site	258

Reflection Time (msec DT)	Depth of Lithologic Contrast (m)	Depth of Impedance Contrast (m)	Average Velocity (km/sec)
90		85-200	1.89
300	285	270	1.90
550 840 <sup>a</sup>	514	514	1.87

<sup>a</sup>Not sampled.

part of the section has yielded a Pliocene/Pleistocene foraminiferal and nannoplankton sequence of a temperate environment with 29-45 meters of Quaternary and 23-56 meters of Pliocene. Upper Miocene was reached at the base of the Tertiary sequence.

#### PALEONTOLOGY

#### **Biostratigraphic Summary**

In the two holes at Site 258, 114 meters of upper Miocene-Recent overlying 411 meters of middle Albian to Santonian sediments were penetrated. The Tertiary part of the section has yielded a Pliocene/Pleistocene foraminiferal and nannoplankton sequence of a temperate environment with 29-45 meters of Quaternary and 23-56 meters of Pliocene. Upper Miocene was reached at the base of the Tertiary sequence.

Benthonic foraminifera indicate a considerable decrease in water depth during lower Pliocene and uppermost Pliocene.

The Cretaceous part of the section consists of a pelagic chalk sequence of Turonian to upper Santonian age. It is rich in well-preserved planktonic foraminifera in the Santonian, but these are scarce and poorly preserved in the Turonian and most of the Coniacian. Radiolaria and nannoplankton are abundant and well preserved throughout this unit. Dark shales of middle Albian to Cenomanian age occur below the chalks. They contain etched assemblages of foraminifera, nannoplankton, and some Radiolaria. These are less abundant, and become poor in the lower part of this unit. The samples at the bottom of the hole are devoid of microfossils. Deposition below the carbonate compensation depth is most probable for the lowermost sediment drilled at this site. At the same time an increase in sedimentation rate and, at the base, an accumulation of slightly coarser detrital material can be observed.

Foraminiferal and nannoplankton assemblages indicate a cool to temperate environment during the Upper Cretaceous.

The Cretaceous part of the section at this site gives an opportunity for correlating biostratigraphic data for the three main groups of planktonic microfossils in a mid to higher latitude area of the Southern Hemisphere.

#### Foraminifera

Neogene: This is the site where the Quaternary section could be divided into Recent and Pleistocene. For this subdivision the criterion suggested by Parker (1973) (see detailed discussion in Chapter 30) was used. The oldest sample which contained *Globigerinoides ruber* (*f. rosea*) is Sample 1-6, 110-112 cm. In the same material also one pink-walled *Globigerina rubescens* was found. This sample was found 8.5 meters below the bottom surface. Thus, if we accept Parker's criterion, the uppermost 8.5 meters of Quaternary sediments are Recent. The underlying 20.5 meters are Pleistocene, and the whole thickness of the Quaternary sequence is 29 meters. However, some doubts about the validity of this criterion are expressed in Chapter 30. The Pleistocene/Recent boundary is therefore indicated with a question mark.

The Quaternary/Pliocene boundary was located utilizing the *Globorotalia truncatulinoides: G. tosaensis* ratio and the extinction of *Globorotalia crotonensis* and *Globigerinoides obliquus, s.l.* 

The limit between upper and middle Pliocene is given, among other criteria, by the extinction of *Globoquadrina altispira* and the appearance of *Globorotalia truncatulinoides*. G. tosaensis, Pulleniatina obliquiloculata, s.s., P. obliquiloculata praecursor, Globigerinoides pyramidalis appear in the lowermost middle Pliocene.

The most characteristic species of the lower Pliocene is *Globorotalia margaritae*.

For the location of the Miocene/Pliocene boundary, the distribution of *Globorotalia inflata s.l.* and *G. crassaformis* was used. The sporadic records of rare or isolated specimens of the former species in the upper Miocene are in marked contrast to the abundant (or frequent) and uninterrupted appearance of specimens of the same species in the Pliocene. More or less the same phenomenon (but not as well pronounced) was observed with respect to *Globorotalia crassaformis*.

The following typical upper Miocene species were encountered somewhat lower than the Miocene/Pliocene boundary: Orbulina suturalis, Globigerinoides amplus, Globigerinita unicava, and Globorotalia petaliformis. It is interesting to note that, as at Sites 253 and 254, these assemblages contain a great number of very small, unidentifiable, planktonic foraminifera.

Quaternary foraminiferal assemblages at Site 258 are typical of the temperate zone. *Globorotalia inflata s.l.* strongly predominates, and *G. menardii s.l.* is recorded sporadically and as isolated individuals.

Site 258 showed a rather enigmatic stratigraphic range for *Globigerinoides sicanus*. For details see Chapter 30, this volume.

No Neogene sediments older than upper Miocene were encountered at this site.

**Cretaceous:** Foraminiferal assemblages in the Upper Cretaceous chalk unit are characterized by a relatively low diversity. Species of the genera *Hedbergella*, *Archaeoglobigerina*, *Globigerinelloides*, and *Heterohelix* strongly predominate over the relatively few species of *Globotruncana* and *Praeglobotruncana*. The age determinations based on these assemblages are generally not as precise as would be desirable for the establishment of a detailed biozonation. In addition, the indurated nature of the chalk in Cores 11-13 has prevented the isolation of satisfactory foraminiferal assemblages.

Well-preserved assemblages of abundant planktonic foraminifera were found in the upper part of the chalk

unit (Hole 258A, Core 8, Section 6 to Hole 258, Core 10, Section 1). Double-keeled species of Globotruncana form, as mentioned, quantitatively a subordinate portion of the assemblages. The presence of G. cf. ventricosa (probably an early evolutionary stage in the development of this species) together with rare G. fornicata manaurensis and the absence of single-keeled forms indicate a Santonian, possibly upper Santonian age for the uppermost part of this unit. The Coniacian/Santonian boundary is tentatively placed between Cores 258-5 and 6, based on the occurrence and relative frequency of G. angusticarinata and G. fornicata manaurensis. The absence of representatives of the G. concavata group, important index forms for the Coniacian-Santonian interval, is probably due to the temperate character of these assemblages, as discussed below.

In the lower part of Core 10, in Core 11, and in the upper part of Core 12, only very few specimens of planktonic foraminifera could be isolated. This interval has tentatively been placed in the Coniacian based on the scattered occurrence of double-keeled *Globotruncana*.

Few specimens of moderately preserved *Praeglobotruncana (P. stephani, P. cf. algeriana, P. helvetica)* were found in highly liquified sediments of Core 12, Section 3 and Core 13, Sections 3 and 4. Although these samples cannot be considered reliable with respect to the true core depth, they indicate that this portion of the chalk unit has to be at least as old as middle to upper Turonian. The Turonian/Coniacian boundary has tentatively been placed above section 3 of Core 12. However, since the Coniacian assemblages above were also collected in a highly liquified sample, this boundary is not very reliable and may eventually be placed slightly higher in the section.

Assemblages of the interstratified clay unit in the lowermost part of Sample 13, CC and in Core 14 are characterized by the association of *Hedbergella planispira*, *H. delrioensis*, *H. simplicissima*, *Globigerinelloides caseyi*, and rare *Schackoina cenomana*. The Cenomanian age of this assemblage is supported by one single specimen of *Rotalipora reideli* in Sample 14, CC.

The assemblage found in Core 15 does not differ essentially from the one found in Core 14. A single occurrence of *Praeglobotruncana delrioensis* in Section 3 of Core 15 is noteworthy. Based on these foraminiferal assemblages, the Albian/Cenomanian boundary cannot be located precisely. *Praeglobotruncana delrioensis* is known in the upper Albian as well as in the Cenomanian.

No planktonic foraminifera occur in Core 16 and most of Core 17. Cores 18 to 20 are characterized by the association of *Hedbergella* sp. aff. *infracretacea*, *H. planispira*, and *Globigerinelloides casey*. This corresponds to the Albian age determined by the nannoplankton. No planktonic foraminifera have been found below Core 20. The assemblages are reduced to few arenaceous and clacareous benthonic forms in Core 21, and to primitive arenaceous psecies in Cores 22 and 23. No ages can be deduced from them. In Cores 24 and 25 no foraminifera have been found at all.

The restricted diversity of the Albian and Cenomanian assemblages, particularly the virtual absence of the genera *Rotalipora* and *Praeglobotruncana*, are interpreted to indicate a cool environment during this time span. The diversity is slightly higher in the Turonian and Coniacian, and particularly in the Santonian, where an increase in the abundance of globotruncanid species is observed. These species, however, still constitute a low proportion of the total planktonic assemblages. The absence of such a typical warm-water species as *Globotruncana concavata* also suggests that the site was located in a cool to temperate environment. For further paleoecologic discussion see Chapter 32, this volume.

The distribution of planktonic and benthonic foraminifera suggests a steady decrease of the water depths from levels below the carbonate compensation depth in the lowermost part of the section (Cores 2-5) to depths within and above that depth during the Albian. A further decrease of the water depth must be assumed for the Turonian to Santonian chalk unit. However, the strong domination of planktonic foraminifera over benthonic forms still indicates a distinctly pelagic environment.

### Calcareous Nannoplankton

Stratigraphy: Nannoplankton assemblages of middle Albian through Santonian and upper Miocene through Quaternary age were encountered at this site. The almost continuous Cretaceous sequence includes abundant nannofossils of Albian (Hole 258, Cores 15-22); Cenomanian (Hole 258, Core 14); Turonian (Hole 258, Cores 12, 13); Coniacian (Hole 258, Cores 6-12); and Santonian (Hole 258, Core 5; Hole 258A, Cores 8, 9) ages, with a minor unconformity between the Cenomanian and the Turonian. The Santonian/Miocene unconformity was recovered in Hole 258A, Core 8, Section 6. The age of the sediments in Cores 6 through 8 is based upon nannofossil samples taken from pebbles and is considered to be upper Miocene. Although many Pliocene foraminifera were encountered, these foraminiferal samples often had to be taken from liquified parts, which apparently were downhole contaminated. As at Sites 251 and 255 an overlap of the uppermost occurring Discoaster sp. and the lowermost occurring Gephyrocapsa sp. is observed. Again the foraminifera are considered to be of Pliocene age; in fact, the Pliocene/Quaternary boundary as determined by foraminifera lies 17 meters above the extinction of discoasters. In most of the Neogene samples reworked Upper Cretaceous and upper Eocene nannofossils have been found.

**Preservation:** Most of the Neogene nannofossils are well preserved. Some signs of overgrowth are observed in the lower Pliocene and upper Miocene samples. The Cretaceous assemblages from the Turonian to Santonian chalk sequence are well preserved to moderately overgrown. Downwards, increasing dissolution effects are observed in the Cenomanian to middle Albian clays. The lowermost, strongly etched nannofossils are found in Core 23.

**Paleoecology:** The preservation of the nannofossils in the Cretaceous sequence indicates a continuous shallowing of this site from middle Albian through Santonian time. The species compositions of the assemblages point to a transitional to cool environment, although a typically Tethyan species (*Lithraphidites alatus*) has been found in Core 14 (Cenomanian). The depositional depth at Site 258 during the Santonian was greater than that of the Gingin Chalk in the Perth Basin, based upon the absence of *Tetralithus obscurus*, *T. ovalis*, and the scarcity of *Lucianorhabdus cayeuxi*. All three are restricted to nearshore or shallow-water paleo-environment and are abundant in the Gingin Chalk. The Neogene assemblages indicate temperate paleo-temperatures.

## SEDIMENTATION RATES

The sedimentation rate in the Pliocene-Quaternary was about 14-17 m/m.y., depending on the interpretation of the Pliocene/Miocene boundary. The disconformity encountered in Hole 258A, Core 8 encompasses a period of at least 66 m.y. The sedimentation rate during the Santonian and Coniacian was 12.6 m/m.y. The relatively slow sedimentation during the Turonian and Cenomanian (1.8 m/m.y.) might be due to a disconformity between the upper Turonian and the upper Cenomanian. During the Albian the sedimentation rate was at least as high as 45 m/m.y.

## SUMMARY AND CONCLUSIONS

## Summary of Results

Site 258 is situated on the northern slope of the Naturaliste Plateau. The water depth is 2793 meters, and seismic profiles show the sediment thickness to be about 0.85 sec DT. Two holes were drilled. The first penetrated 525 meters before rough weather forced us to suspend operations for about 15 hr. A second hole was drilled to 123.5 meters in order to fill in some gaps in the sampling of the upper parts of the section. The oldest precisely datable sediment reached was middle Albian, found at 444 meters in Hole 258.

Five lithologic units can be recognized. From the surface down to 114 meters is a sequence of white and gray soft oozes ranging from late Miocene to Recent in age. Most of the section is Plio-Pleistocene. Towards the bottom of this unit, especially in the Miocene, planktonic and shallow-water (slope facies) benthonic foraminifera appear. Reworked littoral forms also are found. Unit 2, from 114 to 263 meters, is silicified limestones and chalks overlying micarb (recrystallized) chalks (Subunit 2b) ranging from Cenomanian to Santonian in age. The boundary between Units I and 2 is sharp and well defined. Representatives of all three microfossil groups are abundant and well preserved in Unit 2. Unit 3, 263-285 meters is a transitional unit with interbeds of chalk and the dark ferruginous clays of Unit 4 (285-514 m). The clays of Unit 4 are almost black and contain many diagenetic minerals (alkali feldspar, siderite, zeolites, etc.). This unit is middle and late Albian in age. Unit 5 is a Lower Cretaceous (undefined range) sequence of glauconitic sands and muddy silts, well sorted, and in many cases showing graded bedding. The sediments become coarser downwards. There is common, but as yet indeterminate, macrofossil debris and many deep-water benthonic foraminifera. The planktonic foraminifera are represented by only one species.

Sedimentation rates were high in the Early Cretaceous and in the Late Cretaceous and Tertiary but clearly very much lower in the Cenomanian-Turonian transition beds of Unit 3. The unconformity between the Santonian and the Miocene represents a gap in sedimentation of at least 66 m.y.

## **Preliminary Conclusions**

Drilling at Site 258 did not penetrate deep enough for us to resolve the continental or oceanic nature of the crust here. We can say, however, that deep-water marine sediments have been accumulating here since at least middle Albian times. Unit 5 accumulated below the carbonate compensation depth and in cold water (highlatitude) conditions as suggested by the low foraminiferal species diversity. Unit 4 accumulated within the lysocline and in more temperate conditions. Thus, there is a history of gradual shoaling of the sea floor or deepening of the carbonate compensation depth through the Cretaceous. The Upper Cretaceous flora and fauna are essentially similar to those seen at Broken Ridge in the Upper Cretaceous limestone unit, which may or may not be significant.

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APPENDIX A Grain-Size Determinations for Site 258

Core, Section, Top of	Subbottom				
Interval	Depth	Sand	Silt	Clay	
(cm)	(m)	(%)	(%)	(%)	Classification
Hole 258					
1-2, 90	2.4	22.4	50.1	27.5	Sand-silt-clay
1-2,90	2.4	22.4	50.1	27.5	Sand-silt-clay
1-5,44	6.4	46.4	23.4	30.2	Sand-silt-clay
2-2,80	11.8	4.0	17.6	78.5	Clay
3-1,90	48.4	9.9	33.9	56.2	Silty clay
3-3,90	51.4	8.9	32.2	58.9	Silty clay
3-5,90	51.4	16.8	37.2	46.0	Silty clay
4-1, 130	86.8	40.6	37.6	21.8	Sand-silt-clay
5-2,90	125.9	7.0	21.1	71.9	Silty clay
6-2,90	144.9	0.4	28.4	71.2	Silty clay
6-5,90	149.4	0.5	28.3	71.3	Silty clay
7-2,90	154.4	1.3	36.2	62.6	Silty clay
9-1,81	181.3	0.2	15.9	83.9	Clay
10-2, 71	201.7	0.2	11.0	88.9	Clay
10-2, 84	201.8	0.6	33.0	66.4	Silty clay
12-2, 75	236.8	0.1	29.4	70.5	Silty clay
12-4,86	239.9	0.0	31.6	68.4	Silty clay
12-6,91	242.9	0.4	45.1	54.5	Silty clay
13-2,45	255.4	0.1	35.2	64.7	Silty clay
14-1, 3	263.0	0.7	29.1	70.2	Silty clay
15-2, 135	284.9	0.0	19.9	80.1	Clay
15-5, 120	289.2	0.0	28.6	71.4	Silty clay
16-2, 100	303.5	0.0	20.0	80.0	Clay
16-5,90	307.9	0.0	26.0	74.0	Silty clay
17-2,90	322.4	4.2	25.6	70.2	Silty clay
17-5, 18	326.2	0.4	20.1	79.5	Clay
18-2, 101	341.5	0.0	13.2	86.8	Clay
18-4,93	344.4	0.0	9.6	90.4	Clay
20-1, 101	378.0	0.0	16.1	83.9	Clay
21-1,60	406.1	0.0	16.7	83.3	Clay
21-3,90	409.4	0.0	17.6	82.4	Clay
22-2, 100	437.5	0.1	16.8	83.1	Clay
22-5,88	441.9	0.1	3.2	96.7	Clay
23-2, 88	474.4	0.0	12.1	87.9	Clay
24-2, 128	512.8	0.1	6.7	93.2	Clay
Hole 258A					
1-1, 14	.1	27.2	39.5	33.4	Sand-silt-clay
1A-2,90	2.4	29.8	34.2	36.0	Sand-silt-clay
1-2,90	2.4	29.8	34.2	36.0	Sand-silt-clay
2-2, 15	20.6	26.4	24.7	48.9	Sand-silt-clay
3-2,90	31.9	29.7	27.3	43.0	Sand-silt-clay
3-5, 100	35.5	22.2	27.7	50.1	Sand-silt-clay
3-2,90	30.9	13.2	27.2	59.6	Silty clay
4-5, 100	45.0	23.7	31.4	45.0	Sand-silt-clay
5-2,90	49.9	12.6	33.5	53.9	Silty clay
5-5,90	54.4	9.8	38.8	51.4	Silty clay
6-3, 122	70.7	4.3	30.4	65.3	Silty clay
6-5,90	73.4	. 7.6	35.5	56.9	Silty clay
7-3,90	98.9	21.9	53.6	24.5	Sand-silt-clay
7-5,90	101.9	33.9	51.8	14.3	Sandy silt
8-5,30	110.8	60.7	25.1	14.2	Silty sand
9-2, 118	116.7	26.0	21.4	52.6	Sand-silt-clay
9-5,70	120.7	6.2	29.5	64.3	

Core, Section Top of	Sub bottom	Total	Organic			
Interval (cm)	Depth (m)	Carbon (%)	Carbon (%)	CaCO <sub>3</sub> (%)		
(cm)	(11)	(%)	(70)	(70)		
Hole 258						
1-2,88	2.38	10.5	2.5	67		
1-5,88	6.88	10.6	1.0	80		
2-2, 85	11.85	11.2	1.3	83		
3-1, 78	48.28	11.2	0.2	92 77		
3-3,88	51.38 54.38	11.2 10.9	1.9 1.9	75		
3-5,88 4-1,130	86.80	8.6	0.1	70		
이상에 가지 않아야 한 것이 같아요.	125.88	9.3	0.1	74		
5-2, 88 6-2, 88	144.88	10.7	1.2	79		
6-5, 88	149.38	10.7	1.2	78		
7-2,88	154.38	10.3	1.0	77		
9-1, 80	181.30	7.8	0.4	62		
10-2, 70	201.70	7.3	0.3	59		
11-2, 75	217.75	5.2	0.1	42		
12-2, 74	236.74	8.1	0.1	67		
12-4, 85	239.85	8.8	0.1	72		
12-6, 89	242.89	6.2	0.1	51		
13-2, 44	255.44	8.1	0.1	67		
13-4,90	258.90	5.7	0.3	45		
14-1, 103	264.03	0.1	0.1	0		
15-2, 140	284.90	1.2	1.1	1		
15-5, 119	289.19	1.6	1.5	1		
16-2, 103	303.53	1.8	1.2	5		
16-5,90	307.90	1.3	1.1	2		
17-2,88	322.38	1.1	0.8	3		
17-5, 72	326.72	2.7	1.3	11		
18-2, 100	341.50	1.6	0.5	10		
18-4,90	344.40	2.5	1.0	13		
20-1, 107	378.07	1.7	0.8	7		
21-1,41	405.91	1.6	1.1	4		
21-3,88	409.38	2.1	1.5	5		
22-2, 99	437.49	3.3	2.6	6		
22-5, 91	441.91	3.6	2.7	8		
23-2,90	474.40	1.6	1.6	0		
24-2,95	512.45	1.3	1.3	0		
24-5, 50	516.50	2.5	0.5	16		
25-2, 85	521.85	0.3	0.3	0		
25-4, 55	524.55	0.2	0.3	0		
Hole 258A						
1-1,4	.04	10.8	0.1	89		
3-2, 88	30.88	10.9	0.1	90		
3-5,99	35.49	10.4	0.1	86		
4-4,80	43.30	10.9	0.1	90		
5-5,88	54.38	11.0	0.1	91		
6-3, 126	70.76	10.7	0.1	88		
6-5,88	73.38	11.0	0.1	91		
7-3,88	98.88	8.6	0.1	71		
7-5,88	101.88	8.6	0.1	70		
8-5, 29	110.79	9.5	0.1	79		
9-2, 117	116.67	8.2	0.1	68		
9-5,68	120.68	9.8	0.1	81		

-

Core	Cored Interval Below Sea Floor (m)	Sample Depth Below Sea Floor (m)	Diff.	Amor.	Calc.	Arag.	Quar.	Cris.	K-Fe.	Plag.	Kaol.	Mica	Mont.	Trid.	Clin.	Hema.	Pyri.	Gyps.	Bari.	Hali.	Magn.	U-8 <sup>a</sup>
Bulk S	Samples																					
1	0.0-9.5	6.5	63.1	42.4	76.3	16.1	1.8	-	3.7	144	<u></u>	23	-	:	-	-	2.0		<u>н</u> е		-	-
3	47.5-57.0	48.4	57.2	33.2	95.4		4.6	-	-		-		-	-	-	-	-		-	-	—	_
4	85.5-95.0	86.8	63.7	43.3	78.5	-	14.3		5.9	-	1.3		1000	-	$\sim - 1$	-	-	-	=		_	-
6	142.5-152.0	149.7	56.9	32.6	83.2		0.7	3.1			-		10.2		$\rightarrow$		2.7		-	-	-	—
10	199.5-206.0	201.6	60.1	37.7	69.1	-	0.6	8.0	-	+	-	+	18.0	0.5	-		2.3	-	1.4			—
12	234.5-244.0	240.3	77.0	64.1	22.8		1.6	40.8	-				30.2	2.2		-	0.7		1.8	-	-	
14	263.0-272.5	263.0	83.0	73.4		-	2.4	55.9		2.0	2.5	2.5	11.0	4.1	16.8	1	0.4	1000	2.3	5.5		-
		263.5	82.7	73.0	-	_	3.8	46.9	-	-	3.5	2.3	19.1	0.7	19.4	-			4.3	-	-	_
15	282.0-291.5	284.7	79.5	67.9	6.4	-	3.7	20.0	-	-	6.9	4.6	55.6	0.6	1.2	-	—	-	1.0	100		1
16	301.0-310.5	308.1	73.7	58.9	18.3	-	3.8	6.5	-	-	9.2	3.8	56.2	-	2.2	-	-	-		_	- 22	
17	320.0-329.0	327.4	80.7	69.8	10.1		4.3	27.4	-		6.4	6.7	43.8	1.3	$\rightarrow$ :	-	-	-	$\rightarrow$			-
18	339.0-348.5	341.5	83.7	74.5	8.6	1.77	2.7	59.5			3.5	3.9	17.9	3.2	0.7			1.77	-	-		-
21	405.5-415.0	405.7	79.6	68.1	-	-	12.0	39.2	1.7	-	4.7	7.4	30.8	4.1	c := c	-		$\sim$	-	_	-	$\sim$
22	435.0-444.5	441.7	77.1	64.2	1000	-	17.8	-		-	5.1	4.6	71.3	-	$\sim -1$		1.2	200	-			=
23	472.0-481.5	474.4	76.1	62.7	_	-	5.5		-		16.1	6.5	71.9		-	-	1000	50		-		
24	510.0-519.5	516.5	81.6	71.2	57.6	-	1.4	-	6.3		5.8		23.3				5.6	-	$\rightarrow$	-		-
25	519.5-525.0	521.9	90.6	85.3	-		3.3	-	15.6	-	57.2	-1	13.0		$(-1)^{-1}$	10.9	-	-	$\rightarrow$	-		Р
2-20µ	Fraction																					
1	0.0-9.5	6.5	80.0	68.7	-	-	41.1	-	22.5	6.8	7.5	11.4	-	-	4.5	-	6.3	-		-		-
3	47.5-57.0	48.4	78.9	67.0	-	1000	57.3		15.8	5.3	7.4	11.1	-	-	3.1			$\sim -1$	÷		100	
4	85.5-95.0	86.8	77.6	65.0	4	-	55.9	_	20.3		11.9	8.9	-	<u> 222</u>	2.9	-	_	-				_
6	142.5-152.0	149.7	71.4	55.2	$\leftrightarrow$	-	29.6	9.2	3.3	1.4		14.7	10.0	2.7	8.1	-	6.2	-	15.0	-		$\sim - \sim$
10	199.5-206.0	201.6	80.3	69.3	-	-	7.4	34.4	-		_	4.5	38.2	3.7	0.9	-	0.9	-	10.0	-		-
12	234.5-244.0	240.3	79.7	68.3		-	7.3	64.1	1.5	1.0	$\rightarrow$	4.5	12.5	3.3	0.8	-	-	_	5.1	-	-	-
14	263.0-272.5	263.0	70.7	54.2	-		3.6	57.3	-		1.5	1.8	-	4.6	22.0	-	0.8	$\sim -1$	8.3	-		-
11	20010 21210	263.5	70.7	54.3	-	_	6.3	34.1	_		2.0	2.3	100	2.7	41.5	-	2.5		8.6	1.5	122	_
15	282.0-291.5	284.7	68.0	50.0	-	1	14.6	-	2.8		18.2	6.8	27.8	_	28.0	_	_		1.7	-	-	-
16	301.0-310.5	308.1	77.8	65.4	-	-	14.3	23.2	3.2	-	13.8	9.7	31.5	_	1.5	-	_	_	2.9	_	_	_
17	320.0-329.0	327.4	79.3	67.7		1	13.3	17.8	4.6		16.0	16.2	26.7	2.1	1.5	-	1.1	_	2.3			_
18	339.0-348.5	341.5	80.2	69.1	_		9.1	65.9	3.1	0.9	6.3	6.7	-	3.1	2.3	_	0.6	-	2.0	-	-	
	405.5-415.0	405.7	74.8	60.6	-	_	23.2	24.9	7.5	1.5	9.0	9.9	22.5	-	2.5	2	0.0		1.5		1	-
21				53.7			32.9	24.9	11.0	1.5	9.0	9.9	22.5	12		_	4.8	- 27	1.5			
22	435.0-444.5	441.7	70.4		-	-				1.7	1.202.002	12.4	25.8 39.8	_	-		4.8					
23	472.0-481.5	474.4	67.8	49.8	_	1	11.3		5.0		25.5			-				-	57-5 10-1		0.00	2
24	510.0-519.5	516.5	80.7	69.8	_	_	8.5	-	31.7		4.4	1.9	44.3	-		10.0	9.2	_	-		22.2	
25	519.5-525.0	521.9	84.9	76.3	-		6.1	-	31.4	$\sim$	11.6		7.8	-		19.8	-	-	220	1	23.2	Р

APPENDIX C X-Ray Analyses for Site 258

Core	Cored Interval Below Sea Floor (m)	Sample Depth Below Sea Floor (m)	Diff.	Amor.	Calc.	Arag.	Quar.	Cris.	K-Fe.	Plag.	Kaol.	Mica	Mont.	Trid.	Clin.	Hema.	Pyri.	Gyps.	Bari.	Hali.	Magn.	$U-8^{a}$
<2µ	Fraction																					
1	0.0-9.5	6.5	92.2	87.7	-	-	9.5	-	-	~	42.0	14.0	30.6	-	-	-	3.9	-	-	-	-	5 <u>-</u>
3	47.5-57.0	48.4	86.9	79.5		-	11.8	$\sim - 1$	-	-	48.0	17.5	16.0				-	6.7	$\sim - 1$	100		
4	85.5-95.0	86.8	85.1	76.8		-	7.4	-	-	$\simeq$	36.2	14.9	39.0	_	-		—	2.5	_		-	
6	142.5-152.0	149.7	85.1	76.7	-	$\rightarrow$	3.5	38.9	-	-	-	4.7	38.4	2.0	1.2		0.9	-	7.4	2.9	-	
10	199.5-206.0	201.6	82.4	72.5	-	-	3.9	42.5	-	-	-	2.6	28.3	1.5	-	-	0.3	-	10.9	10.1	-	
12	234.5-244.0	240.3	81.7	71.4		-	0.3	50.6	-	-		2.4	37.9	1.7	-	<u></u> 5	0.4	-	1.8	4.9	-	-
14	263.0-272.5	263.0	85.0	76.6		-	2.0	64.0	-	-	1.7	1.4	19.8	5.8	0.5	<del></del>	-	-	1.4	3.6	-	-
		263.5	89.0	82.9		-	5.5	64.9	-	_	2.2	_	15.2	6.5	_	_	0.6		3.0	2.0	-	
15	282.0-291.5	284.7	79.4	67.9		-	12.5	10.3		-	13.5	4.0	40.2	-	6.5		-	-	4.4	8.6	-	-
16	301.0-310.5	308.5	84.0	74.9		-	5.0	30.0	-	-	6.7	15.2	32.2		-				2.2	8.7	-	-
17	320.0-329.0	327.4	81.8	71.6		-	8.7	42.5		$\sim$	4.9	3.0	29.8	1.7	-		-	-	1.7	7.7	-	-
18	339.0-348.5	341.5	84.3	75.5		-	1.9	67.0	-		1.7	0.8	17.3	4.0	$\sim - 1$		-	-	0.9	6.5	-	-
21	405.5-415.0	405.7	80.6	69.7	-	-	11.0	38.8	-	-	2.5	14.4	10.4	-	-		_	_	_	22.8		100
22	435.0-444.5	441.7	84.7	76.1	-	-	43.8	=7	-	_	8.2	4.6	12.7		_			-	_	30.6	-	-
23	472.0-481.5	474.4	80.2	69.1		-	8.9		-	-	19.0	6.6	52.2	-	$\sim$	-		-		13.4	-	
24	510.0-519.5	516.5	94.2	90.0	-		522			-	12.4	11.8	59.2		1	1.5	1.1.1	-	_	16.6	<u></u>	-
25	519.5-525.0	521.9	93.0	89.0	-	-	2.1	$\sim 10^{-1}$	-	-	59.5	-	21.4	-		11.9	-	-	-	5.0	-	-

APPENDIX C - Continued

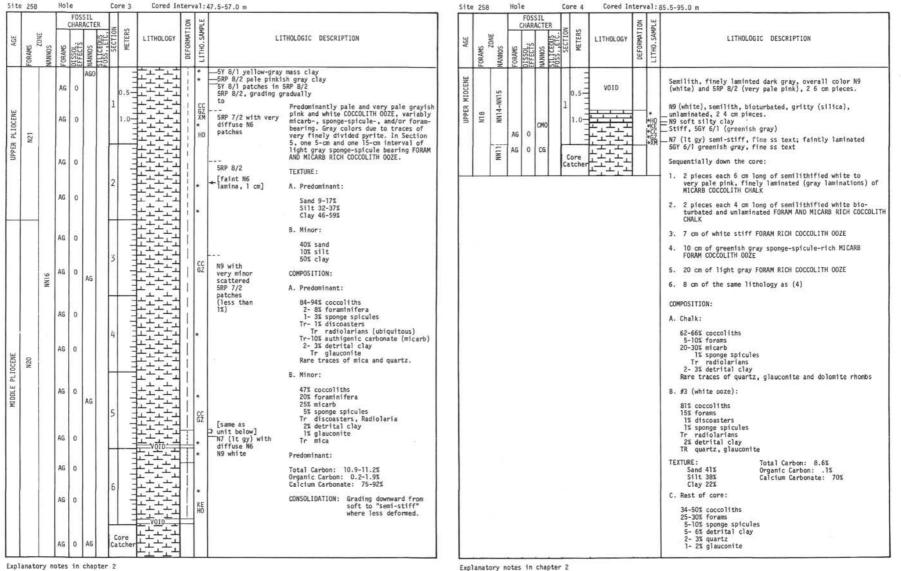
<sup>a</sup>Narrow peaks at 2.753Å and 3.006Å among others. P = present.

Site 25	•	Hole	_		LO	ore 1 Cored In	iter	val:	0.0-9.5 m		510	e 25	8	Но	1.7		Cor	*e 2	Cored Int	erva	1:9	.5-19.0 m	
AGE FORAMS	ZONE NANNOS	C	EFFECTS HAL	ACTER	FOSS., ETC. SECTION	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LIT	ITHOLOGIC DESCRIPTION	AGE	FORAMS	ZONE		FOSS CHARA 10SSOL	CTER	FOSS. ETC. SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	u	THOLOGIC DESCRIPTION
UATERNARY N22-N23	NN20	AG AG		AG AG AG	1 2 3 4 5 6	1.0.5.1444444444444444444444444444444444		* * * * * * * *	with gray and TEXT Pred. N7 (lt gy); few COMP chunks 5YR 8/1 A. P (pink) 8 N7 B. M N7 B. M Stretched 1 5YR 8/1 chunk] FR 8/1 chunk] Stretched Calc Calc	ht gray FORAM-BEARING COCCOLITH 002E, th minor pinkish gray patches and yellowish y intervals of FORAM RCH COCCOLITH 002E FORAM COCCOLITH 002E. ITURE: Sand 22-46% Silt 23-50% Clay 28-30% POSITION: Predominant: 82-90% coccoliths 1- 8% forams 2- 3% diatoms Tr radiolarians Tr sponge spicules 5- 6% detrital clay Minor (FRCO & FCO): 63-90% coccoliths 10-15% forams Tr radiolarians 2- 5% sponge spicules 5% detrital clay Minor (FRCO & FCO): 63-90% coccoliths 10-15% forams Tr radiolarians 2- 5% sponge spicules 5% detrital clay Tr glauconite Rare traces of quartz, mica, and dolomite rhombs tal Carbon: 1.0-5.2.5% IsioLIDATION: Soft.	Curternary		6LW 6LW	AG AG AG AG	0	AG AG AG	1 2 3 Ca	0.5			GZC CC	N7 layer with sharp upper bound., gradstional lower bound.] [2 cm N6 patch] Mixed N7 and 56Y 6/1 It ol gy. Coarser than N7 N7	Except for 28 om interval near the base of the core, the material is lig jary SPONGE AND FORM BEARING OCCOLI 002E with rare dark layers and deform laminae which are colored by traces of very finely divided pyrite. The 28 cm interval in Section 3 consis of SPONGE, MICARB AND FORAM RICH 00CCOLITH 002E and SPONGE RICH MICARB FORAM 002E, greenish gray in color. TEXTURE: A. Predominant (SFBCO): Sand 4.0% Silt 18% Clay 78% B. Minor (MFRCO and SRMFO)(Smear slide 30-55% sand 15-20% silt 25-55% clay COMPOSITION: A. Predominant (SFBCO): 79-88% coccoliths 3 - 8% forams Tr - 2% radiolarians 3 - 5% detrial clay Traces of glauconite Section 3 only: 3% micarb B. Minor (MFRCO and SRMFO): 20-35% forams 6-43% coccoliths 10-20% sponge spicules 1- 2% radiolarians 3 - 5% detrial clay Tr - 1% glauconite 1% diatoms 25-30% micarb 5% detrial clay Tr - 1% glauconite 1% guartz Total Carbon: 11.2% Organic Carbon: 1.3% Calcium Carbonate: 83% CONSOLIDATION: Soft to stiff.

SITE 258

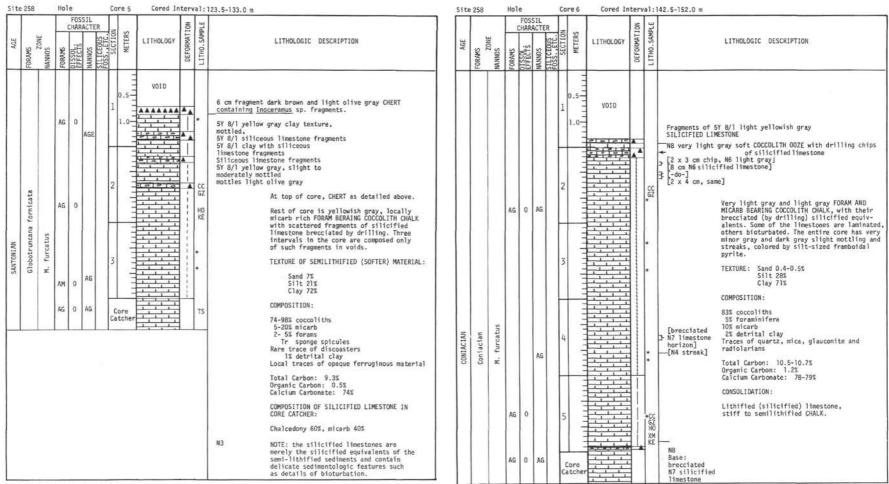
Explanatory notes in chapter 2

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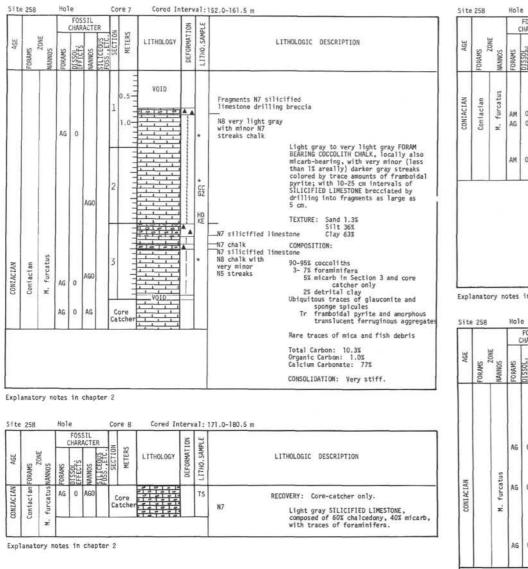


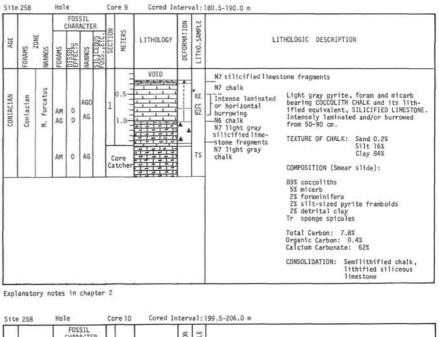
SITE 258

375



Explanatory notes in chapter 2





		_	FOSS HAR/		R	s		LON	MPLE		
AGE	FORAMS ZONE NANNOS	FORAMS	DISSOL.	NANNOS	FOSS, ETC.	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE	LIT	HOLOGIC DESCRIPTION
CONIACIAN	M. furcatus	AG AG AG	0	AG AGO	1	0.5			HANCAR	N7 and N8 limestone 	Light gray and very light gray COCCULTH CHALK and lithified equiv- alent SILICHIED LIMESTONE, typically micarb and foram bearing, locally pyritiferous as well, the pyrite occurring as framboids and radiolarian replacements. Abundant bioturbation. TEXTURE: Sand 0.2-0.6% Silt 11-33% Clay 66-89% COMPOSITION: 88-89% coccoliths 3-5% micarb 5-7% foraminifera 2% detrital clay Local traces of quartz and sponge spicules; locally up c% pyrite. Total Carbon: 0.3% Calcium Carbonate: 59% CONSOLIDATION: Semi-lithified chalk, lithified imestone. N.B.: Silicification and resultant hardness is variable in the lithestones.

SITE 258

Site	258		Hol	е			Co	re 11	Cored In	terv	al:	215.0-225.0 m	
				FOS	SIL	R				NOI	SAMPLE		
AGE	FORAMS	20NE NANNOS	FORAMS	DISSOL. EFFECTS	NANNOS	FOSS., ETC.	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAM	LITHO	LOGIC DESCRI
		5	FP	0	CGO AGO		1	1.0		A.		N8 chalk <u>56</u> Y 6/1 & 56Y 7/1 Same, brecciated N8 chalk N8 brecciated chalk ;48 chalk, frag-	MICARB-RIC locally hi content to the top of very light digitation with depth gray inter nent, atta scolored in and burrow in interva together w

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1 1

Core

Catche

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colored types.

IPTION ICH COCCOLITH CHALK. high enough in foram to be foram bearing; at of the core predominantly ht gray with fewer interons of greenish gray, but th in the core the greenish ervals become more promi-taining 50% of the core in 3. Contacts between the intervals are gradational, ows of one color are found vals of the other color, together with light olive gray ments up to 6 cm Interdigitated 90% burrows in both colors. The burrows are generally horizontal to sub-5GY 6/1 & 10% 5Y 5/1 horizontal. light olive gray ch Same, severely brecci-ated by drilling The whole core is made up of 6-10 cm long pieces, with a few severely Chalk, N8 with many brecciated intervals as indicated. N7 burrows, fragments 6-10 cm long. Section 3: 50% 56Y 6/1 & 50% TEXTURE: Sand Tr-2% Silt 2-4% Clay 94-98% (Smear Slide) N8 chalk alternating pieces 6-8 cm long. Contacts are grada-tional. 5 mm wide sub-COMPOSITION: horizontal burrows of N8 in 5GY 6/1 intervals 71-72% coccoliths and vice versa, with 25% micarb 5Y 6/1 burrows in both 1- 2% foraminifera 2% detrital clay Rare traces of mica, glauconite and sponge spicules Total Carbon: 5.2%

# Organic Carbon: .1% Calcium Carbonate: 42% CONSOLIDATION: Semi-lithified. NOTE: The silicification which has

lithified intervals in the previous cores appears to be totally lacking in Core 11.

Explanatory notes in chapter 2

furcatus

magnificus-M.

×

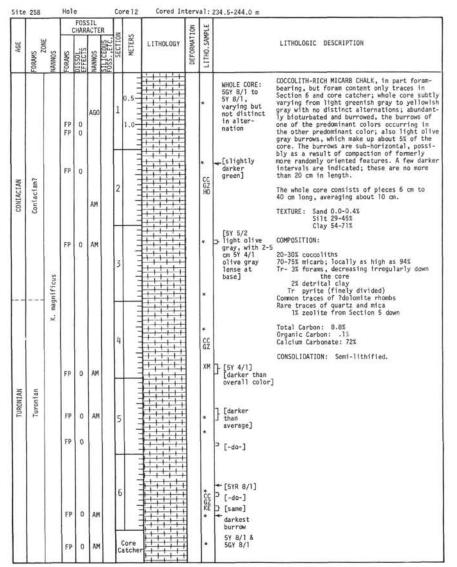
Contactan?

CONTACIAN

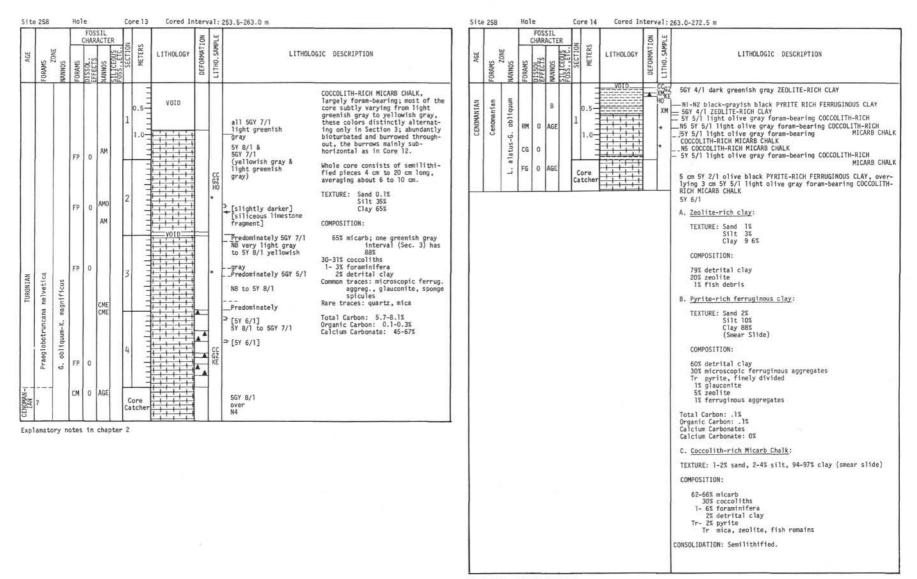
FP 0

CM D AG0

1000



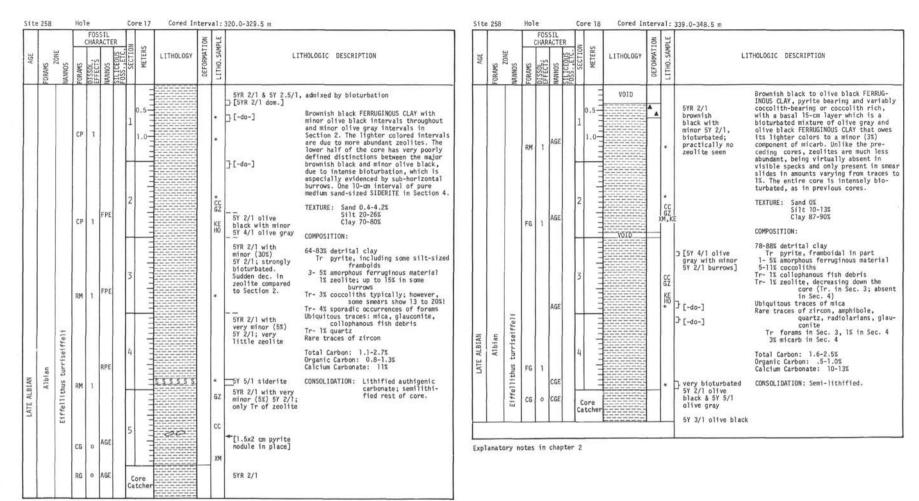
Explanatory notes in chapter 2



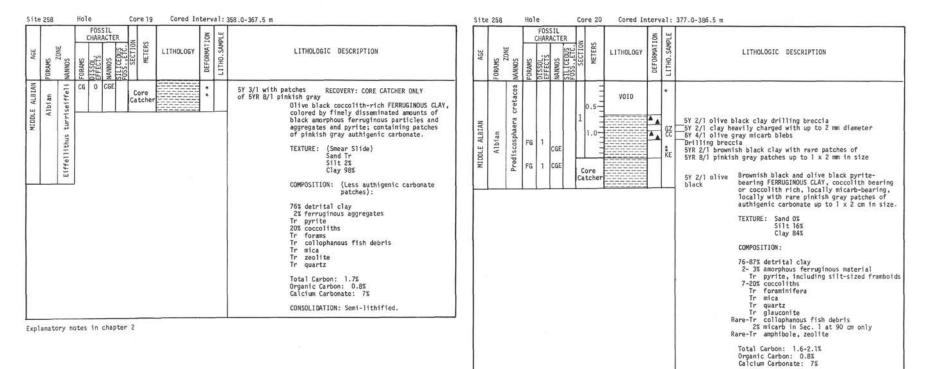
AGE S ZONE S		FOSS	11						282.0-291.5 m		e 258	-	Но			 re 16	corea m	reerv	41. 5	301.0-310.5 m
FORAMS Z( NANNOS		DI SSOL.	ACTER	F05S., ETC.	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE	FORAMS	ZONE	FORAMS		ACTER		LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC DESCRIPTION
LATE ALBIAN Albian-Cenomanian Eiffellithus turriseiffeli L. alatus	FG CG FM FG FG	0	AGE AME CME FPE AME	1 2 3 4 5 6	0.5- 1.0-			· · · · · · · · · · · · · · · · · · ·	<pre>SY 6/1 light olive gray zeolite and micarb bearing _COCCULTH-RICH CLAY "N2 grayish black FERRUGINOUS CLAY SY 6/1 zeolite &amp; micarb bearing COCCULTH CLAY with vertical _J mm N3 burrows 2 cm long Alternating N3 pyrite &amp; micarb bearing COCCULTH RICH CLAY and SY 6/1 light olive gray zeolite &amp; micarb bearing COCCULTH RICH CLAY _SY 6/1 light olive gray and SY 4/1 olive gray zeolite &amp; micarb bearing COCCULTH RICH CLAY SY 6/1 light olive gray and SY 4/1 olive gray zeolite &amp; micarb bearing COCCULTH RICH CLAY _SY 6/1 greenish gray zeolite &amp; micarb bearing COCCULTH SY 3/1 olive gray to SY 2/1 olive black RICH CLAY SY 6/1 greening COCCULTH RICH CLAY _SY 2/1 olive black pyrite &amp; zeo;ite bearing COCCULTH RICH CLAY SY 2/1 olive black pyrite &amp; zeo;ite bearing COCCULTH RICH CLAY _sy 2/1 olive black pyrite &amp; zeo;ite bearing COCCULTH RICH CLAY _sy 2/1 olive black zeolite bearing FERRUGINOUS CLAY NOTE: Speckled with fire sand-sized zeolites, which are also concentrated in burrows.</pre>	LATE ALBIAN	Albian-fenomențan	Eiffellithus turriseiffeli	RG	D	CPE	0.5	VOID		KE SZC HO * * CCZZ KE MY *	<ul> <li>14 om piece</li> <li>550 6/1</li> <li>Caving</li> <li>SY 2/1 = 303</li> <li>and/or foram bearing; locally (core catcher)</li> <li>cocolith-rich, intensely bioturbated and</li> <li>burrowed. The burrows tend to contain concentrations of white zeolite specks, which occur in lesser abundance scattered throughout the material. The burrows are invariably sub-horizontal and up to 7.5 mm wide.</li> <li>TEXTURE: Sand 0%</li> <li>Silt 20-26%</li> <li>Clay 74-80%</li> <li>COMPOSITION:</li> <li>69-87% detrital clay (average 81%; but 58% in core catcher)</li> <li>3- 5% amorph. microscopic ferrug. material Tr pyrite, including framboids</li> <li>1- 5% zeolite</li> <li>Tr- 1% mica</li> <li>2% micarb</li> <li>Rare traces to 1% occurrences of quartz, also fish debris</li> <li>Rare traces to 1% occurrences of quartz.</li> <li>also fish debris</li> <li>CONSOLIDATION: Semi-lithified.</li> </ul>

Explanatory notes in chapter 2

SITE 258

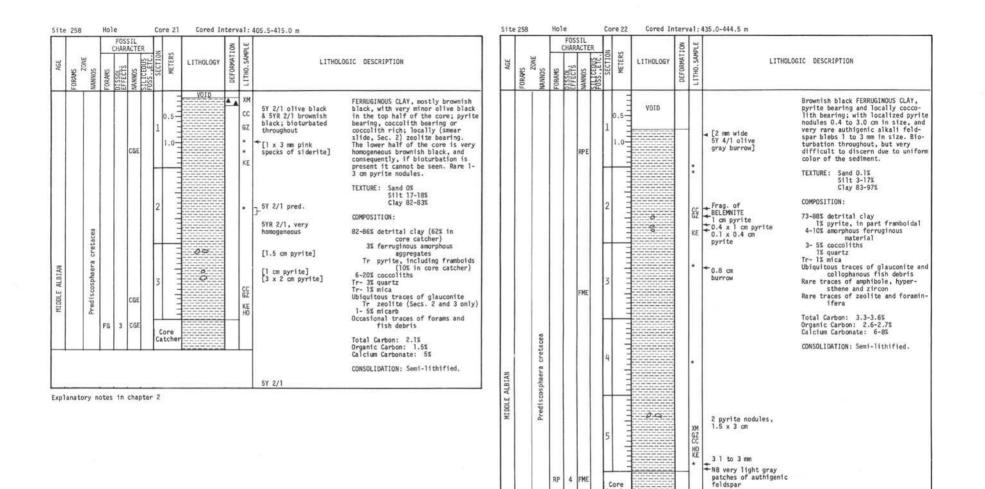


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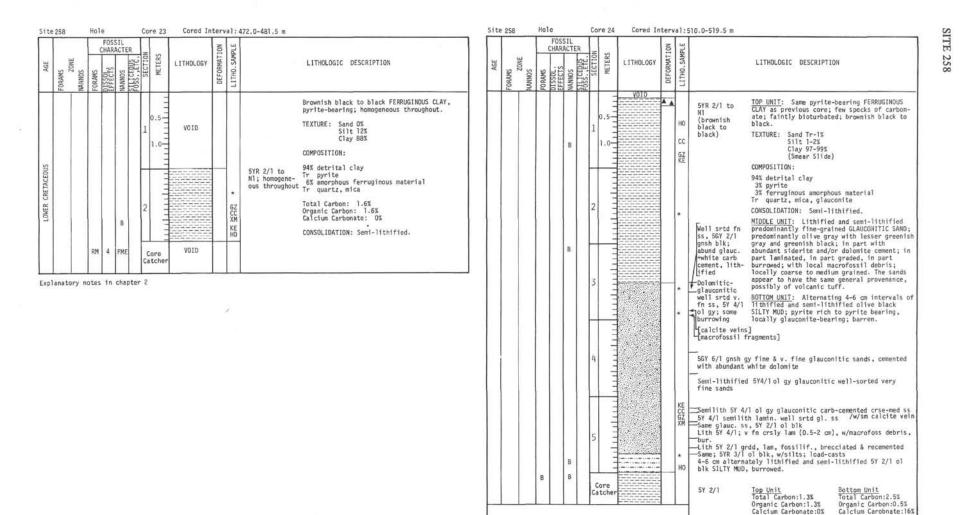
CONSOLIDATION: Semi-lithified.

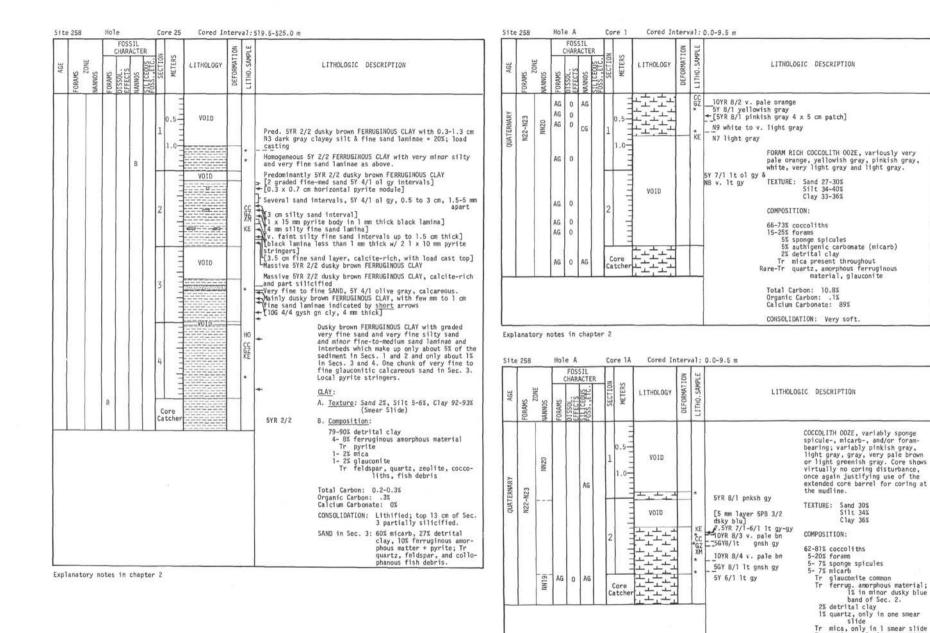


RP 4

Core Catche

N1





SITE 258

Tr pyrite, only 1 smear slide CONSOLIDATION: Very soft.

N8 1t gy

Explanatory notes in chapter 2

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ite 258	1	Hole	-	_		Co	re 2	_	C	ored	In	terv	/a1:	19.0-28.5 m			Site	258		Ho1e	10	Core :	3	Cored	Inter	val::	28.5-38.0 m	
AGE FORAMS ZONE NANNOS	NANNUS	FORAMS	EFFECTS AN	CTE	-	SECTION	METERS		.17	101.0	GY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION			AGE	FORMAS	NANNOS	CH	ARACTE NANNOS	SECTION		ITHOLOG	DEFORMATION	LITHO. SAMPLE	LITHOL	OGIC DESCRIPTION
NITENARRY REVARRY REVEALERVARRY	6 NN	AG AG AG AG AG AG AG		AG		1 2 3 Cca	ore				[4 [4]	*	*	Light gray, light i and minor dark gree COCCOLIT 002E, sym ticarb-bearing and bearing. TEXTURE: Sand 26% Silt 25% Clay 49% TEXTURE: Sand 26% Silt 25% Clay 49% COMPOSITION: 7.5YR 7/0 B4-88% coccoliths 1 - 7% forams 2 - 4% sponge spic 5% 4/1 dk gnsh gy 8 56% 7/1 gnsh gy N7 to N8	nish gray nge spicule locally for les rruginous tes	e- and	UPPER PLIOCENE QUATERVARY	N21 N22-N23	6LMN	AG A	) ) ) ) ) ) ) ) ) ) ) ) ) ) ) ) ) ) )	10.5	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~			* CCC GZ	2.5Y 7.5/0 lt gy 2.5Y 7.5/0 2.5Y 7.5/0 2.5Y 7.5/0 1 gy ←[2 x 5 cm 2.5Y 7.5/0 1 gy ←[5GY 5/1 3 x 2 cm patch] 2.5Y 7.5/0 [6 x 3 cm 5Y 8/1.5] =25GY 5/1 gnsh gy 2.5Y 7.5/0 [6 x 3 cm 5Y 8/1.5] =[typical 5Y 6/2 1 t ol gy mottle, 1 x 2 cm 5GY 8/1 + 56 8/1	COCCOLITH OOZE, predominantly light gray, with very minor light gray; yellowish white and greenish gray; imottles 1 x 2 on; sponge spicule bearing, and variably micarb- and/ or foram-bearing. TEXTURE: Sand 13-30% Silt 27-28% Clay 43-60% COMPOSITION: 62-89% forams 2 - 5% micarb 3 - 8% sponge spicules 2 - 3% detrital clay Tr flauconite Tr ferrug, aggreg. in Sec. 6 and core catcher, less common traces of mica One trace occurrence of radiolarians. Total Carbon. 1% CoNSOLIDATION: Soft.

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**SITE 258** 

Site 258	Hole A Core 4 Cored Interval: 38.0-47.5 m	258 Hole A Core 5 Cored Interval: 47.5-57.0 m	
AGE FORAMS ZONE	FOSSIL CHARACTER SOUNDAN SOUNDAN Soundary Sounda	SINVEGO	C DESCRIPTION
MIDDLE PLIOCENE N20	AG         0         AG         0.5         V010           1         0.5         0.5         0.5         0.5         0.5           1         1.0         1.0         1.0         1.0         1.0         1.0           AG         0         AG         0.5         0.5         0.5         0.5         0.5           AG         0         AG         0.5         0.5         0.5         0.5         0.5           AG         0         AG         0.5         0.5         0.5         0.5         0.0         0.5         0.0         0.5         0.0         0.5         0.0         0.5         0.0         0.5         0.0	$ \begin{array}{c c c c c c c c c c c c c c c c c c c $	And 10-13% Silt 34-39% Silt 34-39% Silt 34-39% Soliths ms arb of core rital clay zes of feruginous aggregates i of heavy minerals and ite framboids te tramboids te z bughout core, traces of DASIERS; in core catcher, 1%. on: 11.0% rbont . 1% rbonate: 91%

0 AG 58 9/1

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Core Catche

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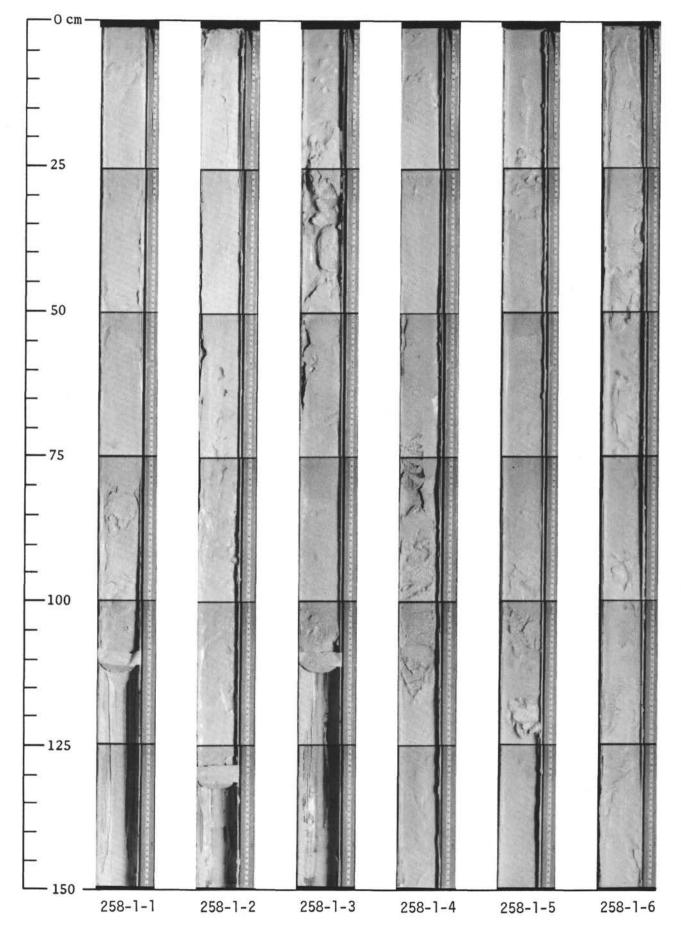
			t 1	055	11	- 1				1	- 1 h -								FOS	SSIL		- 12		1	1 I		
FORAMS	ZONE	NANNOS	FORAMS	ARA	NANNOS	FOSS., ETC.	METERS	LITH	OLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOG	IC DESCRIPTION		AGE	FORAMS ZONE	NANNUS	CHAS	RACT	FOSS., ETC. B	SECT ION METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE		LITHOLOGIC -DESCRIPTION
	2		AG AG AG AG AG AG AG AG AG AG AG AG AG A		AMO AMO AGO	1	0.5 1.0					56Y 6/1 gnsh gy 55 8 7/1 58 9/1, s1 moti w/5P 4/2 1t pri purple] 58 9/1, no motiling	y hish gy ksh /1] se gg, dia.	Predominantly bluish gray C ODZE, typically foram-beart locally micarb rick; with m grayish green COCCOLITH 002 micarb rick and 4-8% Silt 30-36% Clay 57-65% COMPOSITION: 68-99% coccoliths (ave. 80% I% Discos. 1-10% forams (ave. 5%) Tr-20% micarb (ave. 5%) Tr-20% micarb (ave. 5%) Tr-23% gouge spicules (ave 1-4% detrital clay (ave. Tr-1% jupite, in part silt framboids Tr-2% quartz Rare traces of mica Total Carbon: 10.7-11.0% Organic Carbonate: 88-91% CONSOLIDATION: "Semi-stiff	or ;ram ; Tr- 2%) %) sized	UPPER MIOCENE	BIN		0 3. 0 3.	CGO CGO CGO CGO CGO	0	22 33 44			* CCZZKEM * CCZZKEM * *	<pre>[ Mixed 58 9/1] 56 8/1 1 gnsh gy [ moderately mottled w/ 55 9/1, 0.5 cm mottles] 56Y 8/1 </pre>	COCCOLITH 002E. In sections 1, 2 and 3, light greenish gray, foram-bearing and sponge spicule - and micrab-rich. In Sec 4 and 5, greenish gray, foram-, micarb- sponge spicule-rich. SECTIONS 1-3: TEXTURE: Sand 22-34% Silt 52-54% Clay 14-24% COMPOSITION: 73-74% coccoliths 10-12% micarb 10% Sponge spicules 2% detrital clay Traces:glauconite, pyrite One trace of mica One trace of discoasters SECTIONS 4 and 5: TEXTURE: Sand 35% Silt 15% Clay 50% COMPOSITION: 43-44% coccoliths 10% forams 25% sponge spicules 2-3% quartz 5-6% detrital clay Tr- 1% glauconite and discoasters Total Carbon: 1.% COMSOLIDATION: Stiff where less deform

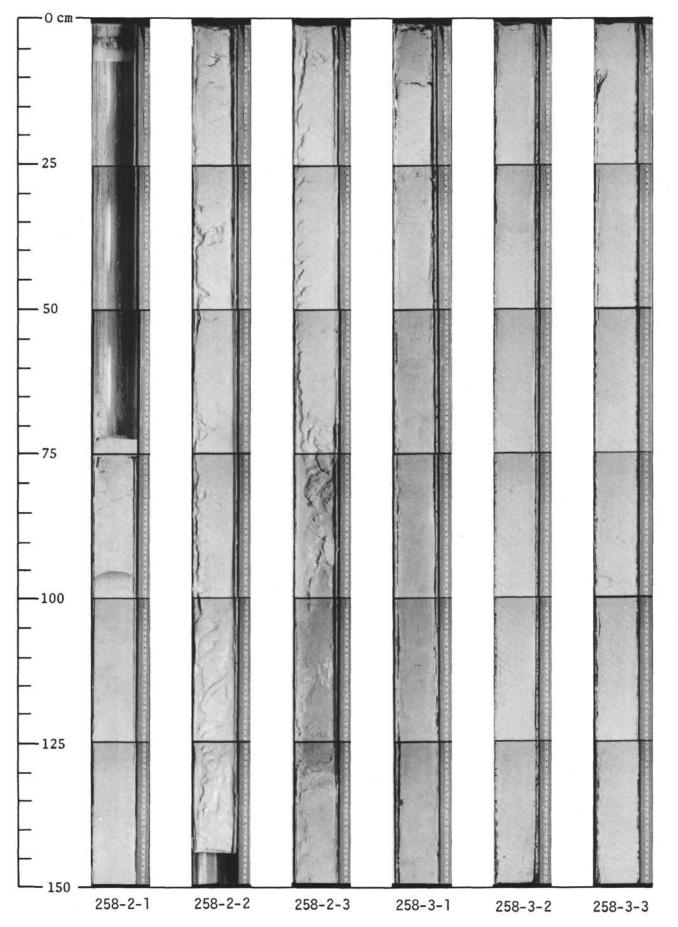
SITE 258

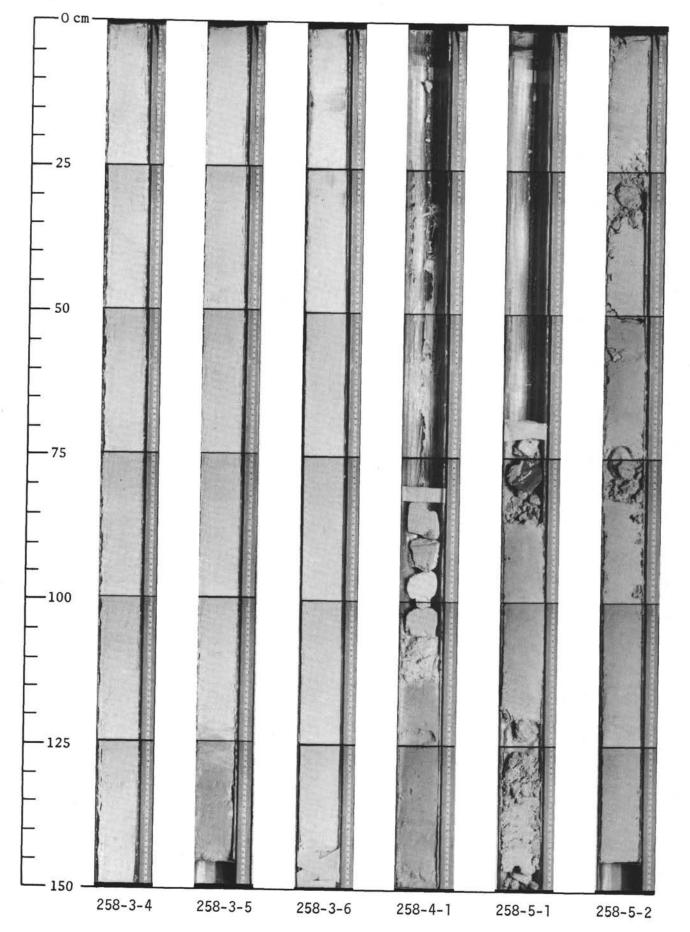
Site 258	Hole A		ore 8	Cored Int	terval:	104.5-114.0 m	Sit	te 25	8	Hol	e A		Cor	e 9	Cored I	nterv	a1:	14.0-123.5 m	
AGE FORAMS ZONE	FOSSIL CHARACT SUCCESSING SUCCESSING SUCCESSIN COCESSIN SUCCESSIN COCESSIN SUCCESSIN SUCCESSIN SUCCESSIN SUC	ER	METERS	LITHOLOGY	DEFORMATION LITHO.SAMPLE	LITHOLOGIC DESCRIPTION	ÅGE	FORAMS	ZONE		FOSS CHARA . TOSSIQ	CTER	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO. SAMPLE	LITHOLOGIC	DESCRIPTION
SANTONIAN UPPER MIOCENE NIA	AG     0-1     AG       AG     0-1     AG       AG     1     AG		0.5				SANTONIZAL					460 460 462 462 462 462 462	2 3 4 5 6 Cot				* KEHOMCCZ GZ * * *	Pr         pr           10YR 8/2         ST           yv. pale orange         da           5GY 8/1         ma           1t gmsh gy         sa           with mixed         th           minor         10YR 8/2           [2-3 cm]         patches           5GY 5/1         gnsh gy]           SGY 5/1         GOI           gnsh gy]         88           4         [2 x 7 cm SILIC           ON         Ra	<pre>ram bearing COCCOLITH 002E, edominantly light greenish gray Secs. 1-3, predominantly very le orange in Sections 4-6. In c. 3, one 2 cm thick layer of llowish gray SILICIFIED LIME- ONE which delicately preserves e bioturbation of the unconsoli- ted lithologically equivalent trfx. Small drill cuttings of the me SILICIFIED LIMESTONE occur roughout the badly deformed core. XTURE: Sand 6-26% Silt 27-30% Clay 53-64% MPOSITION: -91% coccoliths -5% forams 6% micarb 1% detrital clay Tr finely divided pyrite throughouts matter e 1% occurrence of zeolite ry rare trace of glauconite ngle trace occurrences of sponge spicules and collophanous fish debris tal Carbon: .1% Silt Carbon: .1% NSOLIDATION: Stiff ooze; lithified silticified lime- stone.</pre>

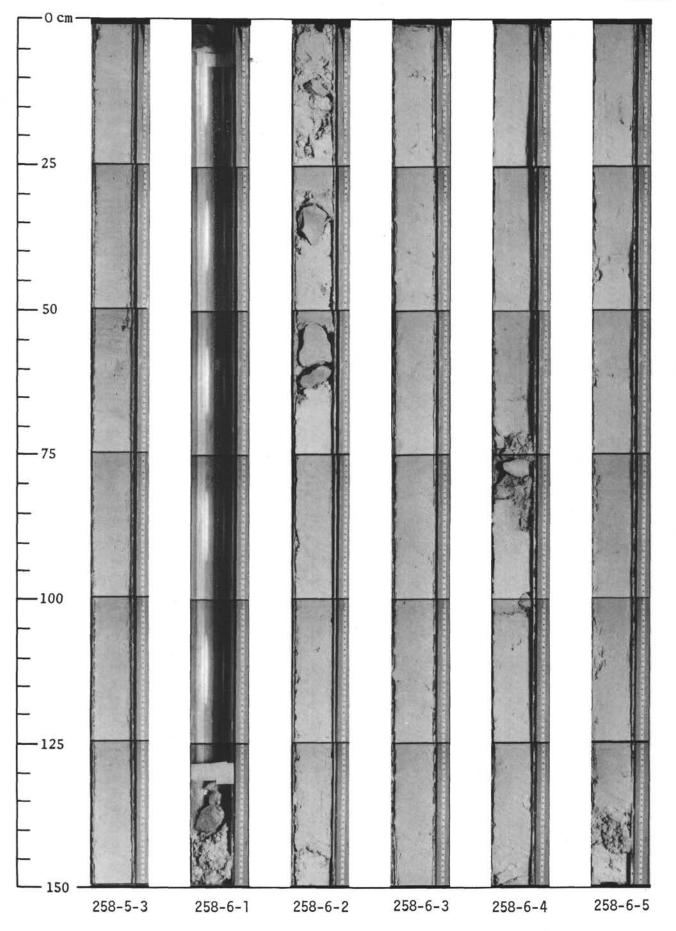
Explanatory notes in chapter 2

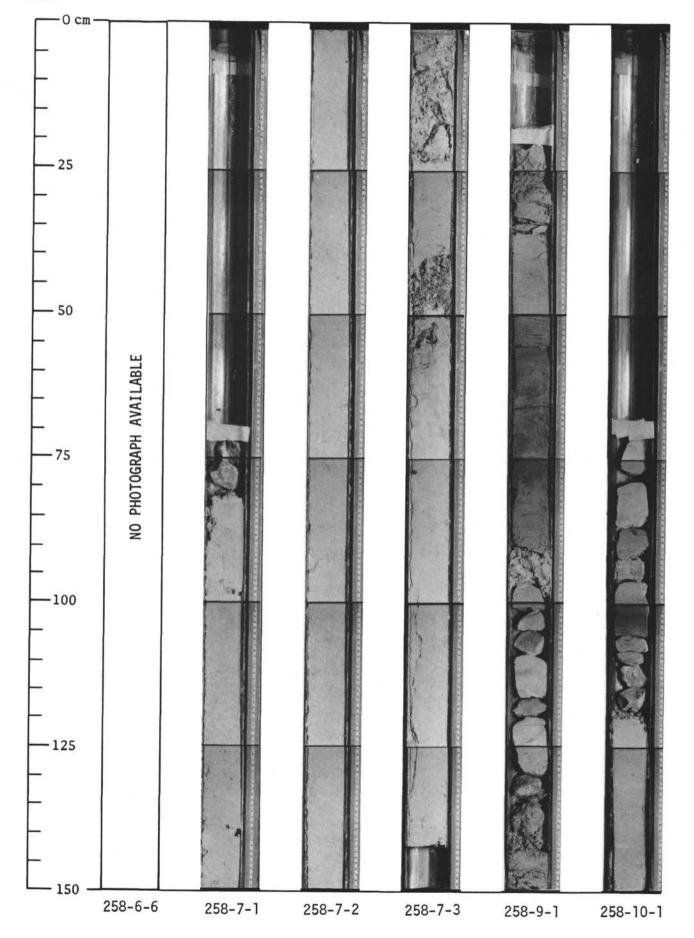
SITE 258

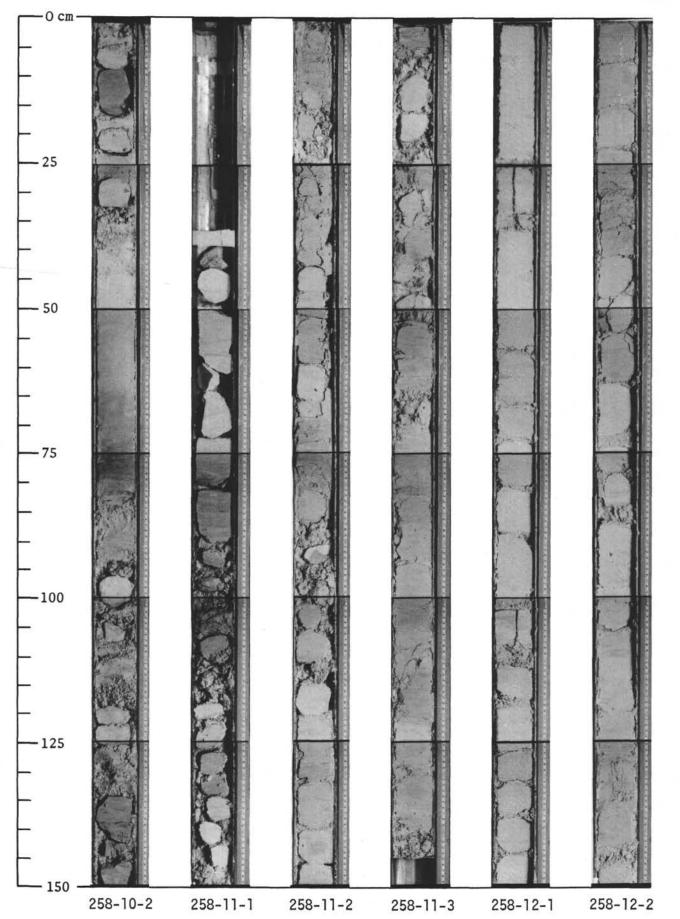


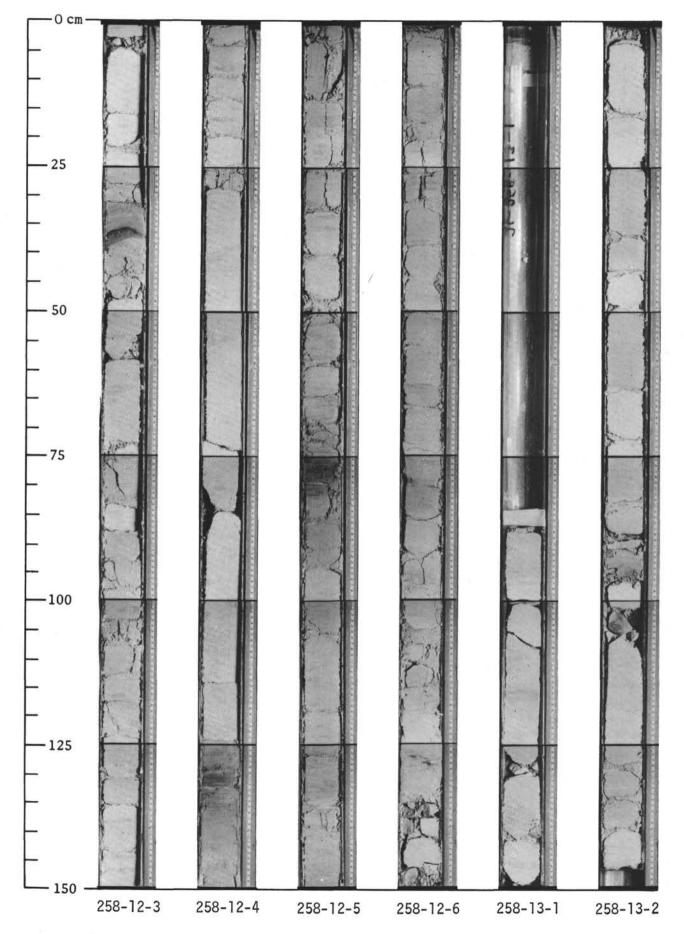


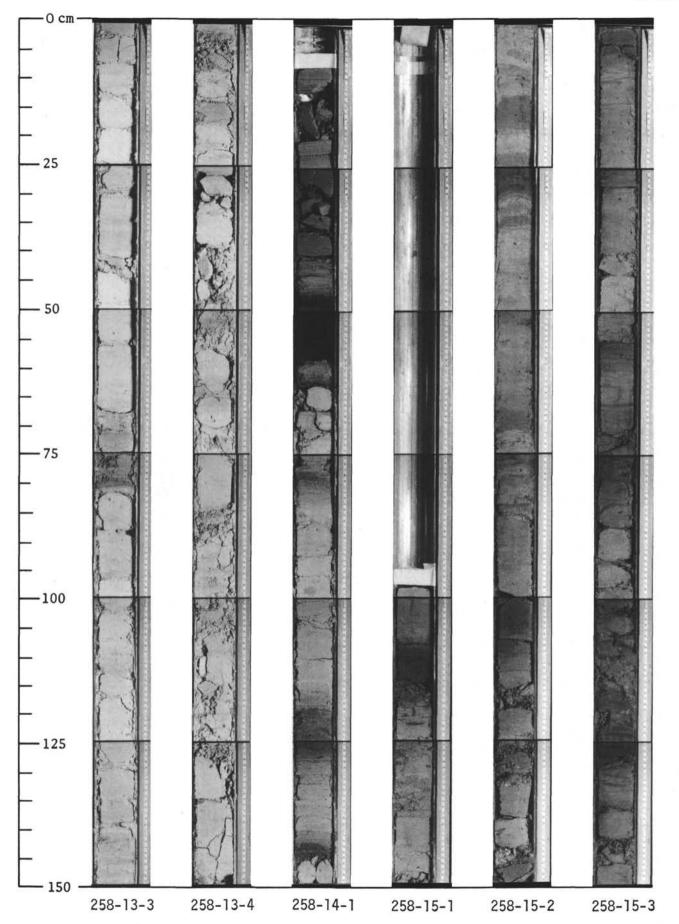


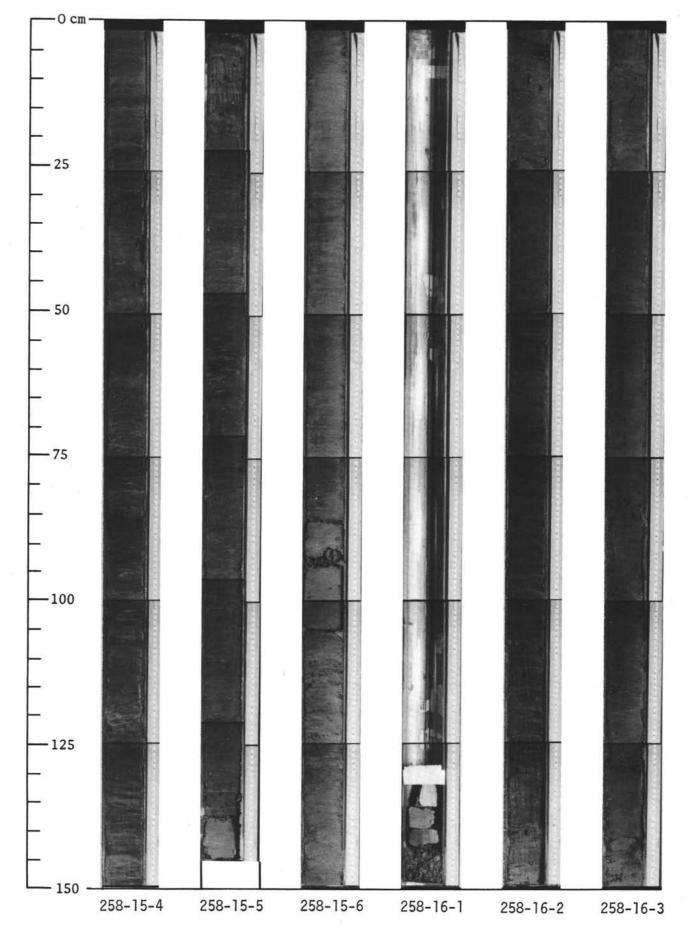


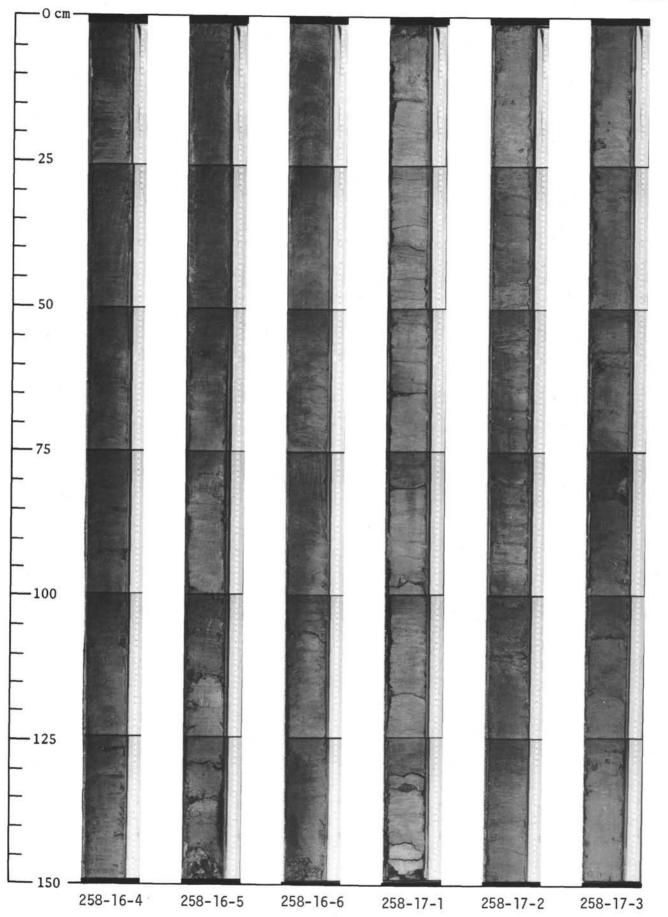


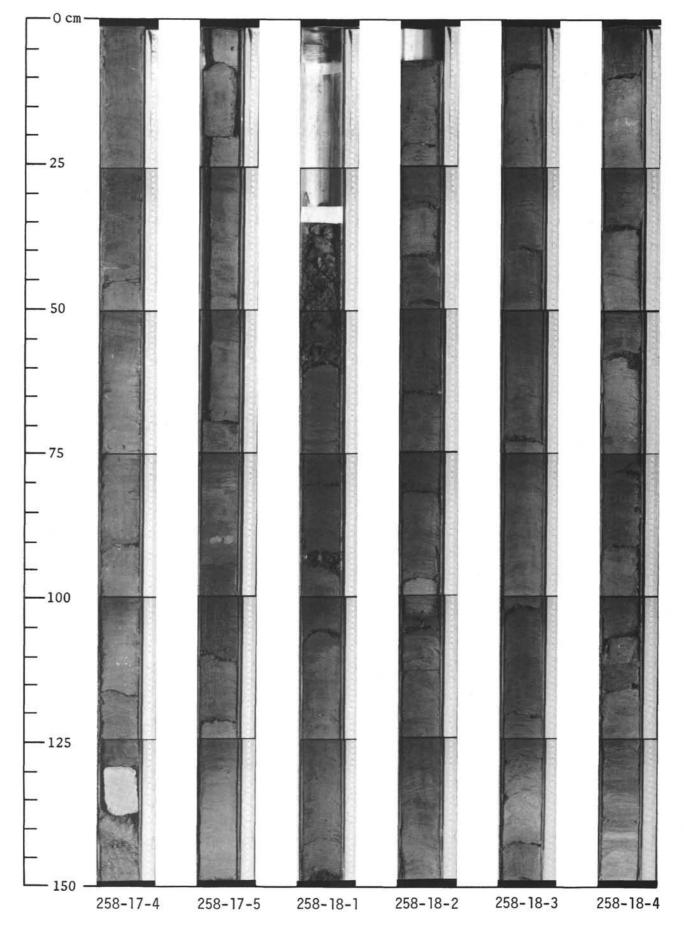


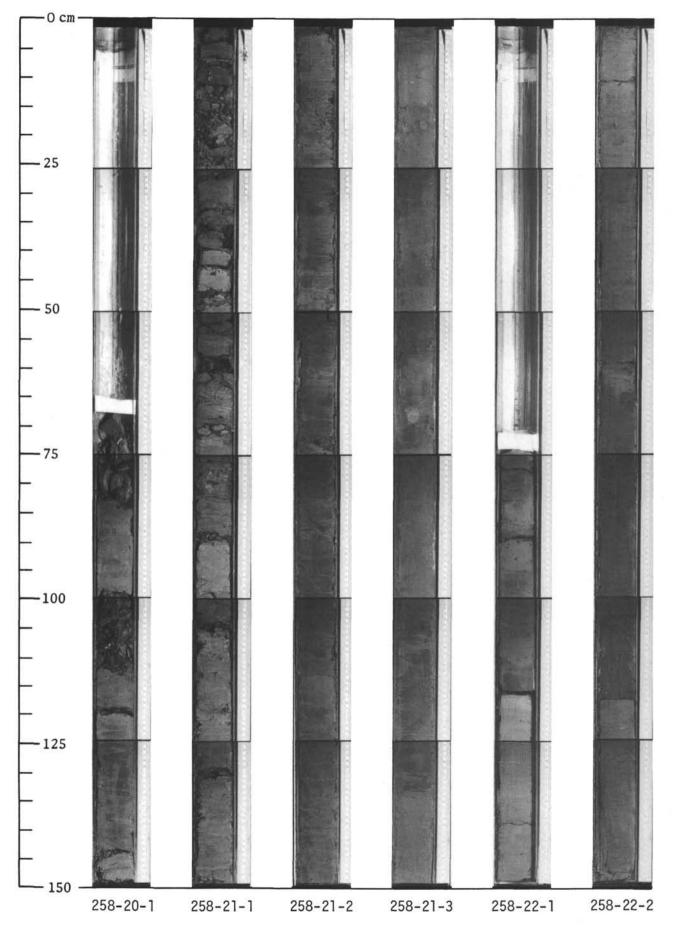


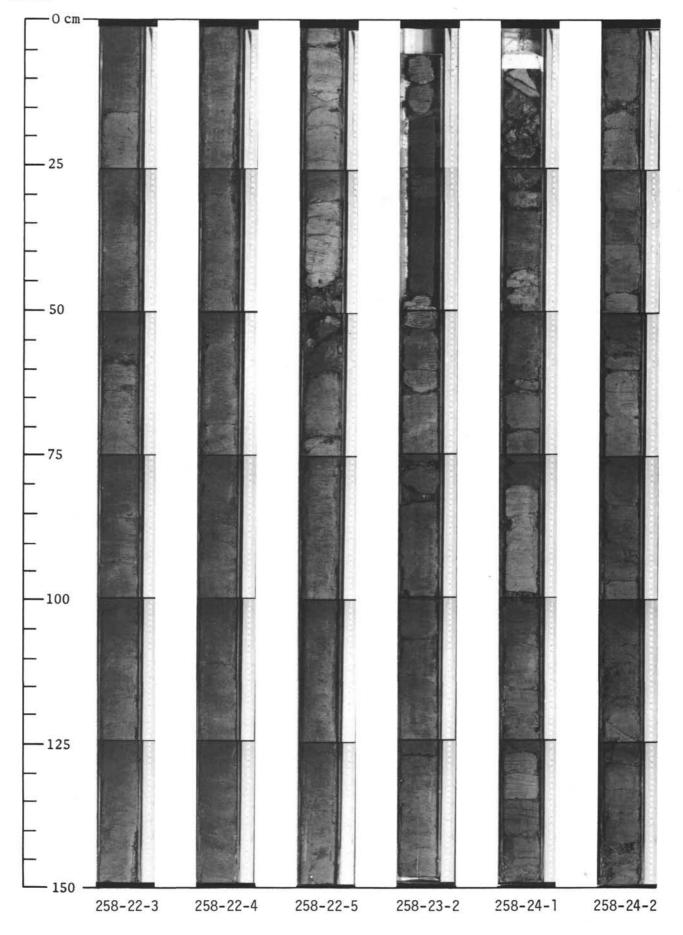


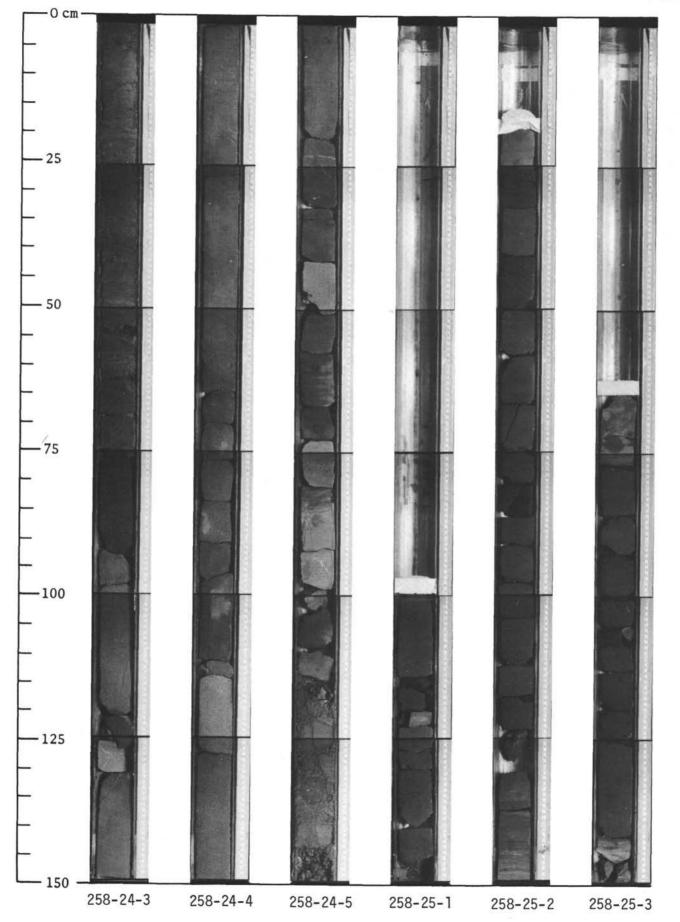


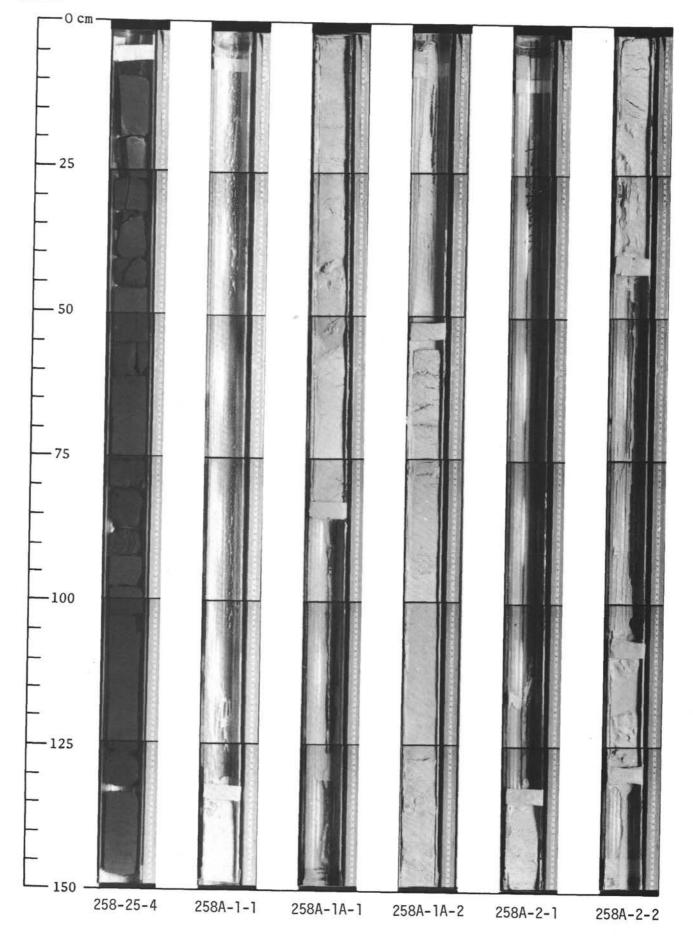


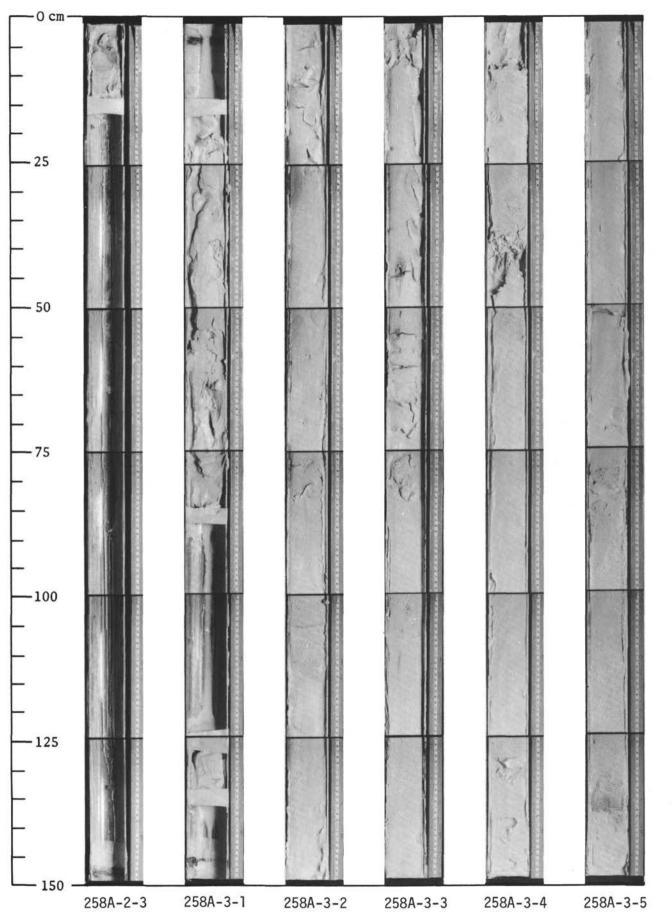


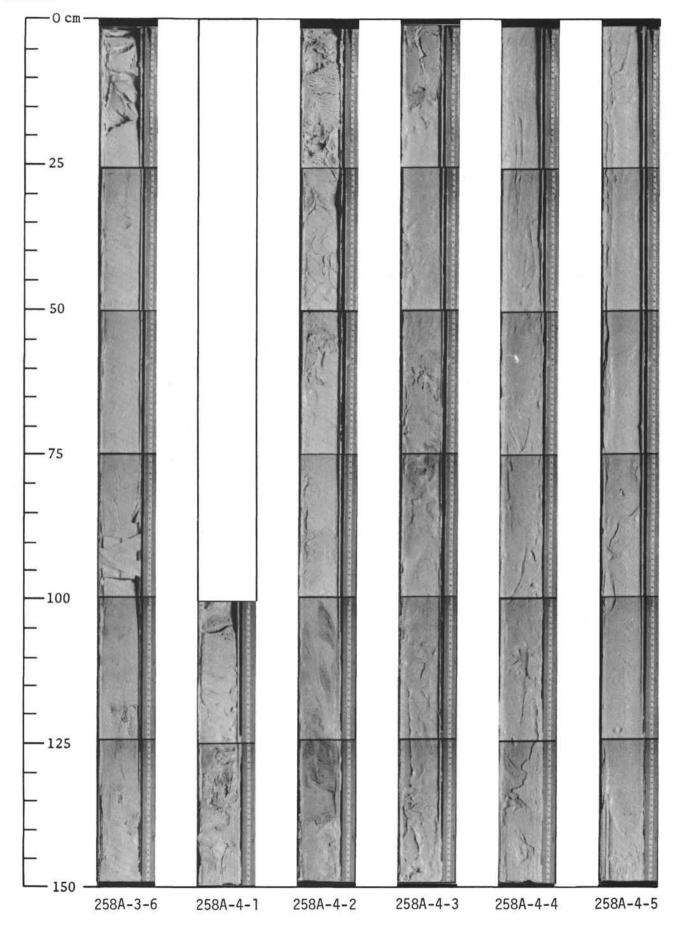


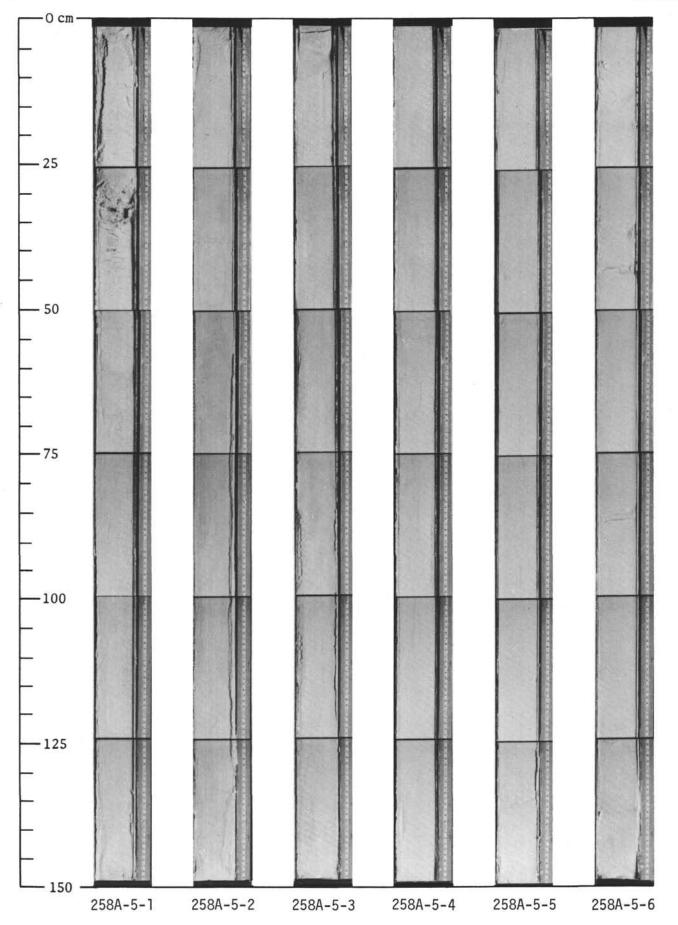


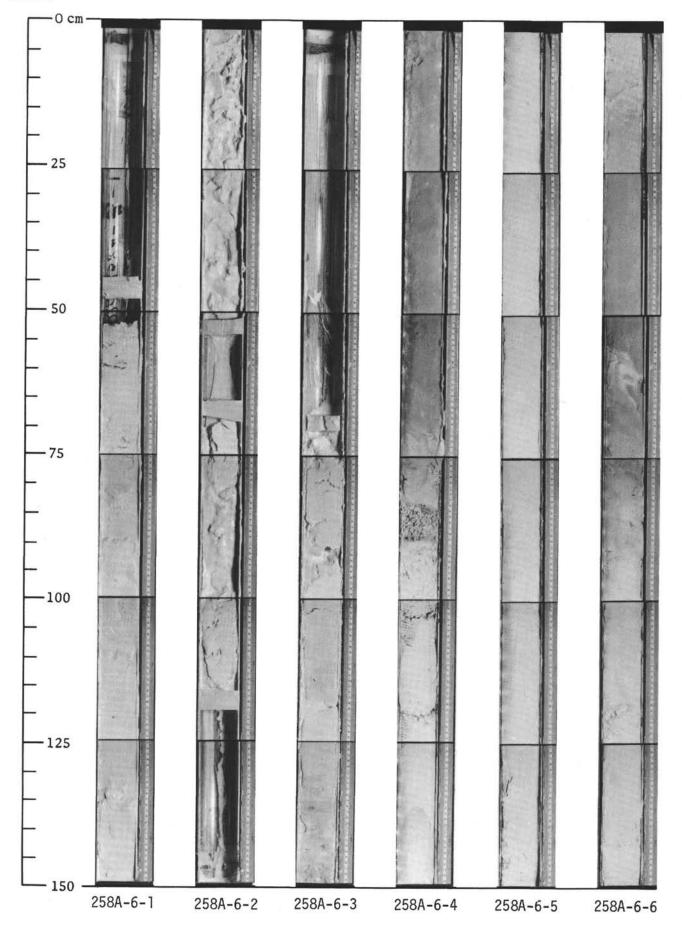


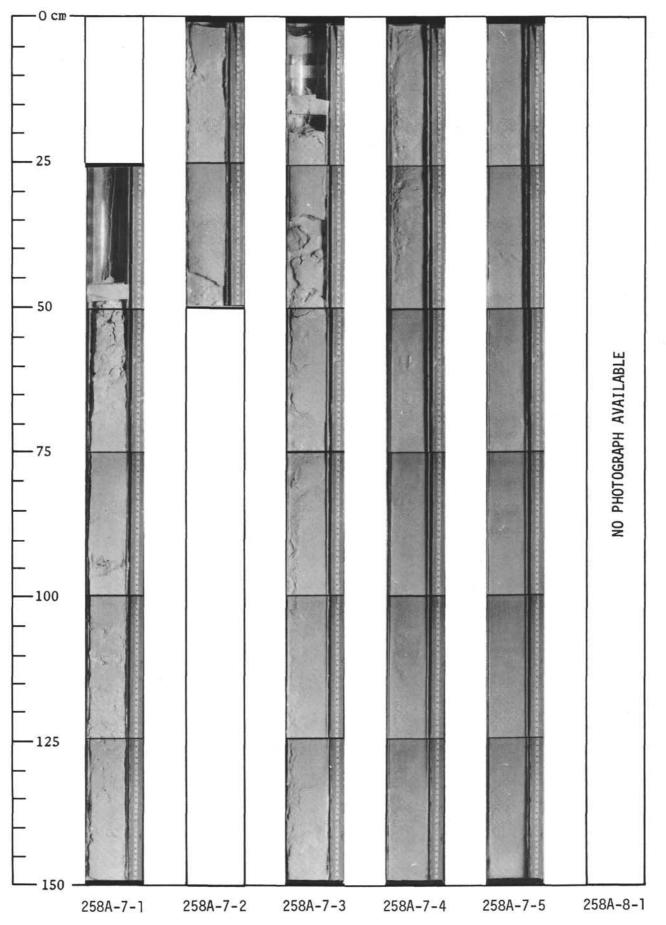




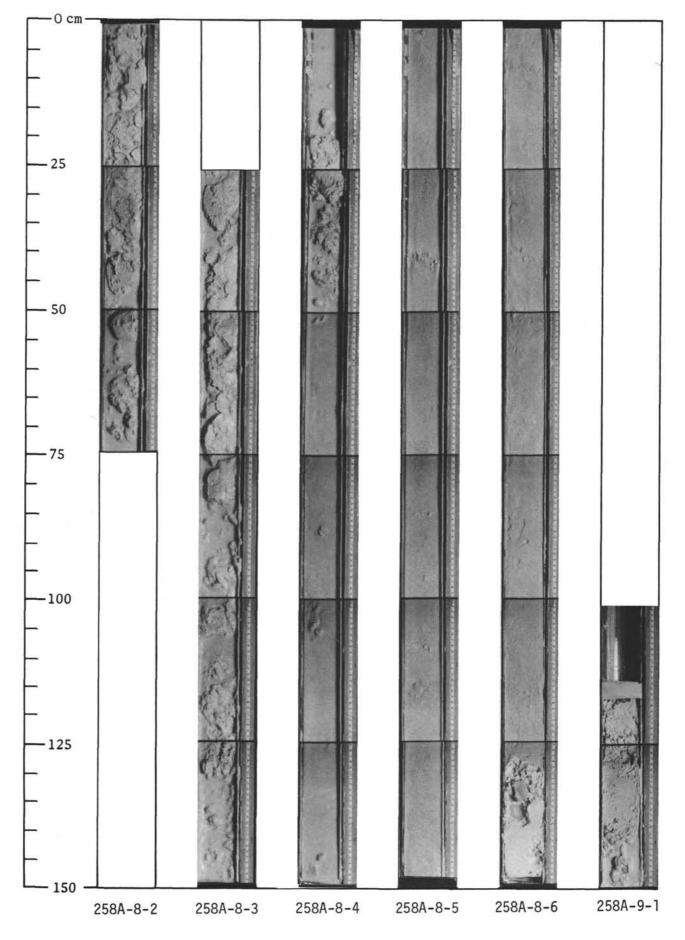




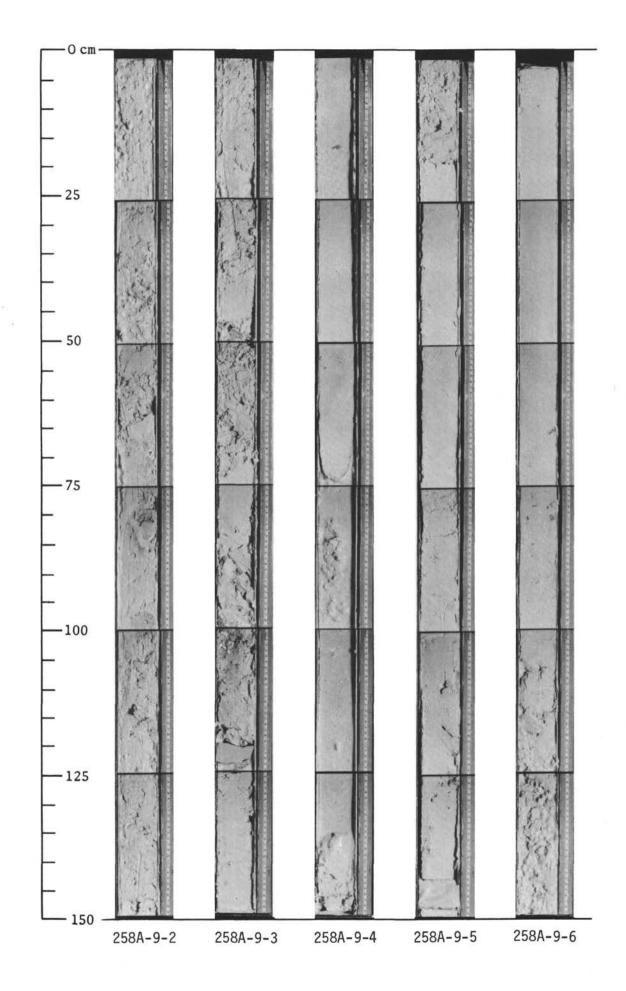




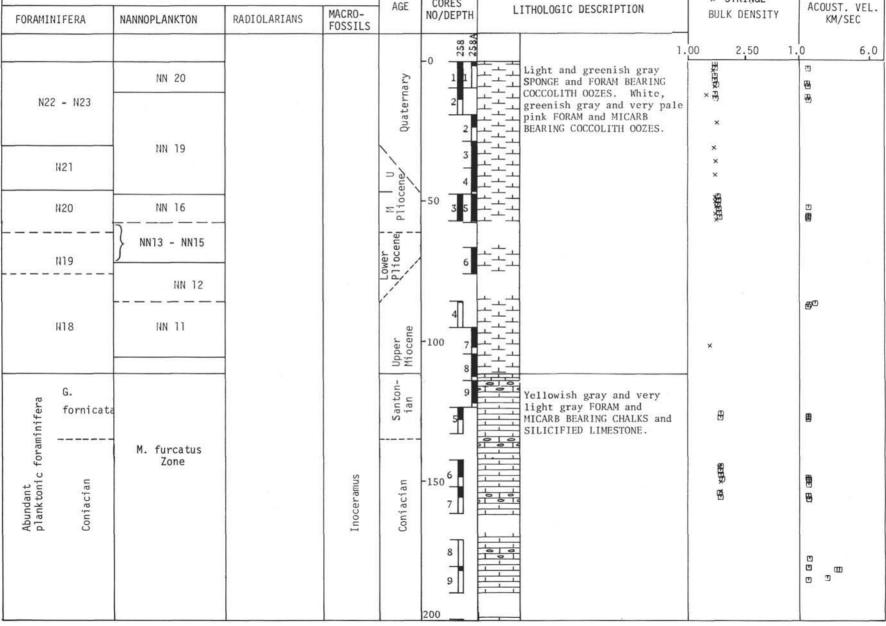
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SUMMARY OF DRILLING RESULTS: SITE 258/0 - 200 m C GRAPE BIOSTRATIGRAPHY × SYRINGE CORES AGE LITHOLOGIC DESCRIPTION MACRO-NO/DEPTH BULK DENSITY RADIOLARIANS NANNOPLANKTON FOSSILS 258/ 1.00 2.50 .0 HATERY FE Light and greenish gray NN 20 Quaternary SPONGE and FORAM BEARING COCCOLITH OOZES. White, greenish gray and very pale pink FORAM and MICARB × BEARING COCCOLITH OOZES. NN 19 × × M L U Pliocene × 50 NN 16 NN13 - NN15



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SITE 258

BIOSTRATIGRAPHY				00055	LITHOLOGIC DESCRIPTION		GRAPE × SYRINGE		-	
FORAMINIFERA	NANNOPLANKTON	RADIOLARIANS	MACRO- FOSSILS	AGE	CORES NO/DEPTH		CRIPTION	BULK DENSITY	ACOUS KM	T. VEL. /SEC
Mostly poor plank- tonic assemblages - d Soniacian Coniacian	M. furcatus Zone			Coniacian	- 200 10	Light greenish	Light greenish gray and yellowish gray COCCOLITH	00 2,50	1.0	6.0
	K. magnificus				11	MICARB CHALK a	nd MICARB	۳.	Ф	۵
	Zone			Turonian	12	+ + + + + + + + +		EEFE	6	
E P. helvetica Cenomanian planktonic Foraminifera	_/M. staurophora'_ G. obliquum Z. _L. alatus Z			Ceno- manian	13	Interstratif greenish gra RICH DETRITA black FERRUG light olive	y ZEOLITE L CLAY, olive INOUS CLAY,	B	Ø	
	Eiffellithus	COCCOLITH DE and light ol COCCOLITH RI COCCOLITH RI COCCOLITH RI COCCOLITH RI	TRITAL CLAY, ive green /	EFEED	•					
Few nonkeeled planktonic foraminifera (Hedbergella, Globigerinello- ides)	turriseiffeli Zone			Late Albian	- 300 16	Electric Brownish black Electric black FERRUGIN CLAY.	and olive OUS DETRITAL	EFEE	₩.	
					17			EEEED	0	
					18 - 350			Ħ	8	
	Prediscosphaera cretacea			Middle Albian	19					
	Zone				20					
				Σ	400					

## SUMMARY OF DRILLING RESULTS: SITE 258/200 - 400 m

SITE 258

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## SUMMARY OF DRILLING RESULTS: SITE 258/400 - 600 m

BIOSTRATIGRAPHY					CORES	1.1	LITHOLOGIC DESCRIPTION		GRAPE × SYRINGE	
FORAMINIFERA	NANNOPLANKTON	RADIOLARIANS	MACRO- FOSSILS	AGE	NO/DEPTH	LITHOLOGIC DESCRIPTION		BULK DE	ENSITY	ACOUST. VEL KM/SEC
Few benthonic foraminifera	Prediscosphaera cretacea Zone			Middle Albian	- 400 🛱 21		[1 to 3 cm diameter pyrite nodules]	1. <u>00 2.</u>	50 1.	0 6.( mo
Rare arenaceous foraminifera			Belemnite		22 - 450	55955 55555	[0.4 to 1.0 cm diameter pyrite nodules]	Đ		C
	W. barnesae C. margereli W. britannica W. communis			Cretaceous	23					23
				Indeterminate	- 500 24 25 - 550		Olive gray and greenish gray fine-grained GLAUCONITE DETRITAL SANDSTONE underlain by dusky brown laminated GLAUCONITIC DETRITAL SILTY CLAY. Several 1 mm thick black laminae in the clay merge into pyrite stringer			00 30 10
	¥.				600					