The Shipboard Scientific Party¹

Position:

146: 15°06.99'N, 69°22.67'W 146A: 15°06.99'N, 69°22.74'W 149: 15°06.25'N, 69°21.85'W.

Water Depth:

146: 3949 meters 146A: 3949 meters 149: 3972 meters.

Penetration:

146: 762 meters 146A: 105 meters 149: 390 meters.

Recovery:

146: 145.6 meters (19%) 146A: 4.6 meters (4.8%) 149: 239.9 meters (57%).

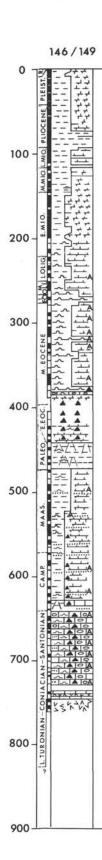
ABSTRACT

Sites 146 and 149 were drilled within 2 km of each other in the central Venezuelan Basin and are discussed as one site.

At Site 149 oozes and chalks of the Cenozoic were cored to 382 meters where hard limestones and cherts of Horizon A'' were encountered (406 m at Site 146). The Early (?) Eocene to early Early Miocene sediments are characterized by abundant Radiolaria.

At Site 146 sediments below Horizon A" were cored to Horizon B", identified as dolerite sills intruded into Coniacian limestones. Cherts and limestones of the Early (?) Eocene overlie noncalcareous Paleocene claystone. The Campanian and Maestrichtian pelagic marl and chalk contrast with the varied lithology of the Santonian which consists of silicified radiolarian limestone, with intercalated basaltic ash, carbonaceous layers, and radiolarian sands.

The upper contact of dolerite with the overlying sediment was not recovered, but the lower contact with an underlying limestone 0.5 meter thick showed a chill zone about 2 cm thick. The upper contact of the underlying sill was also not recovered.



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BACKGROUND

The central Venezuelan Basin was selected for the location of the drill site planned for the recovery of the oldest sediment in the Caribbean Sea. This basin lent itself to this particular objective because, despite the fact that it is surrounded by lands, islands, and ridges, terrigenous sediments are restricted to bathymetric depressions adjacent to sediment sources. As a result, the major part of the basin apart from the small abyssal plains in the deeper regions is blanketed with a uniform cover of acoustically transparent (pelagic) sediments, quite unlike the thick stratified (abyssal plain) sediments of the Colombian Basin west of the Beata Ridge.

The pelagic sediments are about 0.8 to 0.9 sec (reflection time) thick and are characterized by subbottom reflectors, Horizons A" and B" (Ewing et al., 1967, 1968, 1971; Edgar et al., 1971). Horizon A" was identified as an early Eocene chert from examination of piston cores recovered from a fault escarpment on the eastern flank of the Beata Ridge (Talwani et al., 1966; Edgar, 1968). Drilling in the Venezuelan Basin on Leg 4 also recovered chert at the level of Horizon A" and demonstrated the great areal extent of the chert layer.

Horizon B" is a smooth reflector that is similar to sedimentary layers or igneous rock in certain parts of the Pacific (Horizon B') and areas of the Atlantic near inactive continental margins (Horizon B') but quite unlike the typical Atlantic acoustic basement or Layer 2. Seismic refraction studies indicated compressional wave velocities through the layer below B" that range from 3.2 to 5.5 km/sec (Edgar et al., 1971). Rarely did the seismic reflection records show any reflections below B" using standard oceanographic seismic techniques, but processed reflection data collected by Gulf Oil Company's *Gulfrex* (Eaton and Driver, 1969) clearly indicated layered structure below B" in the central Venezuelan Basin.

Site 29 of Leg 4 (Bader, Gerard et al., 1970) was drilled near the central part of the Venezuelan Basin at a point at which the thickness above Horizon A" was less than the maximum. The objectives of this site had been to obtain sedimentological information relevant to the history of the Caribbean and to recover an essentially complete Cretaceous and Tertiary biostratigraphic sequence. The drilling was terminated at Horizon A", which was thought to correspond to unconsolidated radiolarian ooze overlying chert of middle Eocene age, after several unsuccessful attempts had been made to penetrate it. Continuous coring of the site provided a good stratigraphic section for post-Horizon A" time, but a considerable interval (including the entire Oligocene) was missing, and a high degree of solution diminished the biostratigraphic value of the fossil assemblages recovered (Figure 1).

The unrealized objective of Site 29 of recovering Cretaceous or older pelagic sediments in the central Venezuelan Basin remained the prime scientific objective for Site 146. The thickness of sediment between Horizons A'' and B'' implies a middle Mesozoic age for the latter, if accumulation rates have been fairly uniform (Edgar et al., 1971). A further indication that older sediments may be found in the Venezuelan Basin is the widespread occurrence of Early Cretaceous and local Jurassic volcanic rocks found

LEG 4

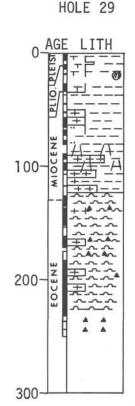


Figure 1. Columnar section of Site 29, Leg 4, showing basic lithology recovered. Drilling was terminated in chert related to Horizon A".

in the island-arc systems of the eastern Greater Antilles and Lesser Antilles, which imply the existence of crust of at least that age on the inside of the arcs.

Site 146/149 was located about 35 km north of Site 29 in shallower water (3949 m versus 4282 m) where the thickness of sediment above Horizon A" was considerably thicker and, presumably, more calcareous and stratigraphically complete (Figures 2 and 3). The site was also in proximity to the Gulfrex seismic line that showed reflectors below Horizon B". Two technological improvements available on Leg 15 which had not been available on Leg 4 were employed for deeper penetration. The first of these was an improved roller bit with tungsten carbide inserts for cutting through the chert and other brittle rocks. The other new capability was that of reentry, which had been tested previously but never used operationally. With this capability, a worn bit could be replaced and the drill string returned to the hole in order to continue drilling at the lowest depth reached by the first bit.

Previously obtained high quality seismic reflection records in the area (*Vema*-26 and *Gulfrex*, which had been the basis for Site 29) showed a uniform area of pelagic sedimentation and precluded the necessity for extensive site surveying by the *Glomar Challenger*. Figure 4 shows the

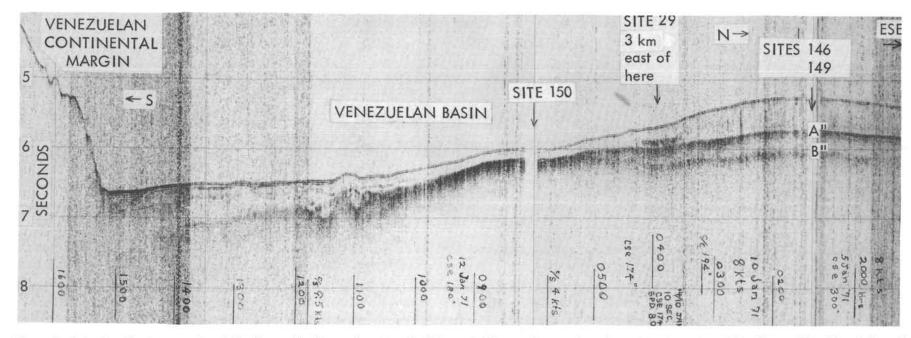
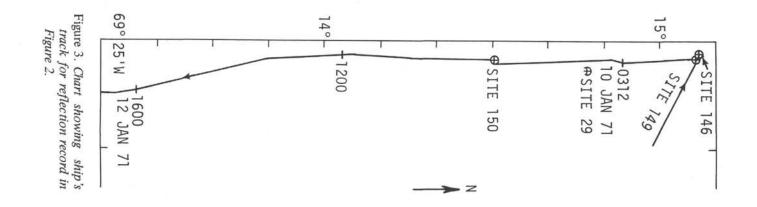


Figure 2. Seismic reflection record made by Glomar Challenger from Site 146/149 to the Venezuelan continental margin and passing within 3 km of Site 29 and through Site 150. Pelagic sediment thickness at Site 146/149, typical for the Venezuelan Basin, then southward to a minimum at Site 150, then becoming thicker toward the abyssal plain at the foot of continental margin.



SITE 146/149

bathmetry in the area of Site 146/149 and Site 29, based on Vema-20 site survey (Bader, Gerard et al., 1970).

OBJECTIVES

Site 146/149 was the primary site on Leg 15, designed to recover a complete sequence of sediments from the Venezuelan Basin for lithologic and biostratigraphic studies and to determine the nature and age of Horizon B". Penetration below Horizon B" would have identified the nature and the age of sub-B" layers that had been noted on the *Gulfrex* records.

OPERATIONS

The ship arrived on Site 146 (Figure 5) at about 0900 hours on 15 December. After dropping two beacons (the first had too weak a signal to be usable), the reentry cone was rigged and lowered over the port side in the manner that had proven successful during the reentry trials six months previously. However, after the cone had been lowered and filled with water (its large opening faced upwards), it was lifted momentarily out of the water by the roll of the ship in the 3 to 4 meter swell, broke its cable, and was lost. A second cone, this time oriented horizontally, was successfully launched at 1530 hours on 16 December and keelhauled into position amidships. Fifty meters of casing were lowered through the cone and secured to it. A 4-cone tungsten carbide bottom roller bit completed the bottom-hole assembly. The assembly was secured to a mechanical release mechanism in the cone (with the bit within the casing), the cone with its casing was freed from the ship, and the remainder of the drill string was made up in the conventional manner. When the casing reached the sea floor, it was washed into the soft upper sediments. With the cone on the sea floor, the release mechanism was actuated with a tool lowered on the sand line, the drill string was released from the cone, and drilling commenced on 17 December. Two spot cores were taken in the upper part of the section, at 96 and 254 meters, as drilling proceeded until the chert was encountered at 406 meters. Continuous coring began at the chert and continued to 701 meters through early Tertiary-late Cretaceous limestones, cherts, siliceous clays, and ash beds. Because the bit had reached 48-hours drilling time (nearly its total life expectancy), the decision was made, at 1000 hours on 21 December, to retrieve the drill string and change the bit rather than risk losing a cone from the bit in the hole. The drill string was reassembled and run down to the ocean floor by the afternoon of 22 December. An Edo rotating sonar transducer, which scans the ocean floor rather like a radar, was lowered into its position immediately beyond the opening of the bit. The ship was maneuvered to within 30 meters of the cone by tracking three sonar reflectors on the cone through the use of a cathode ray display on the bridge. An attempt was then made to position the drill string over the cone by moving it laterally with a jet of water pumped through the drill string and issuing through an opening near the bit. There was no detectable lateral movement and the decision was made to position the drill string by moving the entire ship. At 0658 hours on 22 December the bit was centered and the drill string lowered. Two successive cores (39 to 45 m and 96 to 100 m) were

taken to confirm reentry, but the core barrels returned full of Pleistocene mud, indicating that the bit had missed the core and had started a new hole (146A). This core also showed no trace of the barite mud which had been pumped into the hole before withdrawal. With the realization that the reentry had failed, the drill string was withdrawn to the ship and the jetting device, which was believed to have interfered with water circulation, was removed. The drill string was reassembled and returned to the sea floor by 2100 hours on 24 December. The Edo transducer was lowered, returned for minor repairs, and lowered a second time. The second reentry attempt was successful at 0630 hours on 25 December. Drilling continued and the new bit cut an additional 61 meters of very hard sediments. Drilling terminated 16 meters into a dolerite sill, after penetration rates had slowed to 1 meter per hour. The drill string was withdrawn and the ship departed for Curacao at 1700 hours on 27 December to replace the reentry engineers with geochemists, who would be aboard for Sites 147, 148, and 149. On departure from Site 146 it was agreed that the vessel should return to the area with the geochemists to core the section overlying Horizon A". Sites 147 and 148 were drilled before returning to complete the task (149).

Site 149 was approached in high seas during a four hour interval in which no acceptable satellite fix was obtained. By dead reckoning the site was located and later found to be within 1.5 km of Site 146 (Figure 6). Drilling commenced on 6 January and consisted of continuous coring down to the chert of Horizon A". The uppermost six cores were taken with the extended core barrel and a smaller diameter liner. The remaining cores were cut with conventional core barrel. Drilling disturbance was moderate to severe and many of the cores contained large fractions of displaced, younger sediment. Two cores, 12 and 24, were cut while fresh water containing rhodamine dye was circulated to test the contamination of cores by drilling fluid. Parts of the cores were found to be completely saturated with dye when cut. Because many cores were stored for up to two days before cutting, the entrapped sea water and the ship's vibration turned these into a flocculent soup. Drilling was terminated late on 9 January and the ship was underway for Site 150 early on 10 January.

LITHOLOGY

The columnar section at Site 146/149 can be divided into seven major lithologic units. Two of these units are soft to semiconsolidated oozes that overlie Horizon A"; the remainder, lying below Horizon A", are well consolidated to lithified.

The two ooze units are defined by the presence and absence of Radiolaria. The upper unit (0-198.5 m) is predominantly chalk and marl ooze and is divided into three subunits based on the carbonate content. The uppermost subunit (0-71.85 m) consists of foraminiferal nannoplankton chalk and marl ooze. Average carbonate decreases downward from approximately 50 down to 40 percent. Coarser fossil components are pteropods (upper 9 m only) and planktonic foraminifera, with minor micromollusks, echinoid spines, and benthonic foraminifera. Radiolaria occur very rarely near the top of the unit. Indeterminate black spots (hydrotroilite?) are conspicuous

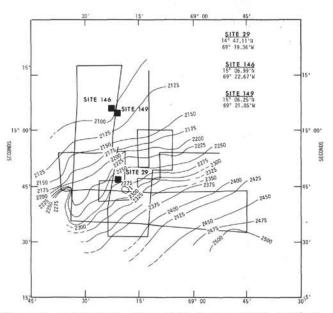
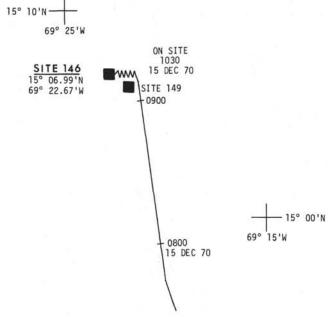
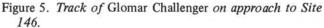


Figure 4. Bathymetric chart of the region of Site 146/149 and Site 29 of Leg 4. Chart was prepared from a survey by the Vema-26 cruise. (Bader, Gerard et al., 1970).





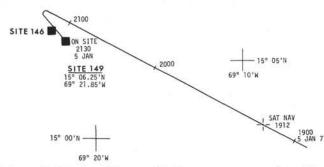


Figure 6. Track of Glomar Challenger on approach to Site 149.

in the lower part of the unit. Volcanic constituents, especially plagioclase and hornblende, are scarce but are found throughout.

A pelagic brown clay (71.85 to 96.45 m) underlies the uppermost subunit with a relatively abrupt upper contact. The clay has a negligible carbonate content, is brownish down to about 87 meters, and varicolored (gray, dark yellowish orange, and light brown) below this. Coarse constituents are limited to rare benthonic foraminifera, very rare and partly dissolved planktonic foraminifera, rare fish debris, and sparse volcanic minerals, especially plagioclase. Native copper was found in two places at about 75 and 85 meters; aggregation in the form of pellet-like shells and cylindrical tubes strongly suggests biological concentration of this element. Dolomite rhombs and anatase dipyramids (both a few microns in length) are persistent throughout this unit. Palygorskite and barite were found.

Below the clay subunit are interbedded clays, and calcareous clays and marl oozes (96.45-135.45 m). The top of this subunit is defined by the first occurrence (drilling) of carbonate below the clay subunit. The lower boundary is established at the first occurrence of chalk which coincides with the last occurrence of clay. The clay is mostly yellowish brown, firm, and contains K-feldspar and carbonate rhombs. The calcareous clays and marls are extensively mottled and contain plagioclase, apatite, and glauconite.

The lowermost subunit (135.45 to 198.50 m) of foraminiferal nannoplankton marl oozes and chalk oozes have been disturbed extensively by drilling which may, in part, be attributed to fresh water having been used for the drilling fluid during coring for the geochemical program. The carbonate content increases from the base of the overlying subunit to about 70 percent at 150 meters, then declines to about 50 percent at the base of the unit. The amount of carbonate fluctuates noticeably in units a fraction of a meter thick. Burrow mottling and black (iron-manganese?) spots are conspicuous in the marl layers. Planktonic foraminifera vary from strongly dissolved to well preserved, but in most instances show noticeable solution effects. Volcanic debris occurs throughout but increases downward, with plagioclase the most persistent element and glass and pumice appearing below 170 meters. The sediments of this unit are compacted to varying degrees but are nowhere lithified. Minor clinoptilolite was found by X-ray diffraction.

The first appearance of a siliceous fauna marks the top of Unit II, a semi-indurated calcareous radiolarian ooze and radiolarian nannoplankton chalk (198.50 to 382.50 m). The carbonate content of this unit fluctuates widely but, on the average, declines from about 70 percent at the top to about 40 percent near the base. Planktonic foraminifera are almost completely dissolved, but nannoplankton are common and generally well preserved. Volcanic constituents are abundant in this interval and ash beds (Figure 7) are conspicuous below 240 meters. Plagioclase is persistent and clino- and orthopyroxene, apatite, and red hornblende are found. Microtektites (this volume) were found in the core catcher of Core 31 (279 m). The unit can be subdivided into two subunits on the basis of sediment



Figure 7. Radiolarian nannoplankton marl with distinct volcanic ash bed. Oligocene. (149-29-2, 113-125).

components. The upper subunit (198.50 to 270.00 m) is a radiolarian-rich nannoplankton marl. The clay component decreases rapidly at the lower boundary and is essentially absent in the underlying subunit of radiolarian nannoplankton chalk. The only ash bed recovered above Horizon A'' at this site was found at about 250 meters (Figure 7). The first occurrence of chert is at 289 meters, which is 117 meters above the main chert layer comprising Horizon A''. Palygorskite was detected by x-ray diffraction in the three lowest cores.

The first occurrence of chert and limestone of Unit III was found at 382.50 meters at Site 149 but at 406 meters at Site 146. Consequently the exact depth of the contact between the second and third unit, which here identified with Horizon A", is not clearly established. Volcanic clay is a minor component of this unit. Fresh glass and radiolaria are not found; cristobalite (replacements of radiolaria?), clinoptilotite, and montmorillonite comprise the clay units. Planktonic foraminifera are common in the chalk. Palygorskite was detected by x-ray diffraction at 422 meters. The lower boundary of the unit is defined by the presence of green siliceous claystone at 440.65 meters.

Beneath the chalk-chert sequence is a uniform, green siliceous claystone (449 to 477.25 m) which is markedly compacted, showing a conchoidal fracture and a tendency toward shaley fissility. This Unit (III) consists predominantly of cristobalite. Fossils are absent, except for cristobalite pseudomorphs after Radiolaria. Mottling is inconspicuous and flattened. X-ray diffraction showed rhodochrosite at 451 meters, palygorskite at 478 meters, and barite and clinoptilolite throughout.

The next lower major unit (V) is a varicolored nannoplankton chalk and marl sequence (477.25 to 630.30 m) that consists of interbedded neutral-colored chalks and varicolored (pinkish brown, brownish gray, etc.) marls. The average carbonate content increases downward from about 40 to about 80 percent. This unit is characterized by color contrasts and by the ubiquity of mottling and burrowing (Figure 8). Most of the burrows are subhorizontal, including horizontal chevron (Zoophycos) burrows, but many are vertical (Teichichnus), helicoid, or dendritic (Chondrites) in three dimensions. Interbedded sandy radiolarian layers, of possible local turbiditic origin, containing subordinate planktonic foraminifera, with sharp lower contacts and with or without grading, occur below 534 meters. In the marl layers radiolarians are absent and planktonic foraminifera present but badly preserved. Chert (Figure 9; Frontispiece) occurs at 557 meters (151 m below the main chert horizon). An occurrence of garnet and Melson, this volume), which may be (Donnelly authigenic, was noted at 517 meters. Lithification is moderate throughout this unit, increasing noticeably downward. The downward increase in carbonate distinguishes two subunits: brownish gray and olive gray marl (477.25 to 521.35 m), and a subunit of varicolored chalks (521.35 to 630.30 m) which is dominantly pinkish brown above, light gray in the middle, and pinkish gray below. Volcanic ash beds occur throughout this unit, but are more abundant downward. They contain a few volcanic crystals with exceptional beds that are rich in biotite, quartz, alkali feldspar, apatite, and zircon but are commonly almost pure montmorillonite. Barite is conspicuous throughout. Planktonic foraminifera are scarce and poorly preserved.

The next lower unit (VI) is radiolarian limestone (630.30 m to 721.00 m) which is characterized by a higher degree of lithification, more abundant cherts and ash layers, microfractures, and occurrences of pyrite. The carbonate content of this unit fluctuates widely between about 50 and 80 percent. The upper part is a fairly homogeneous, light gray limestone (630.30 to 692 m) and the lower part is a varicolored limestone with contrasting lithologies, including white radiolarian sands, very dark carbonaceous layers (Figure 10), and conspicuous ash layers. The organic-rich lavers, restricted to the Turonian-Coniacian-Santonian, were analyzed for their organic carbon; the results are shown in Table 1. Sapropelite was defined by Olausson (1960) as a sediment containing more than 2.5 percent organic carbon; three of the samples analyzed classify as sapropelitic. Burrowing is conspicuous throughout (Figure 11), especially in the less carbonate-rich beds. X-ray diffraction shows minor barite, clinoptilolite, and palygorskite.

Immediately above the upper of two dolerites is a varied lithic subunit (721,00 to 738,30 m) differing from the overlying subunit in the presence of conspicuous graded beds of altered, greenish basaltic ash (Figure 12). The ash is interbedded with limestones and radiolarian sands (Figure 13), some graded like those in the above unit. Below this unit is an upper dolerite sill (738.30 to 739.63 m) whose upper contact was not recovered and whose lower contact is a chilled border about 2 cm thick. Below this dolerite is 0.5 meters of highly recrystallized limestone crowded with



Figure 9. Radiolarian limestone showing sharp contact of replacement chert and limestone. Campanian. (146-28-3, 32-45).



Figure 8. Nannoplankton chalk and marl showing a dark volcanic clay bed and pale chalk whose contrasting colors emphasize the extent of burrowing. Maestrichtian. (146-16-6, 9-35).

Figure 10. Radiolarian limestone showing layer rich in organic carbon. Note flattened burrows. Turonian to Santonian. (146-38-1, 88-101).

TABLE 1 Total Carbon, Organic Carbon, and CaCO₃ Per Cent for Samples Taken from Dark Carbonaceous Layer (Turonian to Santonian), Site 146

| Sample | Total C (%) | Organic C (%) | CaCO3 (%) | | |
|-----------|----------------|------------------|--------------|--|--|
| 36R-2-87 | 8.1 | 2.5 | 48 | | |
| 36R-2-138 | 7.7 | 6.3 | 12 | | |
| 38R-1-106 | 14.5 | 11.2 | 28 | | |
| 39R-1-92 | 0.2 | 0.1 | 1 | | |

planktonic foraminifera. The limestone consists of calcite, a very well-crystallized montmorillonite, and some pyroxene (evidently metamorphic diopside).

The lower dolerite sill was penetrated from 740 to 762 meters; more than half of the drilled interval was recovered, but the upper contact was not. Grain size variations within the recovered material were minimal. The dolerite appears to be quite fresh and moderately fractured and only slightly veined.

PHYSICAL PROPERTIES

Wet-bulk Density, Water Content, and Porosity

Discussion

Wet-bulk density and porosity were measured by two methods aboard the Glomar Challenger: Gamma Ray Attenuation Porosity Evaluator (GRAPE) and individual sample volume-weight measurements (sample data are the enclosed dots on core and hole plots). Water content was determined by weight-weight relationships. These methods are described and discussed in Appendix I. In general, the data have a precision of ±5 per cent. The GRAPE data should be viewed with caution as it results from continuous diameter-scanning along the entire length of an unopened core which includes undisturbed sediment, disturbed sediment, and drilling slurries. The Appendix contains precautionary discussions concerning interpretation of these data and equations for diameter corrections for punch cores in the upper 45 meters of this site and for hard core samples which have a smaller diameter than the core liner.

Some stiff sediments or rocks are cored without plastic flowage of the sample and, since the drill bit has a smaller diameter than the core liner (2.60-in internal diameter), a space between the liner and the core results. The remaining space may be filled with a drilling slurry, highly disturbed sediment, or, in some cases, air. A problem arises here because a 2.60-inch diameter is assumed in the density calculation.

Where necessary the GRAPE data of Site 146 were adjusted (G.C.D. [GRAPE corrected diameter] data are the dotted density and porosity lines on the hole and core plots) for varying diameters of the core. Diameters are generally smaller than 2.60 inches by about 12% in cores recovered from 494 to 680 meters, 13 to 14% from 680 to 705 meters, respectively, 20% from 705 to 740 meters, and 14% from 740 to 760 meters.

Density adjustments for varying diameters in the data below 494 meters at Site 146 were made by observing the



Figure 11. Burrowed limestone, with fish bone (lower center), recrystallized fossils, and flattened, highly organic burrows. Bar is 1 mm long. Turonian to Santonian. (146-36, CC).

GRAPE data and determining the typical densities of the disturbed sediments or drilling slurries between the hard rock sample. The typical density was about 1.2 g/cc, but values ranged from 1.0 to 1.35 g/cc, except for air around the limestone and basalt cores at the bottom of Hole 146. However, this "slurry" density is for the most part an assumption, and the reader may wish to recalculate the data using other densities (the formula is in Chapter 13). Only maximum density trends and minimum porosity trends were recalculated (dotted lines in core and hole plots) as lesser densities and greater porosities are suspected of being drill-disturbed sediments, smaller diameters, or fractured rocks. These adjusted data are only approximations.

Results

In general, wet bulk densities irregularly increase (1.4 to 1.8 g/cc) then decrease (1.8 to 1.5 g/cc) between the surface and Horizon A", while between Horizons A" and B", densities were about 2 g/cc (1.9 to 2.2 g/cc) with a slight decrease to 1.8 g/cc (1.8 to 2.3 g/cc) near Horizon B". Horizons A" and B" were cored at 406 and 739 meters, respectively, at Site 146.

Wet bulk densities of the Tertiary calcareous ooze, marl, and clay above Horizon A" irregularly increase from the surface value of 1.4 g/cc (76% porosity) to 1.6 g/cc (64%) at 10 meters, and then irregularly increase to approximately 1.8 g/cc (50%) at 200 meters. Below 200 meters, densities irregularly decrease to about 1.4 g/cc (68%) at 375 meters,



Figure 12. Interbedded limestone, basaltic ash, and clay showing graded bedding, current scour, and flattened burrows. Turonian to Santonian. (146-39R-2, 65-84).

which corresponds with the occurrence of radiolarian ooze and radiolarian chalk. Also, below 362 meters, silicification and occurrences of chert become more common resulting in densities of about 2.1 g/cc.

Between Horizons A" and B" (406 to 739 m at Site 146), Cretaceous chalk, marl, and radiolarian limestone (and chert) have typical maximum densities of about 2 g/cc (40 to 20%) with a range from 1.9 to 2.3 g/cc (16%). Some chert nodules and layers have densities near 2.2 to 2.3 g/cc. The radiolarian limestone near Horizon B" has a slightly lower density of 1.8 to 1.9 g/cc (range of 1.8 to 2.3 g/cc). The diabase which formed Horizon B" has densities of 2.68 to 2.88 g/cc.

Porosity plots are slightly different from the density plots as the composition of the sediment or rock varies greatly. An example is opaline radiolarian tests, which are

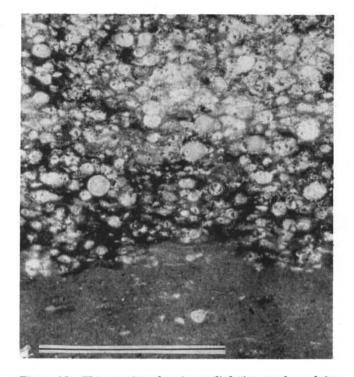


Figure 13. Thin section showing radiolarian sand overlying clay with a sharply defined contact. No clear size grading is apparent in this section, but other radiolarian sands are graded. Bar is 1 mm long. Turonian to Santonian. (146-41R-1, 104-106).

very light weight (2.2 g/cc) and have a very delicate, fine structure, versus the heavier calcite (2.7 g/cc) foraminifer tests or nannofossil plates. Therefore siliceous and calcareous sediments may weigh the same but have different porosities.

As a generality, porosity between the surface and Horizon A" varies irregularly between 44 and 70%, decreasing then increasing with depth. Between Horizons A" and B" porosity varies irregularly (8 to 42%) but appears to decrease from about 36 to 22%. Some chert nodules and thin layers are hard and dense with essentially zero porosity.

In more detail, porosities in the Tertiary sediments above Horizon A" decreased from 76 to 65% from zero to 10 meters depth and from 65% to 45 to 50% within 10 to 200 meters depth. These sediments are chalks and chalk ooze, marl and marl ooze, and clays. However, porosities in radiolarian-rich chalks and oozes from 200 to 375 meters increased from 50% to 65% respectively.

The Cretaceous sediments and rocks between Horizons A'' and B'' typically have a porosity of 36% between 495 and 577 meters. Below 577 meters, porosity irregularly varies (9 to 42%; 0% for chert nodules and layers), but generally decreases to about 22 to 32% at 660 meters depth, and then typically increases to 36% (with a low of 12%) at 720 meters. The porosity between Horizons A'' and B'' inversely correlates with the sound velocity (see section on Sound Velocity for Plot of Porosity).

Water contents varied irregularly between the surface and Horizon A". They ranged from 53 to 28% at the surface. From 10 to 200 meters below the sea floor, water contents irregularly decrease from 42 to 30%, while from 200 to 300 meters they increase irregularly to about 50%. This increase appears to be related in part to the radiolarian content of the sediment. In general, sediments containing radiolarians have lower densities and higher porosities and water contents than comparable calcareous sediments.

Sound Velocity

Introduction

Velocities at which sound travels through sediment and rock core samples were measured by the Hamilton Frame technique, which is discussed in Appendix I. This method has a precision of ± 1.1 percent.

These measurements were taken several days after the cores were split; however, the sediments were stored in sealed D-tubes containing wet towels and refrigerated $(2-5^{\circ}C)$ in a humid atmosphere to minimize moisture loss. If these sediments or rocks lost pore water, then their measured velocities would be significantly lower (10-20%) than velocities obtained from other seismic data; however, the data are in good agreement.

Only sediments and rocks which appeared to be physically undisturbed had velocity measurements made. These velocities were measured parallel to the bedding planes, unless otherwise noted in the tables, and are reported at laboratory pressures and temperatures (21.6 to 25.8°C; Table 2).

Results

Sound velocities measured at Site 146 range from 1.51 to 5.34 km/sec. Clay and chalk ooze in the upper part of the hole (98-103 m and 254-263 m) have typical velocities of 1.51 and 1.63 km/sec, respectively, while the lithified clays, chalks, marls, and radiolarian limestone in the lower part of the hole have typical velocities ranging from 1.95 to 2.95 km/sec. Velocities in reddish brown chert fragments range from 4.70 to 5.34 km/sec with an average of 5.09 km/sec. A diabase sample from a sill at 739 meters below the sea floor propagated sound at 4.86 km/sec.

Velocities were measured parallel and perpendicular to the bedding planes of Oligocene and Cretaceous samples. Oligocene, light gray radiolarian nannofossil chalk ooze recovered from 262 meters has velocities of 1.57 and 1.58 km/sec perpendicular and parallel to the bedding, respectively, while Maestrichtian light gray nannofossil chalk from 522 meters has velocities of 1.90 km/sec parallel to the bedding and 1.81 km/sec perpendicular to the bedding. These two spot checks are not sufficient data to draw any definite conclusions.

Discussion

Major lithologic sections were stratigraphically distinguished by their relative and uniform velocity differences. Average velocities were calculated for these sections, but velocities for the chert nodules (or layered fragments) were not averaged in these groupings as it is not known what thicknesses of the chert are in situ, and they might thus skew the average from the modal velocities (Figure 14). The sound velocity-stratigraphy is displayed in Table 3. In general, velocity variations correlated with different lithologic sections as defined by the geologists, except in the bottom of the hole where velocities were controlled less by sediment type and more by cementation.

According to averaged core velocities, air-gun records, and drilling rates, the reflector Horizon A" appears to be the hard rock (foraminiferal nannofossil chalk and chert) at 406 meters (at Site 146) below the sea floor and the reflector Horizon B" is probably the diabase at 739 meters depth. It seems unusual that there were not other reflectors between Horizon A" and Horizon B", where varying velocities were recorded. It is possible that such zones would be reflectors with differing reflection frequencies, and it is also possible that these velocity changes in situ are in part gradational.

The sedimentary section above Horizon A" (0-406 m at Site 146), calculated from the seismic reflection travel time and drilled depth, has an average velocity of 1.60 to 1.65 km/sec, which is in agreement with the meager laboratory sound velocity data in this section. From the laboratory sound velocity data the section between Horizons A" and B" has an average velocity of 2.44 km/sec (not including chert). This average velocity agrees with the average velocity of 2.47 km/sec (including chert) obtained from the seismic air-gun reflection times from Horizons A" and B" and their drilling depths.

The dolerite velocity of 4.86 km/sec is in good agreement with J. Ewing's (personal communication) sonobuoy data (Figure 14) and with previous refraction data. A summary of previous refraction and wide-angle reflection data is presented in Edgar et al (1971).

Considering the precision of all the data involved, it seems unusual that the data agree so well, and, therefore, the following discussion of their agreement and the following manipulation of data may not be warranted but is of interest.

If it is assumed that there is a 3- to 4-meter cumulative thickness of chert between Horizons A" and B", then the sample velocity data would theoretically match Edgar's air-gun-derived velocity for this section of 2.47 km/sec. In order to match Ewing's sonobuoy sound velocity average between A" and B" of 2.59 km/sec, which was taken 6 miles to the north $(15^{\circ}20'N, 69^{\circ}50'W)$, the shipboard laboratory velocity data would have to include about a 15-meter section of chert, which does not seem probable at Site 146. Of course Ewing's geologic section may not be identical to the one at Site 146 (facies may change rapidly as suggested by Site 150), and factors such as temperature and pressure must be taken into consideration.

Negative temperature corrections would be significant in the high-porosity surficial sediments where the bottom water is about 4°C, but below here, the temperature would increase with depth. However, these negative corrections would be counterbalanced by positive corrections resulting from increasing pressure (and decrease in porosity), which is about 395 bars at the sea floor at Site 146 and 455 bars at Horizon A".

In general one would expect in situ sound velocities to be higher because of pressure corrections. At Site 146, pressures at Horizons A" and B" would be about 455 and 530 bars, respectively (calculated from sample densities).

SITE 146/149

TABLE 2 Hamilton Frame Sonic Velocities, Site 146

| Core | Section | Upper Interval ^a (cm) | Depth in Hole (m) | Velocity ^b (m/sec) | Temperature (°C) | Remarks |
|------|---------|--|-------------------------|----------------------------------|---------------------|---|
| 1 | 1 | 125.0 | 97.25 | 1511 | 24.8 | Clay. |
| 1 | 2 | 95.0 | 98.45 | 1527 | 24.8 | Clay. |
| 1 | 3 3 | 25.0 | 99.25 | 1482 | 25.4 | Clay. |
| 1 | 3 | 25.0 | 99.25 | 1467 | 25.8 | Clay. |
| 1 | 4 | 95.0 | 101.45 | 1501 | 24.8 | Clay. |
| 1 | 6 | 101.0 | 104.51 | 1539 | 24.9 | Clay. |
| 2 | 4 | 54.0 | 259.04 | 1623 | 23.8 | Foram nanno chalk ooze, clay and radrich. |
| 2 | 4 | 85.0 | 259.35 | 1678 | 23.5 | Foram nanno chalk ooze, clay and radrich. |
| 2 | 5 | 90.0 | 260.90 | 1574 | 23.5 | Nanno chalk ooze, with radiolarians. |
| 2 | 6 | 54.0 | 262.04 | 1577 | 23.5 | Nanno chalk ooze, with radiolarians. |
| 2 | 6 | 75.0 | 262.25 | 1563 | 24.7 | Nanno chalk ooze, with radiolarians. |
| 5 | 1 | 25.0 | 422.25 | 2140 | 24.5 | Foram nanno chalk, clay-rich. |
| 5 | 1 | 60.0 | 422.60 | 1702 | 24.5 | Foram nanno chalk, clay-rich. |
| 5 | 2 | 96.0 | 424.46 | 5216 | 25.5 | Chert, brown. |
| 5 | 1 | 96.0 | 424.46 | 5345 | 25.5 | Chert, brown. |
| 5 | 1 | 100.0 | 414.50 | 1928 | 25.5 | Foram nanno chalk, clay-rich. |
| | | | | | | and a war - H |
| 6 | 1 | 63.0 | 431.63 | 1928 | 25.1 | Foram nanno chalk. |
| 6 | 7 | 0.0 | 440.00 | 2025 | 24.1 | Foram nanno chalk. |
| 7 | 1 | 25.0 | 440.25 | 3198 | 24.3 | Foram nanno chalk (light). |
| 7 | 1 | 60.0 | 440.60 | 2477 | 25.3 | Siliceous claystone (dark). |
| 7 | 1 | 131.0 | 441.31 | 2697 | 24.3 | Siliceous claystone. |
| 8 | 1 | 96.0 | 449.96 | 2251 | 24.2 | Silicaous claustone |
| 3 | 1 | | 449.96 | 1927 | 24.2 | Siliceous claystone. |
| | | 78.0 | | | | Siliceous claystone. |
| 9 | 1 | 64.0 | 458.64 | 2033 | 24.2 | Siliceous claystone. |
| 9 | 2 | 55.0 | 460.05 | 2225 | 24.2 | Siliceous claystone. |
| 10 | 2 | 76.0 | 469.26 | 1977 | 24.2 | Siliceous claystone |
| 11 | 1 | 107.0 | 477.07 | 2158 | 24.2 | Siliceous claystone. |
| 11 | 2 | 22.0 | 477.72 | 1768 | 23.5 | Nanno marl. |
| 11 | 2 | 75.0 | 478.25 | 1693 | 23.5 | Nanno marl. |
| 13 | 1 | 100.0 | 495.00 | 1869 | 23.5 | Foram nanno marl, reddish gray. |
| 13 | 2 | 18.0 | 495.68 | 1805 | 23.5 | Foram nanno mari, gray, burrowed. |
| 13 | 3 | 125.0 | 498.25 | 1808 | 22.8 | Foram nanno mari, gray, burlowed. |
| 13 | 4 | 20.0 | 498.20 | 1815 | 22.0 | Foram nanno marl, gray. |
| 13 | 4 | 90.0 | 499.40 | 1813 | 22.8 | Foram nanno marl, gray. |
| 13 | 5 | 35.0 | 500.35 | 1832 | 22.2 | Foram nanno marl, reddish gray. |
| | | | | State and a second | | |
| 14 | 1 | 110.0 | 504.10 | 2502 | 22.2 | Foram nanno marl, gray. |
| 14 | 2 | 38.0 | 504.88 | 1921 | 22.2 | Foram nanno marl, light gray. |
| 14 | 3 | 78.0 | 506.78 | 1961 | 22.2 | Foram nanno marl, gray. |
| 14 | 4 | 78.0 | 508.28 | 2045 | 22.2 | Foram nanno marl, gray. |
| 15 | 1 | 132.0 | 513.32 | 1998 | 22.2 | Nanno marl, gray with light gray band. |
| 15 | 2 | 9.0 | 513.59 | 2064 | 21.8 | Nanno marl, light gray. |
| 15 | 3 | 0.0 | 515.00 | 1957 | 21.8 | Nanno marl, light gray. |
| 15 | 4 | 8.0 | 516.58 | 2015 | 21.6 | Nanno marl, light gray. |
| 15 | 5 | 15.0 | 518.15 | 2027 | 21.6 | Nanno marl, gray. |
| 15 | 6 | 60.0 | 520.10 | 1831 | 21.8 | Nanno marl, light gray. |
| 16 | 1 | 119.0 | 522.19 | 1903 | 21.8 | Foram nanno chalk, light gray. |
| 16 | 1 | 119.0 | 522.19 | 1809 | 21.8 | Foram nanno chalk, perpendicular to bedding |
| 16 | 2 | 120.0 | 523.70 | 1916 | 21.8 | Foram nanno chalk, gray. |
| 16 | 3 | 144.0 | 525.44 | 2116 | 21.8 | Foram nanno chalk, light gray. |
| 16 | 4 | 76.0 | 526.26 | 1515 | 22.1 | Waxy volcanic clay, low density. |
| 16 | 5 | 65.0 | 527.65 | 1969 | 22.1 | Foram nanno marl, gray. |
| 16 | 6 | 101.0 | 529.51 | 1931 | 22.1 | Foram nanno marl, gray. |
| | | | | | | |
| 17 | 1 | 84.0 | 530.84 | 1935 | 22.1 | Foram nanno marl, gray. |
| 17 | 2 | 52.0 | 532.02 | 1967 | 22.1 | Foram nanno marl, gray. |
| 17 | 3 | 128.0 | 534.28 | 2047 | 22.1 | Foram nanno marl, gray. |
| 17 | 4 | 25.0 | 534.75 | 1980 | 22.0 | Foram nanno chalk, gray, bedded. |
| 17 | 4 | 33.0 | 534.83 | 2059 | 22.0 | Foram nanno marl, bedded. |
| 17 | 5 | 48.0 | 536.48 | 2138 | 22.0 | Foram nanno marl. |
| 17 | 6 | 14.0 | 537.64 | 1812 | 22.0 | Volcanic clay. |
| 17 | 6 | 36.0 | 537.86 | 2014 | 22.0 | Foram nanno chalk, light gray. |

| Core | Section | Upper Interval ^a (cm) | Depth in Hole (m) | Velocity ^b (m/sec) | Temperature (°C) | Remarks | | |
|----------|---------|--|-------------------------|----------------------------------|---------------------|--|--|--|
| - | | 201010 | 10000 | | | | | |
| 18 | 1 | 30.0 | 539.30 | 1922 | 22.0 | Foram nanno chalk, light gray. | | |
| 18 | 2 | 70.0 | 541.20 | 2008 | 22.1 | Foram nanno marl. | | |
| 18 | 3 | 100.0 | 543.0 | 2059 | 22.1 | Foram nanno marl, dark gray (bedded). | | |
| 18 | 3 | 147.0 | 543.47 | 2290 | 22.1 | Nanno foram chalk, silty, light gray. | | |
| 18 | 4 | 37.0 | 543.87 | 1963 | 22.1 22.1 | Foram nanno marl, dark gray. | | |
| 18 | 4 | 78.0 146.0 | 544.28 | 2136 | 25.2 | Foram nanno chalk, light gray. | | |
| 18 18 | 5 | .66.0 | 546.46 547.16 | 2012 2102 | 25.2 | Foram nanno chalk, gray. Foram nanno chalk, gray. | | |
| 19 | 1 | 102.0 | 549.02 | 2702 | 25.2 | Nanno foram chalk, sandy. | | |
| 19 | 2 | 28.0 | 549.02 | 2702 | 25.2 | Nanno foram chalk, sandy. | | |
| 19 | 3 | 87.0 | 551.87 | 2381 | 25.2 | Foram nanno chalk, gray. | | |
| 19 | 4 | 101.0 | 553.51 | 2204 | 25.2 | Foram nanno chalk, gray. | | |
| 20 | 1 | 7.0 | 557.07 | 1924 | 25.4 | Foram nanno chalk, white, No. 3. | | |
| 20 | 2 | 82.0 | 559.32 | 2000 | 25.2 | Foram nanno chalk, gray. | | |
| 20 | 3 | 62.0 | 560.62 | 1847 | 25.2 | Foram nanno chalk, top, gray. | | |
| 20 | 3 | 62.0 | 560.62 | 2227 | 25.2 | Foram nanno chalk, bottom, white. | | |
| 22 | 1 | 72.0 | 575.72 | 1963 | 25.2 | Foram nanno chalk. | | |
| 23 | 1 | 16.0 | 584.16 | 2128 | 25.6 | Nanno chalk. | | |
| 23 | 2 | 5.0 | 585.55 | 2061 | 25.6 | Nanno chalk. | | |
| 23 | 2 | 45.0 | 585.95 | 1816 | 25.8 | Nanno chalk. | | |
| 23 | 2 | 58.0 | 586.08 | 2182 | | | | |
| 23 | 2 | 96.0 | 586.46 | 1815 | 25.6 | Clay, volcanic, waxy. | | |
| 23 | 3 | 4.0 | 587.04 | 1720 | 25.6 | Clay, volcanic, waxy. | | |
| 23 | 3 | 60.0 | 587.60 | 2217 | 25.8 | Nanno chalk. | | |
| 23 | 3 | 108.0 | 588.08 | 5179 | 25.8 | Chert, grayish red. | | |
| 23 | 3 | 108.0 | 588.08 | 2970 | 25.8 | Silicified nanno chalk, very pale red. | | |
| 24 | 1 | 63.0 | 593.63 | 2308 | 25.8 | Nanno marl. | | |
| 24 | 1 | 140.0 | 594.40 | 5086 | 25.8 | Chert, brown. | | |
| 25 | 2 | 84.0 | 604.34 | 2540 | 25.8 | Nanno marl; bedding. | | |
| 26 | 1 | 121.0 | 612.21 | 2072 | 24.2 | Nanno chalk, reddish gray. | | |
| 26 | 2 | 68.0 | 613.18 | 2953 | 24.2 | Nanno chalk, hard, reddish gray. | | |
| 27 | 1 | 89.0 | 620.89 | 4727 | 24.2 | Chert, grayish red, waxy luster, conchoidal. | | |
| 27 | 1 | 92.0 | 620.92 | 3141 | 24.2 | Nanno limestone, white between chert. | | |
| 27 | 2 | 71.0 | 622.21 | 3101 | 24.2 | Nanno limestone. | | |
| 27 | 2 | 130.0 | 622.80 | 4698 | 24.3 | Chert, red. | | |
| 27 | 2 | 130.0 | 622.80 | 4672 | - 24.2 | Chert, red. | | |
| 28 | 1 | 138.0 | 630.38 | 3719 | 24.2 | Radiolarian limestone, partly silicified, near chert | | |
| 28 | 2 | 96.0 | 631.46 | 2681 | 22.4 | Radiolarian limestone, partly silicified. | | |
| 28 28 | 3 | 40.0 114.0 | 632.44 633.14 | 5092 3386 | 22.0 22.0 | Chert, brown. Radiolarian limestone; partly silicified. | | |
| 29 | 1 | 118.0 | 639.18 | 2510 | 22.0 | Radiolarian limestone; partly silicified. | | |
| | | | | | | | | |
| 30 30 | 2 3 | 58.0 65.0 | 649.08 650.65 | 2581 2661 | 22.5 22.0 | Radiolarian limestone; partly silicified. Radiolarian limestone; partly silicified. | | |
| 30 | 4 | 89.0 | 652.39 | 2766 | 22.0 | Radiolarian limestone; partly silicified. | | |
| 31 | 1 | 118.0 | 657.18 | 5116 | 22.0 | Chert, red. | | |
| 31 | 2 | 44.0 | 657.94 | 2777 | 22.1 | Foram radiolarian limestone, partly silicified. | | |
| 31 | 3 | 77.0 | 659.77 | 2624 | 22.1 | Foram radiolarian limestone, partly silicified. | | |
| 31 | 4 | 0.0 | 660.50 | 2454 | 37.7.101 | , p, | | |
| 31 | 4 | 104.0 | 661.54 | 3162 | 22.1 | Foram radiolarian limestone, partly silicified. | | |
| 32 | 1 | 83.0 | 665.83 | 2666 | 22.1 | Radiolarian limestone. | | |
| 32 | 2 | 148.0 | 667.98 | 3203 | 22.3 | Radiolarian limestone. | | |
| 33 | 1 | 130.0 | 675.30 | 3344 | 22.3 | Radiolarian limestone. | | |
| 34 | 1 | 75.0 | 683.75 | 2859 | 22.3 | Radiolarian limestone. | | |
| 35 | 1 | 95.0 | 692.95 | 3043 | 22.1 | Radiolarian limestone. | | |
| 35 | î | 136.0 | 693.36 | 3813 | 22.3 | Radiolarian limestone. | | |
| 35 | 2 | 62.0 | 694.12 | 3257 | 22.3 | Radiolarian limestone. | | |
| 36 | 2 | 130.0 | 703.80 | 2662 | 22.3 | Foram radiolarian limestone. | | |
| 38 | 1 | 86.0 | 714.86 | 2706 | 22.3 | Radiolarian limestone. | | |
| 39 | 1 | 88.0 | 719.88 | 3689 | 22.3 | Volcanic ash. | | |
| 39 | î | 88.0 | 719.88 | 3647 | 22.3 | Volcanic ash. | | |
| | | | | | | | | |

TABLE 2 - Continued

| Core | Section | Upper Interval ^a (cm) | Depth in Hole (m) | Velocity ^b (m/sec) | Temperature (°C) | Remarks |
|------|---------|--|-------------------------|----------------------------------|---------------------|-----------------------|
| 39 | 1 | 136.0 | 720.36 | 3049 | 22.3 | Volcanic ash, banded. |
| 39 | 2 | 112.0 | 721.62 | 2883 | 22.1 | Volcanic ash. |
| 41 | 1 | 105.0 | 738.05 | 3082 | 22.1 | Foram limestone. |
| 41 | 1 | 140.0 | 738.40 | 4859 | 22.1 | Diabase. |

TABLE 2 – Continued

^aUpper limit of a 3-cm sample interval.

^bVelocities were measured parallel to bedding unless noted otherwise.

^cThe few samples whose velocity was measured through the core liner and where a "D"-shaped lucite block was used to obtain liner thickness and travel time.

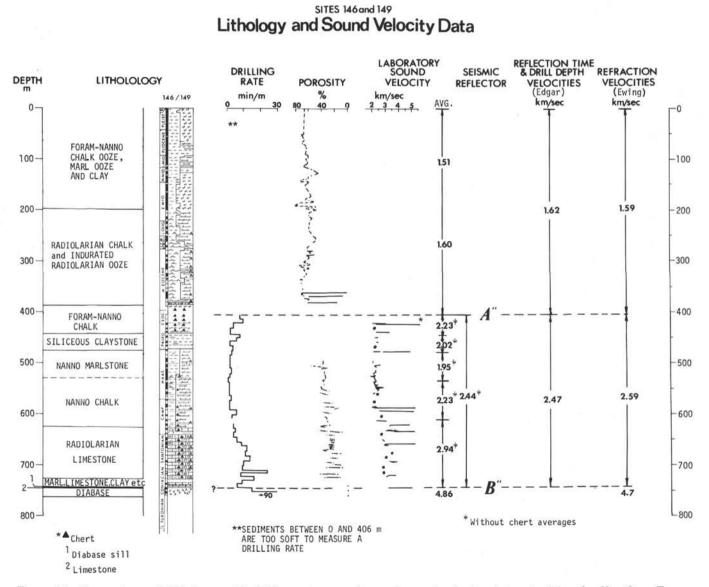


Figure 14. Comparison of lithology with drilling rate, porosity, and sound velocity determined by the Hamilton Frame method onboard Glomar Challenger by calculation from drilling depth and reflection time and from sonobuoy velocities (J. Ewing, personal communication).

| Avg. Vpa (km/sec) | Rock | Depth (m) | | |
|----------------------|--|--------------|--|--|
| 1.51 | Unlithified clay | 98-105 | | |
| 1.60 | Chalk ooze | 254-263 | | |
| 2.23 | Foraminiferal nannofossil chalk | 406-440 | | |
| 2.02 | Green siliceous claystone | 440-478 | | |
| 1.95 | Nannofossil marl | 478-529 | | |
| 2.23 | Varicolored nannofossil chalk | 529-612 | | |
| 2.94 | Chalk, radiolarian limestone, variegated limestone, marlstone, and claystone | 612-737 | | |
| 4.86 | Diabase | 737-738 | | |

TABLE 3 Major Lithologic Sections Stratigraphically Distinguished by their Velocity Characteristics

^aWithout chert data.

Velocity changes versus pressure variations for many rocks, according to Press (1966) and Christensen (1970), are very sensitive to pressure changes from 1 to 500 bars. However, all dolerite samples listed by Press (1966, Table 9-2) suggest little change of sound velocity within these pressures. See the Schrieber and Fox chapter in this volume on rock velocities relative to pressure.

Other problems may arise, in comparing these differing types of sound velocity data, from acoustic anisotrophy in these rocks. Seismic reflection velocity records occur in a direction perpendicular to stratification, shipboard Hamilton Frame velocity measurements were predominately parallel to stratification and Ewing's are interval and refracted velocities.

Natural Gamma Radiation

Natural gamma ray emissions were counted for a period of 1.25 min at 7.62 cm (3 in) intervals along the core, with a counting precision of about ± 100 counts. Sediment disturbance, porosity control, methods, and equipment are discussed in the Appendix.

Cores 1 through 6 at Site 149 were recovered in the extended core barrel which has a 2.25-inch internal diameter compared to the normal 2.60 inches. The volumetric difference per unit length between the two core sizes relative to the smaller core is 33.7%, therefore the natural gamma ray counts of 2.25-inch cores were increased by 33.7% so that they may be compared to the normal 2.60 inch data.

Natural gamma radiation at Sites 146 and 149 roughly corresponds to the different types of sediments or rocks, which range from Recent to Cretaceous and are from 0 to 760 meters below the sea floor. Gamma ray emissions range from 0 to 2100 counts/1.25 min period at 7.62-cm core intervals with typical high values for clays and marls, intermediate values for chalk, and the lower values for radiolarian-rich chalks, radiolarian-rich limestone, and radiolarian ooze. The diabase emitted the least radiation.

Pliocene-Miocene brown clays emitted natural gamma radiation ranging from 500 to 1900 counts with typical counts of about 1500, while Paleocene green siliceous claystone emitted 400 to 700 counts. Pleistocene-Miocene nannofossil and foraminiferal chalk and marl ooze ranged from about 500 to 1300 with typical counts of 900,

Eocene radiolarian ooze ranged from 200 to 400, and Cretaceous radiolarian limestone from 100 to 400. In general, natural gamma ray emissions from mixtures of these sediment end members were related to their mixed proportions. The lowest counts of 100 are emitted from the diabase.

Penetrometer

Needle penetration tests were made at Sites 147 and 149 with a 1-mm diameter needle. Sediment disturbance, methods, and equipment are discussed in the Appendix. CP on the graph stands for complete penetration.

Penetration ranged from 26 mm to zero with penetration decreasing with increasing depth below the sea floor to 230 meters. Between 0 and 30 meters below the sediment surface, needle penetration in Pleistocene-Pliocene gray orange nannofossil chalk and marl ooze decreased from 17 to 9 mm with a range of 9 to 26 mm. Pleistocene-Pliocene gray nannofossil foraminifera marl ooze and foraminifera nannofossil marl recovered from 30 to 70 meters below the sea floor were typically penetrated 10 mm with a range of 6 to 15 mm, while penetrations were less in Pliocene-Miocene brown clay and marls from 70 to 110 meters, which have typical penetration values of about 3 mm (1.5 to 5 mm limits). From 110 meters to 140 meters, penetration was typically 2 to 3 mm with narrower limits (1 to 3 mm), while below 140 meters, in Oligocene and older sedimentary rock, penetration was usually within 0 to 3 mm regardless of lithology (chalks and marls, and radiolarian chalks).

BIOSTRATIGRAPHY

The uppermost lithologic unit recognized at Site 149 consists of foraminiferal nannoplankton chalk and marl ooze and is about 75 meters thick. The first core was taken mostly above the sea floor so that only the core catcher contained sediment. The sample contained abundant pteropods, planktonic foraminifera and calcareous nannofossils. The planktonic foraminifera belong to the Globorotalia fimbriata Subzone of the Globorotalia truncatulinoides truncatulinoides Zone and the calcareous nannoplankton belong to the Emiliania huxleyi Zone; the sample is Holocene. Core 2 (1-10 m below the sediment surface) sampled somewhat older, definitely Pleistocene sediments. The planktonic foraminifera of the upper two sections belong to the Globigerina calida calida Subzone of the Globorotalia truncatulinoides truncatulinoides Zone. Samples of the Globigerina bermudezi Subzone of the Globorotalia truncatulinoides truncatulinoides Zone, which lies between the Globorotalia fimbriata Subzone and the Globigerina calida calida Subzone and is considered to be latest Pleistocene, were not recovered. Calcareous nannoplankton assemblages from the top of Core 2 also suggest that the youngest Pleistocene was not samples. At this site the lowest occurrences of Globigerina calida calida and Emiliania huxleyi both lie in the lower part of Section 2 or the upper part of Section 3 of Core 2. The lower sections of Core 2 and all of the material in Cores 3, 4, and 5 belong to the Globorotalia hessi Subzone of the Globorotalia truncatulinoides truncatulinoides Zone according to the planktonic foraminifera or the *Gephyrocapsa oceanica* Zone using calcareous nannoplankton fossils. The highest occurrence of *Pseudoemiliania lacunosa* lies in the lower part of Section 2 or the upper part of Section 3 of Core 3, so that if the terminology of Gartner is used, only the upper two sections of Core 3 and lower four sections of Core 2 belong to the *Gephyrocapsa* Zone, lower sections and cores to the middle of Core 6 belonging to his *Pseudoemiliania* Zone. Very rare, well-preserved radiolarians are present in Core 2 but are not found in deeper samples from this unit.

The planktonic foraminiferal assemblages from Cores 6 to 11 are generally mixed, the sediments being disturbed by drilling; calcareous nannofossil assemblages from this same interval are also often mixed, but some small undisturbed pieces of the sediment yield normal assemblages in the expected stratigraphic order. Core 6 contains a highly condensed Gephyrocapsa oceanica-Gephyrocapsa caribbeanica-Discoaster brouweri-Discoaster pentaradiatus Zone sequence in less than 6 meters, probably a result of drilling disturbance. The highest occurrence of Discoaster brouweri has been used to determine the Pliocene-Pleistocene boundary between samples from Sections 3 and 4 of Core 6. Sections 1 to 5 of Core 7 are mixed in such a way as to suggest slumping of material from higher in the hole. Only the lowest section (6) of Core 7 contains an apparently indigenous assemblage, belonging to the Discoaster pentaradiatus Zone. Similarly, the first three sections of Core 8 contain mixed assemblages, but apparently indigenous assemblages of the Discoaster surculus Zone are present in Sections 4 and 6 of this core. In Core 9 it is possible to select undisturbed samples for calcareous nannofossils, particularly from the lowest section (6).

The interval from 75 to 138 meters (Cores 10-16) is represented by brown clay with some interbedded marl containing calcareous pelagic fossils. As noted above, the planktonic foraminiferal assemblages from Cores 10 and 11 appear to be mixed. In these cores the planktonic foraminifera are often damaged by solution effects. The calcareous nannoplankton assemblages have also been attacked and are particularly poor in the upper sections of Core 10. In Core 11 the Reticulofenestra pseudoumbilica and Discoaster asymmetricus zones are present in Sections 1 to 3 and 4 to 5 (to 115 cm) but the earlier Pliocene Ceratolithus rugosus and Early Pliocene-Late Miocene Ceratalithus tricorniculatus zones are missing, the lowest part of Section 5 (129 cm) belonging to the Discoaster quinqueramus Zone of the Late Miocene. Section 5 of Core 11 is marly brown clay, and it may be that the calcareous nannofossils representing the latest Miocene-early Pliocene have been completely dissolved. The assemblages which bracket the barren interval between 115 cm and 129 cm are not necessarily in situ, so it is possible that the real thickness of the Early Pliocene-Late Miocene might be greater than the apparent 14-cm interval observed in the core.

The planktonic foraminiferal assemblages throughout the remainder of this unit are moderately to severely damaged by solution. The *Globorotalia acostaensis* Zone of the Late Miocene and the *Globorotalia menardii*, *Globorotalia mayeri*, *Globorotalia fohsi lobata*, *Globoro*- talia fohsi fohsi, and Globorotalia fohsi peripheroronda zones of the Middle Miocene are all recognizable in Cores 12 to 16. However, the Globorotalia fohsi robusta Zone, which should intervene between the Globorotalia fohsi lobata and Globorotalia mayeri zones in the Caribbean was not detectable. The calcareous nannofossil sequence of Discoaster quinqueramus-Sphenolithus heteromorphus zones is complete, so that the absence of the Globorotalia fohsi robusta Zone may be due to the lack of the defining species in the impoverished planktonic foraminiferal assemblages from the lower sections of Core 15.

The subjacent lithic unit recognized at Site 149 is foraminiferal nannoplankton chalk which extends from 138 to 195 meters (Cores 17-22). In cores 17 to 20 the planktonic foraminiferal assemblages are generally well preserved or only slightly to moderately affected by dissolution and represent the Globorotalia fohsi peripheroronda and Praeorbulina glomerosa of the Middle Miocene and Globigerinatella insueta Zone of the late Early Miocene. These cores contain calcareous nannoplankton assemblages belonging to the Sphenolithus heteromorphus and Helicopontosphaera ampliaperta zones. Preservation of calcareous fossils of both groups is notably poorer in Cores 21 and 22 through Section 6. The Globigerinita stainforthi and Globigerinita dissimilis zones (mid-early Miocene) are not distinguishable and are lumped together. The Sphenolithus belemnos Zone can be recognized in sediment at 94 to 96 cm in Section 2 of Core 21. The remainder of Core 21 and Core 22 through Section 6 belongs to the Discoaster druggi Zone. Core 22 contains very rare, poorly preserved fragments of radiolaria, inadequate for zonal assignment.

The fourth lithic unit recognized at Site 149 is an indurated calcareous radiolarian ooze and radiolarian nannoplankton chalk. The core catcher of Core 22 and all subjacent cores (through 43) recovered at Site 149 belong to this unit which extends from 195 meters to the bottom of the hole at 390 meters.

The core catcher of Core 22 and subjacent cores through Section 3 of Core 27 are grayish green radiolarian chalk. Planktonic foraminiferal assemblages are only slightly affected by solution in the higher part of this unit and become progressively more damaged toward the bottom of the unit. The boundary between the Globigerinoides primordius Zone (earliest Miocene) and Globorotalia kugleri Zone (latest Oligocene) is between Cores 25 and 26. In the case of the calcareous nannofossils, the base of the Discoaster druggii Zone lies between samples from Sections 3 and 4 of Core 23. Radiolarians are few to common in Cores 23 to 27 and somewhat dissolved. Core 23 belongs to the Lychnocanoma elongata Zone as do subjacent samples to the lower part of Core 26. The base of Core 26, all of Core 27, and part or all of Core 28 belong to the Dorcadospyris ateuchus Zone.

The sediment from Section 3 of Core 27 through Section 2 of Core 30 is a yellowish gray radiolarian chalk. Planktonic foraminifera are sparse and solution effects have reduced the assemblages to only a few resistant species. Material in Cores 27 and 28 is tentatively referred to the *Globorotalia kugleri* Zone and Cores 29 and 30 to the *Globorotalia opima opima* Zone. Calcareous nannoplankton assemblages in this interval are also affected by solution, but it is still possible to recognize the sequence Triquetrorhabdulus carinatus (Core 27, Section 3–Core 29, Section 1), Sphenolithus ciperoensis Zone (Core 29, Section 2), Sphenolithus distentus Zone (Core 29, Section 3), and Sphenolithus predistentus Zone (Core 30), although secondary species must be used to determine zone boundaries in the absence of the defining species. Nevertheless, this is good evidence that the Late and Middle Oligocene are represented by the sediments in this interval. The radiolarians are few to common and have been somewhat affected by dissolution. The base of the Dorcadospyris ateuchus Zone lies in the lower part of Core 28 or upper part of Core 29. The base of the Theocyrtis tuberosa Zone, which may approximate the Eocene-Oligocene boundary, lies between Cores 30 and 31.

Cores 31 to 34 (Section 3) consist of radiolarian chalk ooze. Planktonic foraminifera are virtually absent although debris is present, indicating that they have been removed through dissolution. Calcareous nannofossils are very sparse in Core 31, somewhat more diverse but still very poor assemblages occur in Cores 32 to 34. Samples from Core 34 are referable to the Middle Eocene Discoaster saipanensis Zone. The radiolarians in these cores are abundant and well preserved, but the assemblages in Cores 31 and 32 contain a high proportion of reworked specimens. As a consequence it is difficult to recognize the bases of the Late Eocene Thyrsocyrtis bromia and Podocyrtis goethana zones which must lie in these cores. The base of the late Middle Eocene Podocyrtis chalara Zone is probably in the lower part of Core 32, and subjacent cores belong to the Podocvrtis mitra Zone.

Core 34, Section 3, through Core 43 (301-390 m) are radiolarian chalk and ooze. Planktonic foraminifera are virtually absent. Calcareous nannofossil assemblages have also been strongly attacked by solution, but it appears possible to recognize the Discoaster tani nodifer Zone in Cores 35 and 36, the Chiphragmalithus alatus Zone in Cores 37 to 39, and the Discoaster sublodoensis Zone in Cores 40 to 41; all belong to the Middle Eocene. No calcareous nannofossils were encountered in Cores 42 and 43. The Radiolaria are abundant and well preserved. The base of the Podocyrtis mitra Zone is between the bottom of Core 36 and the lower part of Core 37. The base of the Podocyrtis ampla Zone is between Cores 38 and 39. The base of the Thyrsocyrtis triacantha Zone is between Cores 42 and 43, and Core 43 belongs to the Theocampe mongolfieri Zone. All of the radiolarian zones are considered to belong to the Middle Eocene.

Two spot cores were taken in post-Eocene sediments at Site 146. In the first core taken at 96 to 105 cm, Sections 1 through 6 were devoid of foraminifera and Radiolaria, but contain fish debris and a poor calcareous nannofossil assemblage of the *Discoaster quinqueramus* Zone (Late Miocene). The core catcher sample from Core 1 contains rich planktonic foraminiferal and calcareous nannofossil assemblages belonging to the Late Miocene *Neogloboquadrina dutertrei* and *Discoaster quinqueramus* zones respectively. Correlation by depth of this core with Site 149 would suggest that this level should belong to the *Globorotalia acostaensis* Zone, and the sediment sampled may represent material slumped from slightly higher in the hole.

Core 2 was taken between 254 and 263 meters below the sediment surface. The upper four sections contain rich planktonic foraminiferal assemblages assigned to the Early Miocene Globigerinoides primoridius Zone. Calcareous nannofossils are also abundant in these sections and belong to the Discoaster druggi Zone. Sections 5 and 6 are devoid of planktonic foraminifera but contain thick-walled benthonic foraminifera and rich calcareous nannofossil assemblages of the early Miocene, part of the Triquetrorhabdulus carinatus Zone. Radiolarians are abundant and moderately well preserved. Those from the upper half of the core belong in the Lychnocanoma elongata Zone and those from the lower part in the Dorcadospyris ateuchus Zone, but, because the core is disturbed, the transition between these zones cannot be investigated in detail. Correlation by depth with Site 149 indicates that this material may be slumped from a level some 40 meters higher.

Continuous coring at Site 146 was initiated at a depth of 406 meters. Cores 3 to 8 belong to a lithified foraminiferal nannoplankton chalk with chert which extends from 382 to 441 meters below the sediment surface. Cores 3 and 4 yielded mostly chert fragments with no foraminifera or calcareous nannoplankton. Radiolarians are rare to common, moderately well to poorly preserved. Core 4 evidently belongs in the Buryella clinata Zone. Cores 5 to 7 contain some levels with planktonic foraminiferal assemblages adequate for dating, although all have been affected to some degree by dissolution. All of these assemblages are assigned to the new Globorotalia edgari Zone of earliest Eocene age. Calcareous nannofossil assemblages from rock fragments in the upper part of Core 5 belong to the Discoaster diastypus Zone: those from deeper levels in Core 5 and from Cores 6 and 7 belong to the Discoaster multiradiatus Zone. A meager assemblage from the top of Core 8 may belong to either the Heliothius riedeli or Discoaster gemmeus Zone (Paleocene). Radiolaria are as above, moderately well to poorly preserved, some being partially dissolved, thus silicified, and others perhaps zeolitized. An assemblage from Core 7 belongs to the Bekoma bidarfensis Zone.

A monotonous green siliceous clay was found in Cores 9, 10, and the upper part of 11. (441-477 m). Foraminifera and calcareous nannofossils are absent, and the Radiolaria belong to that part of the Paleocene left unzoned in the Leg 10 Initial Report.

A calcareous layer at 128 to 138 cm in Section 1 of Core 11 contains planktonic foraminifera belonging to three distinct assemblages occurring in different lenses. These assemblages belong to the *Globorotalia pseudobulloides* and the *Globigerina eugubina* zones of the earliest Paleocene, and a *Globotruncana* Zone of the Late Cretaceous. No calcareous nannoplankton were found in this material, but there are many small calcite crystals which may indicate complete recrystallization of the nannofossils. Cretaceous calcareous nannofossils appear immediately below this layer and the Cretaceous-Tertiary boundary is indicated at 138 cm in Section 1 of Core 11. The highest assemblage at 140 cm is extremely poor, but 3 cm below, specimens become more abundant. Moderately good diverse assemblages of Late Cretaceous forms appear in Section 2 of Core 11. The radiolarians in Core 11 (476 to approximately 480 m) have what is commonly regarded as a "Cretaceous aspect," but they might equally well be early Paleocene since no well-dated assemblages of that age are known.

The sequence from 477 to 630 meters (Cores 12-27) is a varicolored nannoplankton chalk and marl. Planktonic foraminifera are abundant and the assemblages diverse. although almost all have been strongly damaged by dissolution. Core 12 did not recover any material for analysis. Core 13 belongs to the Abathomphalus mayaroensis Zone, Core 14 and the top of Core 15 to the Globotruncana contusa Zone, the remainder of Core 15 and subjacent material to near the base of Core 18 to the Globotruncana gansseri Zone, and the base of Core 18 through the base of Core 21 to the Globotruncana "tricarinata" Zone, all assigned to the Maestrichtian. Calcareous nannofossil assemblages in this interval are moderately well to poorly preserved and all show dissolution effects. Species diversity is generally low and it is impossible to establish a fine zonation. Core 22 contains relatively well-preserved planktonic foraminifera of the Globotruncana calcarata Zone (Late Campanian). Cores 23 to 27 contain progressively more corroded planktonic foraminifera belonging to the Globotruncana elevata Zone (Early Campanian). In this same interval the diversity of calcareous nannofossil assemblages decreases markedly.

From 630 to 738 meters the sequence consists of radiolarian limestone. Planktonic foraminifera are strongly affected by dissolution and the diversity of assemblages is considerably reduced. Nevertheless, it is possible to recognize the *Globotruncana concavata carinata* Zone (Cores 34-38), indicating the Santonian. The basal calcareous sediments (Cores 39-41) contain the assemblage of the *Globotruncana schneegansi* Zone of Late Turonian. Calcareous nannofossil assemblages in this unit tend to be very poor.

Practically all samples from Cores 12 to 41 (from 485 to approximately 740 m) contain rare, moderately wellpreserved (partially dissolved) radiolarians representing a small number of taxa. Many samples (particularly those from sandy layers) from cores below Core 14 contain, in addition, common to abundant, poorly preserved (usually silicified) radiolarians representing large numbers of taxa. These results, together with the fact that an orderly sequence of foraminiferal faunas is found in the cores, suggest that the sparse, better preserved radiolarians are autochthonous, and that the floods of poorly preserved specimens were introduced by some mechanism such as turbidity currents, which carried material not significantly older than that being deposited at this site. This explanation requires a pattern of paleoecological conditions such that the water column at this site was inhabited by a restricted radiolarian assemblage, while at the presumably not-far-distant source of the transported component, a much more diverse fauna flourished.

CONCLUSIONS

Cores recovered from Site 146/149 give a record of apparently uninterrupted sedimentation for the last 80 m.y. Although total recovery was about 50 per cent, only 2 out of 86 core barrels were completely empty. Thus, this site affords an exceptionally fine opportunity for studying a long record of sedimentary history in the Caribbean area.

As a result of the coring and measurement of sound velocities and densities of the rocks of this site, we feel that we can identify Horizons A" and B" with a high degree of confidence. Horizon A" correlates with the chalk/limestone and chert interface separating lithic units 2 and 3. Horizon B" is identified here with the upper dolerite sill (B" was found to be basalt or dolerite in sediments of approximately the same age at four other sites of Leg 15). The latter conclusion is of extreme importance: the nearly ubiquitous distribution of Horizon B" in the Venezuelan Basin and its occurrence in much of the Colombian Basin implies a wide areal extent to this igneous episode.

The earliest lithic unit is extremely varied and shows striking contrasts between successive layers. Although dominantly a radiolarian limestone, there are, especially near the base, intercalated basaltic ash beds, carbonaceous layers, and radiolarian sands. Glauconite and pyrite are present. This diversity in lithologies implies strongly that the topography of the sea floor was irregular, with high and low areas in juxtaposition. The reduced, organic-rich sediment, the relatively high phosphorous contents (Donnelly, Mineralogy and Chemistry of Noncalcareous Sediments, this volume), and the presence of glauconite and pyrite all suggest sedimentation on a bottom with limited water circulation. The intercalations of radiolarian sands indicate incursions of detritus from nearby topographic highs (see Biostratigraphic Summary); the sorting (with or without grading), sharp bottom contacts, and the relatively well-preserved Radiolaria and planktonic foraminifera in a sedimentary environment below the lysocline all indicate that these beds represent rapidly transported sediment from topographic high areas within the basin.

The lower radiolarian limestones pass upward into Campanian-Maestrichtian chalks and marls and younger chalks, oozes, marls, and clays. The gradual disappearance of the intercalated radiolarian sands and the blurring together of lithologies results from the gradual subduing of sea floor topography through sedimentation. The entire post-Cretaceous section is characterized by the near absence of juxtaposed contrasting lithologies, except that ash beds are recognizable as late as Oligocene at Site 149.

The entire post-Santonian section represents nearly constant pelagic sedimentation, with a variable volcanic admixture, under conditions dominated by changing patterns in productivity of calcareous and siliceous organisms and by the relative movement of the depth of calcium carbonate compensation above and below the sea bottom at the site. Thus, the Paleocene and Late Miocene-Early Pliocene intervals are marked by complete solution of calcareous fossils. During most of the remaining times, the site was at about the level of the depth of calcium carbonate compensation, and the planktonic foraminifera were more or less dissolved. During brief intervals the planktonic foraminifera escaped solution. A significant accumulation of siliceous fossils is seen from the late Cretaceous through the late Eocene, reaching a maximum during the latter unit of time and later falling to a low level through the remainder of the Cenozoic Era.

The history of volcanic activity shown by the sediments at Site 146/149 is a complete and interesting one (discussed more fully in Donnelly, Circum-Caribbean Explosive Volcanic Activity . . . , this volume). The basaltic episode at the base of the section actually consisted of several eruptions, represented by basaltic ash layers at Site 146 in Cores 39 and 41. The volcanic clays which form impressively thick layers in the later Cretaceous section (many are several cm thick) are of special interest. Several contain minerals typical of subalkalic, highly explosive eruptions of the continental margins (quartz, biotite, alkali feldspar, apatite, and zircon), but most consist of glass which has completely devitrified to montmorillonite. If these ashes are of continental origin, their thickness at relatively great distances from possible sources is remarkable; igneous activity of that age is found only in the Greater Antilles and, to a limited extent, along the Venezuelan coast. However, similar ashes outcrop in Puerto Rico, Jamaica, Curacao, and Bonaire. The possibility that they may represent eruptions from within the basin cannot be completely dismissed; similar volcanic materials are known from the east and southeastern Pacific.

After a diminution of volcanic activity during the late Maestrichtian-Paleocene, there was a sharp renewal of activity in the middle Eocene. This material is entirely pumice of a calc-alkaline origin, and the probable source was the Lesser Antilles. This activity diminished to a minimum during the middle Miocene but apparently was renewed in the late Pliocene to Recent (Site 148).

The occurrence of chert at Horizon A" provides some valuable insights into the origin of chert layers of this age in many parts of the ocean. The analysis of Donnelly and Nalli (this volume) shows that the present non-chert silica component is not exceptionally great at the level of the main cherts, but was much greater at a higher level at which fresh opaline radiolarian and fresh glass shards were found. The chert represents a recrystallization, or lithification, of opal and volcanic glass, and its coincidence with chalks (in which calcitic fossil debris is recrystallized to sparry calcite) indicates that the problem of chert formation is one of the post-deposition chemical environment, not the rate of supply of silica to the sea floor.

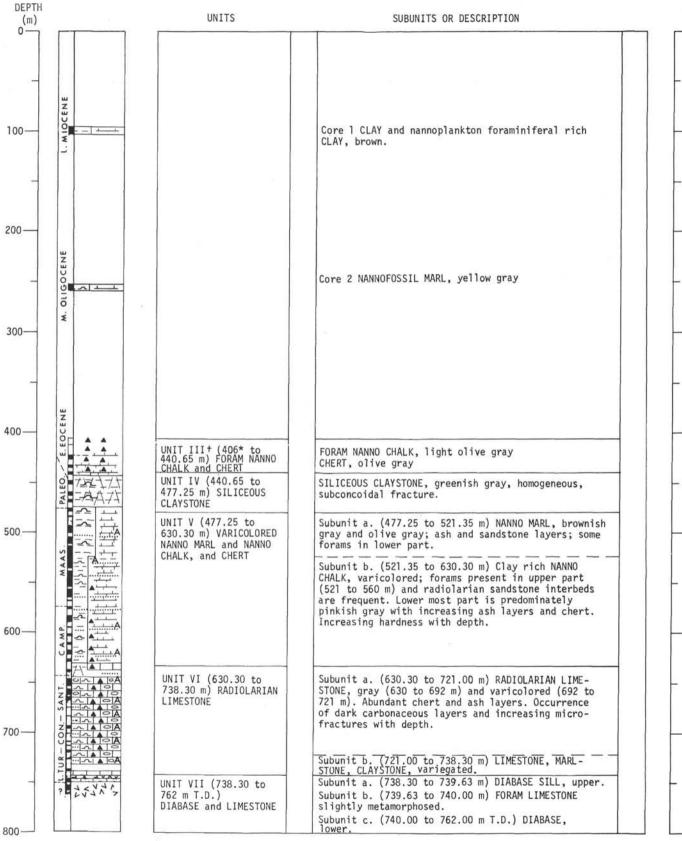
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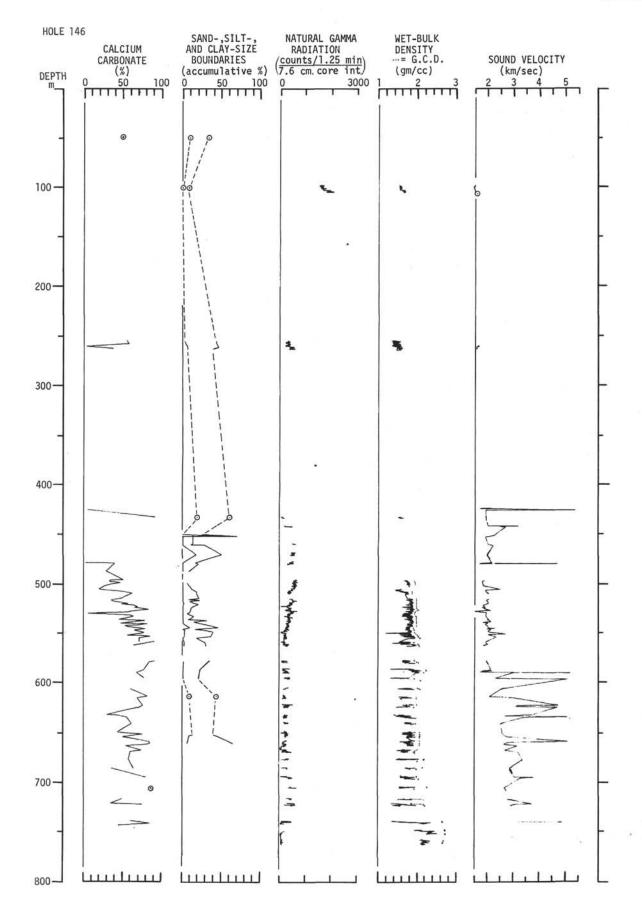


HOLE 146

LITHOLOGY

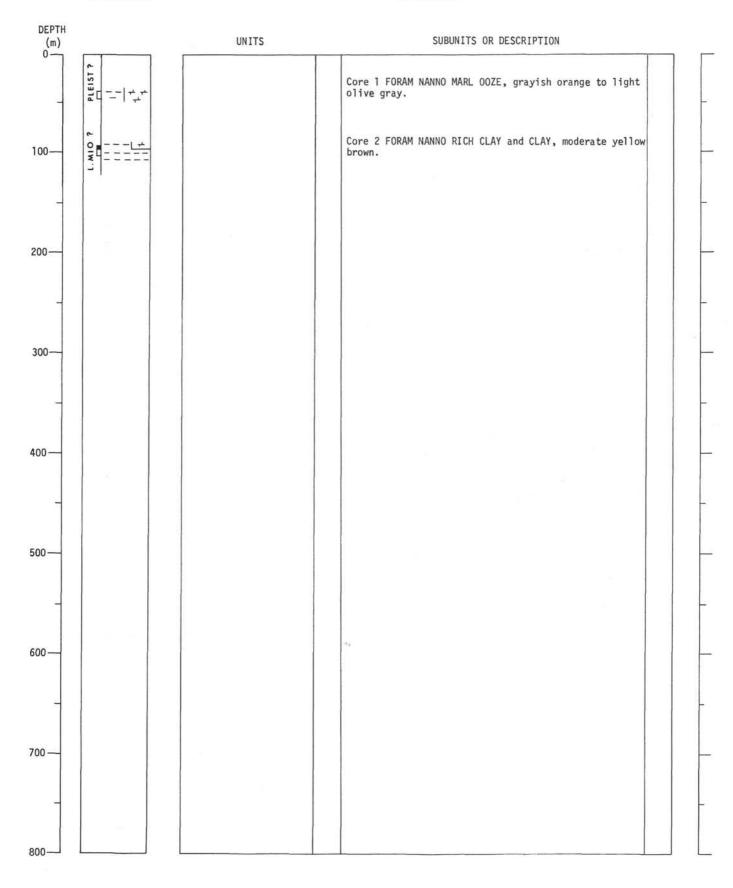


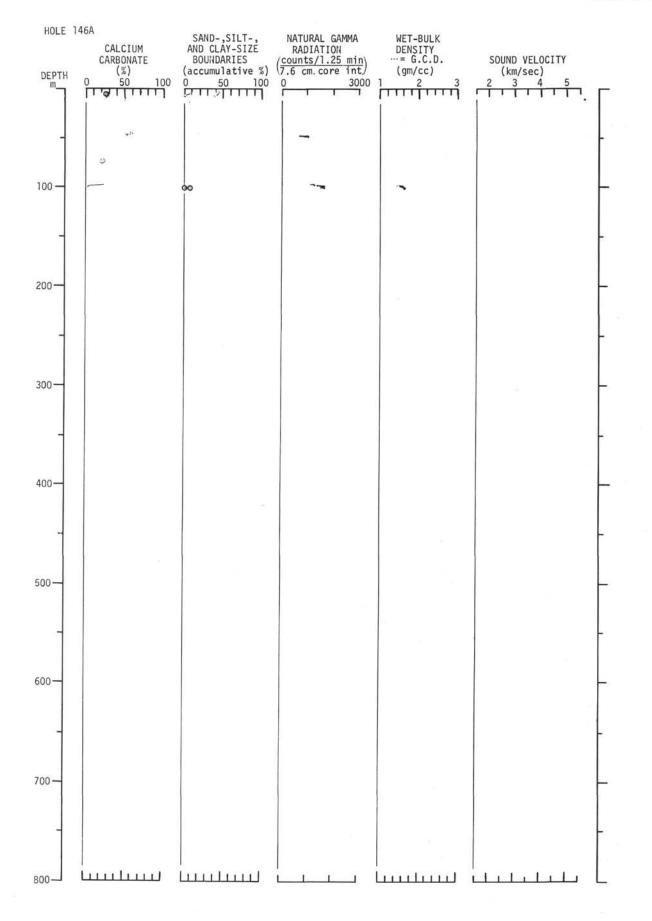
⁺See Site 149 for Lithologic Unit I through V. *Slightly higher in Site 149.



HOLE 146A

LITHOLOGY

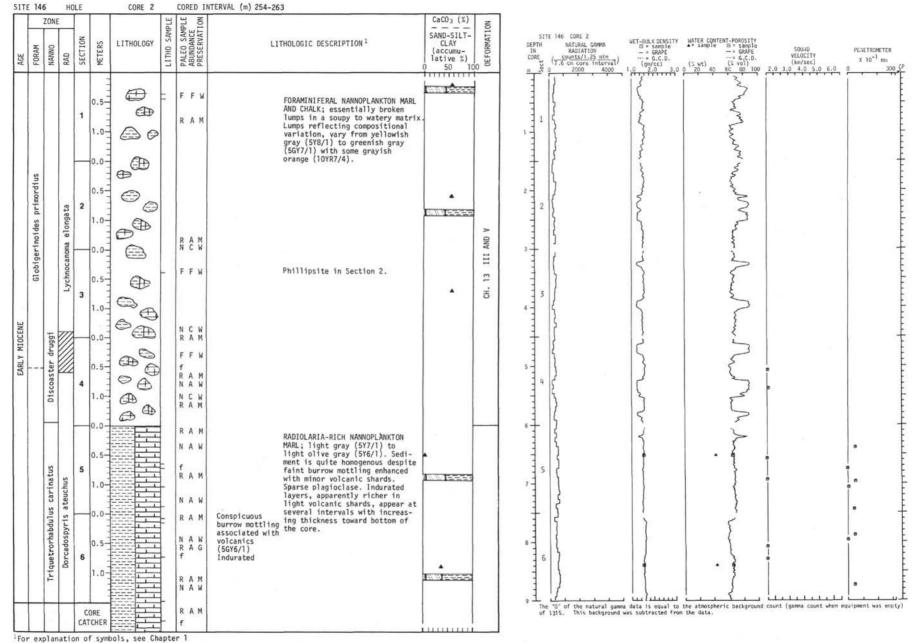




| SITE | 146 | ł | OLE | | CORE 1 | | | NTERVAL (m) 96-105 | | | |
|--------|-----|---------------|------------------|---|-----------|--------------|---|---|---|------------|--|
| AGE | Т | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (% SAND-SILI CLAY (accumu- lative % 0 50 | DEFORMATIO | SITE 146 CORE 1 MET-BULK DENSITY WATER CONTENT - POROSITY DEPTH NATURAL GAMMA D = sample A = sample D = sample IN ANDIATION |
| LATE 1 | | qu'inqueranus | 3 4 5 6 | 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 0.5 1.0 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0 | | | f n fn n fn n ffn N N FRAW | CLAY; predominantly pale yellow- ish brown (10YR8/2) to dark yellowish brown (10YR8/4) with irregular shade changes due to mottling and secondary distur- bances from drilling. Sparse plagioclase and hornblende. Section 1 is mixed with light olive gray (5Y5/2) and contains displaced nannoplankton. Remainder of core is barren except for core catcher which is rich in planktonic foraminifers, quartz, chlorite, and kaolinite. | | E4 13 111 | The start of the start of the start of the the transfer of a start of the start |

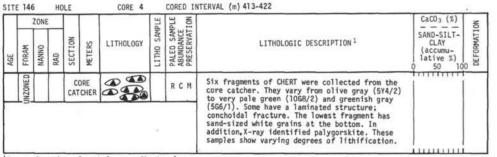
SITE 146/149

·For explanation of symbols, see Chapter 1

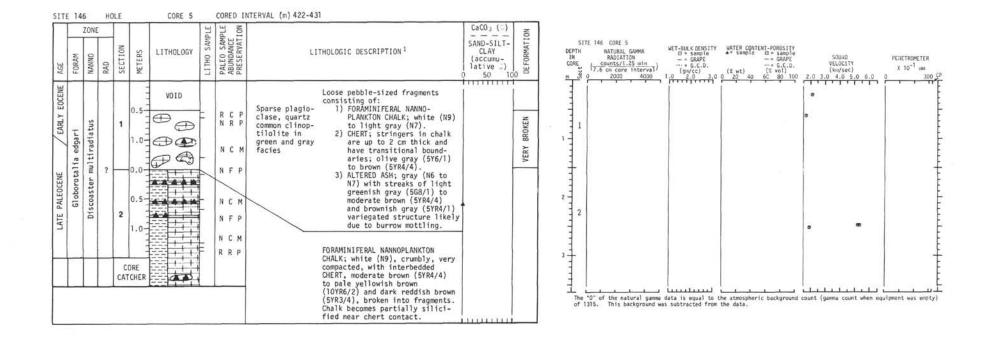


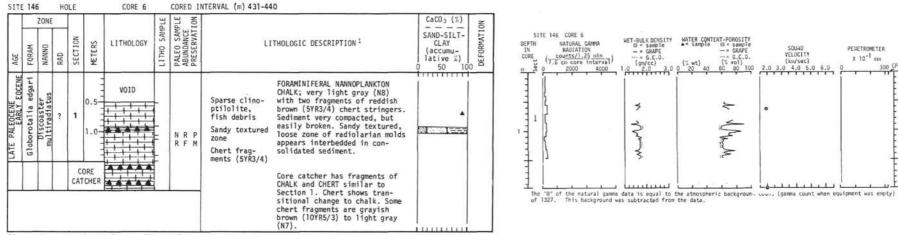
SITE 146/149

| ITE | 14 | 6 | HC | HOLE | | CORE 3 | _ | | NTERVAL (m) 406-413 | | | | | |
|--------------|----|------|-----|---------|-------------|-----------|--------------|---|--|--|--|--|--|--|
| AGE FORAM | | ZONE | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 100 | | | | |
| | | | | | DRE CHER | | | RFM | Only two fragments were recovered from the core catcher: CHERT; dark gray (5Y4/1) with conchoidal fracture and a potential grayish green (5G5/2) alteration rim. CHERT; dark olive gray (5Y3/2) covered on one side with a thin (0.5 cm) layer of very pale green (1068/2) chert. In thin section, cherts contain bands of lighter colored, somewhat calcareous materials. These bands resemble, in their shape and mutual relationship, the mottled and bioturbated areas of the adjoining limestone. Mainly parallel, well-preserved to recrystallized radiolarians are present. In one case, chalcedony has been observed. Microcrystalline quartz fills chambers. | | | | | |
| | | | | | | | | | Abundant "ghosts" of radiolarians and some probably benthonic foraminifer occurs. Dark cherts contain irregular patches of probably argillaceous material. Matrix is mainly microcrystalline cristobalite. Lighter areas contain some scattered barite crystal | | | | | |



¹For explanation of symbols, see Chapter 1



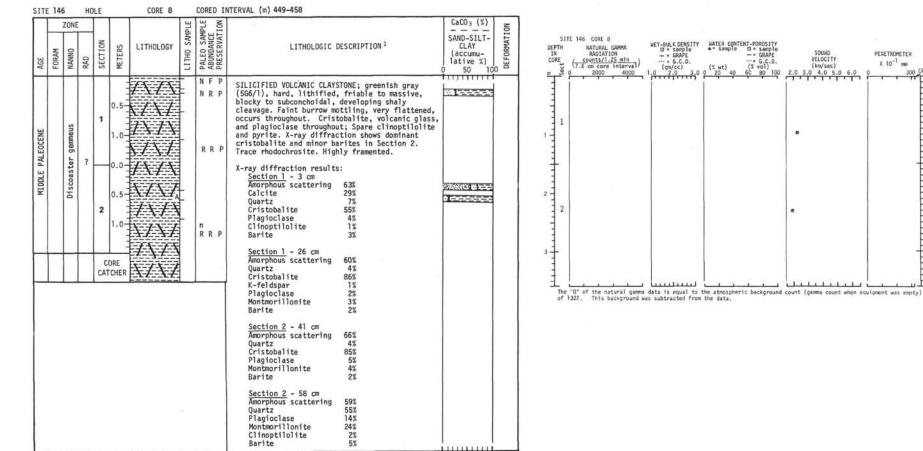


For explanation of symbols, see Chapter 1

| AUC FORAM NANNO RAD SECTION METERS METERS ADOIOHLITI ADOIOHLITI FREERVATION | FURAMINIFERAL NANNOPLANKTON | CaCO ₃ (%) SAND-SILT- CLAY (accumu- 1 ative %) 0 50 100 | SITE 146 CORE 7 WET-BULK DENSITY MATER CONTENT-POROSITY DEPTH NATURAL GAMMA MT-BULK DENSITY MATER CONTENT-POROSITY IN RADIATION |
|--|---|---|---|
| | FURAMINIFERAL NANNOPLANKTON | 11111111111 | m 3 0 2000 4000 1.0 2.0 3.0 0 20 40 6C 80 100 2.0 3.0 4.0 5.0 6.0 0 34 |
| | CHALK; very light gray (N8), indurated, with a sharp basal contact on volcanic clay. Alternating zones of silicified VOLCANIC CLAYSTONE and NANNO- PLANKTON MARL varying from greenish gray (5077/1) to light brownish gray (5077/1) to light burrow mottling occurs at scattered levels. Cristobalite throughout. In general, sparse plagioclase, clinoptilolite, and quartz are present and in Section 1 (60 cm), silicified foraminfers are common. X-ray data show dominant cristobalite. X-ray diffraction results at 145-146 cm. Amorphous scattering 62% Quartz 4% Cristobalite 92% Plagioclase 3% Montmorillonite 1% In Section 1 (145-147 cm) SPICULTIC-RADIOLARIAN CHERT, slightly calcareous, with abundant sponge spicules and well-preserved radiolarians. The chambers are empty or filled by cristobalite. Bioturbated. Detrial quartz. SILICEOUS VOLCANIC CLAYSTONE greenish gray (567/1). | | The "0" of the natural gama data is equal to the atospheric background count (gama count when equipment was emport 1327. This background was subtracted from the data. |

¹For explanation of symbols, see Chapter 1

SITE 146/149



¹For explanation of symbols, see Chapter 1

PENETROMETER

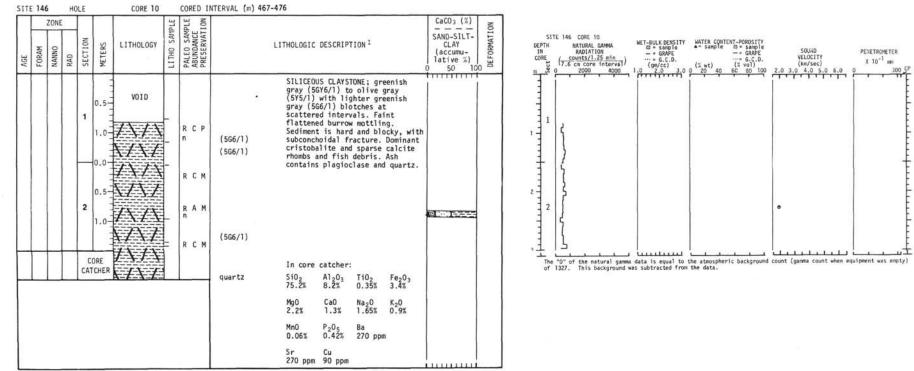
x 10⁻¹ mm

300 CP

| ZONE | | - | ANDIE | PLE | NOI | | CaCO3 (3) 8 | |
|-----------------------|---------|--|------------|-------------|---------------------------|---|--|---|
| FORAM NANNO RAD | SECTION | RTERS CIT | OLOGY OFFI | PALEO SAM | ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ² | SAND-SILT- CLAY (accumu- lative) 0 50 100 | SITE 146 COME 9 MET-BULK DENSITY MATER_CONTENT-PORDSITY DEPTH ANIDAL CAMMA EF sample Sample Sample DEPTH ANIDAL (Sample) EF sample Sample Sample DEPTH ANIDAL (Sample) EF sample SOLAD PEXETROM COME Counts/(JSmin) |
| 2. | | 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0 | | n R R | FM RP CP 0 | 0.5 cm greenish gray (5676/1), blocky, hard, subconchoidal fracture. Sediment oval lense pale greenish gray (5867/2) Scattered cristobalite. X-ray data show K-feldspar. | | The "0" of the natural game data is equal to the atmospheric background count (game count when equipment was of 1327. This background was subtracted from the data. |

- 1

·For explanation of symbols, see Chapter 1



For explanation of symbols, see Chapter 1

| SITE 146 | 6 | HOLE | | CORE 11 | | | NTERVAL (m) 476-485 | | | |
|---|------|------|----------------------------------|-----------|--------------|--|---|---|---------|---|
| AGE FORAM | ZONE | RAD | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | DEFORMA | SITE 146 CORE 11 MET-BULK DENSITY MATER CONTENT-POROSITY DEPTH NATURAL GAMMA DET-BULK DENSITY MATER CONTENT-POROSITY IN RADIATION DESTIN Sample Sample IN RADIATION DESTIN Sample Sample CORE Counts/I/LSP min_ |
| 7 LATE CRETACEOUS EARLY PALEOCENE Gr. bseudobulloides or 7 Globogerina eugubina | | | 0.5- 1.0 0.5 1.0 0.5 | | | P DADAXIN P F DADAXIN R R CEERSTRACK R R | SILICEOUS CLAYSTONE; grayish brown (10YR5/2) to moderate brown (SYR5/4) with slightly lighter blotches due to flattened burrow mottling. Chert stringer appears at bottom. Sediment hard and broken, with cristobalite. Stringer Chert quartz, biotite, plagicalse, barite, and fish debris. X-ray shows polygorskite. INTERBEDDED NANNOPLANKTON MARL AND CLAY; pinkish gray (5Y6/2) Nanno Marl Horizontal burrow mottling associated with color changes, (10YR6/2) light greenish gray (5G8/1) to (5YR8/1) pinkish gray (5Y68/1), occurs throughout. The clays are characteristically the darker zones and the marls the lighter - gradational contacts. X-ray shows K-feldspar and polygorskite Sediment well consolidated and very broken. Upper boundary sharp, with a marlstone separated by a thin claystone layer. X-ray diffraction results: <u>Section 1</u> 63-64cm 113-114cm Amorphous scattering 62% 55% Quartz 8% 69% Cristobalite 76% Plagioclase 6% Palygorskite 12% Montmorillonite 6% 6% Clinoptiolite 2% 9% Barite 2% 5% In thin section the banding is seen to be the result of arg11- laceous and pyrite-rich bio- micrite alternating with sparse highly calcareous biomicrite. Matrix partly silicified. NANNOPLANKTON CHALK (5YR8/1) and RADIOLARIAN CHERT; up to 150µ plagioclase. | | | In the "O" of the natural game data is equal to the atmospheric background count (game count when equipment was empty) of 1387. This background was subtracted from the data. |

¹For explanation of symbols, see Chapter 1

SITE 146/149

| 2 |
|-----|
| F |
| [1] |
| 4 |
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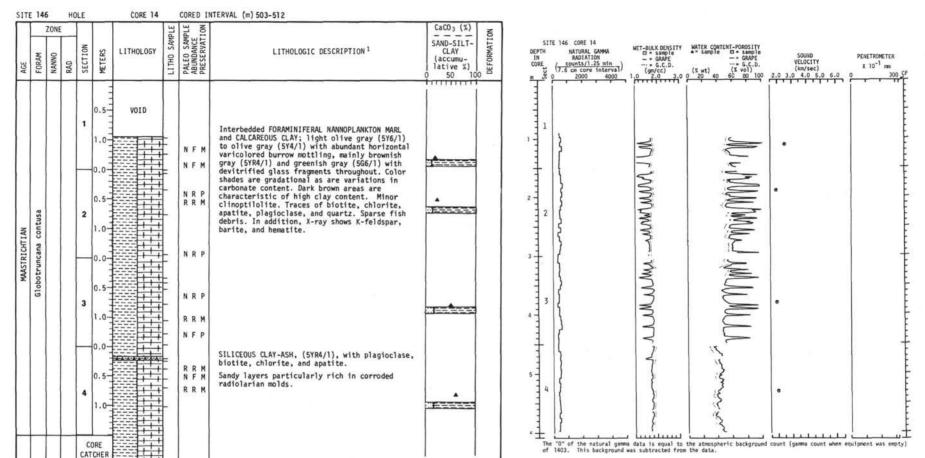
| | _ | ZONE | | | | | SAMPLE | I ON | | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) |
|-------------------|-------|-------|-----|---------|---------------------------|-----------|----------|---|---|--|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SA | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 |
| ? LATE CRETACEOUS | | | | C | 0.5 1.0 DRE CHER | VOID | | RRP | A few pieces of NANNOPLANKTON MARL AND CHALK light brownish gray (5YR6/1), hard and compacted. Lighter colored mottling due to fine burrowing occurs throughout. Some radiolarians and foraminifers, more conspicuous in the upper pieces. Minor biotite, quartz, and carbonate pellets. X-ray diffraction results, 126-127 cm Amorphous scattering 46% Calcite 72% Quartz 26% K-feldspar 2% | 1 |

¹For explanation of symbols, see Chapter 1

| ZONE | 1 I ON | CaCO ₃ (#) 5 |
|---|--|---|
| FORAM NANNO RAD | CALL COLOR CONTRACT | SITE 146 CORE 13 MET-BULK DENSITY MATER CONTENT-POROSITY CLAY DEPTH INTURAL CAMMA MET-BULK DENSITY MATER CONTENT-POROSITY CLAY DEPTH INTURAL CAMMA MET-BULK DENSITY MATER CONTENT-POROSITY IN RADDIATION |
| Abathomphalus mayaroensis Lithraphidites quadratus | 1 0.5 WOID 1 0.0 R R M 1.0 R R M apatite VOLCANIC CLAY about 0.5 cm thick, with sharp basal contact on marl. 0.0 R R M N R P (5YR6/1) NANNOPLANKTON MARL AND CAL-CAREOUS CLAY; predominantly light brownish gray (SYR8/1) and brownish gray (SYR8/1) respectively. The darker zones are which are gradational with lighter marl zones. Abundant burrow mottling in various shades of brown to olive gray (STS/1), pale blue green (SBG7/2), and bluish gray (STS/1), pale blue green (SBG7/1) 0.0 N R M (SB7/1) X-ray diffraction results: 3 N F M (SB7/1) X-ray diffraction results: 0.0 N F M (SB7/1) X-ray diffraction results: 3 N F M Section 3 Section 5 4 N C M Sattering S1% 66% 44% Calcite 72% 14% 77% 0.0 N R R N R M Amorphous Scattering S1% 66% 2% 4 N C M Section 3 Section 5 Statering S1% 66% 2% 0.5 N R P Common chlorite Abundant to dominant quartz and montmorillonite 5% 66% 2% 5 N C M | A B A A A A B A A A B A A A A B A A A A B A A A B A A A A B A A A A B A A |

·For explanation of symbols, see Chapter 1

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¹For explanation of symbols, see Chapter 1

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| LODAL | LUKAM NANNO | Γ | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | SITE 146 CORE 15 DEPTH NATURAL GAMMA IN RADIATION - ERAPE CORE (<u>counts/1.25 min</u>) |
|----------------------|-----------------------|---|---------|--------------|-----------|---------------|---|---|---|--|
| Clobotanoo anoninter | biopotrancana concusa | | 1 | 0.51 | VOID | | NPM | NANNOPLANKTON MARL; predominantly light olive gray (5Y6/1) to light gray (N7), with gradationally intercalated zones from pinkish gray (5Y8/1) to olive gray (5Y5/1). The darker zones typically have higher clay con- tent. Burrow mottling, mostly horizontal, common throughout; usually darker than the sur- rounding sediment. Mottling is variable in shape and size; horizontal chevron-like burrows appear in bottom of Section 6. Interlayering of darker zones fades away toward lower sections. | | |
| | | | | 0.0 | | | R R M R R M N F P | | 1 | |
| | | | | 0.5 | | - | NFM | | | |
| | | | 3 | 1.0- | | | RRM | | | |
| | | | Ľ | 0.0 | | I I I | NRP | | | |
| hand | gansseri | | | 0.5 | | | NRP | | i nic'uluicici. | |
| | | | 4 | 1.0 | | 4 14 14 14 14 | NRP | | | |
| "abota | a lobotruncana | | | 0.0 | | H | R C P/M | | | |
| | | | 5 | 0.5 | | | NFM | Minor quartz, biotite, plagio- | | |
| | | | | 1.0 | | | | clase and apatite. Garnet (?spessartite) in Section 4, 10 cm. Sparse fish debris. | | |
| | | | | 0.0 | | 4 14 14 14 | N F M | In addition, X-ray shows K-feldspar, barite, carbonate rhombs. | 6 | |
| | | | 6 | 0.5 | | Talala | | Horizontal chevron-like | | |
| | | | | 1.0 | | HHH | N C M R R P | burrows. Fossiliferous micrite with rare benchonic foraminifers and rad- iolarians and abundant volcanic fragments. | | |
| | | | | ORE TCHER | | | | rraymencs. | | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was of 1403. This background was subtracted from the data. |

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SITE 146

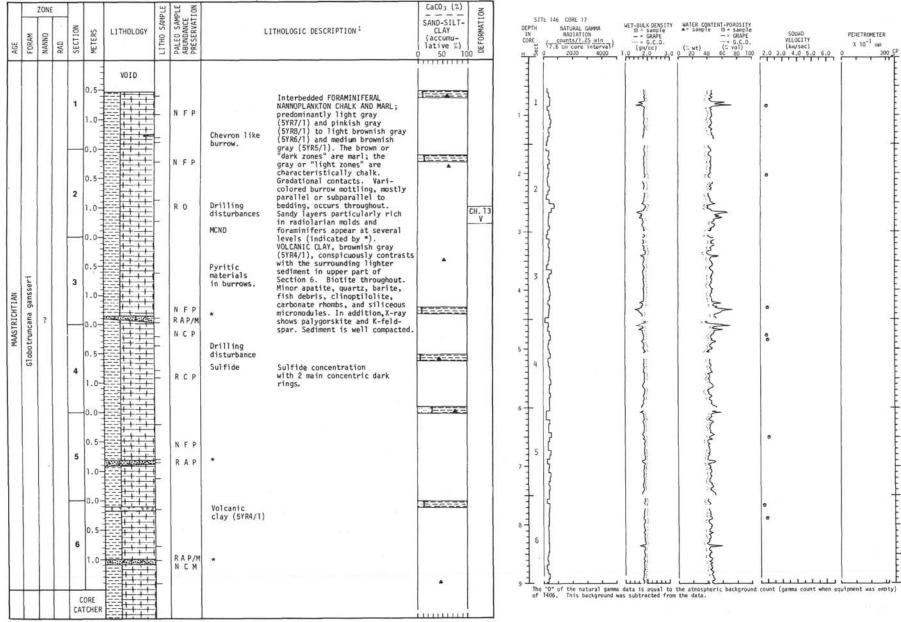
HOLE

CORE 15 CORED INTERVAL (m) 512-521

| ZONE | | | | SAMPLE | ION | | CaCO 3 | (%) NO | | | | |
|-----------------------|---------|--------|-----------|-----------|---|--|---|------------------------|--|---|--|---|
| FORAM NANNO RAD | SECTION | METERS | LITHOLOGY | LITHO SAM | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-S CLA (accu lative 0 50 | mu- e %) DEFORM/ | SITE 146 CORE 16 DEPTH NATURAL GAMAA IN RADIATION CORE <u>counts/1.25 min</u> (7.6 cm core interval) | WET-BULK DENSITY D = sample - = GRAPE = G.C.D. (gm/cc) 1.0 2.0 3.0 | WATER CONTENT-POROSITY * sample B = sample | 500WD VELOCITY (km/sec) 2.0 3.0 4.0 5.0 6. |
| | | 0.5 | VOID | Taria - | | MCND Interbedded FORAMINI NANNOPLANKTON CHALK | AND MARL. | | | }. | 12 | |
| | 1 | 1.0 | | | N F P R C P/M | very light gray (N7) olive gray (5Y6/1) a brownish gray (5Y86/ tively. Darker zones characteristically m | are | * EE1 | , = 1 | 2 | And a | œ |
| | - | 0.0 | | | NFP | apatite gradational with the (5YR3/1) chalky zones. Sparse MCND plagioclase and biot Widespread fish debr | lighter quartz, ite. | . | | } | | |
| | 2 | 0.5 | | | NCP | and clinoptilite. Min bonate rhombs. In ad X-ray shows K-feldsp Several volcanic clay occur. They vary in | nor car- dition, ar and barite. y layers | | 2 2 | | Sec. 1 | |
| | - | 0.0 | | | NCP | and color from dark gray (5YR3/1) to med gray (586/1) olive g and darker shades. motilion shundart th | brownish ium bluish ray (5Y5/1) Burrow | | *† } *† } | | M | 6 |
| gansserf | 3 | 0.5 | | | RRP | Volcanic particularly conspic Clay nearby volcanic clay (5B6/1) darker sediments are burrows. Sandy zone radiolarian molds occ | uous in and layers when filling in rich in | • | 3 | | 12 | |
| | - | 0.0 | | | NFP | lower part of Section volcanic clay at 80 Section 4 has common feldspar, minor clin | n 1. The cm in alkaline optilolite, | | | States and | | |
| 61 obo truncana | 4 | 0.5- | | | N C P R R P | subspherical quartz, biotite, and zircon. 7 cm thick X-ray diffraction re- volcanic | sults. | | s | | 3.5 | |
| | | 1.0- | | | NEP | clay Section (5Y5/1) 123.2- 123.3 c Amorphous Scattering 49% | 37.5- | | | | in the second se | |
| | | 0.0- | | | | Calcite 11% Calcite 11% Quartz 1% K-feldspar 3% Montmorillonite 85% | 2% 2% 97% | | | 7 | 3 | |
| | 5 | 1.0- | | | NFP | MCND=micrónodules | | | 7 5 5 7 4 5 | | | • |
| | - | 0.0 | | | R R M N F P | Volcanic clay (5YR2/1) | | | | \mathbf{n} | ~ | |
| | 6 | 0.5 | | | RO | Volcanic clays (5GY6/1) and (N5) | | 191 | 8 T 6 5 | ~ ~ | W NU | |
| | | 1.0- | | | N F P R O N F P | | | | , <u>1</u>]} | 13 | he atmospheric background the data. | ° |

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¹For explanation of symbols, see Chapter 1



CORED INTERVAL (m) 530-539

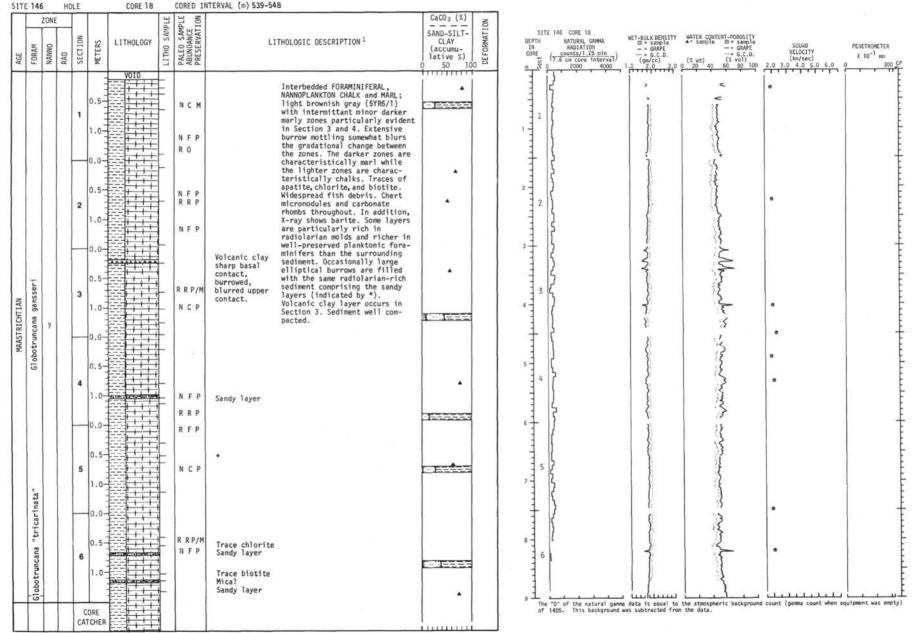
SITE 146/149

SITE 146

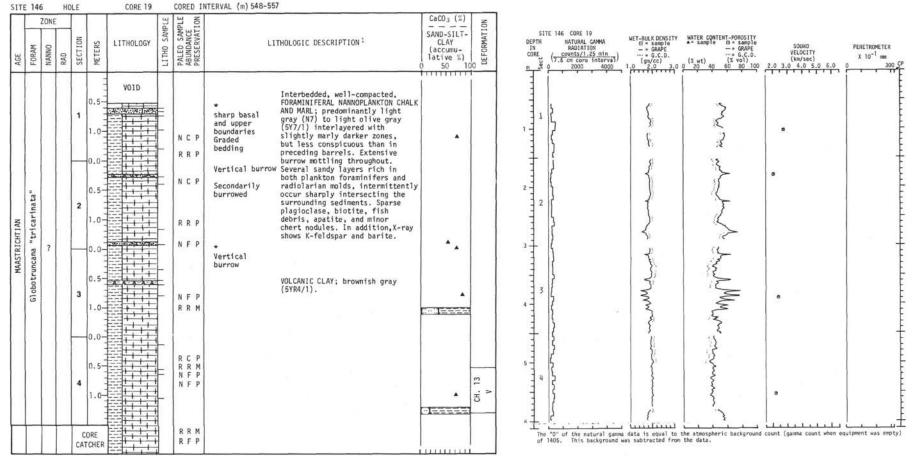
HOLE

CORE 17

| S | |
|----|--|
| T | |
| 1 | |
| 14 | |
| 6 | |
| 14 | |
| 9 | |



For explanation of symbols, see Chapter 1



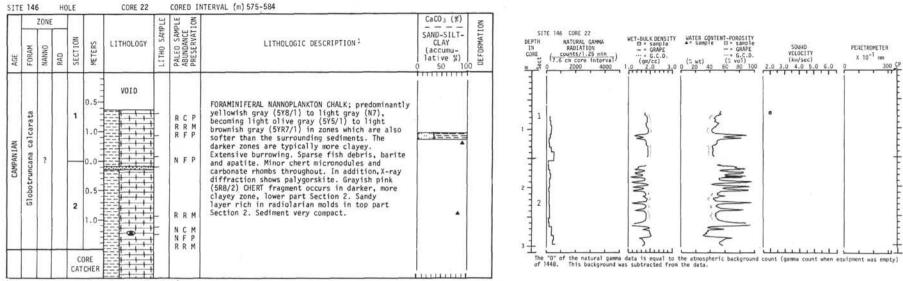
| | | | | | | NTERVAL (m) 557-566 | | _ | | | | | |
|-------------|---------|---------------------------|---|---|---|---|--|---|--|--|---|--|--|
| ZONE RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ¹ | | | 5 17.6 cm core interva | - = GRAPE = G.C.D. (gm/cc) | = GRAPE | SOUHD VELOCITY (km/sec) 2.0 3.0 4.0 5.0 5.0 | РЕНЕТКОМЕТЕЯ X 10 ⁻¹ mm 0 300 СР |
| ? | | 1.0 0.5 1.0 | | | N C P R C P R R P R R P R R P/M N F P R R M | FORAMINIFERAL NANNOPLANKTON CHALK; light gray (N7) to very light gray (N8) with minor inter- bedded light brownish gray (SYR&/1) zones of greater clay content. Extensive burrow motiling throughout. Very compact, but crumbly especially in light gray zones. Trace amounts of biotite, apatite, barite, plagioclase, mica, quartz, and widespread fish debris. Chert micronodules and carbonate rhombs throughout. Chert fragment grayish orange (SYR7/2) with light brown (SYR&/4) nucleus is at top section 1.Sandy layer with predominance of foraminifers and with graded bedding.Zone with abnormally high gamma count contains mainly devirtified volcanic glass with few crystals of plagioclase and biotite. | | | 2 2 2 2 2 3 4 5 5 6 7 7 7 7 7 7 7 7 7 7 7 7 7 | TWW W | the atmospheric background | • | |
| | | 1 NANNO 1 SECTION 3 | NO NUMUN SEBILIZE NVWN 02 0.5 1 0.5 1 0.0 0.5 1 1.0 0.0 0.5 1 0.5 1 0.5 1 1.0 0.0 0.5 0.5 | NO S2 LITHOLOGY 000 0.5 0.5 1 0.5 0.5 0.0 0.5 0.5 1 0.0 0.5 0.5 0.5 0.5 | NHMN SS 32 LITHOLOGY OHLIT 02 0.5 <td< td=""><td>N C P 1 0.5 0.5 N F 0.5 0.5 0.5 N F 1 0.0 0.5 N F 0.5 0.5 0.5 N F 0.0 0.5 0.5 N F 0.0 0.5 0.5 N F 0.0 0.5 0.5 N F 0.1 0.5 0.5 N F 1.0 0.5 0.5 N F 1.0 0.5 0.5 N F 1.0 0.5 0.5 N N 1.0 0.5 0.5 N N 1.0 0.5</td><td><pre>Proceedings of the second second</pre></td><td>0.5 N C P 1 N C P 1 N C P 1 P 1.0 R C P 0.0 R R P 0.5 R R M 0.5</td><td>1 Image: Constraint of the second second</td><td>7 2 CORE N F P 3 0.5 N F P 3 0.5 Sandy layer with predominance of foraminifers and with grade dedding. 7 2 N F P 8 R R P/M 0.5 N F P 8 R R P/M 0.5 N F P 0.5 N F P 0.5 R R P/M 0.5 N F P 0.5 R R P/M 0.5 R R M 0.5 <</td><td>7 2 0.5 0.6 0.7 1 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0</td></td<> <td>1 0.5- 1 N C P greater clark (NNNOPLANKTON CHALK; light gray (N7) to very light gray (N8) with minor inter- bedded light brownish gray (SYR6/1) zones of greater clark content. Extensive burrow motiling throughout. Very compact, but crumbly especially in the protocole Trace anomes of biotite. 1 0.5- 2 R C P 0.0- 0.0- 0.0- 0.0- 0.0- 0.0- 0.0- 0.0</td> <td>7 2 0.5 N F P 0.0 0.5 N F P 0.5 N F P Sady layer with gredomining rays (NB) with light from and with greated bidting. 0.5 N F P Sady layer with gredomining rays (NB) man count contains mainly devirtified volcanic glass with few crystals of plagioclase and biotite. 0.5 N F P R R P/R 0.5 N F P R R P/R 0.5 N F P Sady layer with gredomining rays (NB) disting and with gredomining rays and with rew crystals of plagioclase and biotite. 0.0 0.0 N F P 0.0 R R P R R P 0.0 R R R R R R 0.0 R R R R R 0.0 R R R R R 0.0 R R R R R</td> | N C P 1 0.5 0.5 N F 0.5 0.5 0.5 N F 1 0.0 0.5 N F 0.5 0.5 0.5 N F 0.0 0.5 0.5 N F 0.0 0.5 0.5 N F 0.0 0.5 0.5 N F 0.1 0.5 0.5 N F 1.0 0.5 0.5 N F 1.0 0.5 0.5 N F 1.0 0.5 0.5 N N 1.0 0.5 0.5 N N 1.0 0.5 | <pre>Proceedings of the second second</pre> | 0.5 N C P 1 N C P 1 N C P 1 P 1.0 R C P 0.0 R R P 0.5 R R M 0.5 | 1 Image: Constraint of the second | 7 2 CORE N F P 3 0.5 N F P 3 0.5 Sandy layer with predominance of foraminifers and with grade dedding. 7 2 N F P 8 R R P/M 0.5 N F P 8 R R P/M 0.5 N F P 0.5 N F P 0.5 R R P/M 0.5 N F P 0.5 R R P/M 0.5 R R M 0.5 < | 7 2 0.5 0.6 0.7 1 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 0.0 | 1 0.5- 1 N C P greater clark (NNNOPLANKTON CHALK; light gray (N7) to very light gray (N8) with minor inter- bedded light brownish gray (SYR6/1) zones of greater clark content. Extensive burrow motiling throughout. Very compact, but crumbly especially in the protocole Trace anomes of biotite. 1 0.5- 2 R C P 0.0- 0.0- 0.0- 0.0- 0.0- 0.0- 0.0- 0.0 | 7 2 0.5 N F P 0.0 0.5 N F P 0.5 N F P Sady layer with gredomining rays (NB) with light from and with greated bidting. 0.5 N F P Sady layer with gredomining rays (NB) man count contains mainly devirtified volcanic glass with few crystals of plagioclase and biotite. 0.5 N F P R R P/R 0.5 N F P R R P/R 0.5 N F P Sady layer with gredomining rays (NB) disting and with gredomining rays and with rew crystals of plagioclase and biotite. 0.0 0.0 N F P 0.0 R R P R R P 0.0 R R R R R R 0.0 R R R R R 0.0 R R R R R 0.0 R R R R R |

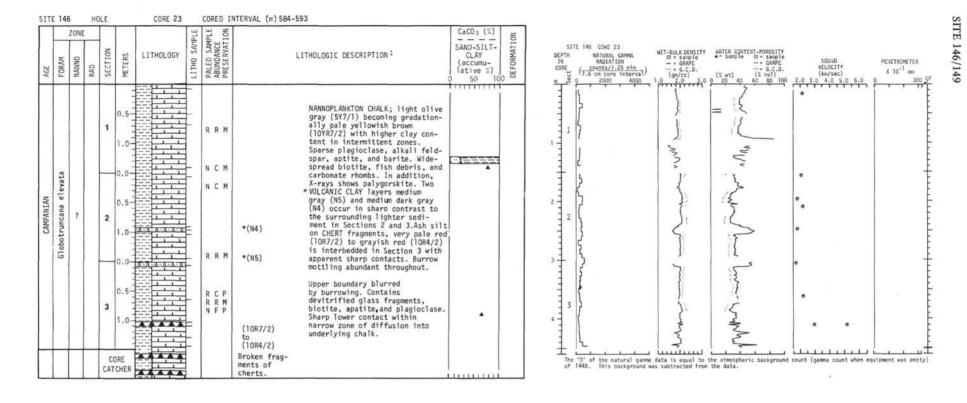
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| | | ZONE | | | | | PLE | I ON | | $\frac{\operatorname{CaCO}_3(\%)}{}$ |
|-----|-------|-------|-----|---------|-------------|-----------|-------------|---|--|---|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPI | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 |
| • | * * | | | |)RE CHER | | | | No recovery. Core catcher consists of FORAMINIFERAL NANNOPLANKTON CHALK, light gray (N7), crumbly and fragmented. | - |

¹For explanation of symbols, see Chapter 1

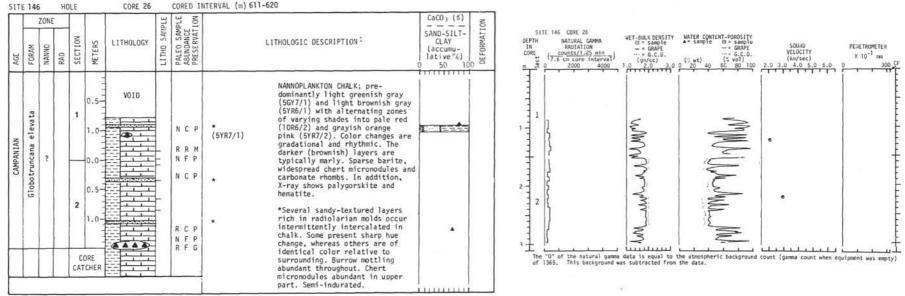
* MAASTRICHTIAN
* * Globotruncana "tricarinata"

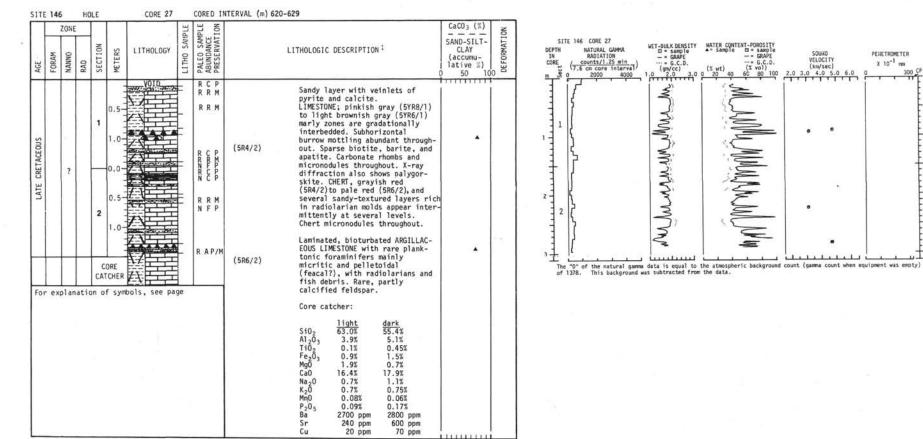




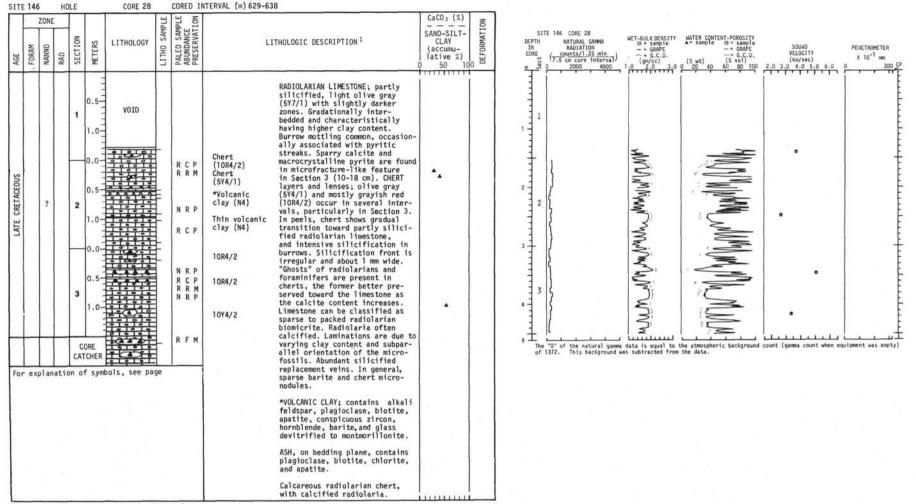
| SITE | Z | ONE | SECTION | TERS | CORE 24 | THO SAMPLE | SAMPLE NUCE VATION | ITERVAL (m) 593- | 602 LITHOLOGIC DESCRIPTION : | CaCO ₃ (%) SAND-SILT- CLAY (accumu- | FORMATION | SITE 146 CUNE 24 DEPTH NATURAL GAMMA BY SAMPLE IN RADIATIONRADE IN RADIATIONRADE SAMPLE SAMPLE IN RADIATIONRADE SOUND PENETROMET | TER |
|---------------|---|-------|---------|-------------------|---------|------------|--|--|--|---|--------------|--|--------|
| CAMPANIAN AGE | + | NANNO | 1 | ₩ 0.5- 1.0- | | | D31A9 N C D W NONUA N C D W NONA N C D W N C D | (5YR4/1) and (5Y4/1) (5Y4/1) (5R6/2) | tively. Lowermost cherts show irregular color banding and may be filling in burrows. Light colored chert is slightly calcareous. Minor quartz, fish debris, barite, chert micro- nødules, and carbonate rhombs throughout. In addition, X-ray diffraction shows K-feldear. | | ⁰ | (km/sec) (1 vt) (1 vt) (km/sec) | empty) |

| | ZON | IE | | | | PLE | APLE FI ON | | | CaCO ₃ (%) | NOI | |
|-----------------|-------|-----|---------|-------------|-----------|-----------|---|---|---|--|----------|--|
| AGE FDRAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAM | PALEO SAMPLE ABUNDANCE PRESERVATION | LI | THOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 10 | | SITE 146 CORE 25 WAT-BULX DENSITY WATER CONTENT-PORDSITY DEPTH NATURAL GAMMA m = sample 4 = sample 1 = sample IN RADIATION |
| EOUS | | | 1 | 0.5 | | | | | Section 1: Very soupy mixture of CHERT cuttings and CHALK, totally disturbed. | | сн. 13 I | |
| LATE CRETACEOUS | ~ | | 2 | 0.0 | | | R C P R R M N F P R R M | (56Y8/1) (56Y8/1) (56Y8/1) Volcanic clay (5YR2/1) (56Y8/1) | NANNOPLANKTON CHALK; predominant- ly light brownish gray (5YR6/1) to pinkish gray (5YR7/1). The darker (brownish) layers are characteristically marly. Minor fish debris, apatite, chert, and barite micronodules. Carbonate rhombs throughout. In addition, X-ray shows K-feldspar. Exten- | | | i with i with i |
| | | | | ORE CHER | 3 | | | | sive burrow mottling throughout, and occasionally associated with contrasting hue change into light greenish gray (56%8/1). Thin lamina of brownish black (5YR2/1) volcanic clay is partly blurred by mottling. Semi- indurated. | | | The '00' of the natural gama data is equal to the atmospheric background count (gamma count when equipment was e of 1365. This background was subtracted from the data. |



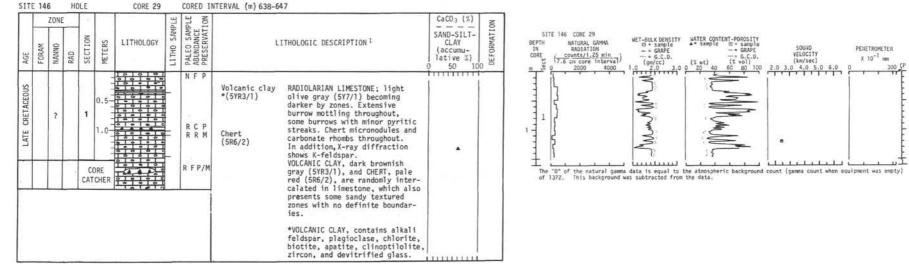


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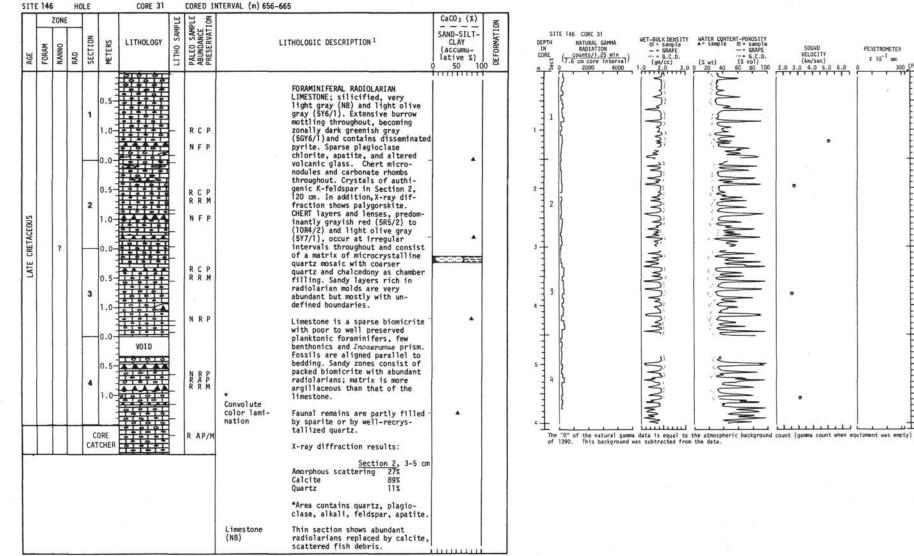
9



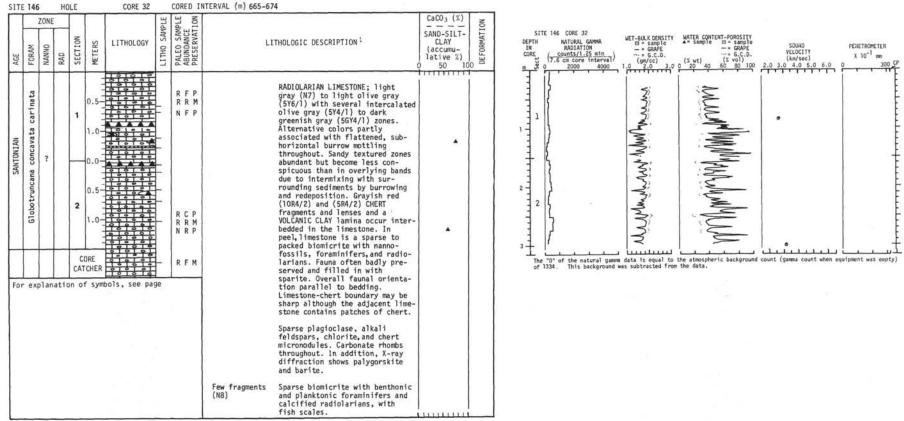
| ITE 146 | | HOLE | | CORE 30 | | | NTERVAL (m) 647-656 | CaCO 1 (%) | T |
|---|------|------|--|---|--------------|---|--|---|--|
| AGE FORAM | ZONE | RAD | NETERC | LITHOLOGY | LITHO SAMPLE | ALEO S BUNDAN RESERV | LITHOLOGIC DESCRIPTION : | SAND-SILT- CLAY (accumu- I Tative %) 0 50 100 | SITE 146 CORE 30 DEPTH NATURAL GAMMA WET-BULK DENSITY MATER CONTENT-POROSITY IN RADIATION |
| SANTONIAN Globotruncanà concavatà carinata | ? | 3 | 1. 0. 1. 0. 0. 0. 1. 0. | • • | | R R M N F P R F M R C P R C P R C P R C P R C P N F P N F P N F P R C R M | <pre>FORAMINIFERAL RADIOLARIAN LIME- STONE: partly silicified, light Chert gray (N7) to light olive gray (5R4/2) of very pale red (10R6/2). Abundant subhorizontal burrow mottling. Chert micronodules and carbonate rhombs throughout. gypsum barie, apatite, and biotite. "Chert In addition, X-ray diffraction (SYR8/1) above part of Section 1. (SYR8/1) above part of Section 1. (SYR8/1) above part of Section 1. (SYR8/1) appear as definite sandy layers whereas others grade into the limestone. Limstone replaced by grayish red (SR4/2) CHERTS and lenses are irregular and surrounded by pinkish gray (SYR8/1) "diffusion" zone. Limestone in acetate peel is fossiliferous micrite to sparse biomicrite, with poorly preserve fauna filled in by sparite. Nearby chert matrix is partly silicified. In the chert, most of the foraminifers and their sparit(c filling are un- silicified. Silicification irregular, following biotur- bation. The radiolarians are Volcanic often calcified. Matrix clay lamina micrite. ragmentation * skeletal fragments. *</pre> | | Monor man Milling . |
| | | | CORE | | IND | RFP | Chert olive gray (5Y4/1) | | The "0" of the natural game data is equal to the atmospheric background count (gamma count when equipment was e of 1367. This background was subtracted from the data. |

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For explanation of symbols, see Chapter 1

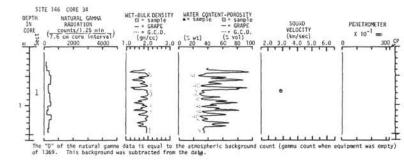


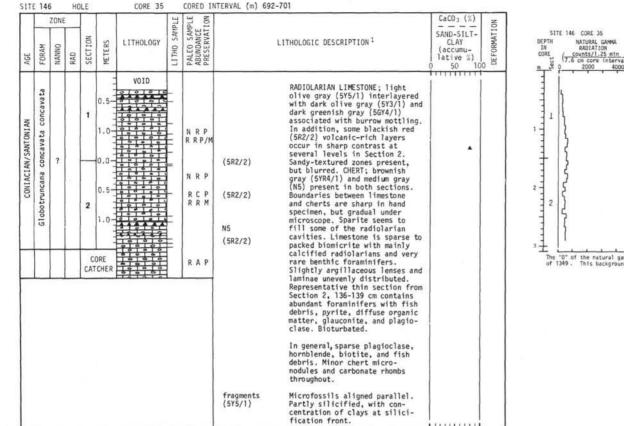
89



| T | ONNE | CELTTON | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | DEFORMA | SITE 146 CORE 33 WET-BULK DENSITY WATER CONTENT-PORDSITY DEPTH NATURAL GAMMA Im sample Im sample Im sample IN RADATION Im sample Im sample Im sample Im sample IN RADATION Im sample Im sample Im sample Sample SOURD PEN CORE Counts/I.25 min_ Im (S.C.D. Im (S.C.D.) VELOCITY VELOCITY X IN Im (S.C.D.) (Smoth) Im (Smoth) < |
|--|------|---------|--------------------|-----------|--------------|---|--|---|---------|---|
| SANTOWIAN Globotruncana concavata carinata | ? | | 0.5 1.0 CORE | | | R R P N R P | RADIOLARIAN LIMESTONE; pre- dominantly light olive gray (577/1) to (576/1) with abundant interlayered dark greenish gray (5674/1) zones, some associated with flattened burrow mottling. calcitic Radiolarian-rich, sandy-textured lenses appear at several levels. Common carbonate rhombs. Trace of chlorite and chert micro- nodules. few fragments (576/1) | | | The "D" of the natural gamma data is equal to the atmospheric background count (gamma count when equipmen of 1342." This background was subtracted from the data. |

| | | ZONE | | | | | SAMPLE | MPLE E | | | $CaCO_3$ (%) |
|---------------------|--------------------------------------|-------|-----|---------|------------|-----------|-----------|---|--|--|---|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAM | PALEO SAMPLE ABUNDANCE PRESERVATION | Ľ | ITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 |
| CONIACIAN/SANTONIAN | Globotruncana concavata concavata | ? | | | 0.5 1.0 | VOID | | N R P R R M | Lithified volcanic clay Slight graded bedding Chlorite Gypsum | RADIOLARIAN LIMESTONE; very light gray (N8) to medium light gray (N6) with laminations in darker shades and dark greenish gray (56Y4/1). CHERT, brownish gray (5YR4/1), is transitional into sandy zones as in previous core. At 51-52: radiolarian limestone fossils calcified. Scattered fish debris. Large barite crystals in burrows. Chemical analysis, Section 1, 51-56 cm. | |
| | | | | | | | | | | Si02 Al203 Ti02 Fe203 Mg0 55.3% 14.9% 0.9% 7.0% 5.9% Ca0 Na20 K20 Mn0 P205 7.4% 2.35% 1.65% 0.02% 0.13% Ba Sr Cu 750ppm 90ppm | |
| | | | | | | | | | No C/C available | Trace chlorite. Trace of gypsum at 103 cm. In addition, X-ray diffraction | |





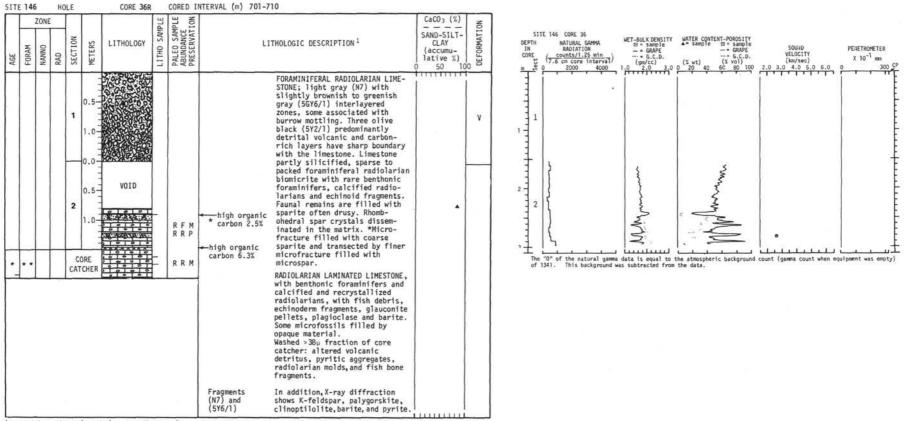
NATURAL GAMMA RADIATION (<u>counts/1.25 min</u> (7.6 cm core interval) WATER CONTENT-POROSITY * sample B = sample - = GRAPE 50020 PENETROMETER VELOCITY (km/sec) 0 4.0 5.0 (% vol) 60 80 10 x 10⁻¹ mm (% wt) 4000 100 2.0 3.0 300 www.www.www.w/www . hunning 1. 11.1 The "O" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1349 . This background was subtracted from the data.

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For explanation of symbols, see Chapter 1

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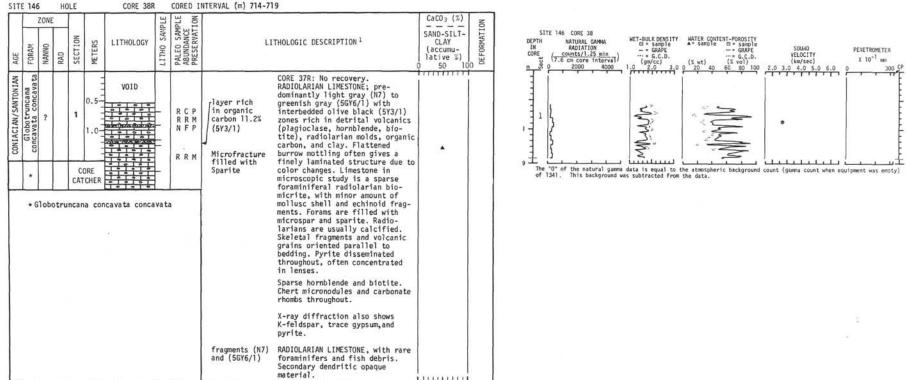
CORE 35 CORED INTERVAL (m) 692-701



Core 37R: No recovery; Cored Interval (m) 710-714.

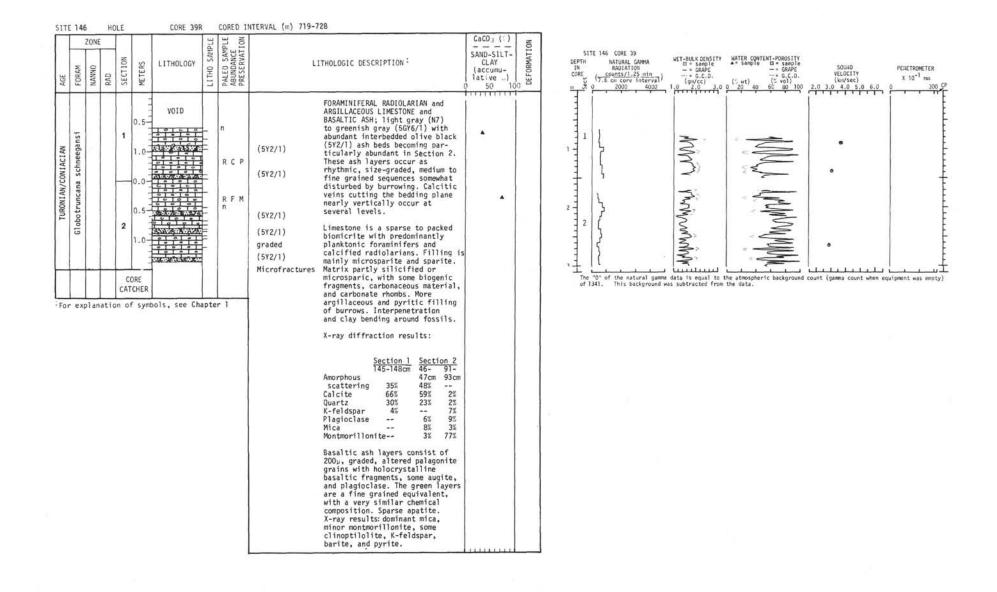
* CONIACIAN/SANTONIAN

* * Globotruncana concavata concavata

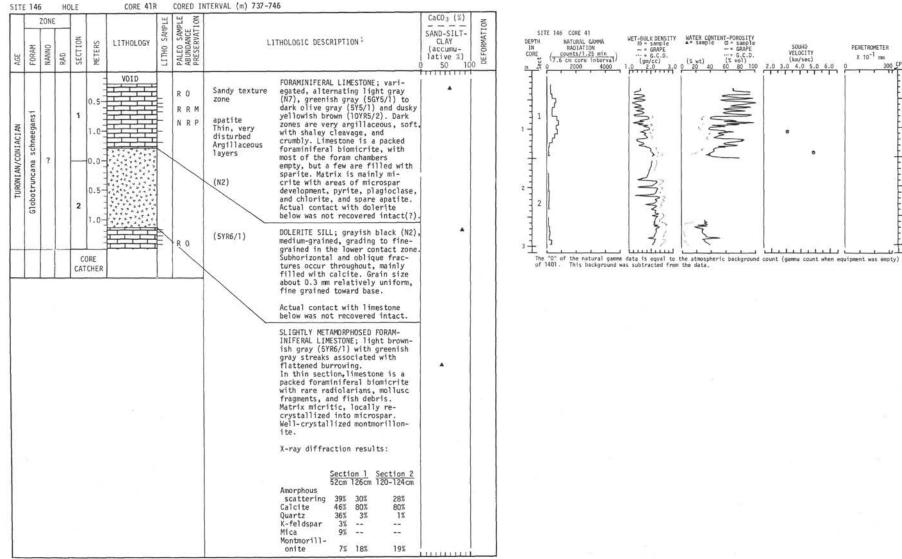


SITE 146/149

¹For explanation of symbols, see Chapter 1



| | | ZONE | | | | | PLE | PLE | | CaCO ₃ (%) | NO |
|-----|-------|-------|-----|-----------|-------------|-----------|--------------|---|--|--|-------------|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LI THO SAMPI | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ¹ | SAND-SILT- CLAY (accumu- lative %) 0 50 10 | DEFORMATION |
| | | | | 1.1.1.2.2 | ORE CHER | | | RRM | LIMESTONE; medium dark gray (N4) with irregular lighter shades of gray to light brownish gray (5YR5/1) with finely sublaminated structure due to flattened burrow mottling. | | |

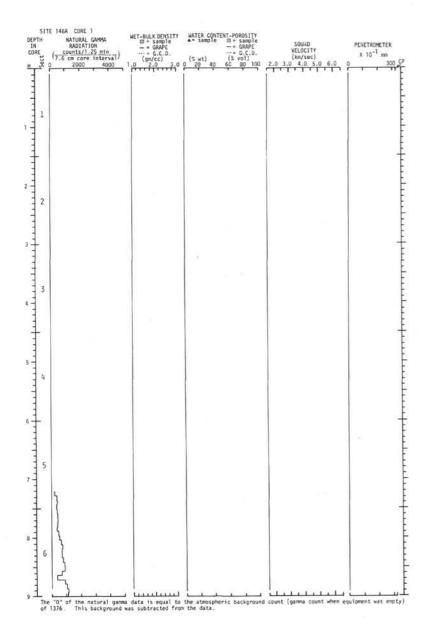


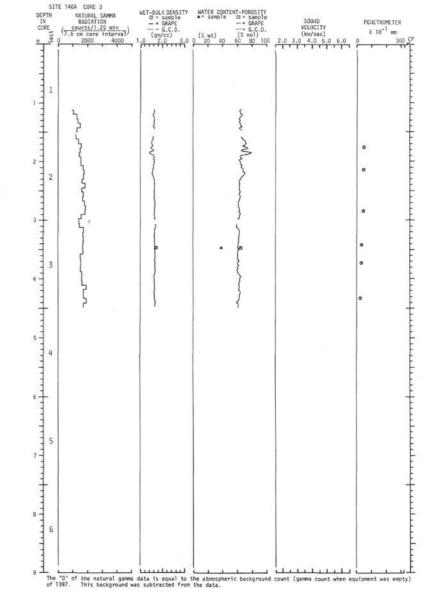
SITE 146/149

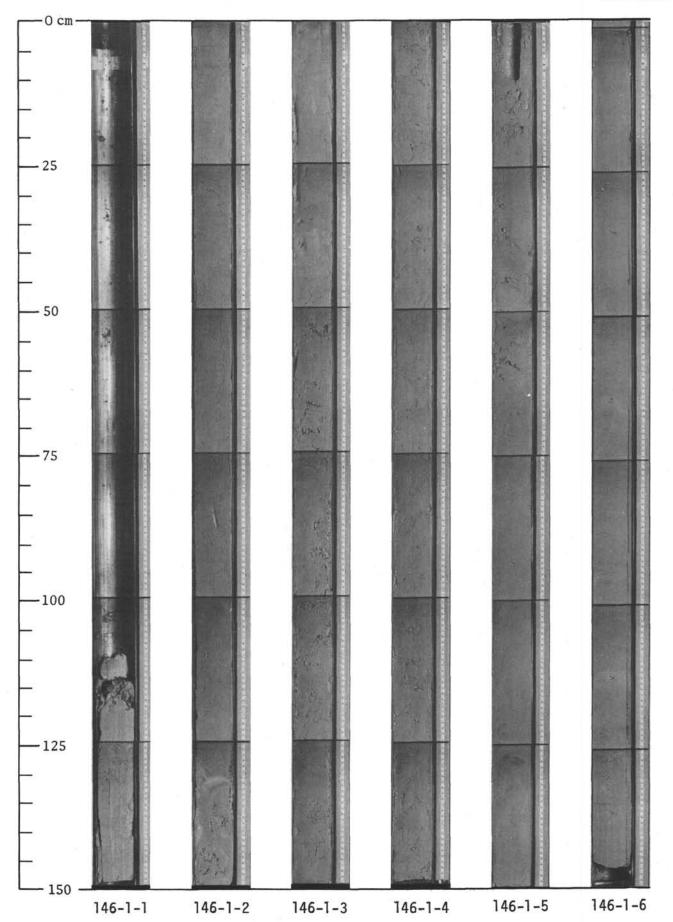
LL

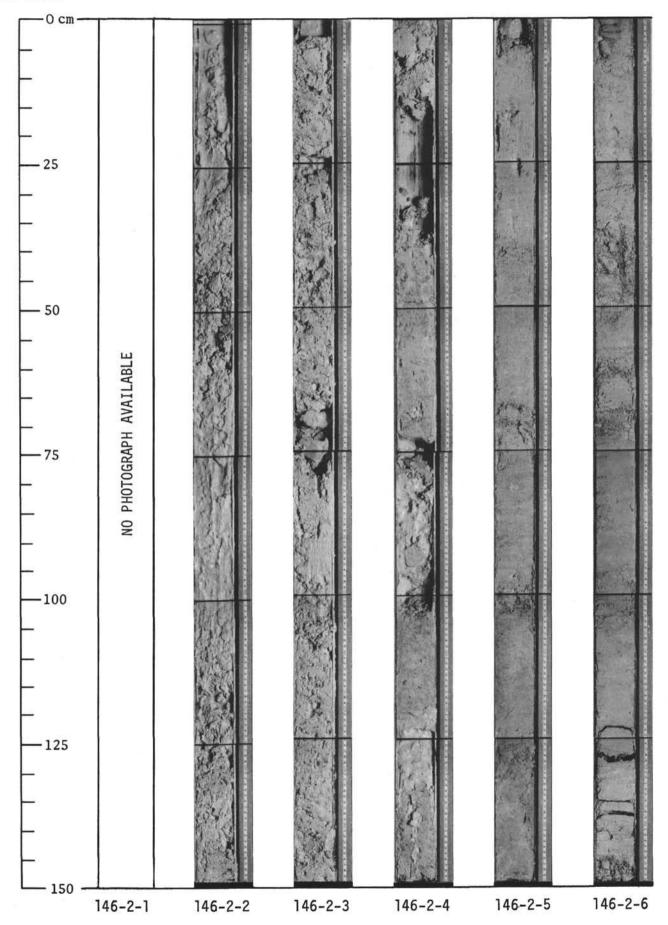
| 1 | ZON | | | | | PLE | PLE | | | CaCO ₃ (%) | NO | |
|-------|-------|-----|---------|------------------------------|-----------|--------------|---|--|--|---|----|---|
| FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | 1 | ITHOLOGIC DESCRIPTION ¹ | SAND-SILT- CLAY (accumu- lative %) 0 50 1 | | SITE 145 CORE 42 WET-BULK DENSITY WATER CONTENT-POROSITY DEPTH NATURAL GAMMA |
| | | | 1 | 0.5- 1.0- 0.5- 1.0- | VOID | | | | DOLERITE; grayish black (N2) with subvertical fractures, some filled with calcitic materials. Conspicuously finer grained (about 0.1 mm) than in barrel 41, very uniform. | | | |
| | | | | 0.5 1.0 | | | | Vertical fracture filled with calcitic material. | DOLERITE | | | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipm of 1366. This background was subtracted from the data. |

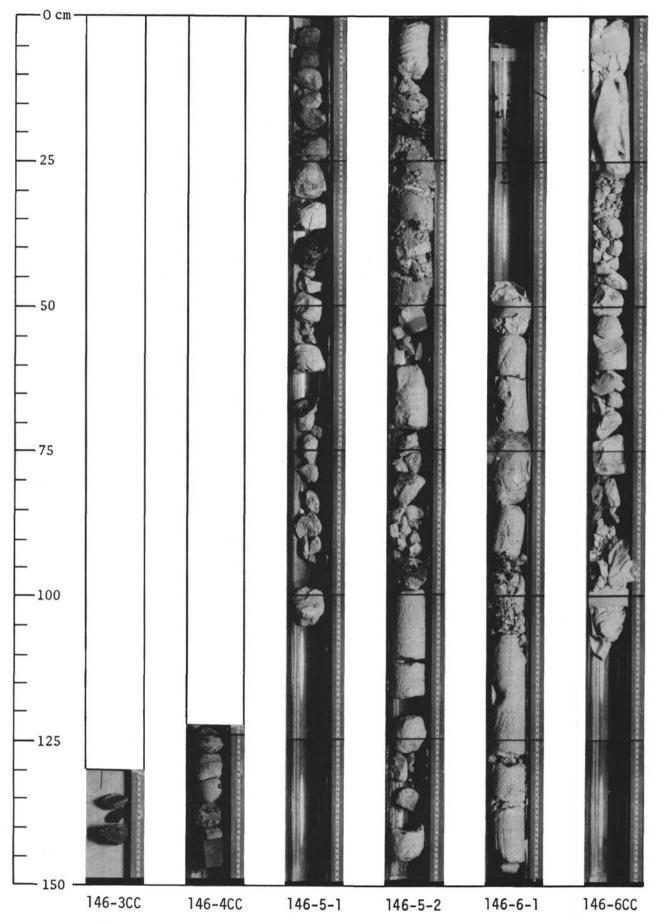
| SITE 14 | 46 | но | LE | | CORE 43 | R | | NTERVAL (m) 755-760 | | |
|--------------|------|----|---|--------|-----------|--------------|---|---|--|--|
| AGE FORAM | ZON | П | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (2) SAND-SILT- CLAY (accumu- lative %) 0 50 100 | SITE 146 CORE 43 DEPTH NATURAL CANMAN, G + SAMPLE CONTENT-POROSITY IN RADIATION G + SAMPLE - SAMPLE SOLID PERETROMETER CORE Counts/1.25 min 6.00 (J. 6. co row finterval) (gr(cc) ((sv1)) (kx/sec) X 10 ⁻¹ mm 6.00 (J. 6. co row finterval) (gr(cc) ((sv1)) (kx/sec) X 10 ⁻¹ mm 6.00 (J. 6. co row finterval) (gr(cc) ((sv1)) (kx/sec) X 10 ⁻¹ mm 6.00 |
| | | | 1 1 2 1 1 2 1 1 0 0 0 0 3 1 1 0 0 0 0 0 4 | | VOID | | | DOLERITE; grayish black (N2) with minor calcitic veining, fine grained, uniform. | | The "0" of the natural game data is equal to the absopheric background court (game count when equippent was empty) |
| | ZONE | | | METERS | CORE 44R | | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ¹ | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative 5) 0 50 100 | |
| | | | COF | | | | | DOLERITE, grayish black (N2). | | |

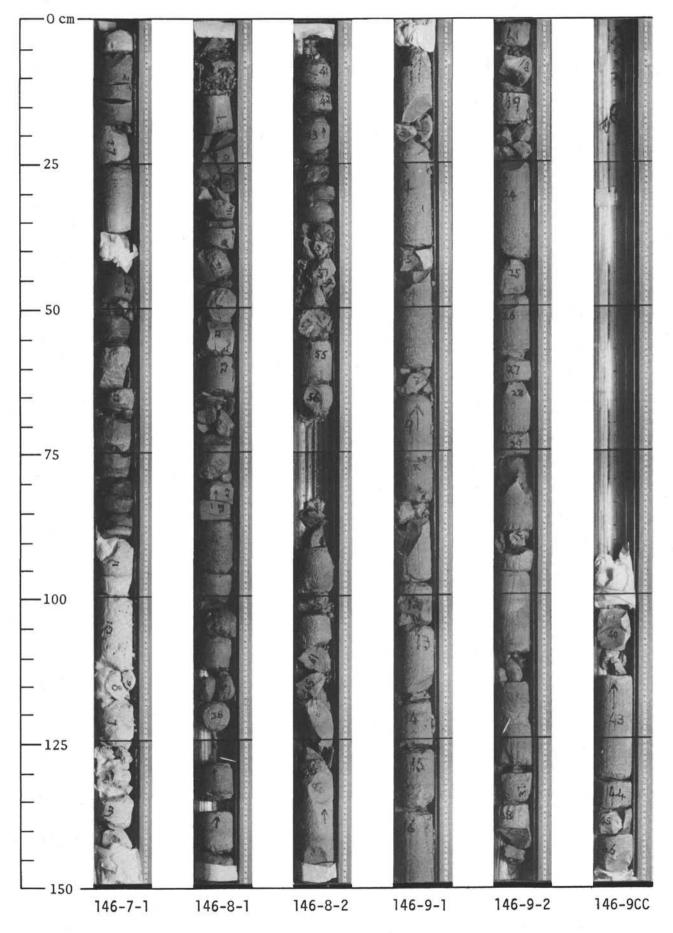


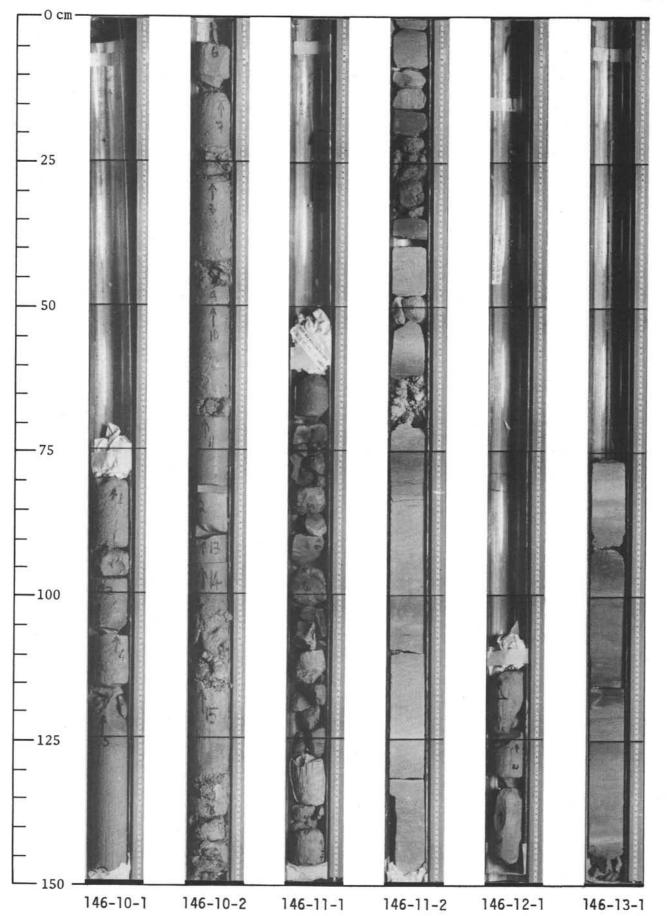


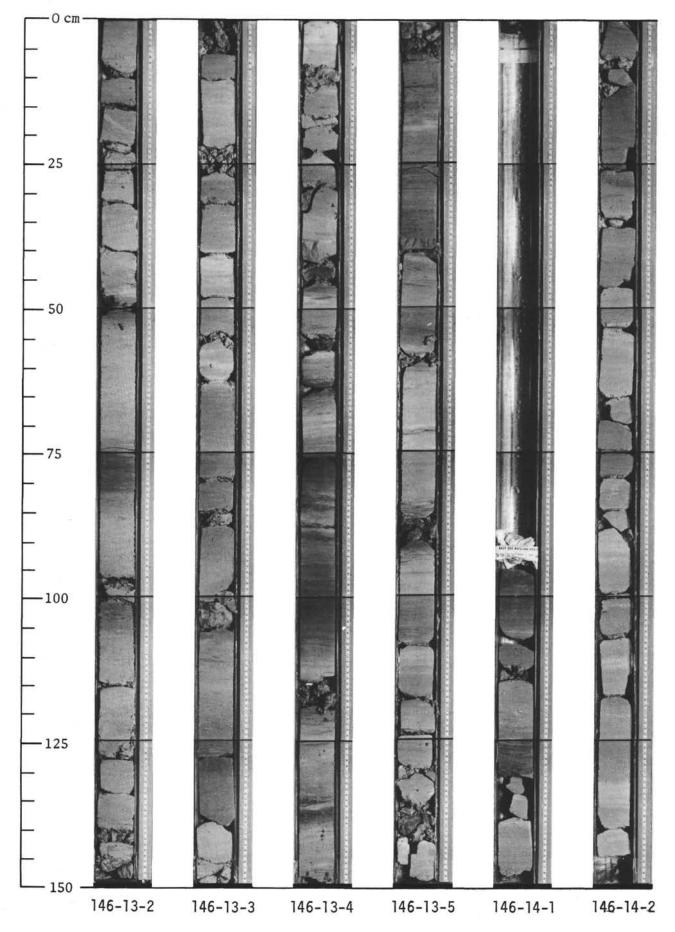




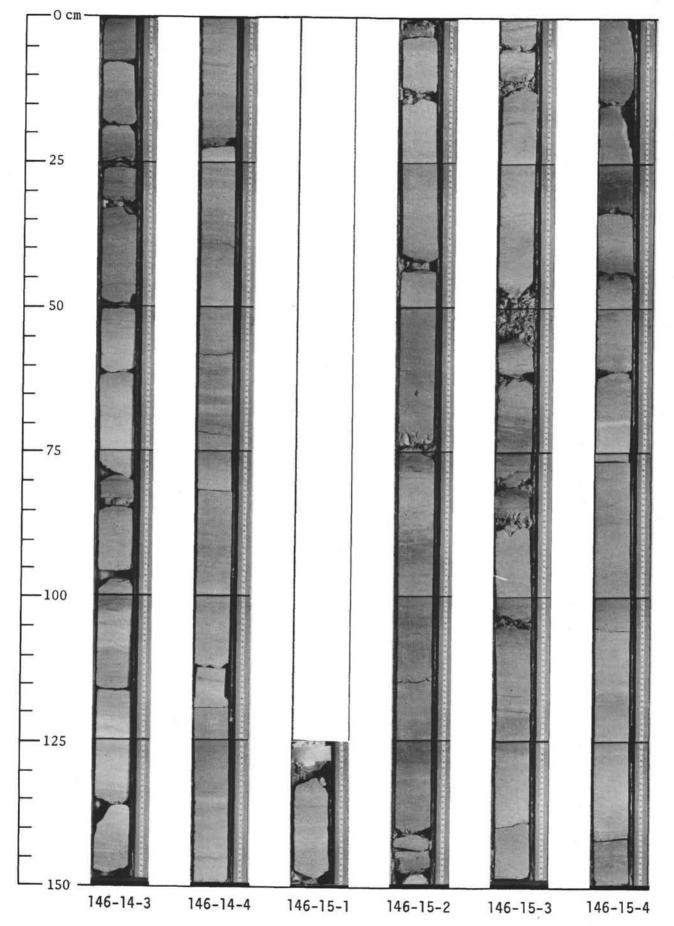


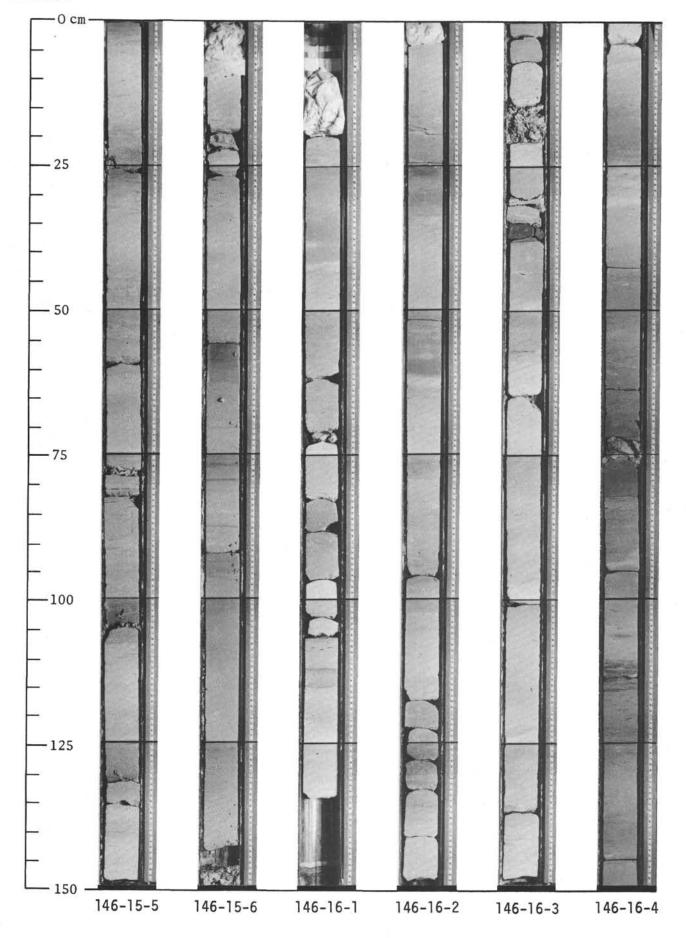


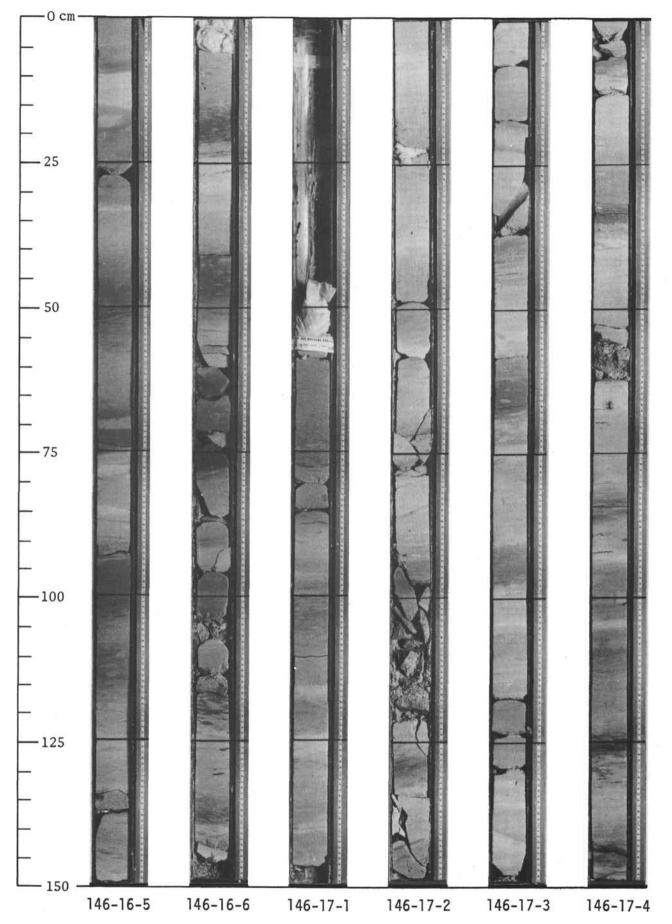


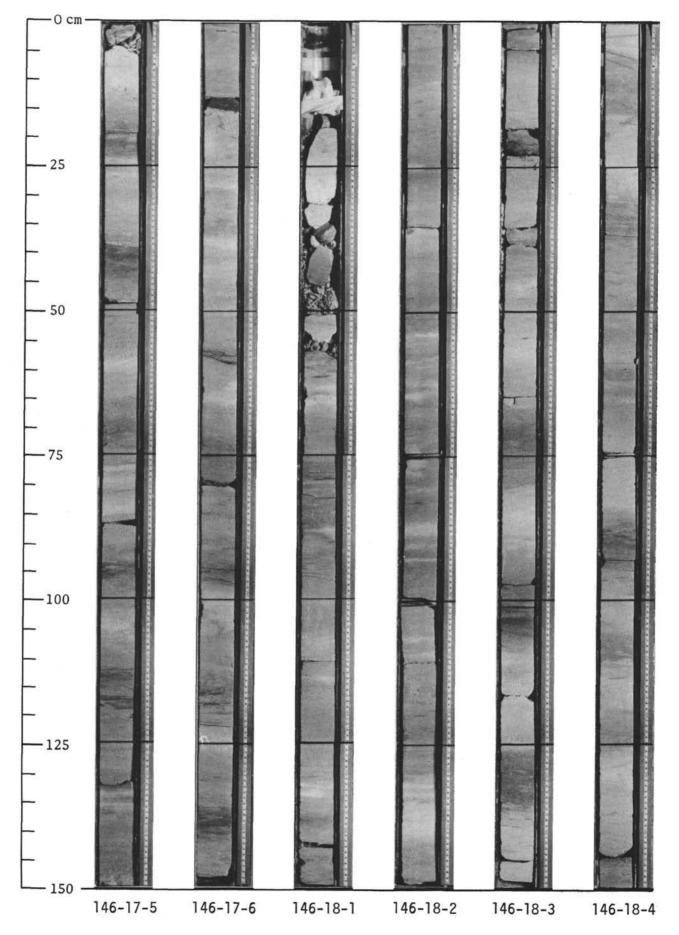


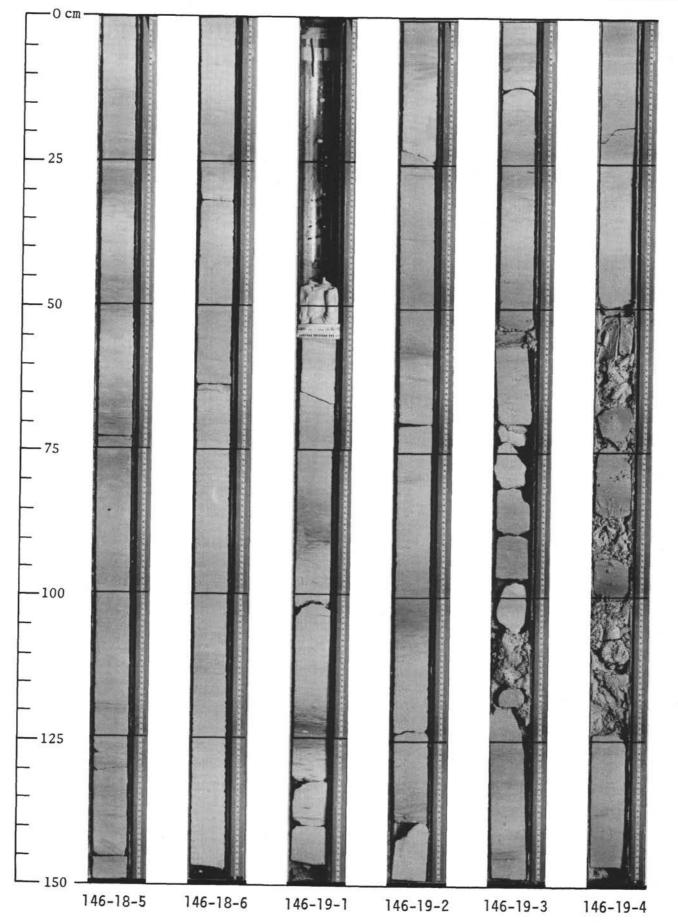
SITE 146/149

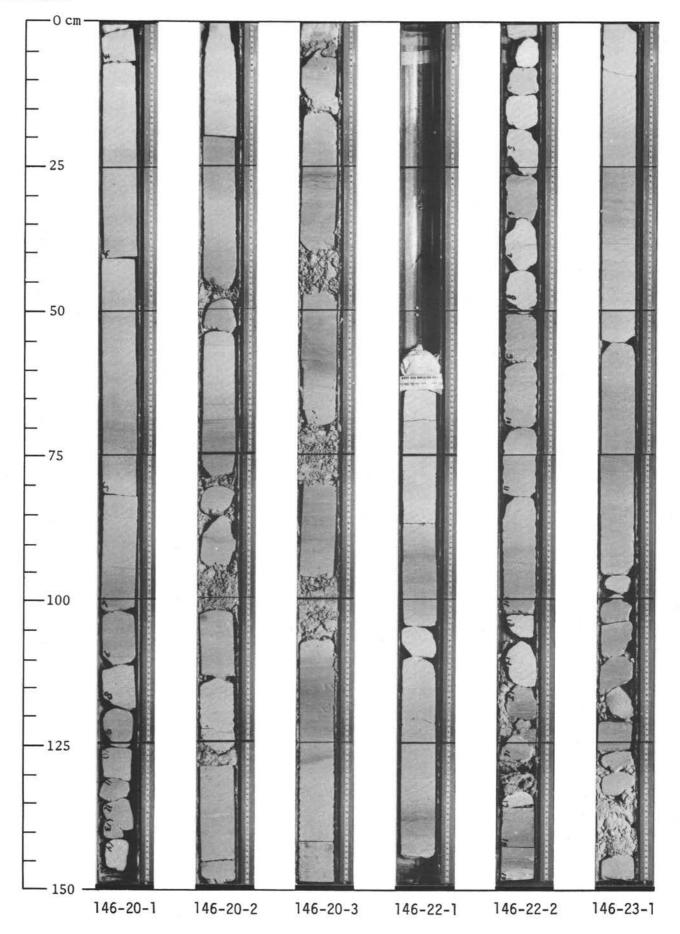




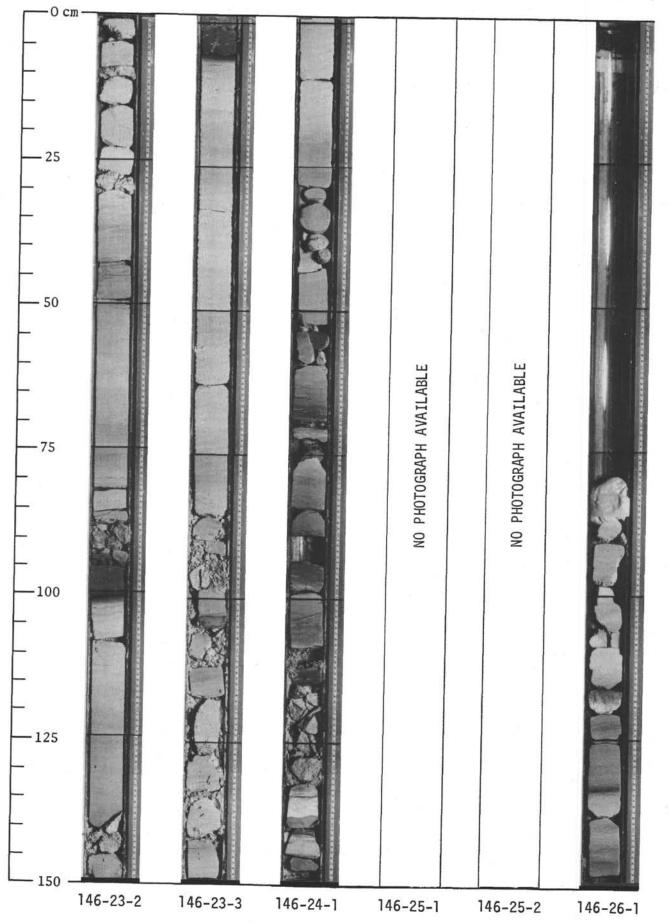


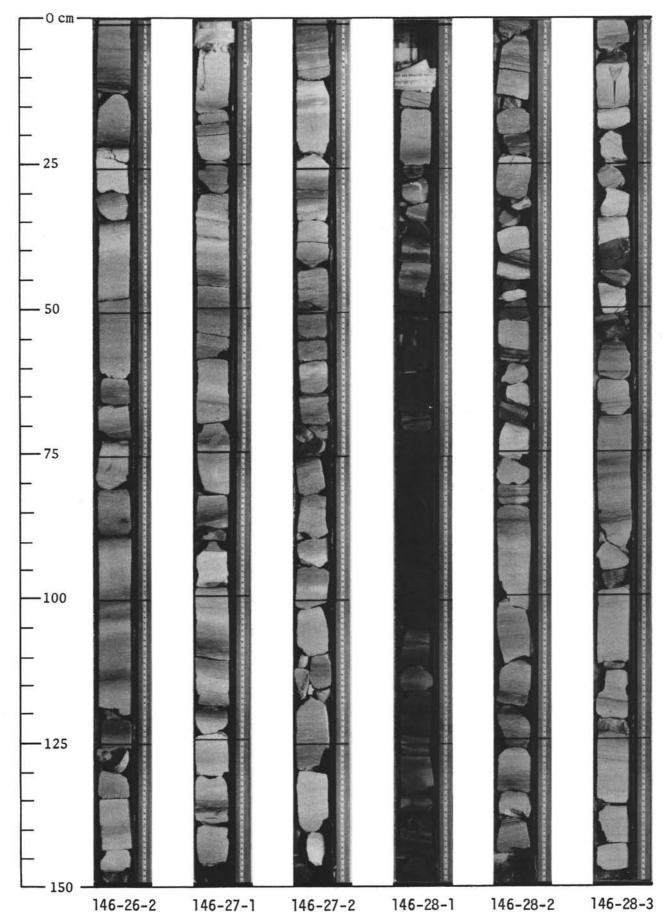


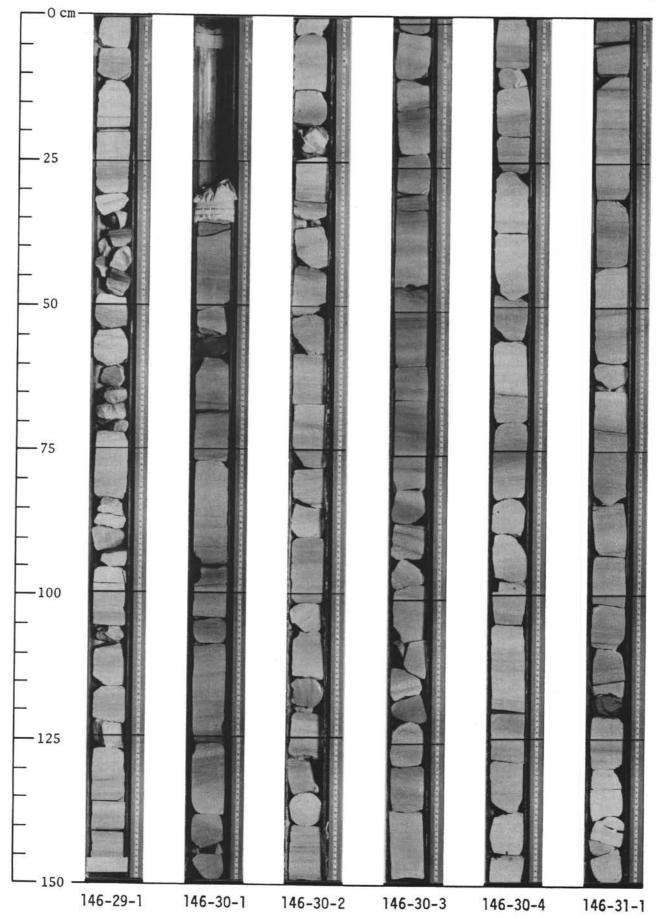


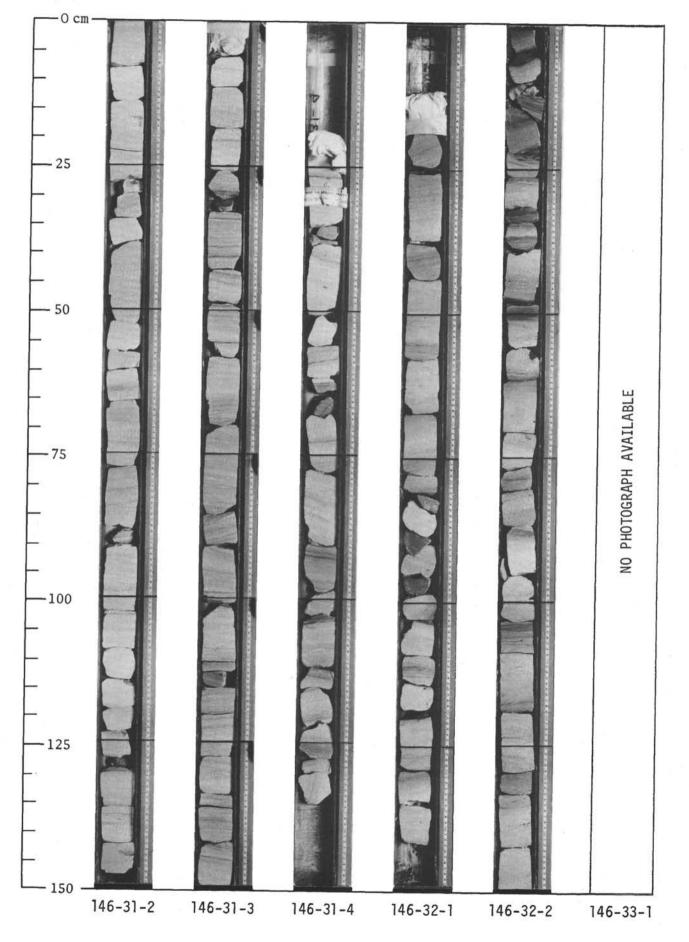


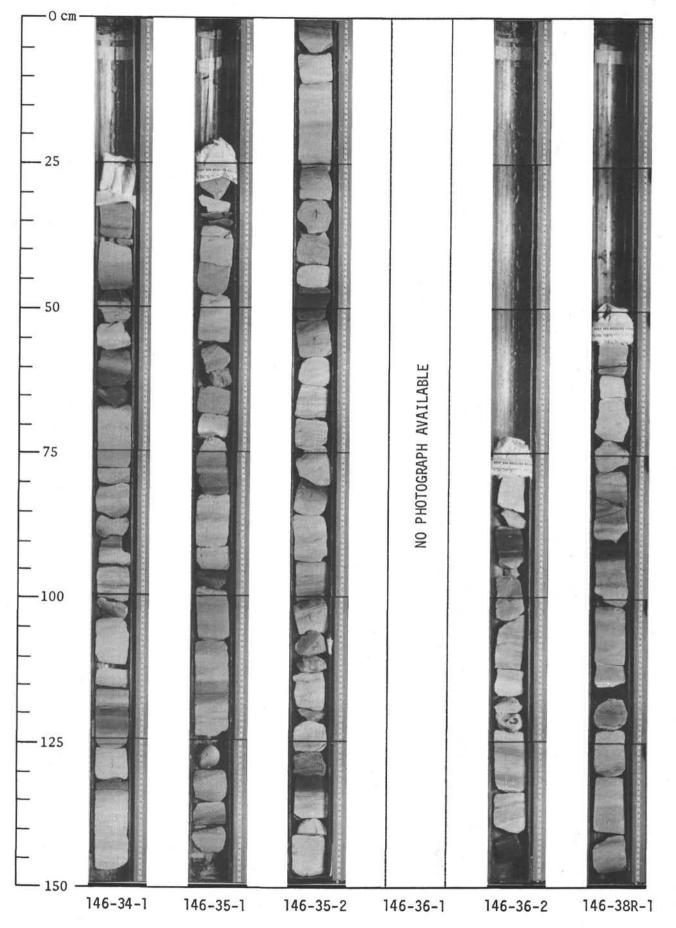


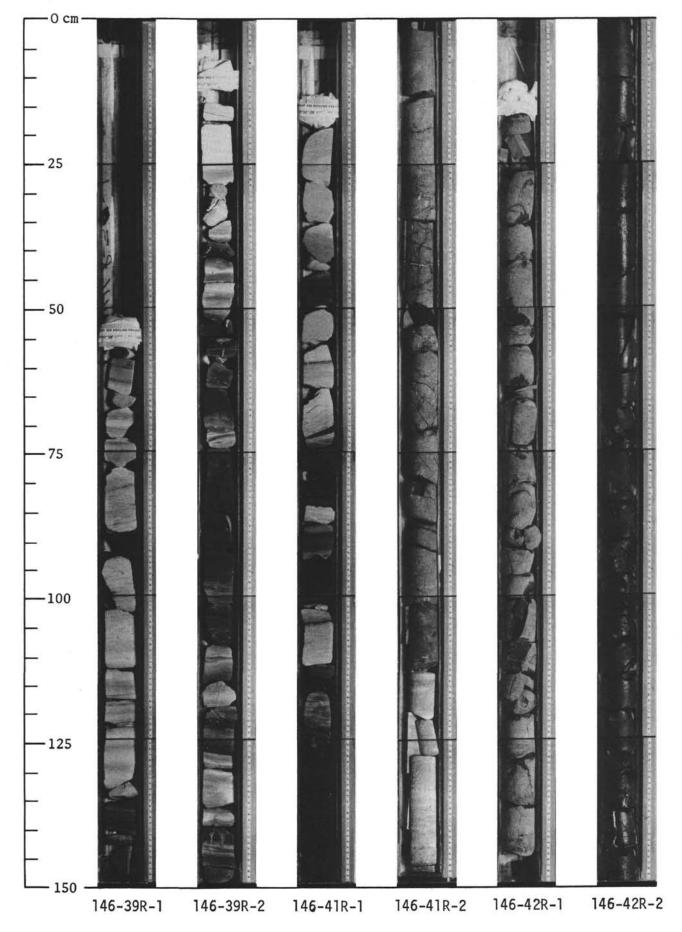


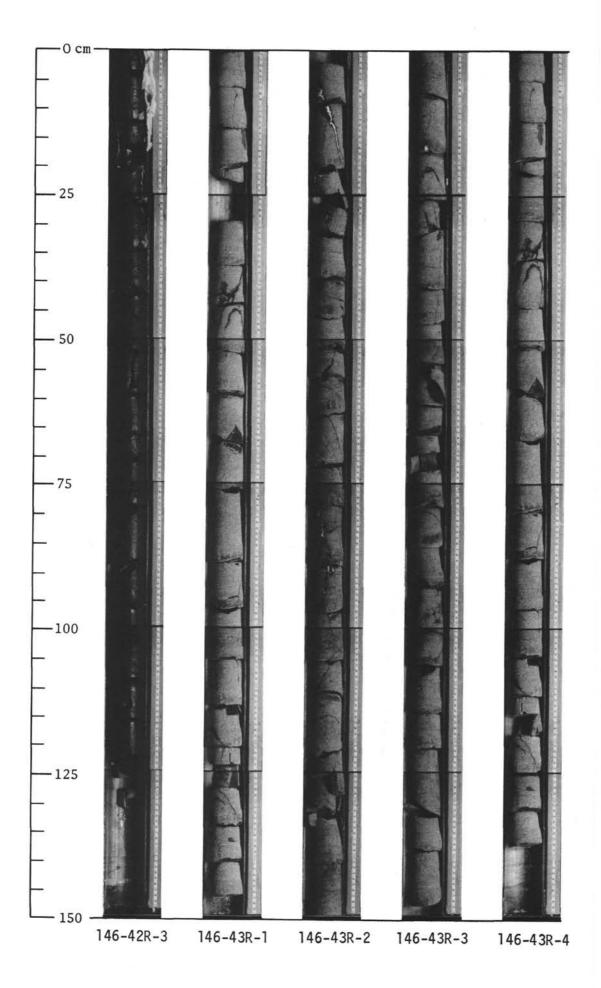








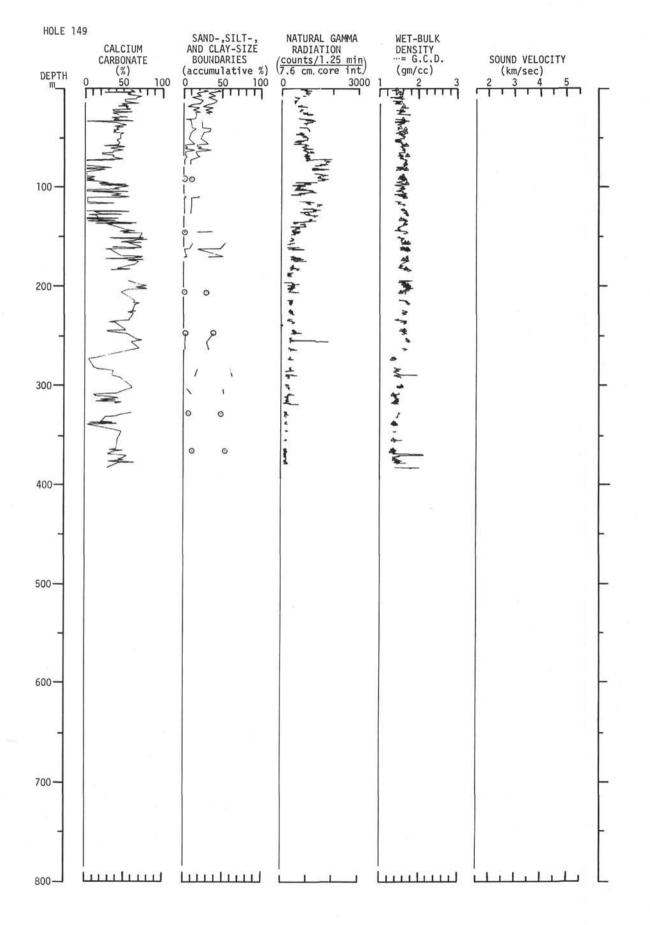




SITE 149

LITHOLOGY

| DEPTH (m) | | UNITS | SUBUNITS OR DESCRIPTION |
|--------------|--------------------|---|---|
| 0 | PLIOCENE PLEIST. R | UNIT I (O to 198.50 m) FORAM NANNO CHALK OOZE, FORAM NANNO MARL OOZE, and CLAY | Subunit a. (0 to 71.85 m) FORAM NANNO CHALK and MARL 00ZE, grayish orange (0 to 30 m) and light olive gray. Pteropods present in upper 9 m. |
| 100 | | | Subunit b. (71.85 to 96.45 m) ZEOLITIC CLAY, brown |
| 100 | M.M.IOK.MIO | | Subunit c. (96.45 to 135.45 m) FORAM NANNO MARL OOZE, bluish gray, dark yellowish orange, and light |
| - | | | brown. Subunit d. (135.45 to 198.50 m) FORAM NANNO CHALK OOZE and FORAM NANNO MARL OOZE, yellowish gray and grayish orange. |
| 200— | 6. MIO | UNIT II (198.50 to | Subunit a. (198.50 to 270.00 m) RADIOLARIAN rich |
| _ | | 382.50 m) RADIOLARIAN NANNO CHALK, RADIOLARIAN RICH NANNO CHALK, and | NANNO MARL; minor radiolarian rich nanno chalk; greenish gray and yellowish gray. |
| | | semi-indurated RADIOLARIAN OOZE | Subunit b. (270.00 to 382.50 m) RADIOLARIAN NANNO |
| 300 — | CENE | | CHALK, light olive gray and greenish gray with interbeds of semi-indurated radiolarian ooze, and frequent ash layers. |
| - | | | |
| | | UNIT III (382.50 to | |
| 400 | | 390 m? T.D.). FORAM NANNO CHALK, and CHERT | FORAM NANNO CHALK, light olive gray. CHERT, olive gray. |
| - | | | |
| 500- | | | |
| | | | |
| - | | | |
| | | | |
| 600 | | | |
| - | | | |
| | | | |
| 700 | | | |
| _ | | | |
| | | | |
| 800 | L | *The upper bound | dary was cored at 406 m at Site 146. |



| | | ZONE | | | | TPLE | APLE | | CaCO ₃ (%) | LON |
|--------|------------------------|------------------|--|--|-----------------------------------|--------------------------|--|---|-----------------------|-----|
| AGE | FORAM | NANNO | | | PALEO SA ABUNDANCE PRESERVA | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 10 | DEFORMATION | | |
| RECENT | Globorotalia fimbriata | Emiliana huxleyi | | | ORE | | FAW | FORAMINIFERAL CHALK OOZE with abundant pteropods. | | |
| | | | | | | | | | | |

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¹For explanation of symbols, see Chapter 1

| 0111 | E 149 | | HULE | | CUKE 2 | - | conco m | TERVAL (III) 1-10 | | | - | | | | | | |
|-------------|--|-----------------------|---------|-------------------------------------|-----------|---------------|----------------------------------|---|---|-------------|-----------------------|------------------|--|--|--|--|--|
| AGE | | ONNO ONNO | SECTION | METERS | LITHOLOGY | LI THO SAMPLE | PALEO ABUND/ PRESEF | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | DEFORMATION | DEP IN COR M | TH | 149 CORE 2 NATURAL CAMMA RADIATION (Counts/1.25 min 7.6 cm core interval) 2000 4000 | WET-BULK DENSITY C = sample - = GRAPE = 6.C.D. (gm/cc) .0 2.0 3.0 1 | MATER CONTENT-POROS #T# * Sample = sample | SOUHD VELOCITY (km/sec) 2.0 3.0 4.0 5.0 6.0 | PENETROMETER X 10 ⁻¹ mm 0 300 <u>CP</u> |
| | | Emiliana huxleyi | 1 | 0.5 1.0 0.5 | | | FAW NCW FAW NCW | FORAMINIFERAL NANNOPLANKTON MARL and CHALK 002E; light yellowish brown (10YR6/2) to grayish orange (10YR7/4) and moderate yellowish brown (10YR5/4). Abundant pteropods in upper sections. Disseminated black spots throughout. Sediment is soft, homogenous, and slightly dis- turbed. | | | 211111 | 1 | | | | | |
| PLEISTOCENE | truncatulinoides (Globorotalia hessi sz) | | 3 | 0.0- | | | N C W F A W F A W F A W | | • • • | | 4 | 3 | | 0 | * Q | | |
| | trunca tul inoides | Gephyrocapsa oceanica | 4 | 1.0- 0.0- 0.5- 1.0- | | | N A W F A W | | A | | 6 | 4 | | 9 | | | 0 |
| | Globorotalia | | | 0.0- 0.5- 1.0- ORE CHER | | - | F A W F A W R R G F A W | | 3 | | | Б The of 1 | 0° of the natural gamma BB6. This background w | data is equal to t as subtracted from | he atnospheric background the data. | count (gamma count when eq | uipment was empty) |

For explanation of symbols, see Chapter 1

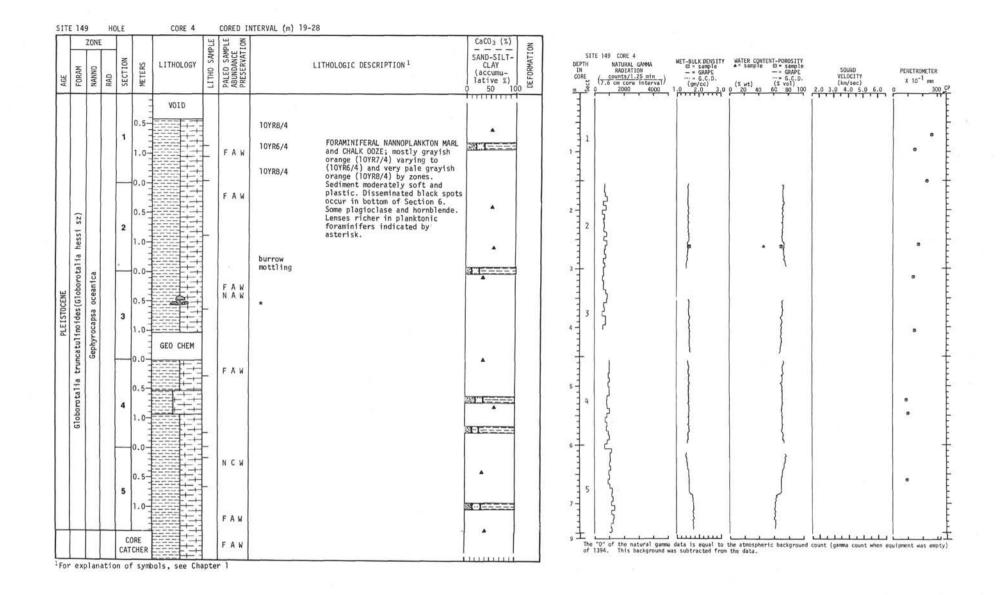
SITE 149

HOLE

CORE 2

CORED INTERVAL (m) 1-10

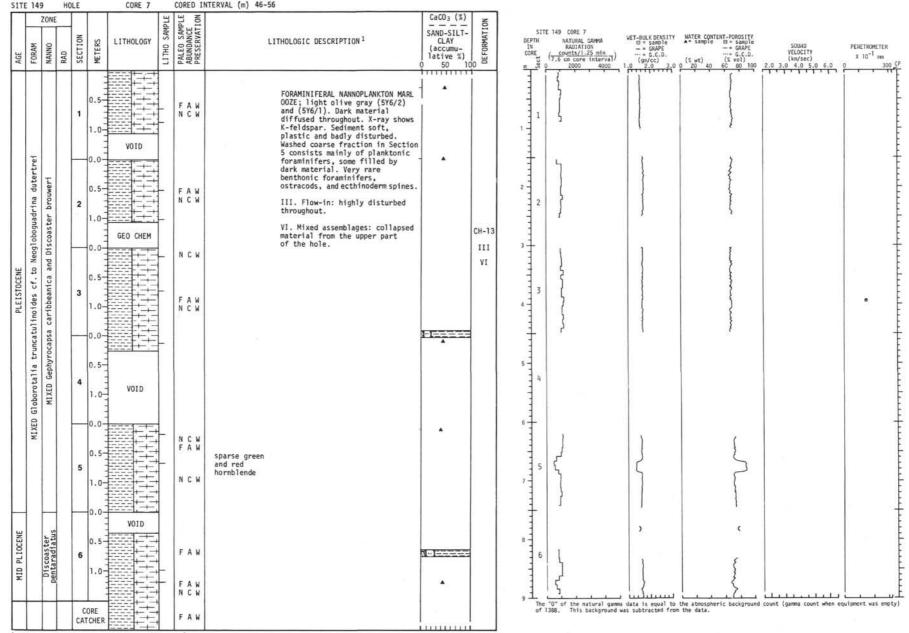
104



| ZONE | NE | _ | | | 2 | I OI | | CaCO ₃ (%) | N | | | | | |
|---|------------------------|-----------|--|-----------|--------------|---|--|---|-----------------------|--------------------------|---|--|-----------|---|
| FORAM NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 1 | DEFORMATION | CORE / counts/1.25 min \ | T-BULK DENSITY D = sample - * GRAPE (gm/cc) 2.0 3.0 | WATER CONTENT-PORC = sample = s: -= G (% wt) (% vc) 0 20 40 6C 8 | APF SUUNU | PENETROMET X 10 ⁻¹ m |
| Globorotalia truncatulinoides(Globorotalia hessi sz) Gephyrocapsa oceanica | nebhl) neebha acamirea | 1 2 3 4 5 | 0.5- 1.0- 0.5- 1.0- 0.5- 1.0- 0.5- 1.0- 0.5- 1.0- 0.5- 1.0- | V01D | | FAW FAW FAW FAW FAW FAW FAW | <pre>FORAMINIFERAL NANNOPLANKTON MARL and CHALK ODZE between grayish orange (10/R7/4) and greenish gray (56/67)) gradually changing in Section 3 to light olive gray (55/22) and greenish gray (55/22) and greenish gray (55/67)). Sediment moderately soft and plastic. Disseminated black spots almost throughout, and abundant at some levels (indicated by *). Some plagio- clase and hormblende occur at a few levels. X-ray also shows K-feldspar. * * *</pre> | | CH-13 III CH-13 | | | · · · · · · · · · · · · · · · · · · · | | 0 0 0 0 0 0 0 0 0 0 0 |

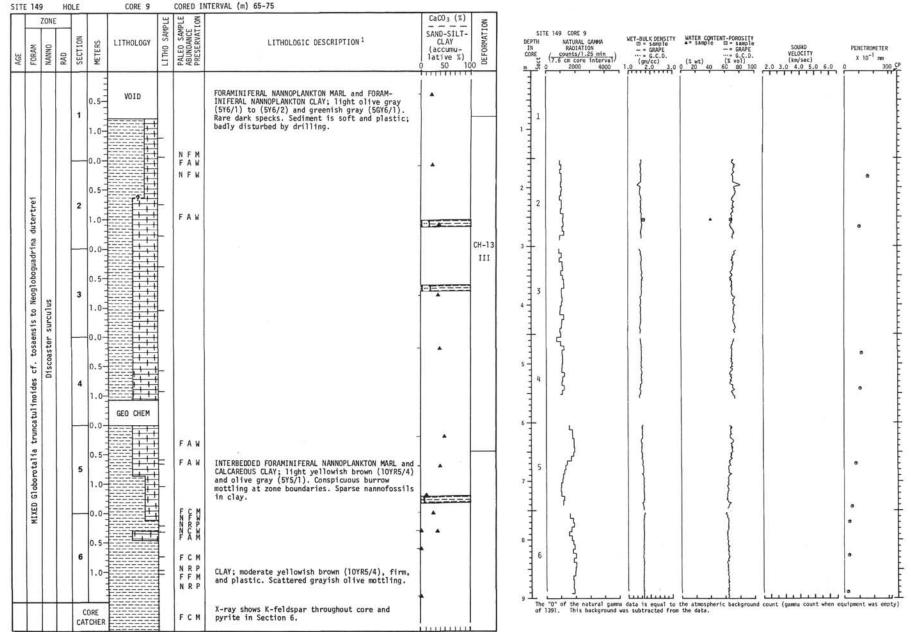
| ITE 1 | | _ | OLE | | CORE 6 | _ | | NTERVAL (m) 37-46 | 1 0.00 (0) | | |
|---|--------------|---|---------|-------------|-----------|--------------|---|--|------------|-----------------------------|---|
| AGE FORAM | ZON | Τ | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | | DEFORMATION | SITE 149 CORE 6 MET-BULK DENSITY WATER CONTENT-POROSITY DEPTH NATURAL GAMMA D= sample D= sample D= sample IN RADATION |
| PLEISTOGENE truncatulinoides cf. tosaensis to Neogloboguadrina dutertrei | Gephyrocapsa | | 3 | 0.5 | V01D | | FAW NCW FAW FAW FAW FAW | FORAMINIFERAL NANNOPLANKTON MARL 002E; light olive gray (SY6/2) to greenish gray (SGY6/1) with intermittent dusky blue (SPB3/2) layering throughout. Some plagioclase and hornblende. X-ray also shows K-feldspar. Sediment moderately firm and plastic. Foraminiferal- rich layer indicated by "+". Black, 1-2 cm diameter pods of sulfide minerals at 90 and 140 cm in Section 6. VI. Mixed assemblages: the section appears Superficially undisturbed as it has good horizontal banding but the fauna and flora indicate mixing with younger material. dusky blue | | CH-13 VI CH-13 TII | |
| LATE PLIOCENE MIXED Globorotalia ti | | | 5 | 0.0 | | | N C W F A W N C W N C W F A W | Dusky yellow (5Y6/4) blotches of oxidized hydrotroilite. | | CH-13 CH-13 III | 6 . 7 . 8 . 9 . 16 . . </td |
| | | | | ORE CHER | | | FAW | | | | of 1391. This background was subtracted from the data. |

For explanation of symbols, see Chapter 1

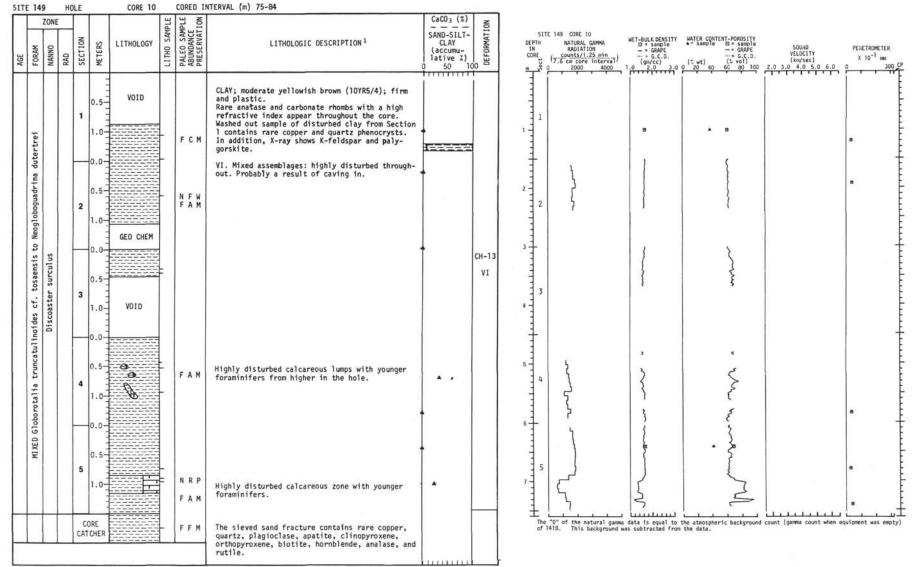


CORED INTERVAL (m) 46-56

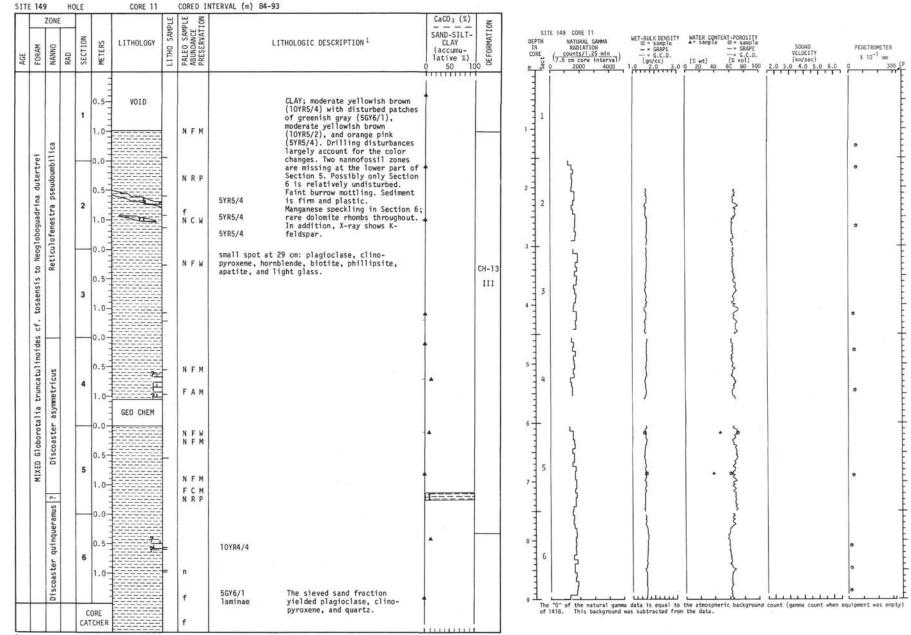
| FORAM | ZON NANNO | Г | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ¹ | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 1 | DEFORM | SITE 149 CORE B MET-BULK DENSITY WATER CONTENT-POROSITY DEPTH NATURAL GAMMA D = sample D = sample D = sample IN RADIATION |
|------------------|----------------------|---|---------|--------|-----------|--------------|---|--|--|--------|---|
| | brouweri | | 1 | 0.5 | | _ | N F M F A W | FORAMINIFERAL NANNOPLANKTON MARL 00ZE; light olive gray (5Y5/2) and (5Y6/2), soft, plastic, and very disturbed. Some plagioclase and disseminated dark material throughout, often dusky yellows. In addition,X-ray diffraction shows K-feldspar. Foraminifers are scarce in Section 6. | kg | | |
| a dutertrei | scoaster | | 2 | 0.5 | | _ | NFM | Lumps of more clayey sediment with smoother texture. III. Flow-in: highly disturbed, homogenized in upper part, with | | CH-13 | |
| Neogloboguadrina | Assemblages of D | | | 1.0 | | | NFM | disturbed layers in lower part. | - | | |
| 4 | 3 | | 3 | 0.5 | | | N F M F A W | | - | 3 111 | |
| cf. tosaensis | ; | | | 0.0 | | | NFW | | | | |
| truncatulinoides | 100 - 01 - 100 - 003 | | 4 | 0.5 | | _ | | | <u> </u> | | |
| Globorotalia tru | | | | 0.0 | GEO CHEM | | FAW | | | СН-13 | |
| MT XFD Globo | | | 5 | 0.5 | | _ | | | 8 0 222 2 | 11 | |
| | Disc | | | -0.0 | | | F A W N F W | | | | |
| | | | 6 | 0.5 | | - | FAW | Distorted, more clayey layers. | | | |
| | | - | | CORE | | - | FAW | fim | - | | 9 The "O" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was en of 1423. This background was subtracted from the data. |



110

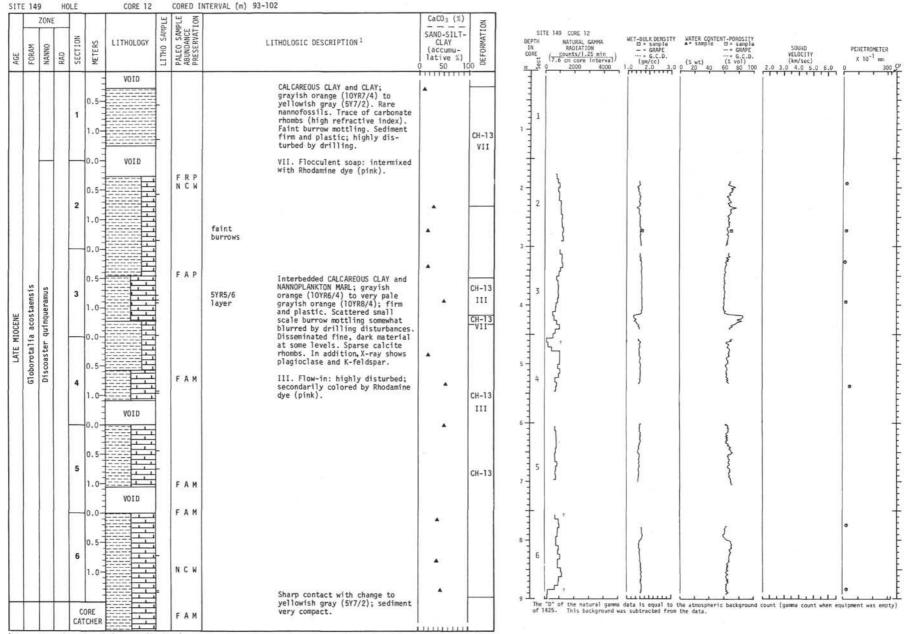


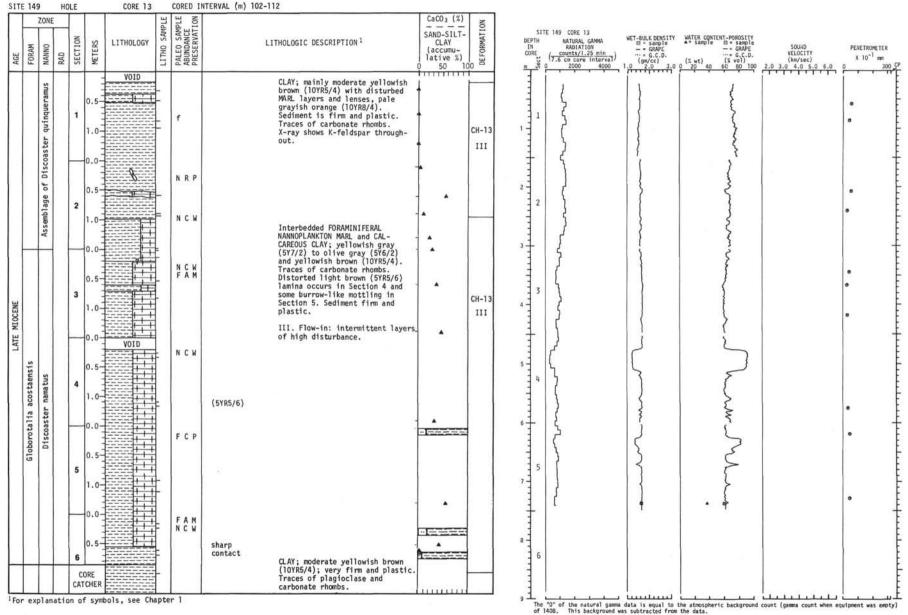
·For explanation of symbols, see Chapter 1



SITE 146/149

For explanation of symbols, see Chapter 1

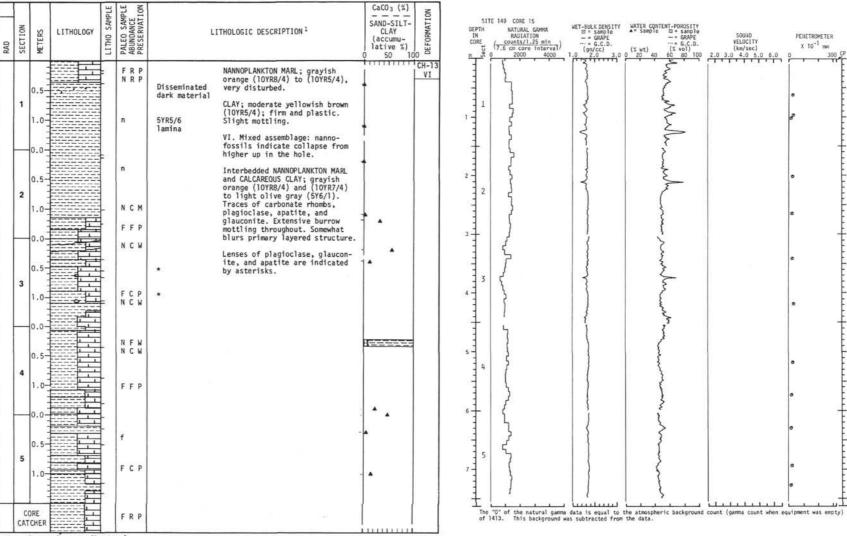




| SITE 1 | 49 | HC | DLE | | CORE 14 | | | NTERVAL (m) 112-121 | | | |
|---|----------|----|---------|--------|-----------|--|---|---|---|-------------------|---|
| AGE FORAM | NANNO | T | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | | SITE 149 CORE 14 WET-BULK DENSITY WATER CONTENT-POROSITY DEPTH NATURAL GAMMA E sample Sound PENETROMETER IN RADATION |
| LATE MIOCENE oborotalia acostaensis | | | 1 | 0.5 | VOID | | N F M F A M | CLAY with minor MARL; mainly grayish orange (10YR7/4) but extensively mixed with varicolor- ed marley lumps due to drilling disturbances. Sparse plagioclase and carbonate rhombs. Sediment is soft and plastic. V. Drilling breccia: varicolored lenses and lumps of marl in- creases toward Section 3. | | | |
| 610 | hama tus | | 2 | 0.5 | | | | | | CH-13 III V | |
| NE ardii | oaster | | | 0.0 | | | NFW | | | | |
| MIDDLE MIOCENE Globorotalia menardii | | | 3 | 1.0 | GEO CHEM | - | N C W F C P n | CLAY; moderate yellowish brown (10YR5/4) becoming slightly grayish yellow green (5GY7/2) toward the bottom. Distinct | | | |
| 610b | | | 4 | 0.5 | | | n | light brown (5YR5/6) lamina occurs in Section 4. Traces of carbonate rhombs and plagioclase. Faint burrow mottling. Sediment is firm to very firm in lower (5YR5/6) part. | | | |
| | | | | DRE | | the state of the s | N F W F F P f | In addition, X-ray diffraction shows K-feldspar scattered throughout. | | | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty of 1414. This background was subtracted from the data. |

For explanation of symbols, see Chapter 1

| S | |
|----|--|
| E | |
| E | |
| - | |
| 46 | |
| 1 | |
| 4 | |



CORED INTERVAL (m) 121-130

116

SITE 149

FORAM

menardii

Globorotalia

mayeri

Globorotalia kugleri

Globorotalia fonsi lobata

Globorotalia fohsi fohsi

Discoaster

MIOCENE

MIDDLE

coal i tus

Catinaster

AGE

ZONE

HOLE

CORE 15

| AGE FORAM NANNO BAOZ | RAD SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | | SITE 149 CORE 16 DEPTH MATURAL GAMMA WET-BULK DENSITY WATER CONTENT-POROSITY IN RADIATION DESDE CORE (-counts/1.23 min) (|
|--|----------------|---|-----------|--------------|--|---|---|-------------|---|
| MIDDLE MIDCENE nda Globorotalia fohsi fohsi Sphenolithus heteromorphus Discoaster exilis | 2 | | V01D | | N C W FR PP N F W F C P F A M F A M | Interbedded NANNOPLANKTON MARL, CALCAREOUS CLAY, and CLAY; moderate yellowish brown (10YR5/4) to yellowish gray (5Y7/2) and olive gray (5Y5/1). Traces of plagioclase and carbonate rhombs throughout. In addition, X-ray diffraction shows K-feldspar in Section 1. Boundaries between layers some- what blurred by extensive burrow mottling throughout. Sediment firm and plastic. Interbedded FORAMINIFERAL NANNOPLANKTON MARL, CHALK, and orange CLAY; yellowish gray (5Y8/1) and layer moderate yellowish brown (10YR5/4). Layering fairly conspicuous despite extensive burrow mottling throughout. Sediment firm and plastic. | | CH-13 VI | |
| fohsi peripheroronda Sphenolithus P | | 0.0 1.0 0.5 1.0 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.0 0.5 0.0 0.5 0.0 0.0 0.0 0 | | | F A M F C M F C M F A M | | · · | | The "O" of the natural game data is equal to the atmospheric background count (game count when equipment was empiricated from the data. |

CORE 16 CORED INTERVAL (m) 130-139

SITE 149

HOLE

| AGE | Т | RAD RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOG | GIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT CLAY (accumu- lative %) 50 | 100 | NCLOWINI TON | DEPTH IN CORE / | 149 CORE 17 NATURAL GAMMA RADIATION counts/1.25 min 7.6 cm core interval 2000 4000 | WET-BULK DENSITY D = sample | WATER CONT * Sample (% wt) 0 20 40 | ENT-POROSITY = sample -= GRAPE = G.C.D. (% vol) <u>6C</u> 80 100 | (km/ 2.0 3.0 4. | CITY (sec) | PENETROM X 10 ⁻¹ | |
|----------------|---------------------------------|---------------------------|---------|----------------------------------|-----------|--------------|---|--|--|---|------|--------------|--|---|--|---|---|--------------------|---------------|--------------------------------|------------|
| MIDDLE MIOCENE | Gioborotalla Tonsi peripreronda | sni | 2 | 0.5- 1.0- 0.5- 1.0- | | | FFP | (ST/) Tota VI.) | egated NANNOPLANKTON MARL; t olive gray (5Y6/1), ish orange (10YR7/4) to rate yellowish brown R5/4) and yellowish gray | | CH-V | | 2 2 | | | | Junior | | | | |
| MIDDLE | Cochooolithine hotococochie | Spheno I thus heteromorph | 3 | 0.5- | | | F C M F A M N C W | NANN(very gray ligh pale Exte | rbedded FORAMINIFERAL OPLANKTON MARL and CHALK; pale orange (10YR8/2) to ish orange (10YR7/4), t olive gray (5Y6/1) and yellowish brown (10YR6/4). nsive burrow mottling | • | | | 4 4 5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 | | | | man harris | | | 0 0 | |
| | glomerosa | | CAT | 0.0 0.5 1.0 0RE CHEF | | | F C M F A M N C W F A M | thro plas Common plagioclase, quartz | ughout. Sediment firm and tic. - | | | | 6 | D ^o of the natural gama 73. This background w | data is equal to as subtracted from | the data. | ric background | count (gamma | count when e | e e quipment was | Lu sempty) |

118

SITE 149

HOLE

SITE 146/149

CORED INTERVAL (m) 139-149

CORE 17

| EVERY MIDCENE 1 1 0.0 0.0 0.0 0.0 0.0 0.0 0. | SAMPLE SAMPLE NCE | NOTICE | CaCO ₃ (%) SAND-SILT- | SLIE 149 CORE 18 VET BUT POINTLY MATER CONTENT_POROTITY |
|--|---|-----------------------------------|--|---|
| EBAK MIOCENE 1 1 1 0.5 1 1 0.5 1 1 0.0 0.0 0.0 0.5 1 1 1.0 0.0 0.0 0.5 1 1 1.0 0.0 0.0 0.5 1 1.0 0.0 0.5 1 1.0 0.0 0.0 0.0 0.0 0.5 0.0 0.5 0.0 0.5 0.0 0.5 0.0 0.5 0.0 0.5 0.5 | LITHO SAMPLE PALEO SAMPLE ABUNDANCE | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- 0 50 100 | DEPTH ANTURAL GAMMA METBULK DENSITY MATER CONTENT-DROSITY IN RADIATION — * GAMPE D* Sample D* Sample Sample CORE (-context)-25 min_1) |
| 4 1 .0 0 .5 1 .0 0 .0 0 .0 | VOID F A N F VOID F A N F | upper 3 sections, due to drilling | | |
| eta angle in the second | GED CHEM | (10YR7/4) | CH-13 III | |

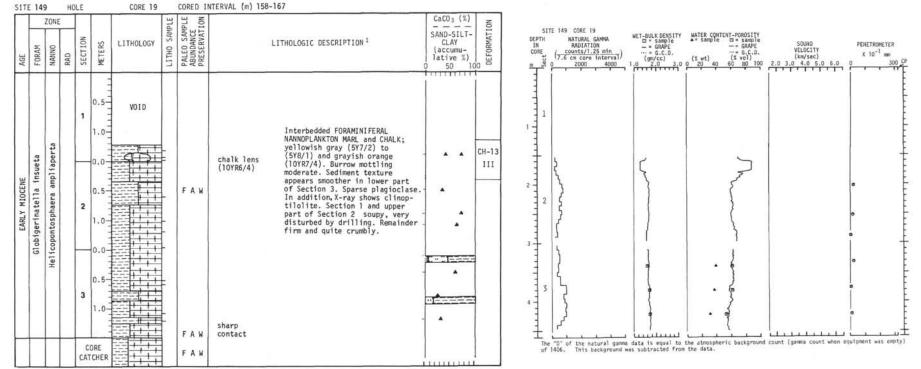
For explanation of symbols, see Chapter 1

SITE 149

HOLE

CORE 18 CORED INTERVAL (m) 149-158

SITE 146/149



For explanation of symbols, see Chapter 1

| 1 | ZONE | | Τ | | | SAMPLE | MPLE E T10N | | CaCO ₃ (% | ION | |
|--|--------------------------------|-----|--|--------|-----------|-----------|---|---|--|--------------|---|
| FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAV | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT CLAY (accumu- lative %) 0 50 | DEFORMA | SITE 149 CORE 20 MET-BULK DENSITY MATER CONTENT-POROSITY DEPTH NATURAL GAMMA DET Sample SOUND PENETROMETER IN RCADATION |
| Globigerinatella | Helicopontosphaera ampliaperta | | 1 1 0 2 1 0 3 1 0 0 3 1 0 4 | | VOID | | F A W F A W N C W | Interbedded FORAMINIFERAL NANNOPLANKTON CHALK, CALCAREOUS CLAY, and MARL; very pale orange (10YR8/2), grayish orange (10YR7/4) to moderate orange pink (SYR8/4) and dark grayish orange (10YR6/4). Some plagio- clase. In addition, X-ray diffraction shows clinoptilolite Layering is somewhat blurred by extensive burrow mottling particularly conspicuous in marly zones. The lighter colored areas tend to be more calcareous than the darker areas, which are typically marly. Sediment firm and slightly crumbly. plagioclase, biotite Dispersed ASH, light brownish gray (SYR6/1) | | CH-13 III | 4 |
| -Globigerinita stainforthi -Globigerinita-dissimilis- | | | 0 5 1 | | | | F A W N C W N F W | | | | |
| 11 | | 0 | COR | | | | FAW | pumice | 1 | | The "O" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was em of 1380. This background was subtracted from the data. |

SITE 149 HOLE

CORE 20 CORED INTERVAL (m) 167-176

SITE 146/149

| ZONE NUMBER LITHOLOGIC DESCRIPTION ¹ CAUCH CA | | ZON | NE | T | | | L L | PLE | | CaCO3 (%) | | | | | | | | | | | |
|--|--------------|--------------|-------|---------|--------|--|---|-------------------------------------|--|---|----|--------------|------|--|-----------|---|---------------------|---------------------------------|-------------|-----------|--------------------|
| S00 1< | FORAM | NANNO | RAD | SECTION | MFTFRC | LITHOLOGY | LITHO SAM | PALEO SAM ABUNDANCE PRESERVAT | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 1 | 00 | DE T C | EPTH | NATURAL GAMMA RADIATION (<u>counts/1.25 min</u>) | - = GRAPE | - = GRAPE | VE (k 2.0 3.0 | ELOCITY km/sec) 4.0 5.0 8 | 6.0 0 | PEHETROME | ETER MIII 30 |
| 1 1 <td>10000</td> <td>Sphenolithus</td> <td></td> <td>_</td> <td>1.0</td> <td>VOID ++++++++++++++++++++++++++++++++++++</td> <td></td> <td>FFM</td> <td>FORAMINIFERAL NANNOPLANKTON CHALK; predominantly grayish orange (10YR7/4) mixed with yellowish gray (5Y7/2) and brown. Varicolored distorted layers as lumps scattered throughout. Some plagioclase.</td> <td></td> <td>CH</td> <td>13 I</td> <td>1</td> <td></td> <td></td> <td>and the second second</td> <td></td> <td></td> <td></td> <td></td> <td></td> | 10000 | Sphenolithus | | _ | 1.0 | VOID ++++++++++++++++++++++++++++++++++++ | | FFM | FORAMINIFERAL NANNOPLANKTON CHALK; predominantly grayish orange (10YR7/4) mixed with yellowish gray (5Y7/2) and brown. Varicolored distorted layers as lumps scattered throughout. Some plagioclase. | | CH | 13 I | 1 | | | and the second | | | | | |
| | stainforthi- | | - DDL | 4 | 1.0 | | ╌╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿╎┿ | FFM | and CHALK; between grayish orange (10YR7/4) and yellowish gray (5Y7/2) to slightly darker shade. Sediment firm and extensively burrowed. | | | 5 | 4 5 | | | my manus manus manus | | | 0 0 0 0 0 0 | | |

122

| 1101 | | CURE 22 | - | | | 1 | | |
|------|---------|--|--|--|---|--|---|---|
| | SECTION | 운 _LITHOLOGY 말 | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION ¹ | SAND-SILT- CLAY (accumu- lative %) 0 50 1 | DEFORMATION | SITE 149 CORE 22 WET-BULK DENSITY HATER CONTENT-PORDSITY DEPTH NATURAL CAMPAS GE = Sample Sample Sample Sample Sample Sample Source Sample Source Sample Source |
| | 1 | | | | grayish orange (10YR7/4), lower part of Section 6 is pale grayish orange (10YR8/4) to light olive gray (5Y7/1) in lowermost part. Light greenish gray blotches appear to be the result of bioturbation. Sparse plagioclase. | mmmm | CH-13 11 CH-13 | |
| - | 2 | 0.0 0.5 V0ID | | | Sediment highly disturbed. I. Watery: stored in a freezer box or unopened. | | сн-13 I | |
| | _ | 0.0 | | | | | CH-13 | |
| | 3 | | | | | | I | |
| | 4 | 0.5- VOID | | | | | | 5 4 |
| | 5 | NOT OPENED | | | | | | |
| | | 1.0- 0.0- IN FREEZER 0.5- BOX | | | | | I I | |
| | 6 | | 111111 | FRP NCW | Sparse radiolarians in layer Disturbed at 96-103 cm of Section 6. top. In general, traces of radiolar- ians in Section 6. | | | |
| | E KAD | 1 2 3 4 5 | E S S S S S S S S S S S S S | E 0 0 0 0 0 0 0 0 0 0 0 0 0 | E SQ LITHOLOGY MUMARY OUTWARY 02 05 SQ LITHOLOGY MUMARY OUTWARY 1 0.5 VOID Image: Constraint of the second s | E Sparse radiolarians in layer at 50 percent of the percent of th | E Signed Stress Sparse radiolarians in layer at 96-103 cm of Section 6. 92 Signed Stress LITHOLOGIC 92 Signed Stress LITHOLOGIC DESCRIPTION 1 1 Stress Stress 1 Stress Stress 1 Stress Stress 0.5 Stress Stress 0.5 VOID Stress 0.5 NOT OPENED Stress 1.0 IN FREEZER Stress 0.5 NOT OPENED Stress 1.0 IN FREEZER Stress 0.0 Stress Stress 0.5 Stress Stress 1.0 Stress Stress 1.0 Stress Stress 1.0 | E Intrologic E Carol (C) SAU-STLT- CLW (LAW (LAW) (LCW) (L |

·For explanation of symbols, see Chapter 1

SITE 149 HOLE

CORE 22 CORED INTERVAL (m) 185-195

SITE 146/149

| | 1.5 |
|--|-----|
| | |
| | |

SITE 149 HOLE CORE 23 CORED INTERVAL (m) 175-204 LITHO SAMPLE PALEO SAMPLE ABUNDANCE PRESERVATION CaCO3 (%) ZONE DEFORMATION SAND-SILT-SITE 149 CORE 23 WATER CONTENT-POROSITY = sample = sample = GRAPE SECTION WET-BULK DENSITY B = sample - = GRAPE METERS LITHOLOGIC DESCRIPTION 1 DEPTH LITHOLOGY CLAY NATURAL GAMMA FORAM NANNO (accumu-RADIATION (<u>counts/1.25 min</u> (7.6 cm core interval) SOUND PENETROMETER RAD CORE VELOCITY (km/sec) AGE lative %) (1 vol) 6C 80 100 x 10⁻¹ mm (% wt) 20 40 50 100 н 2000 4000 2.0 3.0 4.0 5.0 6.0 VOID ++++ +++++ 111 FORAMINIFERAL NANNOPLANKTON CHALK; predominantly grayish orange (10YR7/4) with mixed light greenish gray (5GY8/1) and moderate orange pink (6YR8/4). Color mixed due to extreme drill-0 5 1 Ĥ ing disturbance. Traces of radiolarians and 1 -.0 plagioclase. IV. Fragmentation highly disturbed, with firm CH-13 lumps of plastic sediment. IV druggi + 0.5 Discoaster 2 2 .0. . primordius 3-Interbedded FORAMINIFERAL NANNOPLANKTON CHALK and RADIOLARIA-RICH NANNOPLANKTON CHALK and FAM . bipes NCW And NADLCHARLERATE HEAVED FAILS and light greenish gray (5078/1). Traces of plagioclase and sponge spicules. Faint burrow mottling throughout. Sediment is compact yet crumbly. MIOCENE ٠ . 10.5 RFP Lychnocanium 壹 . Globigerinoides 3 3 EARLY +1 4 -NCW .0 GEO CHEM ł 4 + + NCW + + 5-. carinatus 4 4 .0 FAM -1-Triquetrorhabdulus ٠ 6-5 . NFW 7 -.0-. . Ŧ LILLI hunting FAM CORE The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1427. This background was subtracted from the data. + CATCHER RFM

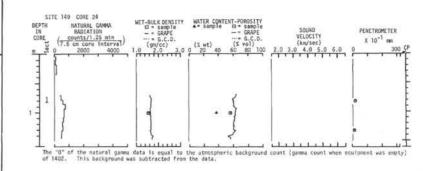
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SITE 146/149

300

For explanation of symbols, see Chapter 1

| ZONE SAND-SILT- CLAY (accumu- lative %) UNIVEN 300 VOID 000 000000000000000000000000000000000000 | | | ZONE | | 1 | | | SAMPLE | APLE FION | | CaCO ₃ (%) |
|--|---------------|-------------------------------|---------------------------------|--------------------|---------|---------|-----------|--------|-----------------------------------|--|-----------------------|
| WIND DUE TO THE PART OF THE PA | AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | | PALEO SA ABUNDANCI PRESERVA | LITHOLOGIC DESCRIPTION 1 | 0 50 100 |
| | EARLY MIOCENE | Globigerinoides primordius | Triquetrorhabdulus carinatus | Lychnocanium bipes | 1 C | 1.0 | VOID | | F R M R F M | RADIOLARIA-RICH FORAMINIFERAL NANNOPLANKTON MARL; grayish yellow green (56Y7/2), compact. Faint scattered burrow mottling. Sparse apatite. | |



| | 1 | ONE | | | | | PLE | | CaCO ₃ (%) |
|---------------|--------------------|----------------------|------------------|---------|--------------|----------|---|--------------------------|---|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | ETHOLOGY | LITHO SAMPLE PALEO SAMPLE ABUNDANCE | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 1 |
| | | | | 1 | 0.5 | | FFNRCN | | |
| (E | primordius | carinatus | elongata | | 0.0- | | RCN | | |
| EARLY MIDCENE | Globigerinoides pr | Triquetrorhabdulus (| Lychnocanoma elc | 2 | 0.5- | | f NCV RCI | | |
| | Globi | Trique | Ly. | 3 | 0.0- | | RCI | 1 | |
| | | | | 3 | 1.0- | | NCV | | |
| | | | | | ORE FCHER | | FAN RCM | | |

SITE 149 CORE 25 $\begin{array}{c} {\sf MET-BULK DENSITY} \\ \begin{array}{c} {\sf METE-BULK DENSITY} \\ {\sf meters} \\ {\sf meters}$ NATURAL GAMMA RADIATION (<u>counts/1.25 min-</u> (7.5 cm core interval) 0 2000 4000 DEPTH SOUHD VELOCITY (km/sec) 2.0 3.0 4.0 5.0 6.0 PENETROMETER IN CORE G.C.D. (% vol) 60 80 100 x 10⁻¹ ms 1.0 300 ----CH-13 . 2 3-3 4 -Ŧ The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1402. This background was subtracted from the data.

DEFORMATION

III

¹For explanation of symbols, see Chapter 1

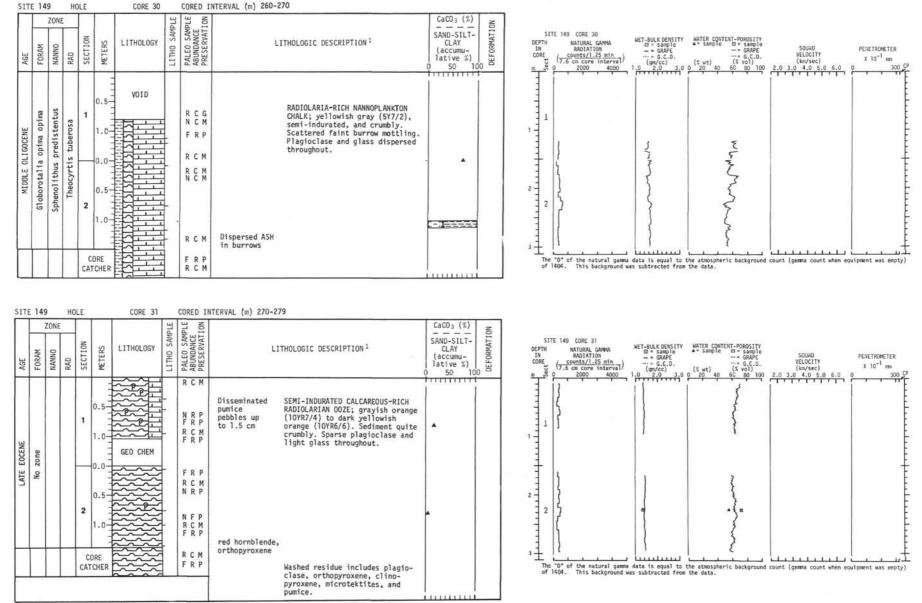
| SITE | 149 | 1 | HOLE | ŝ. | | CORE 26 | | | NTERVAL (m) 223-232 | | | |
|---------------------|--------|--------------|-----------------------|--------------|---|--|------------|--|--|---|-------------|--|
| ω. | T | ONE | 0 CTTON | CI TUN | TERS | LITHOLOGY | THO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- | DEFORMATION | SITE 149 CORE 26 WET-BULK DENSITY MATER CONTENT-POROSITY DEPTH NATURAL GAMMA GF sample A* sample GF sample SOUND PERETROMETER IN RADIATION - GRAPE - GRAPE SOUND PERETROMETER CORF CONTENT/J25 min - GRAPE - GRAPE SOUND PERETROMETER |
| LATE OL IGOCENE AGE | ugleri | us carinatus | Lycnnocanoma elongata | 0. | 5 0 0 5 0 0 5 0 5 5 0 5 5 5 5 5 5 5 5 5 | VOID | TITHO S | R C M F R P R C M R C M F R P R F M | RADIOLARIA-RICH NANNOPLANKTON MARL and CHALK; grayish yellow green (5GY7/2) to moderate grayish yellow green (5GY6/2) and light olive gray (577/1). Scattered, faint burrow mottling, light greenish gray (5GY8/1). Some plagioclase and light volcanic glass. Sediment firm and quite crumbly. | CLAY (accumu- lative %) 0 50 10 | 0 | DEPTH NATURAL CAMA VET = BUK DENSITY VALUE CONSTITUTION SOURD SOURD SOURD PETERMETER CORE 0 Counts/1/25 and (7.6 or core interval) |
| | | | | CORE CORE | | $\frac{1}{2}$ $\frac{1}$ | | N C W R F M R C M F R P | Disseminated pumice | | | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was empty) of 1402. This background was subtracted from the data. |

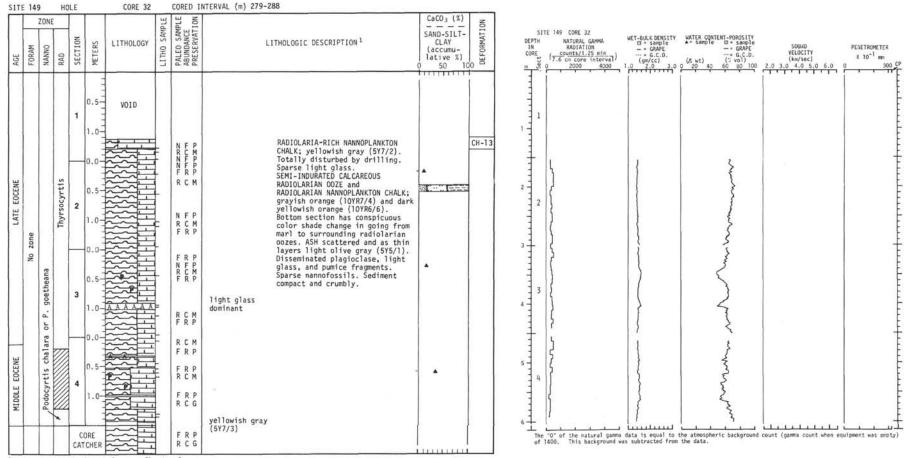
126

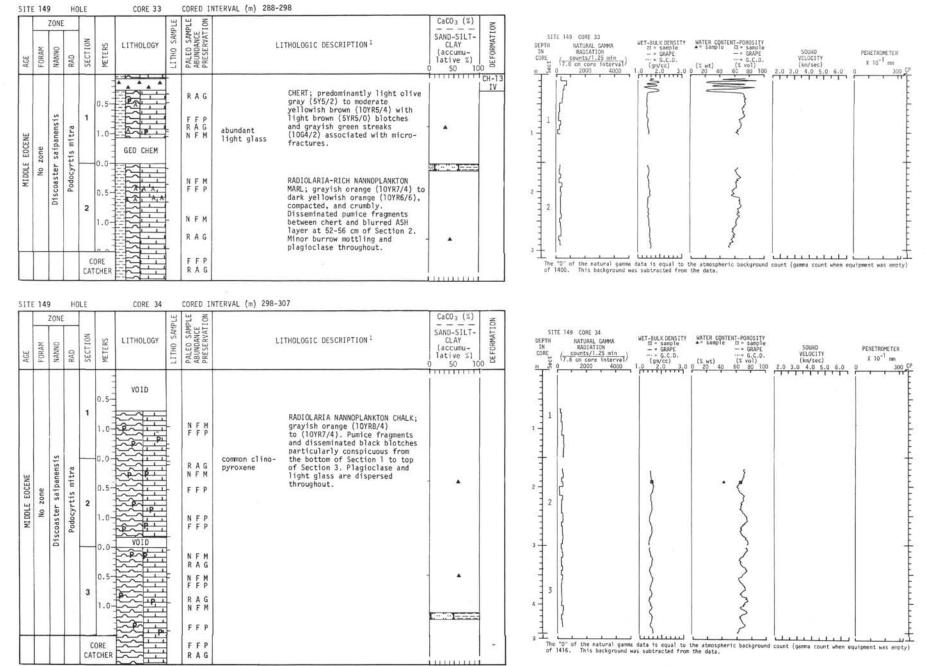
| SITE 14 | ZO | | OLE | | CORE 27 | | | NTERVAL (m) 232-241 | CaCO | | z | |
|--|------------------------------|-----|---------|--|-----------|--------------|--|--|---|--------------------------------------|-----|--|
| AGE FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND- CL/ (accul lativ 0 50 | SILT- AY amu- ve %)) 10 | | SITE 149 CORE 27 NET-BULK DENSITY WATER CONTENT-PORDSITY DEPTH NATURAL GAMMA C=*sample Sample Sample Sample IN RADIATION |
| LATE OLIGOCENE ? Globorotalia kugleri | Triquetrorhabdulus carinatus | | 2 3 | 0.5- 1.0- 0.5- 1.0- 0.5- 1.0- | | | N C W F C M R C M R C M R C M R C M R C M R C M | RADIOLARIA-RICH NANNOPLANKTON MARL; grayish yellow green (5077/2), compact and moderately crumbly. Burrow mottling. Glass and plagioclase dispersed throughout. VII. Flocculent soup secondaril colored by Rhodamine dye. disseminated pumice RADIOLARIA-RICH CALCAREOUS CLAY trace hornblende disseminated pumice | y | | VII | The '0° of the natural game data is equal to the atmospheric background count (game a count when equipment was empty) of 1456. This background was subtracted from the data. |

| ZONE | | |
|--|---|--|
| FORAM NANNO | LITHOLOGY HIT CLAY LITHOLOGIC DESCRIPTION ¹ LITHOLOGIC DESCRIPTION ¹ | WET-BULK DENSITY WATER CONTENT-POROSITY SOUND B1 - sample - sample SOUND |
| Gioborotalia kugleri Triquetrorhabdulus carinatus | VOID RADIOLARIA-RICH NANNOPLANKTON MARL: between yellowish gray (SY7/2) and dusky yellow (SY6/4) to grayish orange (DNR7/4) toward bottom. Sediment very firm and crumbly. Burrow mottling throughout. Dispersed plagioclase and glass shards throughout: sparse apatite and quartz. Disseminated ASH lump indicated by asterisks. 1 N C P R C M N C W R C M N C W R C M . Abundant ASH; grayish blue to medium brownish gray, highly bioturbated. N C W R C M . Abundant ASH; grayish blue to medium brownish gray, highly bioturbated. | m data is equal to the atmospheric background count (gamma count when was subtracted from the data. |

| SITE 14 | 19 н | IOLE | | CORE 29 | | INTERVAL (m) 251 | 260 | | | |
|--|---|----------------|---|---------|---|-----------------------------|---|---|--|-------|
| AGE FORAM | ZONE RAD | SECTION | METERS | | PALEO SAMPLE ABUNDANCE | | ITHOLOGIC DESCRIPTION ¹ | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 10 | Image: Solution SITE 149 CORE 29 MET-BULK DENSITY MATER CONTENT-PORDSITY DEPTH NATURAL GAMMA D= sample Sampl | |
| MIDDLE OLIGOCENE Globorotalia opima | Sphenolithus ciperoensis tuberosa | | 0.5 | | R C N C F C N C C N C C N C C | Dispersed ASH in burrows | RADIOLARIA-RICH NANNOPLANKTON CHALK; yellowish gray (5Y7/2) to greenish gray. Abundant dispersed light ASH. Burrow mottling, sparse light glass, and plagioclase throughout. Sediment firm and crumbly. | | | |
| | ? Sphenolithus distentus Theocyrtis | 3 CC CAT | 0.5-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1 | | N C R C R C R C f | Dispersed ASH | Distinct ASH layer, brownish gray (5YR5/1) in Section 2. Despite burrowing layer presents sharp upper and lower contacts with sloped, graded bedding upward. Abundant light glass, biotite, and zircons; sparse apatite. | | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was en of 1404. This background was subtracted from the data. | mpty) |





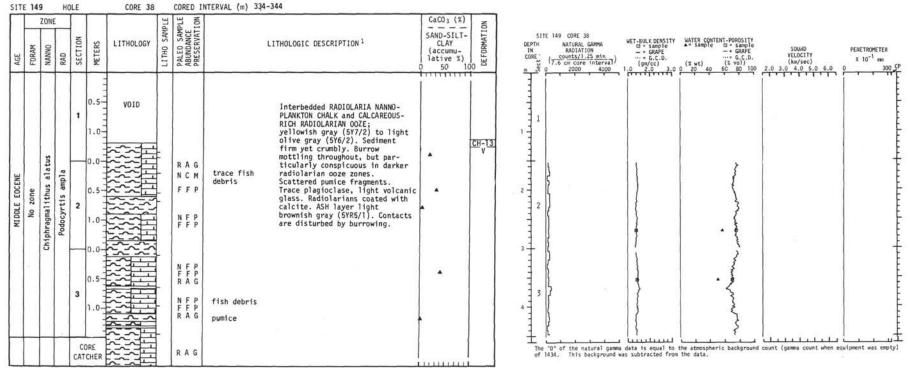


| ZONE | T | | | E LE | NO | | CaCO3 (%) | |
|------------------------------------|------------------|---|------------------|-------------------------------|---|---|---|--|
| FORAM NANNO | RAD | SECTION | 요 LITHOLOGY 뽀 | LITHO SAMPLE | PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 | |
| No zone Discoaster tani nodifer | rodocyrtis mitra | 1 1 1 0 0 1 0 0 0 1 0 0 0 1 0 0 0 1 0 0 0 0 0 0 0 0 0 0 0 0 0 | | | FAGP FPMGFAGP FFAGPAGM FAGPFAG FFF FFAG FFF FFAG FFF | Interbedded RADIOLARIA NANNO- PLANKTON CHALK and SEMI- INDURATED CALCAREOUS-RICH RADIOLARIAN OOZE; yellowish gray (5Y7/2) to (5Y8/2). Burrow mottling and disseminated pumice fragments. Some plagio- clase and sparse light volcanic glass. Radiolaria coated with calcite appear below Section 2. Sediment quite crumbly. | | |
| | | 0 5 1 0 6 1 1 0 0 6 | | - F - F - N - J F | F M F P A G F M F M F M F M F M F M A G Dispersed ASH A G ASH A G ASH ASH ASH ASH ASH ASH ASH ASH | | • | The "0" of the natural game data is equal to the attospheric background count (game count when equipment w of 1425. This background was subtracted from the data. |

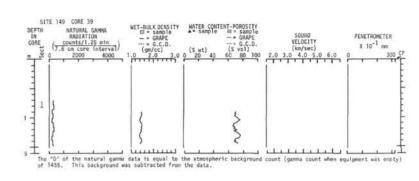
| ITE | | | HOLE | | CORE 36 | | | NTERVAL (m) 316-325 | | _ |
|-----|---------|---|---------|--------------------------------------|-----------|--------------|---|--|--|--|
| AGE | Т | DAN ONNE | CECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) O 50 100 | SITE 149 CORE 36 WET-BULK DENSITY WATER CONTENT-PORDSITY DEPTH NATURAL CAMMA D + sample D = sample IN RADIATION |
| | No zone | f tani nodifer Podocvrtis mitra | 1 | 0.5- 1.0- CORE | V01D | | R F R A F R A R F R A | RADIOLARIA NANNOPLANKTON CHALK and SEMI-INDURATED RADIO- LARIAN 00ZE; yellowish gray (5Y8/2) and light olive gray (5Y6/2) respectively. Abundant volcanic ASN; light brownish sparse light gray (5YR6/1) below 136 cm. glass Some plagioclase and sparse- | | |
| | | | | LORE AT CHER | | | RAG | light glass throughout. Radio- laria coated with calcite. | hummi | The "0" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was e of 1434. This background was subtracted from the data. |
| TE | 149 | ł | HOLE | | CORE 37 | | | NTERVAL (m) 325-334 | | |
| AGE | Т | RAD NAWNU NA | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | CaCO ₃ (%) SAND-SILT- CLAY (accumu- lative %) 0 50 100 | SITE 149 CORE 37 WET-BULK DENSITY WATER_CONTENT-POROSITY SOUND PENETROME DEPTH NATURAL GAMMA ES sample Sample Sample Sample Sample SOUND PENETROME IN RADIATION — G.C.D. … G.C.D. … COME (Exr/s) VELOCITY X.10 ⁻¹ U Q 2000 4000 1.0 G.2.0 3.0 C.2.4 0.0 2.0 3.0 4.0 5.0 6.0 0 |
| | | Unipriagmaiturus alatus Podocyrtis amola | 1 | 1.0- 0.0- 1.0- 0.0- 0.0- | | | RAGMPF FFFMG FFFM FFFM FFFM FFAG NFFAG NFFR NFM | RADIOLARIA NANNOPLANKTON CHALK; very pale orange (10YR8/2), very compact and crumbly. Trace plagioclase. Radiolaria coated with calcite. | | |
| | | | | 0.0- 0.5- 1.0- CORE | | | N R P F F P R A G F F P F F P R A G | Dispersed ASH | • | The "0" of the natural game data is equal to the atnospheric background count (gama count when equipment was of 1434. This background was subtracted from the data. |

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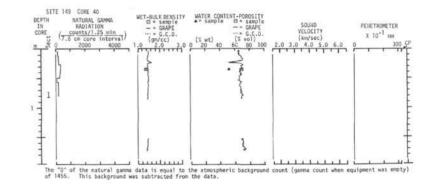
¹For explanation of symbols, see Chapter 1



| | | ZONE | | (| | | SAMPLE | ION | | CaCO ₃ (%) |
|---------------|---------|----------------------------|----------------------------|---------|--------|-----------|-----------|---|---|---|
| AGE | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAM | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 |
| MIDDLE EOCENE | No zone | Chiphragmalithus alatus | Thyrsocyrtis triacantha | 1 | 0.5 | V01D | | R A G N F F P F F P | RADIOLARIA NANNOPLANKTON CHALK; yellowish gray (5Y8/1), compact, and crumbly. | • |
| - | - | 5 | t - | | ORE | | | FFP RAG | pumice fragments | |



| - | - | ONE | _ | | | | SAMPLE | 110N | | | CaCO ₃ (%) |
|---------|-------|---------------------------|-------------------------|---------|--------|-----------|-----------|---|----------|---|---|
| FODAU | FORAM | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAP | PALEO SAMPLE ABUNDANCE PRESERVATION | LITH | IOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) 0 50 100 |
| a 1 - 1 | NO Z | ? Discoaster Sublodoensis | Thyrsocyrtis triacantha | 1 | 0.5 | | | R A G N F P F F P R A G N F F P | debris (| RADIOLARIA NANNOPLANKTON HALK; yellowish gray (SY7/2); semi-lithified and more friable zones interbedded in upper 20 cm of Section 1. | |

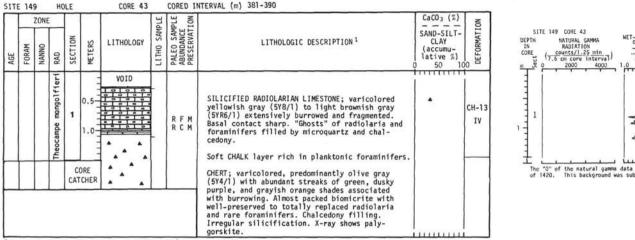


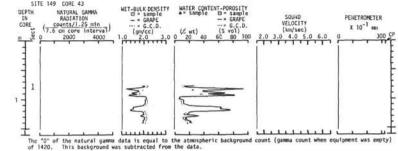
| STIC | | - | HUL | - | _ | CORE 41 | - | | TERTAL (m) SOL ST | | | |
|------|------|-------|-------------------------|---------|------------|--|--------------|---|---|-----------|-------------|--|
| AGE | Т | NANNO | RAD | SECTION | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | | DE FORMATIO | SITE 149 CORE 41 WET-BULK DENSITY MATER CONTENT-POROSITY DEPTH NATURAL GAMMA UP - Sample = Sample = Sample = Sample 1N RADIATION |
| ENE | zone | | Inyrsocyrtis triacantha | | | | | FFR PPG PPG PPG NGP G P FR FA FFA FFA FFA FFA FFA FFA FFA FFA FF | Interbedded SEMI-INDURATED CALCAREOUS-RICH RADIOLARIAN 00ZE and RADIOLARIA NANNOPLANKTON CHALK; pale yellowish gray (5Y8/2) to very pale grayish orange (10YR8/4) and grayish orange (10YR7/4). Sparse plagioclase, light volcanic glass, apatite, diatoms, and fish debris. In addition, X-ray diffraction shows palygorskite. Radiolarians coated with calcite. Sediment quite crumbly with some semi-lithified and more friable zones. Pumice scattered at several levels above volcanic ash of Section 6. Faint burrow mottling. | lative %) | 00 | $\frac{5}{3}$ |
| | | | | 6 C0 | 0.5 1.0 | \`\`\`\`\`\`\`\`\ \`\`\`\`\`\`\`\`\ | | F F P R A G R A G F F P R A G | | | | The "O" of the natural gamma data is equal to the atmospheric background count (gamma count when equipment was emp of 1455. This background was subtracted from the data. |

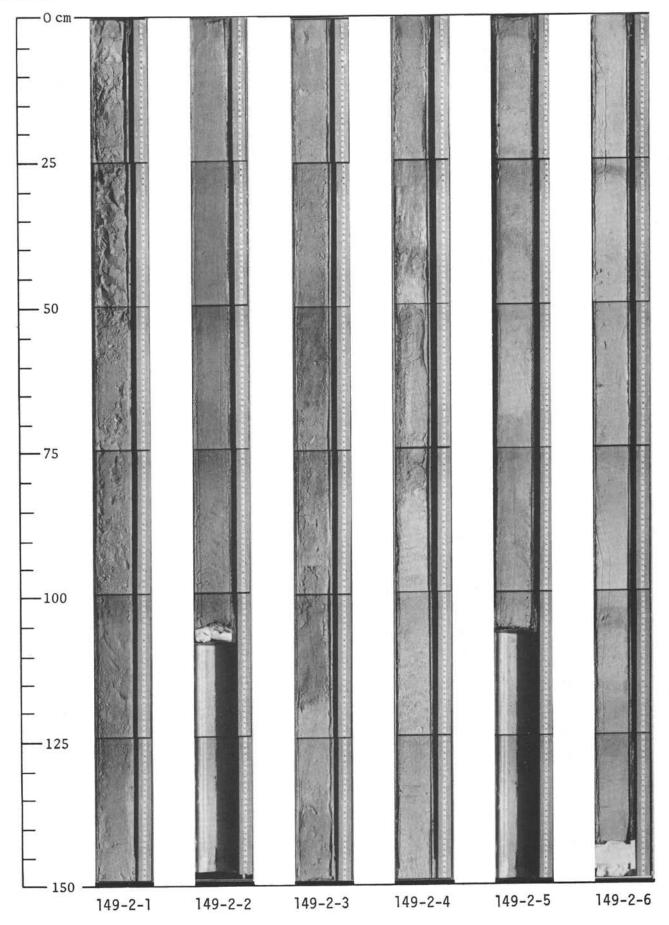
CORE 41 CORED INTERVAL (m) 362-371

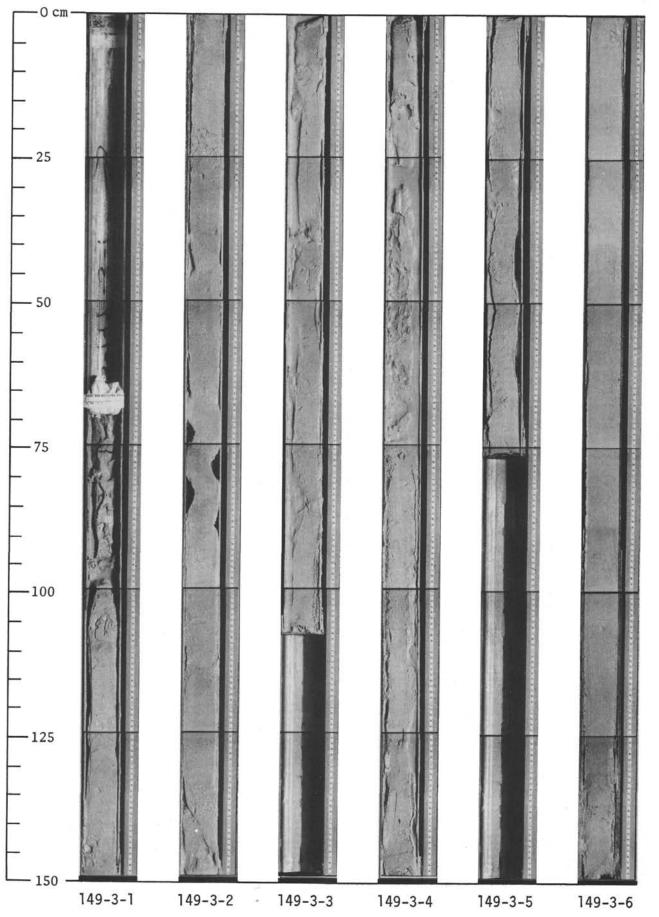
SITE 149 HOLE

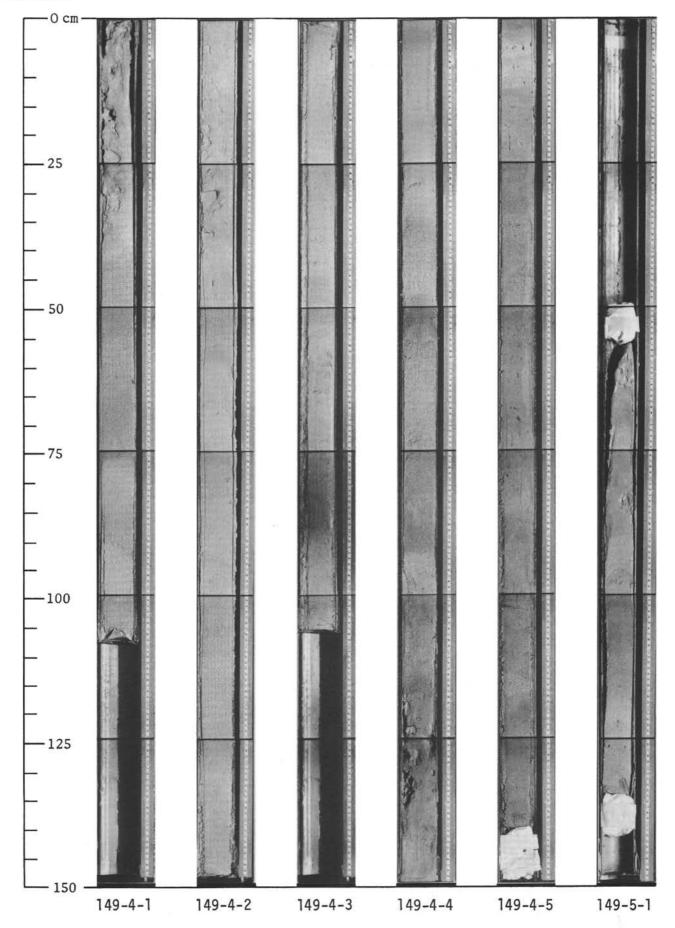
| ITE 149 ZONE | HOLE | T | CORE 42 | | | NTERVAL (m) 371-381 | CaCO ₃ (%) | | |
|-----------------|--------------|--|-----------|--------------|--|--|---|-------------------|--|
| × ç | | METERS | LITHOLOGY | LITHO SAMPLE | PALEO SAMPLE ABUNDANCE PRESERVATION | LITHOLOGIC DESCRIPTION 1 | SAND-SILT- CLAY (accumu- lative %) | DEFORMA | SITE 149 CORE 42 DEPTH NATURAL GAMAA WET-BULK DENSITY WATER CONTENT-POROSITY IN RODATION — GRAPE — GAVE Subit Stample Solution PERETROMETER CORE <u>conts/1.25 min</u> — GRAPE — GAVE VELOCITY ELECTROMETER (7.6 cm core interval) (9%, C) (5 wt) (1 v01) (kn/sec) X 10 ⁻¹ mm 0 2000 4000 1,0 (9%, C) (5 wt) (2 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 m 20 2000 4000 1,0 (9%, C) (5 wt) (2 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (2 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) (2 v03, 0 4, 0 5, 0 6, 0 0 300 M 20 2000 4000 1,0 (9%, C) (5 wt) (1 v01) |
| | Thyrsocyrtis | 0.5 1.0 0.5 1.0 0.5 1.0 0.5 1.0 0.5 0.5 0.5 0.5 0.5 0.5 0.5 0 | | - 4 | R A G F F P R A G F F P F F P F F P R A G F F P R A G F F P | DRILLING SAND of broken fragmen caused by drilling. Interbedded INDURATED RADIO- LARIA NANNOPLANKTON CHALK; yellowish gray (5Y8/1) to grayish orange (10YR7/2), crumbly. Some plagioclase, and sparse light volcanic glass and diatoms. In addition, X-ray shows palygorskite. Few pumice fragments in Section 3. sparse fish debris Sparse fish debris ASH layer disturbed by burrowin sharp basal contact CALCAREOUS-RICH RADIOLARIAN OOZE; carbonate increases downward. CHERT pebble; between light brownish gray (5YR6/1) and light olive gray (5Y6/1). | .5 | <u>CH-13</u> V | The "D" of the natural game data is equal to the attospheric background count (game count when equippent wes experi- |

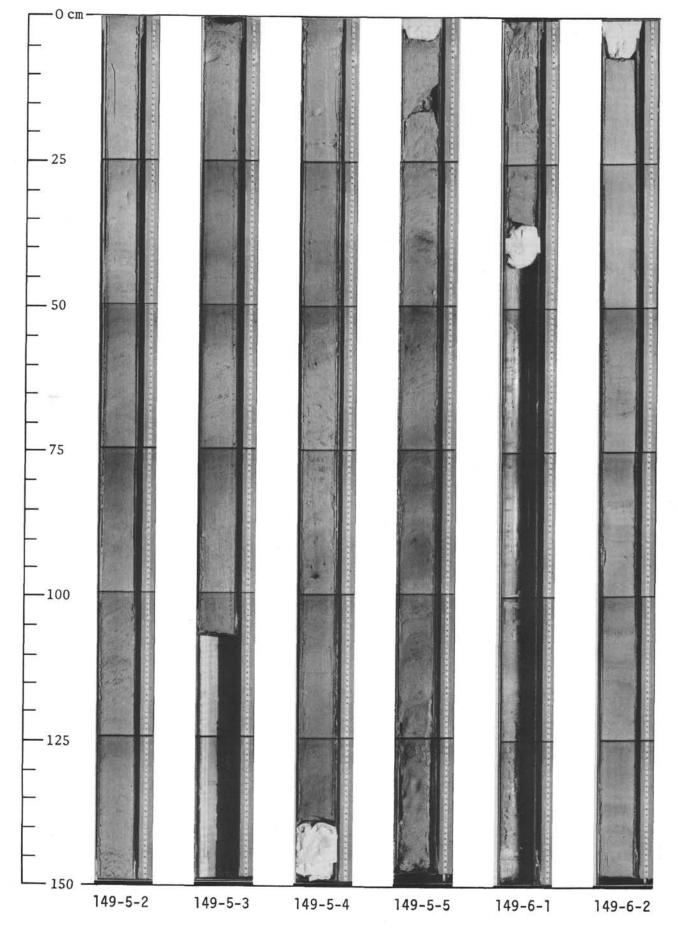


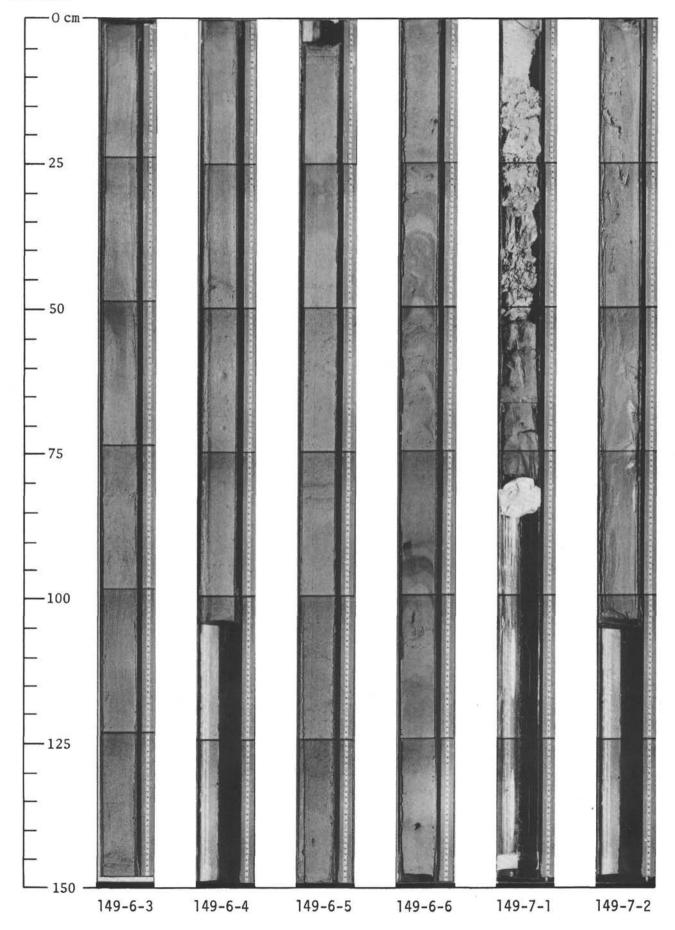


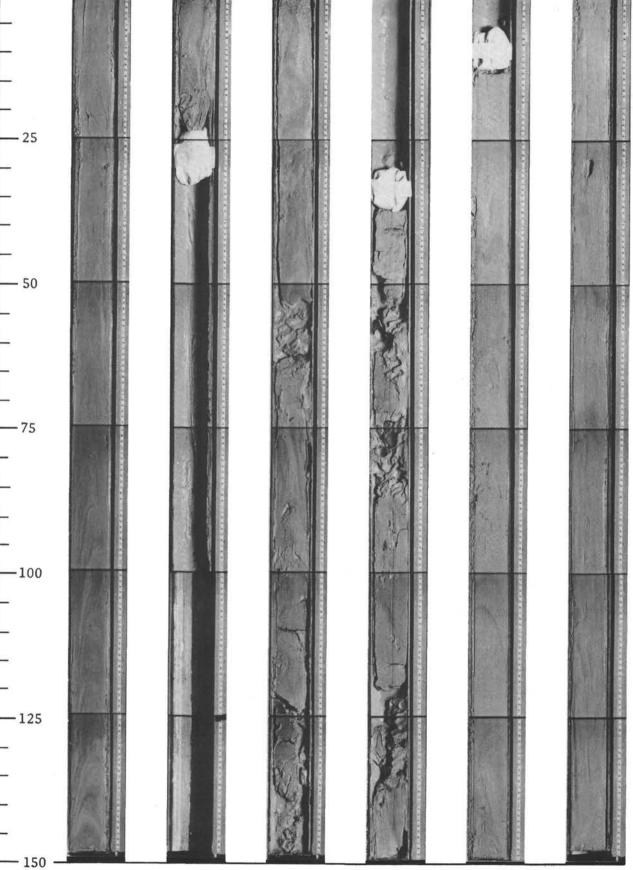












149-7-3

0 cm -

149-7-4

149-7-5 149-7-6

149-8-2

149-8-1

