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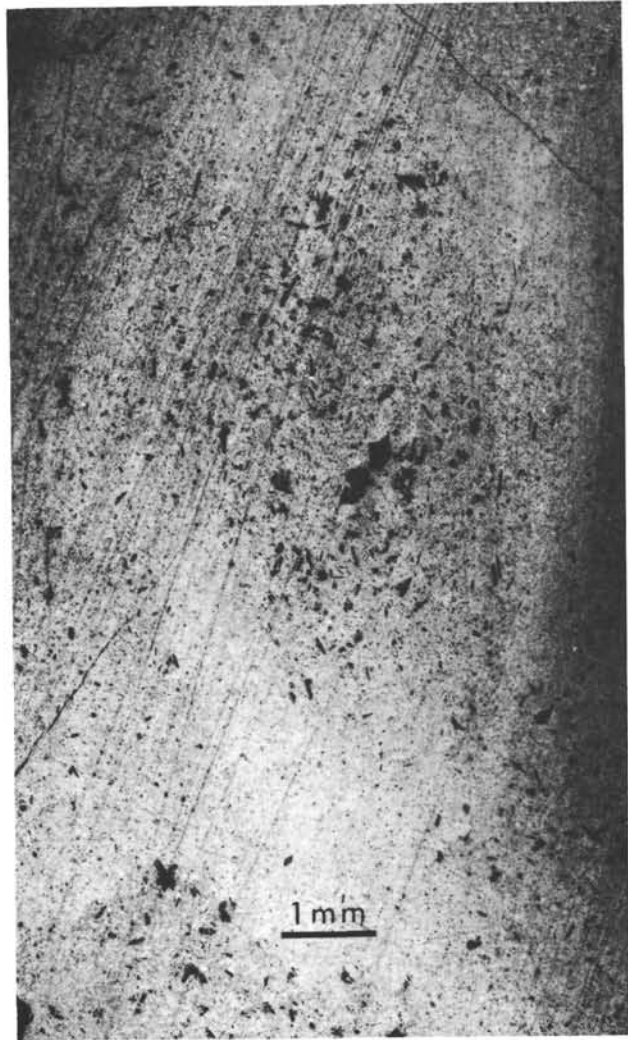


Figure 2. Reverse print of a metabasalt (spilite) from a drill bit sample. The dark patches are porphyritic and glomeroporphyritic clusters of plagioclase accompanied by chlorite pseudomorphs set in a very fine-grained matrix of acicular feldspar, chlorite, and alteration products. Relic crystal outlines suggest that olivine, now replaced by chlorite, was originally present.

26.2. PETROGRAPHY OF ROCKS FROM THE CRESTAL AREA OF THE GORRINGE BANK

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INTRODUCTION

Samples of igneous rock were recovered from the crestal area of Gorrige Bank in 1965 by coring and dredging from the R/V *Robert D. Conrad* of the Lamont-Doherty Geological Observatory. The sampling objective was to obtain materials which would shed light on the composition and geological history of this uplifted part of the Azores-Gibraltar seismic zone. Numerous boulders and rock fragments were obtained in one dredge, RC 9-7 (36°44.6'N, 11°05.8'W; water depth = 104-130 m), and two piston

cores, RC 9-206 (36°40.5'N, 11°04.4'W; water depth = 736 m) and RC 9-208 (36°42'N, 11°07.3'W; water depth = 104 m).

Much of the recovered material was badly weathered, and most of the exposed rock faces were encrusted with carbonate crusts including colonies of byrozoa and coral. Smaller pebbles were often rounded, probably indicating abrasion near an ancient surf-zone on the bank.

Twenty samples were studied in thin section with a petrographic microscope. The analyses are presented in Tables 1 and 2 according to the two igneous rock suites recognized.

TABLE 1
Oceanic Suite

Sample	Rock Type	Major Mineral Composition	Metamorphic Facies
Oceanic Crustal Rocks			
RC9-206-2	Metagabbro	Labradorite, hornblende	Amphibolite
RC9-206-3	Cataclastic metagabbro	Labradorite, diällage, iron oxides	Greenschist amphibolite
RC9-206-5	Cataclastic metagabbro	Labradorite, diällage, apatite, iron oxides	Greenschist amphibolite
RC9-206-6	Cataclastic metagabbro	Labradorite, diällage, apatite, iron oxides	Greenschist amphibolite
RC9-206-11	Cataclastic metagabbro	Labradorite, diällage, apatite, iron oxides	Greenschist amphibolite
RC9-206-13	Cataclastic metaanorthosite (gabbro fragment)	Labradorite, diällage	Greenschist amphibolite
RC9-208-1	Cataclastic metagabbro (anorthosite)	Labradorite	Amphibolite
Ultramafic (Mantle?) Rocks			
RC9-206-12	Amphibole-biotite pyroxene	Augite, aegirine-augite, brown hornblende, biotite, apatite, interstitial zeolites	Amphibolite
RC9-208-250 ₂	Pyroxene bearing hornblendite	Brown hornblende, augite, interstitial apatite, sphene, zeolite(?)	Amphibolite

OCEANIC SUITE

The first group of rocks (7 of 20 specimens) is characterized by cataclastic metagabbro which is composed of primary labradorite, diällage, and iron oxides with varying amounts of secondary amphibole (blue green hornblende, and/or actinolite), chlorite, and in some samples, quartz and epidote. Most of the metagabbro samples exhibit cataclastic textures. These metagabbro samples are similar to the metagabbro sample cored at Site 120 on the north flank of Goringe Bank (described in Chapter 26.1) as well as the metagabbros recovered from the bounding walls of transform faults (see Miyashiro *et al.*, 1971), and they are thought to compose the oceanic crust (Layer 3).

Two of the twenty samples are ultramafic specimens. One is an amphibole-biotite pyroxenite characterized by augite, aegirine-augite, brown hornblende, biotite, apatite, and interstitial zeolites. The other is a pyroxene-bearing hornblendite (Figure 1). The occurrence of actinolite and blue green hornblende, which mantle the primary brown

TABLE 2
Alkaline Suite^a

Sample	Rock Type	Major Mineral Composition	Metamorphic Facies
Alkaline Intermediate Series			
RC9-208-2	Aegirine phonolite	Aegirine, sanidine, nepheline, analcite, iron oxides	Zeolite
RC9-208-4	Aegirine phonolite	Aegirine, sanidine, nepheline, zeolites	Zeolite
RC9-208-5	Sanidine phonolite	Sanidine, nepheline, aegirine, nosean, iron oxides, zeolites	Zeolite
RC9-208-6	Nosean phonolite	Nosean(?), sanidine, nepheline, aegirine, zeolites, andesite	Zeolite
RC9-208-8	Sanidine phonolite	Sanidine, nepheline, aegirine, analcite, zeolites	Zeolite
RC9-208-250 ₄	Sanidine-nosean phonolite	Sanidine, nosean, nepheline, aegirine, analcite(?), zeolites	Zeolite
Alkaline Mafic Series			
RC9-208-3	Mafic phonolite	Sanidine, analcite, nosean(?), aegirine-augite, aegirine, zeolites	Zeolite
RC9-206-1	Mafic phonolite	Aegirine, sanidine, nepheline, nosean, zeolite, kaolinite	Zeolite
RC9-208-250 ₁	Mafic phonolite	Nosean, aegirine, nepheline, sanidine, biotite, anorthoclase	Zeolite
RC9-208-250 ₃	Welded tuff (or devitrified lava)	Nosean, hornblende, biotite, pyroxene, leucite(?), aegirine-augite, aegirine, sanidine, anorthoclase, glassy matrix	Zeolite
RC9-206-7	Alkalic metabasalt	Altered olivine, augite, plagioclase, interstitial feldspatoid (nosean:), melilite(?)	Zeolite

^a>10 per cent feldspathoids.

hornblende and pyroxene, indicates amphibolite facies metamorphism; the occurrence of zeolite in veins and interstitial zeolite suggests later stage retrogressive zeolite metamorphism. Ultramafic rocks of this type have been recovered in dredge hauls from the base of transform fault escarpments (Dmitriev *et al.*, 1971; Bonatti *et al.*, 1971). The ultramafic rocks recovered from transform faults are interpreted as representing the lower part of oceanic crust or upper mantle(?).

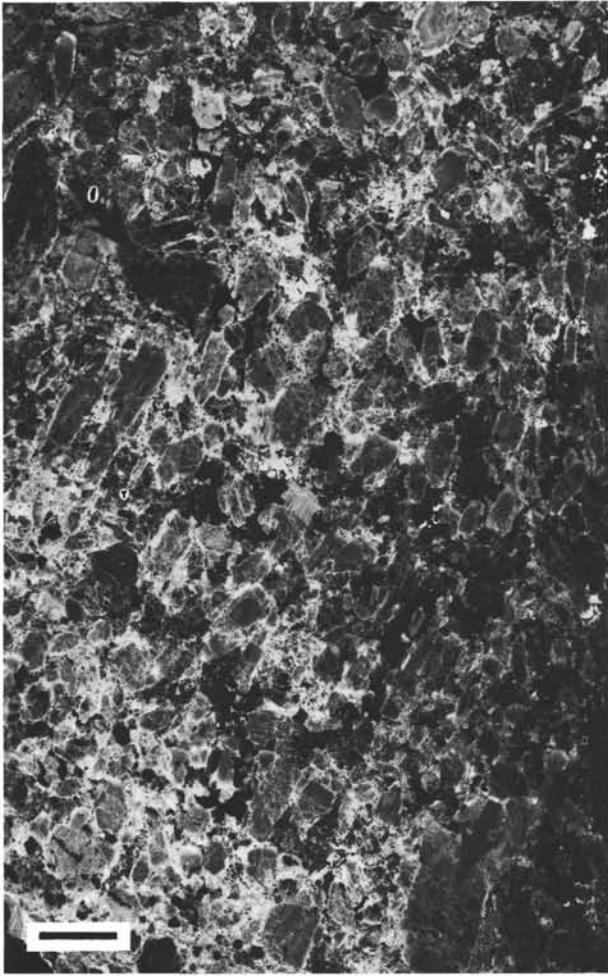


Figure 1. Reverse print enlargement of a thin section of pyroxene bearing hornblendite (RC 9-208-250₂); revealing a medium-grained fabric and brecciated appearance. The larger crystals (arrows) are brown amphibole of an edenite composition. Scale bar represents 1 millimeter.

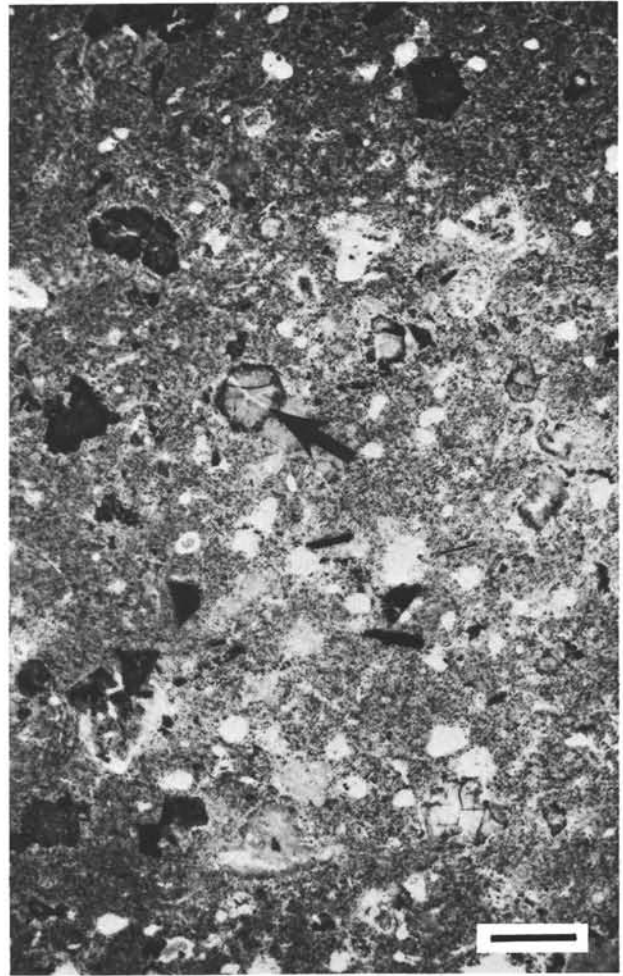


Figure 2. Reverse print enlargement of a thin section of mafic phonolite (RC 9-208-250₁), containing large crystals of nosean (arrow) and numerous thin laths of sanidine. Many of the more mafic minerals are replaced by dark green pyroxene and iron oxides. Biotite is also replaced and has inclusions of zircon without a pleocroic halo. Scale bar represents 1 millimeter.

ALKALINE SUITE

The second group of samples (11 of 20 specimens) is a suite of silica deficient, alkaline rocks (Figure 2) and is characterized by the occurrence of pyroxene (aegirine), feldspar (sanidine), and feldspathoids (nepheline and/or nosean) as the principal minerals and, depending on the sample, lesser amounts of analcite, hornblende, pyroxene (augite, or aegirine-augite), and iron oxides. The occurrence of zeolite in the matrix and in veins in all the samples suggests zeolite facies metamorphism. The occurrence of such undersaturated alkaline rocks in the cores and a dredge located on the crest of Gorringle Bank is unusual because this igneous rock type (phonolite) is not characteristic of the rock types typically recovered from the ocean floor (dredge hauls invariably recover one or a combination of the following rock types: tholeiitic basalt, alkaline basalt, metabasalt (zeolite and greenschist facies), gabbro, metagabbro (greenschist and amphibolite facies), and serpen-

tinized periodotite. Although phonolite rocks occur in minor amounts in a late stage differentiate in association with alkaline basalts on some volcanic oceanic islands (e.g., Ascension and St. Helena), phonolites are more generally found in association with igneous bodies located behind subducting zones (e.g., Andes).

The occurrence of metagabbros in bottom samples from the crest of Gorringle Bank, a part of the Azores-Gibraltar Ridge, is not surprising and in fact these rocks suggest that the oceanic crust (Layer 3) is exposed along a portion of Gorringle Bank. The occurrence of a suite of intermediate and mafic alkaline rocks composed of feldspathoids, however, is not similar to rocks previously sampled from the ocean basin, and this suite is certainly not compatible with the rock types thought to be generated along accreting plate margins of the Mid-Oceanic Ridge System. Rather,

these feldspathoid rich rocks are characteristic of igneous rocks found behind some subducting margins. Although the feldspathoidal phonolites could represent later stage differentiates associated with an alkaline basalt seamount on the Azores-Gibraltar Ridge, we interpret the occurrence of the feldspathoidal phonolites as reflecting a short and rapid subducting phase which accompanied the deformation of the Horseshoe Abyssal Plain sediments and uplift of the Goringe Bank. As to the age of the intrusions, we have noted welded tuffs containing xenoliths of mafic phonolite in direct association with baked skeletal debris of former neritic fauna. This evidence would place the intrusion of the phonolite at a time when the bank was already shallow. According to the facies sequence of pelagic sediment recovered at Site 120, the transition there from a deep oceanic sea bed commenced during or after the Early Miocene, and was essentially completed by the Early Pliocene.

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