

5. SITES 102-103-104 – BLAKE-BAHAMA OUTER RIDGE (NORTHERN END)

The Shipboard Scientific Party¹

INTRODUCTION

Geophysicists studying the deep-sea area off the Blake Plateau during the late 1950's concluded that the Blake-Bahama Outer Ridge is a foundered chain of islands and reefs structurally connected to the Cape Fear Arch off the Carolinas, (Hersey *et al.*, 1959). During the early 1960's scientists of the Lamont-Doherty Geological Observatory made continuous seismic reflection profiles over the ridge and showed that it is a vast accumulation of sediments lying on a horizontal reflecting surface (Layer A) which in turn lies above the rough acoustic basement reflector (Ewing and Ewing, 1964). This discovery led to a number of concepts attempting to explain the processes of formation of such huge sediment bodies in the deep sea.

Ewing *et al.* (1966) suggested that much of the material forming the ridge had been eroded from the Blake Plateau by the Gulf Stream. Andrews (1967) suggested that the ridge was formed by an enormous slump from a proximal up-slope position, in this case from the Blake Plateau region.

Heezen and Hollister (1964) and Heezen *et al.* (1966) proposed that the ridge, as well as many other fine-grained sediment bodies in the deep sea, owes its shape and structure to the effects of mud-carrying, contour-following deep currents ("contour currents") associated with deep thermohaline circulation.

Bryan (1970) and Markl *et al.* (1970) combined these concepts, suggesting that interaction of turbid flow from the Blake Plateau with the southward-flowing contour current initiated the development of the ridge which subsequently was enlarged and perhaps somewhat reshaped by further deposition and erosion associated with the contour currents.

The Blake-Bahama Outer Ridge (Figure 1) includes deposits of several shapes and sizes, from dunes a few kilometers long and tens of meters in height, to ridges many hundreds of kilometers long and thousands of meters high (Markl *et al.*, 1970).

The principal objective of Sites 102, 103 and 104 was to determine the geologic history of the Blake-Bahama Outer Ridge by investigating its structure, composition, rate of growth, and the provenance of constituent sediments.

OPERATIONS

The ship arrived at Site 102 on the crest of the Blake-Bahama Outer Ridge during the late hours of April 30, and drilling was commenced at 1120 hours on May 1. A light-set diamond bit was chosen for this hole. The bottom hole assembly was the same as that used for most of the holes of this leg, consisting mainly of bit and outer core barrel, eight drill collars, and three bumper subs.

The sediment was soft and penetration progressed smoothly for the first three cores. Drill string rotation and pumping were required below a depth of 100 meters, but there was no difficulty in penetration until a depth of about 620 meters was reached. At this depth, which correlates well with the prominent conformable reflector at 0.62 second below bottom in the profiler record (Figure 2a), the sediment became much firmer and drier, and penetration slowed severely. After more than twelve hours drilling in this zone with a penetration rate of only 3 m/hr, the hole was abandoned. There had been some speculation that the bit might have clogged or the seals were blown out of a bumper sub, rather than that the sediment caused difficult penetration; but when the string was recovered no equipment problem was found. Therefore, the sharp decrease in penetration rate must indicate a significant change in lithology at the depth of the prominent seismic reflector at 0.62 second. By comparing the data from this hole and Hole 103 (Figure 2b), it is obvious that the reflector is not a time boundary and, therefore, must result from diagenetic processes.

Most of the cores at this site contained substantial amounts of organic material and were gassy, usually partially extruding themselves from the plastic liners and creating gaps in the remaining sections. Chromatograph measurements indicate that the gas is principally methane with only a small fraction (usually less than 0.5 per cent) of ethane.

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Hole 103 was drilled on the southwest flank of the Blake-Bahama Outer Ridge at a location where seismic

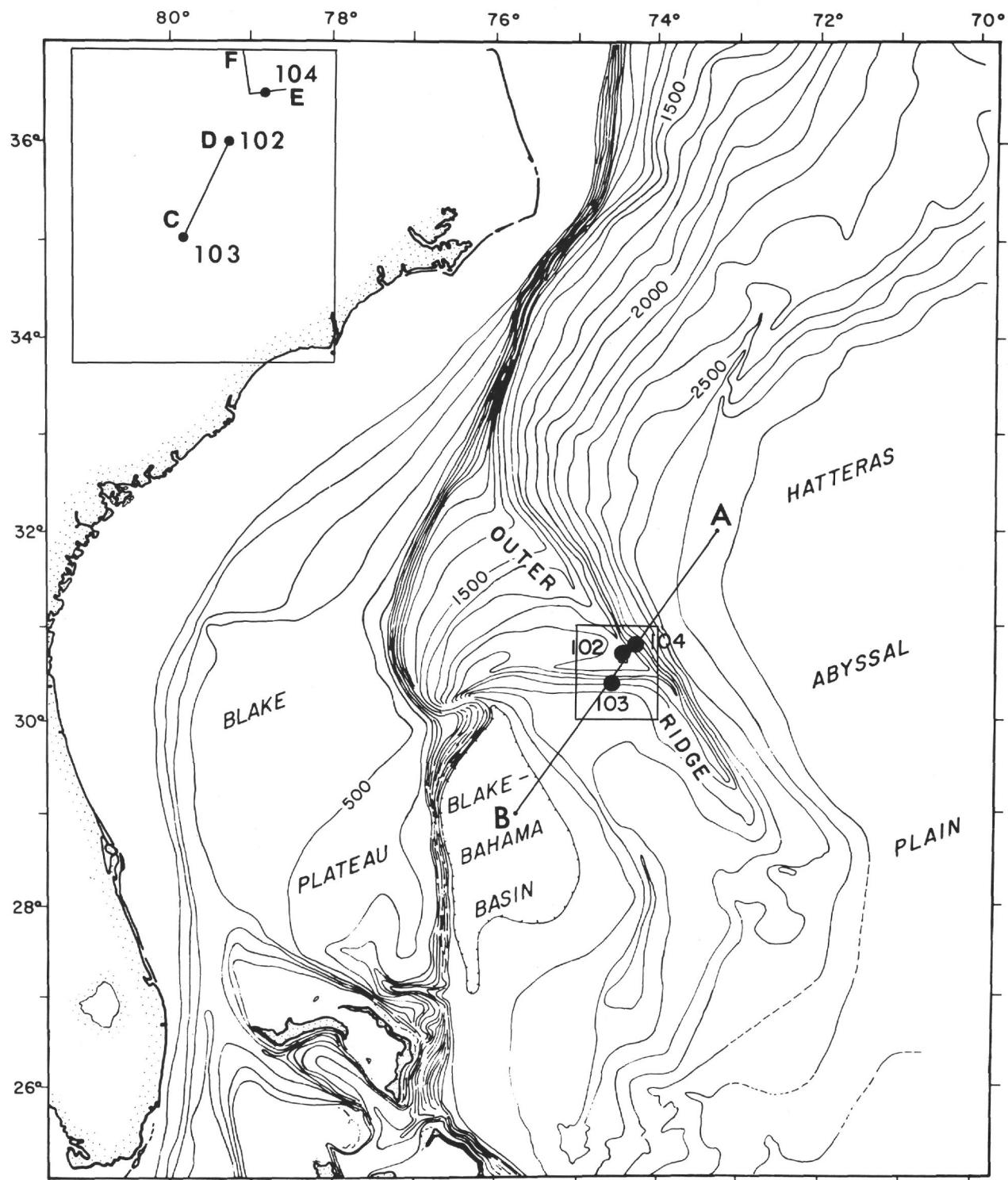


Figure 1. Bathymetry of the Blake-Bahama Outer Ridge system (after Mankl, et al, 1970). Track segments locate seismic profiler sections shown in Figures 2a, 2b and 2c.

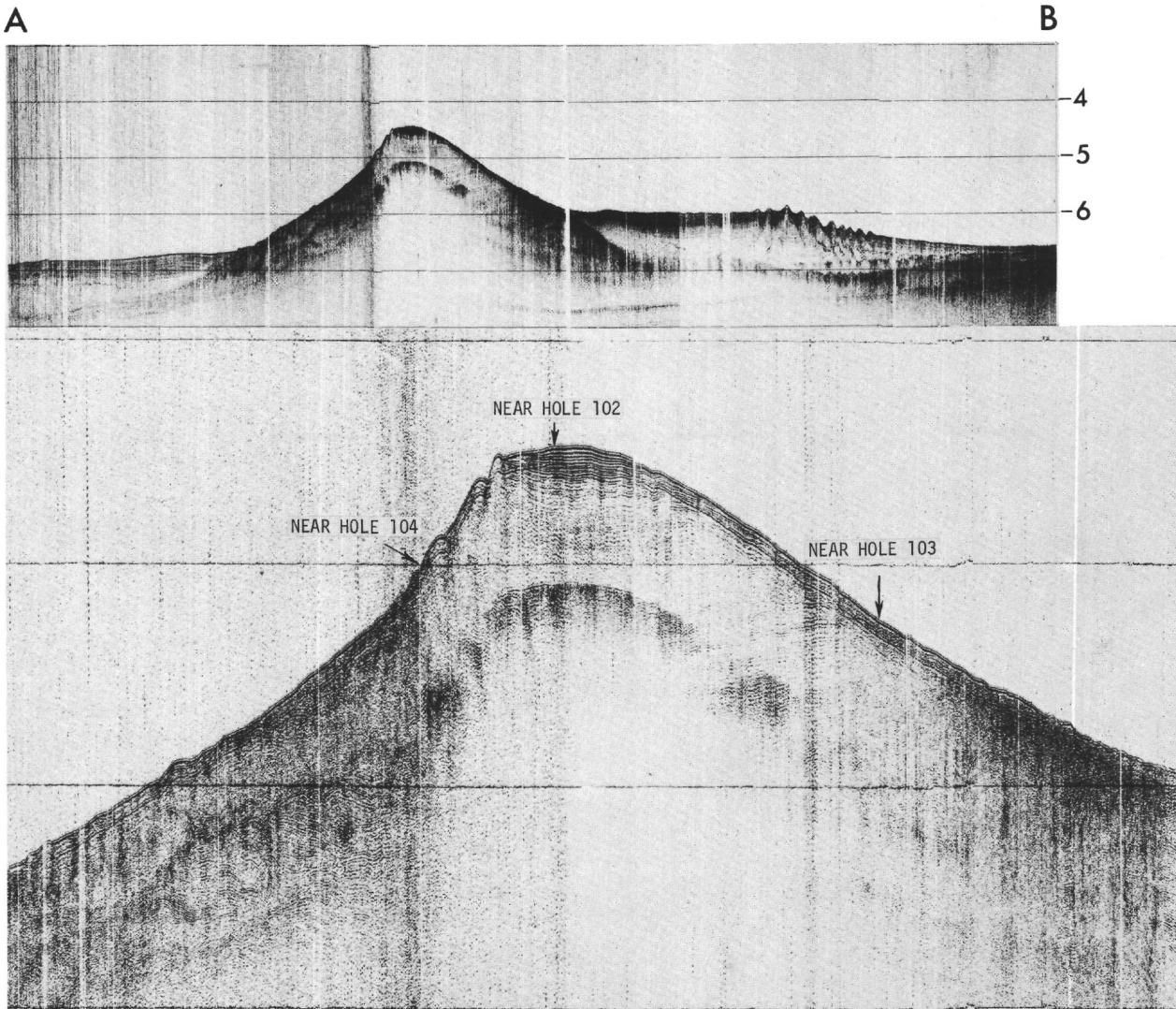


Figure 2a. Vema 21 Seismic profiler section across the Blake-Bahama Outer Ridge, (crestal region expanded below). See Figure 1 for location.

profiler data indicated an outcropping of the deeper beds penetrated in Hole 102 (Figure 2b). Drilling commenced at 1400 hours on 5 May, 1970. The second core, taken at a subbottom depth of 40 meters, brought up late Miocene sediment which had not been reached in the crest hole until a depth of more than 500 meters.

After the difficulty experienced at Hole 102 in drilling the firmer sediment with the diamond bit, a drag bit was chosen to drill this hole. It appeared to perform well, and during the drilled interval above the deepest core, penetrated firm clay at a good rate. Unfortunately, after Core 7 was taken the hole began to show signs of serious collapse and had to be abandoned.

The sediment samples were very similar to those at Hole 102, consisting mainly of silty clay with appreciable amounts of organic material. All of the cores

contained a substantial amount of gas, consisting primarily of methane with a trace of ethane.

After the collapse of Hole 103, enough time remained for about 36 hours more drilling, and a decision was made to drill a third hole (Hole 104) on the northeast flank of the ridge (Figure 2c) at approximately the same water depth as Hole 103, drilled on the southwest flank. Drilling commenced at 1400 hours on 7 May, 1970, in a water depth of 3821 meters. The 3-cone carbide button roller bit used for Hole 99A was employed again.

After only 3 or 4 meters of penetration, several hard, apparently thin beds were encountered, requiring bit rotation and a substantial amount of weight. The hard material, some of which was recovered in Core 1, appears as a finely laminated calcite "crust" covered with calcite-filled burrow-like structures.

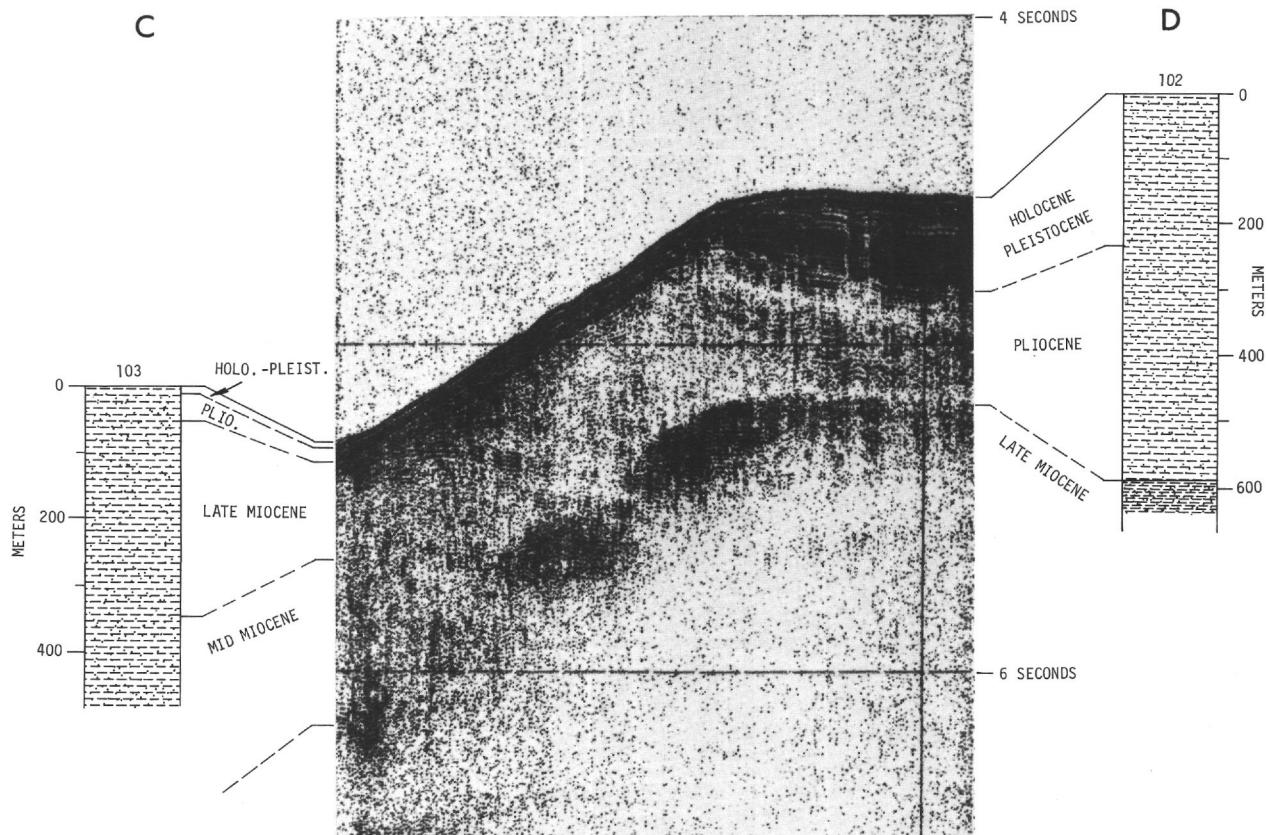


Figure 2b. *Lithology and seismic stratigraphy at Site 102 (right) and Site 103 (left).*

Most of the cored section consisted of hemipelagic silty mud similar to that found in Holes 102 and 103. The deepest core, taken at a depth corresponding to that of the prominent conformable reflector at 0.62 second (Figure 2c) contained approximately 15 centimeters of very hard ankerite. This layer or nodule required several minutes to drill through and was as hard drilling as any of the cherts, limestones, or basalts cored in previous holes during this leg. The material underneath the ankerite appeared to be little different from that above it. Unfortunately, time had run out at this point, and drilling was terminated in order for the ship to make its scheduled arrival at Norfolk.

The situation with respect to gas in the sediments was about the same as at Holes 102 and 103. The chromatograph showed only a trace of ethane. The first three cores of Hole 104 had a strong odor of hydrogen sulfide (H_2S), but this diminished rapidly in the deeper part of the section.

STRATIGRAPHY

Biostratigraphy

Assemblages of abundant calcareous nannoplankton, planktonic foraminifera, dinoflagellates, Radiolaria,

and diatoms of Holocene to early middle Miocene age were encountered at these sites. Approximately 600 meters of middle Miocene, 300 meters of late Miocene, 350 meters of Pliocene and 220 meters of Pleistocene sediments are present.

The foraminiferal and nannoplankton assemblages contain the same tropical/subtropical components that have been observed in worldwide equatorial latitudes. They correspond well with the biostratigraphic zones established by Blow (1969), Martini and Worsley (1970), and the nannofossil zonation of Milow used on Leg 9. The composite section of Sites 102, 103 and 104 contains each of Zones N. 23 to N. 10/9 as defined by Blow (1969).

Two major stratigraphic boundaries (according to foraminiferal data) are documented; the Miocene-Pliocene in Core 2 at Site 103, and Pliocene-Pleistocene in Core 8 at Site 102. In addition, three major unconformities are present; one at Site 104 (Core 1), and two at Site 103 (within Core 1).

Pollen grains and spores occur in all the examined samples. Pleistocene grains consist mainly of pine,

— 4 SECONDS

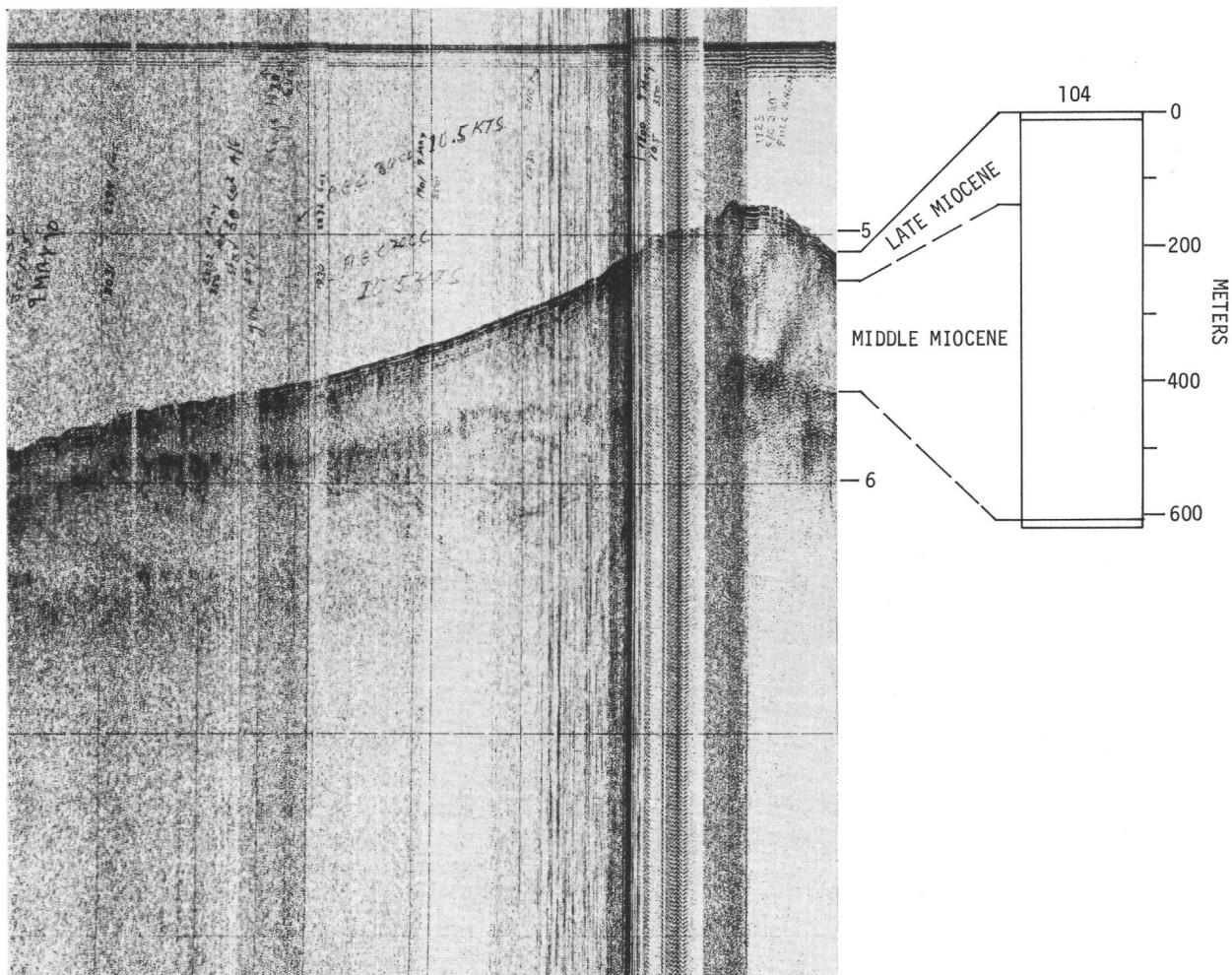


Figure 2c. Lithology and seismic stratigraphy at Site 104.

scattered spruce and pteridophyte spores, which suggests transportation from a northerly source. Reworked grains of Cretaceous and Carboniferous ages are also present. In the upper and middle Miocene sediments, the pollen assemblages are more varied, containing appreciable amounts of oak pollen, grasses, sedges, alder and hickory; this suggests either a climate warmer than the Pleistocene in the same northerly direction, or increased stream discharge and marine currents, or change in source area. Reworked grains of Cretaceous age are also more common, some with excellent preservation, which suggests possible contributions from the Atlantic Coastal Plain or deep-sea deposits. The foraminifers offer little evidence with regard to source of sediment, except that some sublittoral forms are usually minor constituents of the foraminiferal assemblages. This indicates displacement from shallow water.

Foraminifera

Refer to the chapter titled "Neogene Foraminiferal Biostratigraphy" for a detailed analysis of the planktonic foraminiferal assemblages at Sites 102, 103 and 104.

Calcareous Nannoplankton

Coring at Site 102 recovered sediment ranging in age from late Miocene to Quaternary. The first core, from 9 meters, contained a typical Pleistocene nannoplankton assemblage. The second core, from 27 meters, is also Pleistocene and represents perhaps a glacial interval judging from the common to abundant concentration of large specimens of *Coccolithus pelagicus*, a species that has a definite cold-water preference in modern oceans. Cores 3 to 8 were likewise assigned to

the Quaternary. The core catcher of Core 9 contains a few specimens of *Discoaster brouweri* and thus, was, assigned to the latest Pliocene N. 21 Zone of Blow. Core 10 contains an assemblage indicative of the lower part of N. 21. Core 12 was placed within Zone N. 20. Cores 13, 14 and 15 were assigned to the early Pliocene Zone N. 19, and Core 16 to either very low in Zone N. 19 or possibly to the upper part of Zone N. 18. Cores 17, 18 and 19 were considered to be late Miocene, Zone N. 18.

Coring at Site 103 recovered sediments ranging in age from Holocene at the sea floor to middle Miocene at a depth of 449 meters. Core 1 extends from the surface to a depth of 8 meters, and the bottom of the core contains an early Pliocene assemblage. In superposed sequence, Cores 2 through 7 recovered nannoplankton assemblages of late Miocene in Core 2 to middle Miocene in Core 7. In all but the lower two cores, nannoplankton are present in abundance and are well preserved. Core 6 contains only a few forms, and in Core 7 they are very rare. In these two cores solution effects are commonly observed on the foraminifera.

Cores from Site 104 recovered sediments as young as Quaternary at the top and as old as earliest middle Miocene at a depth of 617 meters. The top of the first core contains a Quaternary assemblage, but the sample from the core catcher was assigned to the late Miocene (probably Zone N. 17 of Blow). The remaining cores (2 through 10) revealed a gradual increase in age through the middle Miocene, terminating in the earliest middle Miocene *Sphenolithus heteromorphus* Zone (N. 9/10).

Dinoflagellates

Dinoflagellate cysts were recovered from all samples examined at Site 102. *Opercudinium centrocarpum* (Deflandre and Cookson), *Tectatodinium pellitum* Wall, and species of *Spiniferites* are common in the Quaternary assemblages of Cores 1 through 6. *Achomosphaera ramulifera* (Deflandre) occurs in the Pliocene cores, in addition to those species found also in the Quaternary. At Site 103, *Hystrichosphaeropsis obscurum* new species was found in the cores of Miocene age. The Miocene cores contain a diversified assemblage of pollen grains, including those assigned to oak, hickory, Compositae and grasses. Reworked grains are scattered through the samples, including those of Cretaceous and Carboniferous ages.

Hystrichosphaeropsis obscurum occurs in the Miocene cores of Site 104. *Pentadinium taeniagerum* Gerlach occurs in Core 10 at Site 104.

Rate of Sediment Accumulation

During late Miocene time, sediment at Site 102 accumulated at a rate of 2.5 cm/1000 yr. and increased to 13.7 cm/1000 yr. during the Pleistocene.

At Site 103 approximately 200 meters of late middle and early late Miocene sediment accumulated at a rate of 19.3 cm/1000 yr., and continued at this rate to the end of the Miocene. Rates through the early and middle Pliocene were about 2.6 cm/1000 yr.

At Site 104 middle Miocene sediments accumulated at a rate (19.5 cm/1000 yr.) similar to that calculated for the late middle and early late Miocene at Site 103. During the remainder of late Miocene time, a considerable drop in sedimentation rate (1.8 cm/1000 yr.) occurred at Site 104. In other words, during the late middle and early late Miocene the rate of accumulation at Site 103 (west flank of ridge) was approximately twice that at Site 104 (east flank of ridge). This suggests perhaps that the greatest amount of sediment was being supplied to the southwestern portion of the ridge during late Miocene time.

Lithology

At Sites 102, 103 and 104—on the crest (3425 meters), southwest flank (3964 meters) and northeast flank (3811 meters), respectively, of the Blake-Bahama Outer Ridge—a total of 1727 meters of dark greenish-gray hemipelagic silty mud was penetrated; 192 meters of this material were recovered in 36 cores. Sediment age ranges from middle Miocene to Holocene. All three sites have in common the late Miocene foraminiferal biostratigraphic Zone N. 17 as well as correlatable sediment types as determined by smear slide analysis. They will be treated together as one sequence representing the uppermost kilometer of ridge crest sediment.

Three main variations in the hemipelagic sediments are distinguished (1) siliceous, calcareous, and heavy mineral-rich Pleistocene sediments deposited at a rate of about 14 cm/1000 yr., (2) non-siliceous, siderite-pyrite-glaucnionite-rich Pliocene and late Miocene sediments deposited at a rate of about 9 cm/1000 yr., (3) siliceous, noncalcareous, siderite-rich middle Miocene sediments deposited at a rate of about 19 cm/1000 yr.

Pleistocene Sediments (N. 22 - N. 23)

Approximately 220 meters of the Pleistocene was penetrated at Site 102; 72 meters of this were cored and 52 meters were recovered. The accumulation rate of almost 14 cm/1000 yr. is extremely high for a deep-sea environment, which is rarely over 5 cm/1000 yr. Less than 5 meters of the surface cores at Sites 103 and 104 were Pleistocene and, thus, the following discussion deals entirely with material recovered in Cores 1 through 8 at Site 102.

The upper 10 meters of sediment at Site 102 resemble the sediment recovered in piston cores taken on the

crest of the Blake-Bahama Outer Ridge by Lamont-Doherty scientists (Heezen *et al.*, 1966). The upper half meter is Holocene yellow-brown and gray, laminated, silty, foraminiferal ooze. Smear slide examinations reveal an abundance of clay minerals, inorganic silt-sized detritus and organic debris, foraminifera and coccoliths, aragonite needles, and a flood of radiolarians, diatoms, and sponge spicules. Siliceous microfossils are almost totally lacking in the underlying Pliocene sediment. The abundant inorganic detrital silt laminae contain abundant heavy minerals.

Below the first few meters, the color becomes dark gray to grayish-brown; black patches of iron sulfide become quite abundant. Methane gas with traces of ethane becomes detectable below about 10 meters below bottom. Below a depth of about 100 meters, a few layers of rose-gray and mottled yellow-green silty clay occur. Siderite, dolomite and aragonite are found sporadically, and burrows are often filled with inorganic light-colored quartz silt.

A few hard, well-indurated crusts of limy mud containing cemented wormlike burrows and burrow-fillings were found at Site 104. Rather large fragments of hard lithographic limestone were found immediately below the crust. The lower limit of the fragments coincides with a major unconformity that separates the early Pleistocene from the late Miocene. Similar crusty surfaces have been found in modern sediments where strong bottom currents scour the sea floor thereby temporarily preventing deposition. It is the absence of Pliocene sediments at Site 104 which indicates recent periods of scouring or nondeposition on the east flank of the Blake-Bahama Outer Ridge.

Pliocene—Late Miocene Sediment (N. 21 – N. 15)

At Sites 102, 103 and 104, a total of 687 meters of Pliocene (N. 18/19 – N 21) to late Miocene (N. 14/15) sediment was penetrated; 83 meters of this was recovered in 18 cores. This late Miocene sediment is very similar to the Pliocene material and thus the two will be treated as one lithologic unit. It is homogeneous, dark greenish-gray, silty hemipelagic mud, occasionally mottled with lighter olive-gray hues.

The distinction between Pleistocene and Pliocene-late Miocene sediments lies in minor mineralogic differences revealed by smear-slide analysis and an abundance of siliceous material. Within the Pleistocene, siderite, pyrite and glauconite make their first consistent appearance. They increase in amount with depth in the section. Many of the burrows are found to be filled with siderite silt. Aragonite needles and heavy minerals, common in the Pleistocene, become rare to absent in the older sediment. In the Pleistocene deposits radiolarians, sponge spicules and diatoms are very abundant, whereas in the Pliocene-late Miocene they are rare.

A significant change in mineralogy and palynology also occurs near the boundary between the late Miocene and middle Miocene. This boundary, penetrated at Sites 103 and 104, was not reached at Site 102 and thus the following discussion is based on smear slides obtained from these two correlatable sites.

Middle Miocene Sediment (N. 14 - N. 9)

A change at the late-middle Miocene boundary is a gradual decrease in abundance of heavy minerals. Foraminiferal tests and nannoplankton are almost totally lacking in smear slides of middle Miocene sediment. However, their decrease has been rather gradual since about middle Pliocene, and may reflect progressive dissolution of carbonate as a function of time and/or depth of burial. At the bottom of the oldest sediment recovered (Core 10, Site 104, Zone N. 10/9), a bed or nodule of ankerite was cored.

Below the middle Miocene-late Miocene boundary an increase in abundance of siliceous organisms occurs. Also, near this boundary an abrupt change in pollen and spore assemblages is noted. Pleistocene and Pliocene sediment contains spruce and fir pollen that appears to have been transported from northerly sources, whereas the middle Miocene sediment contains more temperate forms such as alder, hickory, oak, grasses and sedges that may have been derived from more proximal sources and thus have a more restricted transportational history.

DISCUSSION AND CONCLUSIONS

Cores recovered from these sites can be correlated on the basis of one common biostratigraphic zone (N. 17) as well as on the basis of similar sediment composition. Zone N. 16 is present in Holes 102 and 103. The nature of the sediments and the attitude of reflectors recorded on seismic reflection profiles (Markl *et al.*, 1970) suggest that at least the upper 600 meters of the northern end of the Blake-Bahama Outer Ridge is a large accumulation of rapidly-deposited, hemipelagic, silty carbonaceous clay of Pliocene-Pleistocene age that has been transported to this region from the north by southerly-flowing bottom currents. Erosion has occurred along both flanks of the ridge. The underlying 600 meters of late and middle Miocene hemipelagic sediment recovered at Site 104 appear to have been derived from nearer sources. The upper 150 meters of sediment recovered at Site 102 appears to have been deposited as a conformable lens of Pleistocene hemipelagic silty clay on top of the Pliocene and middle Miocene accumulations.

The western flank of the ridge in the vicinity of Site 103 represents a continuation of the straight, east-west oriented slope running from the vicinity of "Nose" of the Blake Plateau to the ridge crest at approximately

30° 30' N. The unconformities apparent in the seismic profile near Site 103 must be related to the same processes that have formed this east-west linear slope. The unconformities and disturbed reflectors seen in profiles over the eastern flank of the Blake-Bahama Outer Ridge in the vicinity of Site 104 indicate erosion.

As discussed in a previous chapter, the Tertiary sediments at Site 101 at the southern end of the Blake-Bahama Outer Ridge have been transported from the north by southerly-flowing currents and, subsequently, penecontemporaneously sculpted and shaped by the Western Boundary Undercurrent (Heezen, Hollister and Ruddiman, 1966).

At Hole 102, the drilling-rate graph shows only one important break — that at a depth of 620 meters. In the appropriate region of the seismic record, there is only one reflector that could reasonably correlate with the drilling break; that is, the one at 0.62 second which is conformable to the topography and appears to cut across bedding planes. Such a correlation indicates an average sound velocity of 2.0 km/sec in the upper 600 meters of section. It is possible to divide the section into two layers and compute average velocities of 1.9 and 2.3 km/sec, respectively—as shown in the drilling-rate graph (Figure 3).

Drilling at Hole 103 did not achieve sufficient penetration to permit a velocity determination to be made.

At Hole 104, penetration was made at a uniform rate down to 615 meters where a very hard layer, or nodule, of ankerite was cored. Unfortunately, there was not enough time to drill deeper, so we cannot be certain whether the ankerite was an isolated nodule or part of an ankerite-rich zone. In any case, this drilling break is the only observed event to correlate with the reflector at 0.61 second. The correlation indicates an average sound velocity of 2.16 km/sec for the upper 600 meters, a value rather close to that computed for the section at Hole 102.

A sound velocity greater than 2 km/sec seems to be somewhat high for the hemipelagic mud cored at these sites, particularly when we consider that most of the material contained significant amounts of methane. However, if we assume a lower velocity for this material, we are faced with having to accept the fact that we drilled through the prominent reflector without noticing it and then encountered a significant drilling break underneath the reflector. This may indicate that it is the homogeneous zone beneath the reflector that resisted drilling, and the reflectivity of the zone just above it is due to factors not apparent in the normal examination of the cores.

Alternatively, we can accept the interval velocity of about 2 km/sec as a valid measurement and attribute the rather high velocity to the presence of the methane. Stoll *et al.* (1971) have pointed out that under usual deep sea conditions, the pressure and temperature in the uppermost few hundred meters of sediment are such that methane would exist as a hydrate—an ice-like substance. They further made laboratory measurements in which the speed of sound in water-saturated sand was observed to increase from about 1.7 km/sec to more than 2.5 km/sec after methane had been bubbled through the sample for several hours at 2° to 3° C and 800 to 1000 psi pressure. Although this may account for the abnormal sound velocity, this is not an obvious explanation of the reflector itself. The fact that it is so conformable to the sea floor and cuts across time lines ranging from mid-Miocene to Quaternary suggests that it is caused by a diagenetic process with a short time constant. It seems reasonable to assume that temperature would change much more rapidly than pressure with increasing depth in the sediment, so the diagenesis appears to have a strong tendency to follow an isotherm and, therefore, may be directly or indirectly related to the presence of gas.

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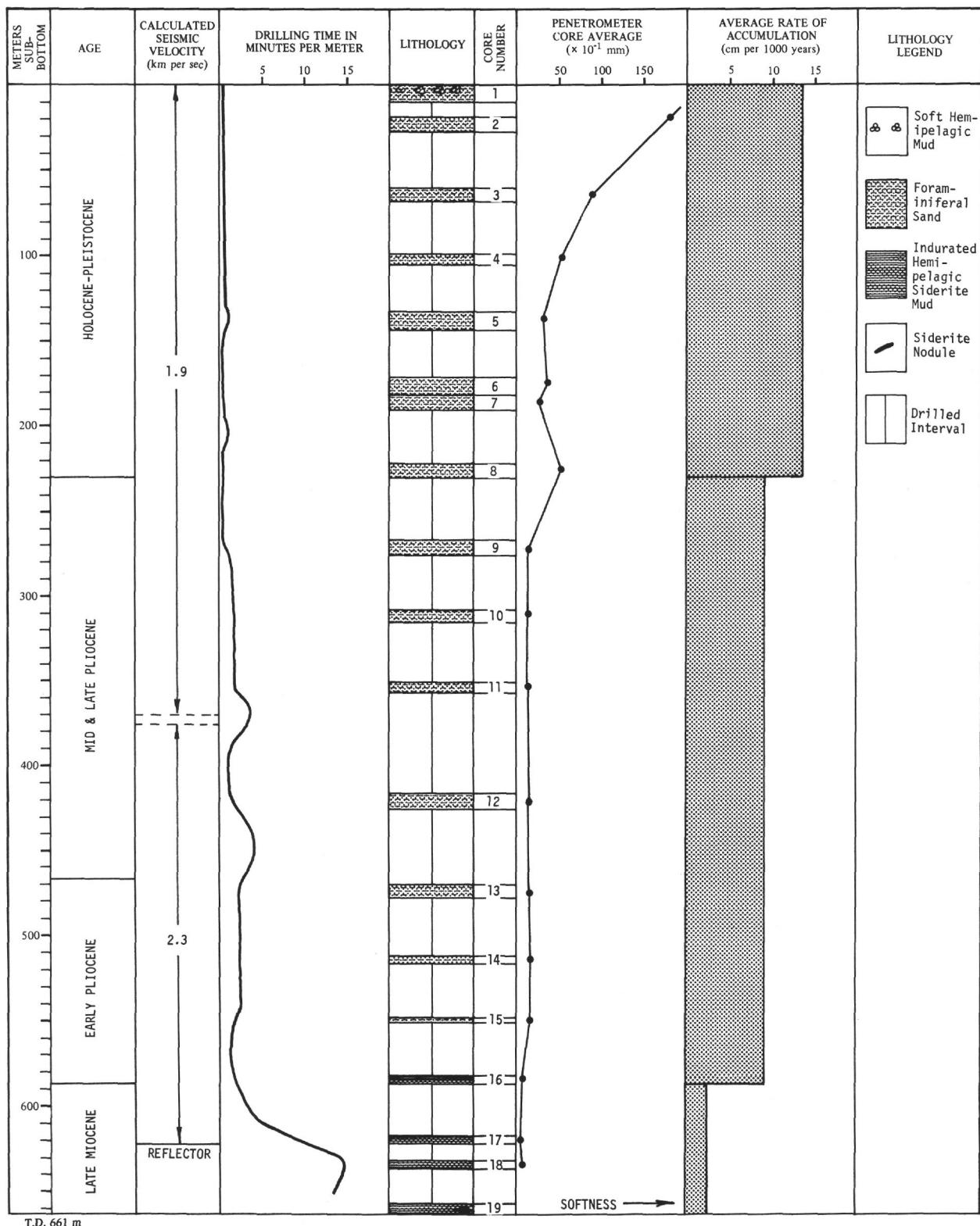


Figure 3a. Site 102 summary chart

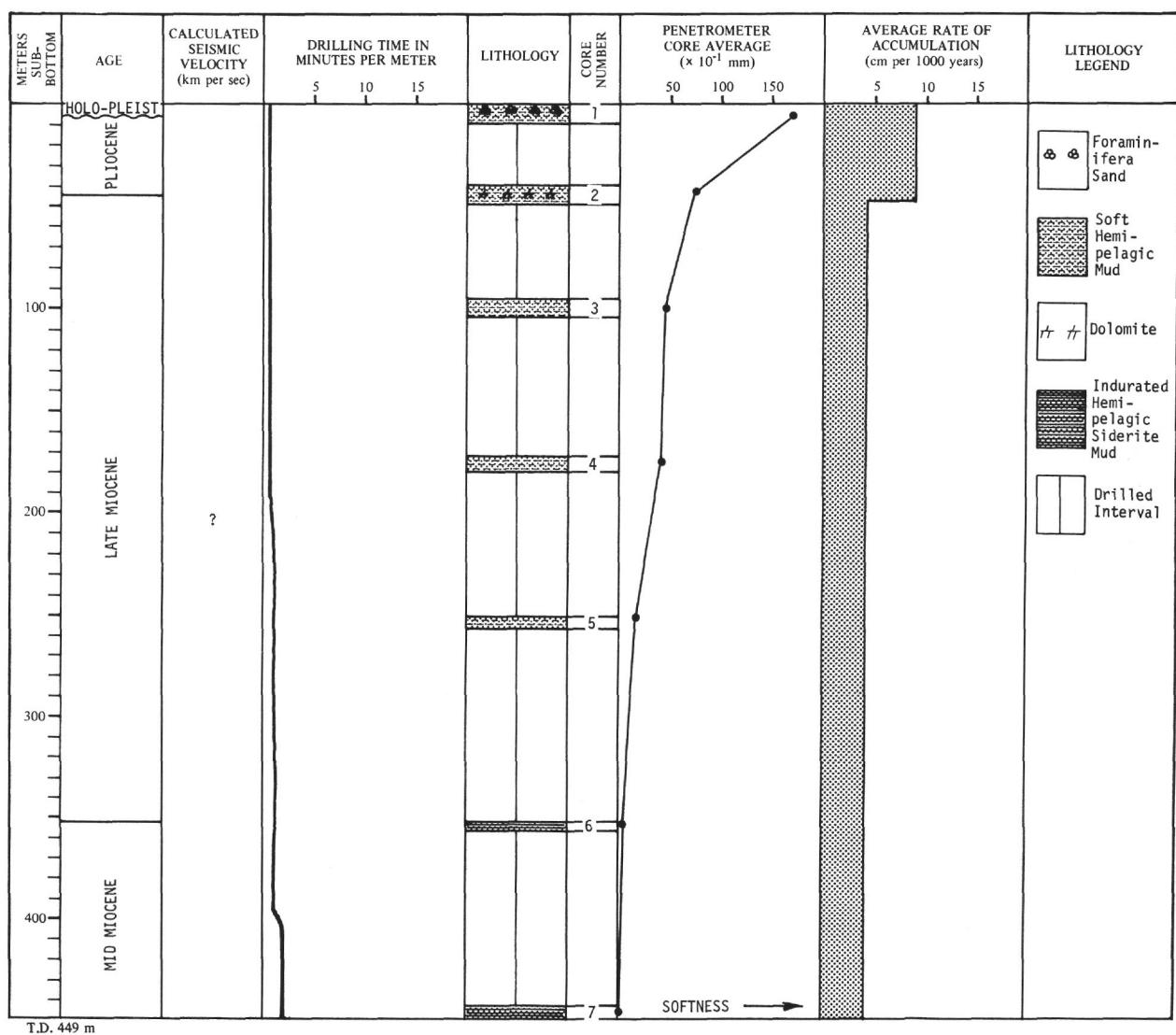


Figure 3b. Site 103 summary chart

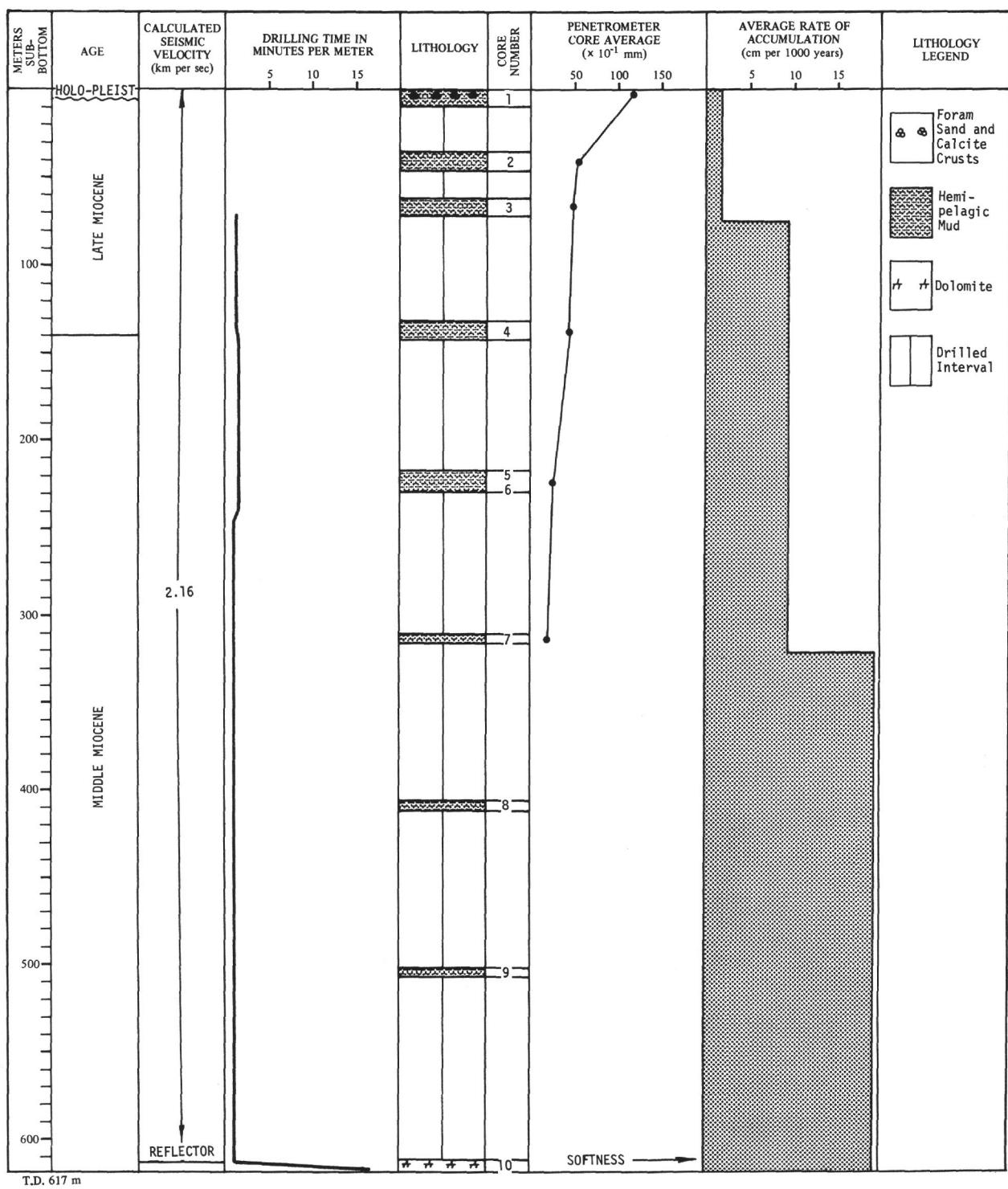


Figure 3c. Site 104 summary chart

Hole 102

Latitude: 30°43.93'N.

Longitude: 74°27.14'W.

Water depth: 3426 meters (drill pipe); 3414 meters (PDR)

Core No.	Interval Cored (meters) ^a					Lithology	Age			
	Depth	Amount	Recovery	Subbottom Depth			Foraminifera	Nannoplankton	Dinoflagellates	
							Holocene-Late Pleistocene	Quaternary		
1	3436-3445	9	9	9	Gray hemipelagic mud	Late Pleistocene				
(Drilled)	(3445-3454)	(9)		(18)						
2	3454-3463	9	8.7	27	Gray hemipelagic mud	Late Pleistocene	Early Quaternary	Quaternary		
(Drilled)	(3463-3494)	(31)		(58)						
3	3494-3503	9	1.7	67	Gray hemipelagic mud	Late Pleistocene	Early Quaternary	Quaternary		
(Drilled)	(3503-3532)	(29)		(96)						
4	3532-3541	9	4.3	105	Gray hemipelagic mud	Late Pleistocene	Early Quaternary	Quaternary		
(Drilled)	(3541-3569)	(28)		(133)						
5	3569-3578	9	9	142	Gray hemipelagic mud	Late Pleistocene	Early Quaternary	Quaternary		
(Drilled)	(3578-3608)	(30)		(172)						
6	3608-3617	9	9	181	Gray hemipelagic mud	Early Pleistocene	Early Quaternary	Quaternary		
7	3617-3626	9	9	190	Gray hemipelagic mud	Early Pleistocene	Early Quaternary	Quaternary		
(Drilled)	(3626-3655)	(29)		(119)						
8	3655-3664	9	1.8	228	Gray hemipelagic mud	Early Pleistocene-Late Pliocene	Early Quaternary	Quaternary		
(Drilled)	(3664-3702)	(38)		(266)						
9	3702-3711	9	9	275	Gray hemipelagic mud	Late Pliocene		Pliocene		
(Drilled)	(3711-3742)	(31)		(306)						
10	3742-3751	9	9	315	Gray hemipelagic mud	Late Pliocene		Pliocene		

Figure 4a. Core Summary table, Site 102.

Core No.	Interval Cored (meters) ^a				Lithology	Age			
	Depth	Amount	Recovery	Subbottom Depth		Foraminifera		Nannoplankton	
								Dinoflagellates	
(Drilled) (3751-3789)	(38)			(353)					
11	3789-3793	4	3	357	Gray hemipelagic mud	Late Pliocene	Middle Pliocene	Pliocene	
(Drilled) (3793-3855)	(62)			(419)					
12	3855-3859	4	4.5	423	Gray hemipelagic mud	Late Pliocene	Middle Pliocene	Pliocene	
(Drilled) (3859-3909)	(50)			(473)					
13	3909-3912	3	4.4	476	Gray hemipelagic mud	← Early Pliocene →		Pliocene	
(Drilled) (3912-3948)	(36)			(512)					
14	3948-3949	1	1.4	513	Gray hemipelagic mud	← Early Pliocene →		Pliocene	
(Drilled) (3949-3984)	(35)			(548)					
15	3984-3985	1	1	549	Gray hemipelagic mud	← Early Pliocene →		Pliocene	
(Drilled) (3985-4020)	(35)			(584)					
16	4020-4021	1	0.7	585	Gray hemipelagic mud	Early Pliocene-Late Miocene	Early Pliocene	Pliocene	
(Drilled) (4021-4054)	(33)			(618)					
17	4054-4055	1	1.5	619	Gray hemipelagic mud	← Late Miocene →		Miocene	
(Drilled) (4055-4070)	(15)			(634)					
18	4070-4072	2	3.4	636	Gray hemipelagic mud	← Late Miocene →		Miocene	
(Drilled) (4072-4095)	(23)			(659)					
19	4095-4097	2	1.5	661	Gray hemipelagic mud	← Late Miocene →		Miocene	

^aAll intervals are measured by drill pipe from the derrick floor which is 10 meters above water surface.

Figure 4a. Core Summary table, Site 102 (Cont)

Hole 103

Latitude: 30°27.08'N.

Longitude: 74°34.99'W.

Water depth: 3964 meters (drill pipe); 3992 meters (PDR)

Core No.	Interval Cored (meters) ^a				Lithology	Age		
	Depth	Amount	Recovery	Subbottom Depth		Foraminifera	Nannoplankton	Dinoflagellates
1	3974-3982	8	9	8	Hemipelagic mud	Holocene-Late Pleistocene-Early Pliocene	Early Pliocene	
(Drilled)	(3982-4012)	(30)		(38)				
2	4012-4021	9	9	47	Hemipelagic mud with dolomite-rich layer	Early Pliocene Late Miocene	Late Miocene	
(Drilled)	(4021-4068)	(47)		(94)				
3	4068-4077	9	9	103	Hemipelagic mud	← Late Miocene →		Miocene
(Drilled)	(4077-4144)	(67)		(270)				
4	4144-4153	9	6.8	179	Hemipelagic mud	← Late Miocene →		Miocene
(Drilled)	(4153-4221)	(68)		(247)				
5	4221-4230	9	2.5	256	Hemipelagic mud	← Late Miocene →		Miocene
(Drilled)	(4230-4317)	(87)		(343)				
6	4317-4326	9	1.4	352	Hemipelagic mud	← Late Miocene →		Miocene
(Drilled)	(4326-4414)	(88)		(440)				
7	4414-4423	9	0.25	449	Dark gray shale	Late Miocene	Middle Miocene	

^aAll intervals are measured by drill pipe from the derrick floor which is 10 meters above water surface.

Figure 4b. Core Summary table, Site 103.

Hole 104

Latitude: 30°49.65'N.

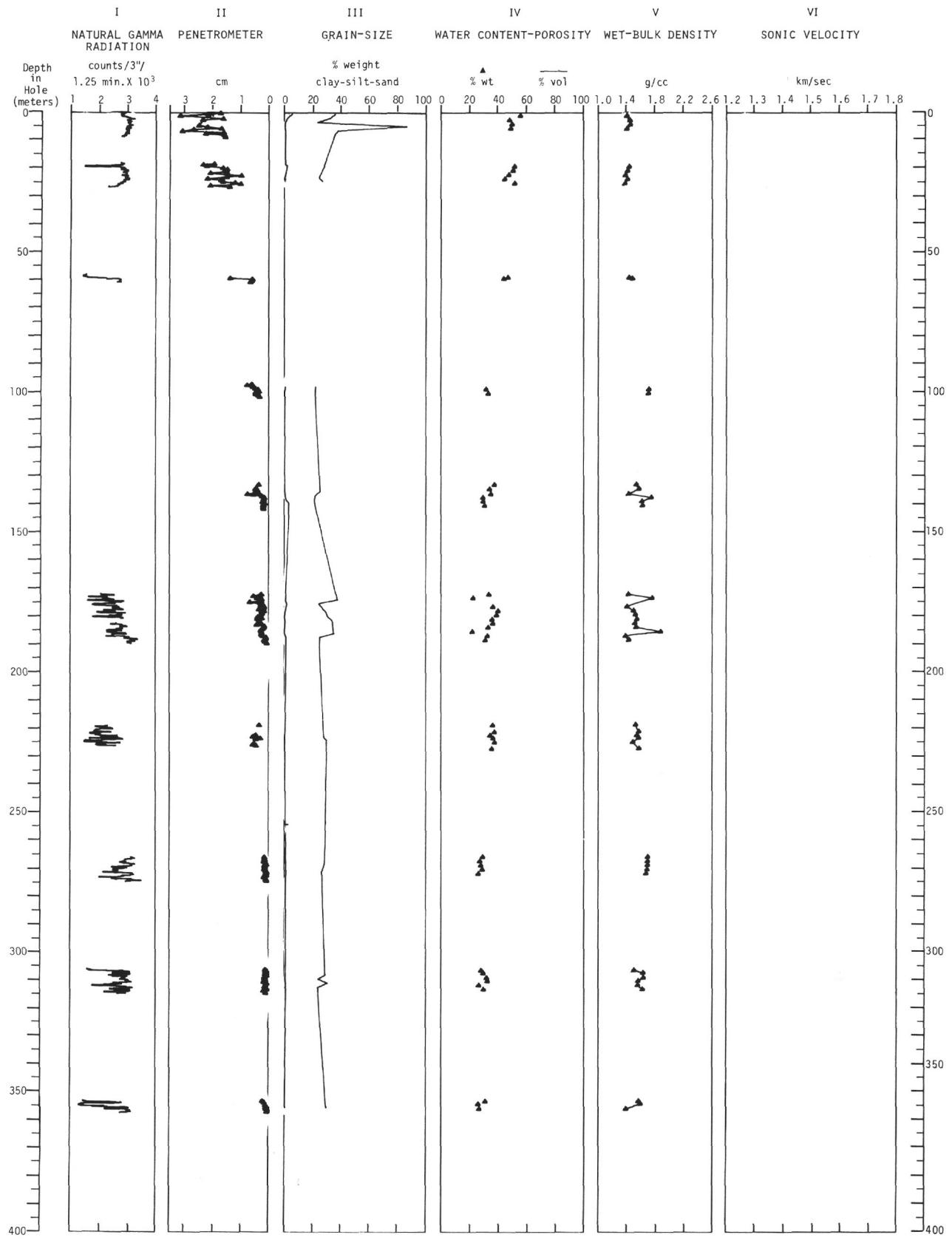
Longitude: 74°19.64'W.

Water depth: 3811 meters (drill pipe); 3833 meters (PDR)

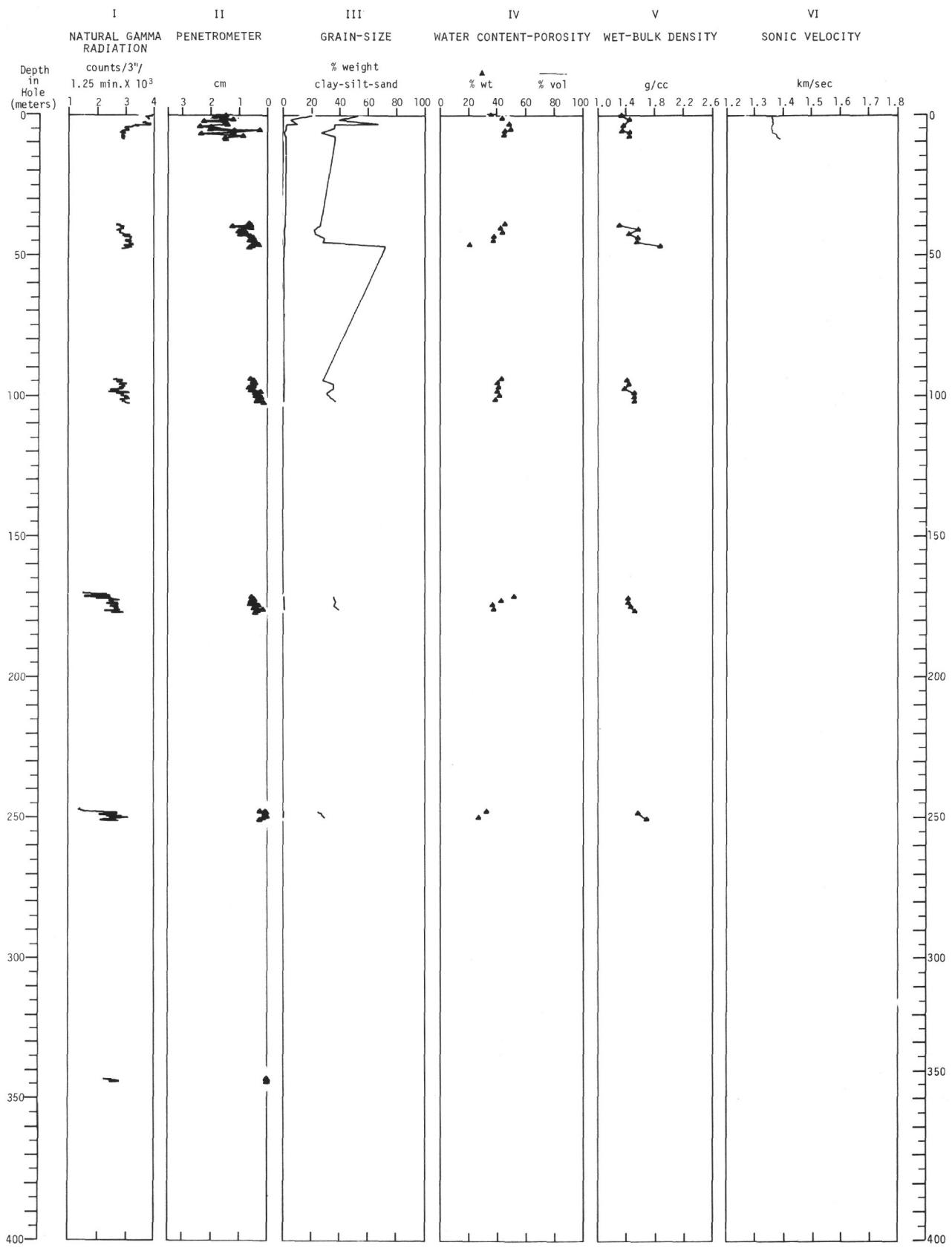
Core No.	Interval Cored (meters) ^a					Lithology	Age		
	Depth	Amount	Recovery	Subbottom Depth	Foraminifera		Nannoplankton	Dinoflagellates	
1	3821-3830	9	9	9	Hemipelagic mud worm-burrow crust	Holocene- Pleistocene- Late Miocene	Late Miocene		
(Drilled)	(3830-3857)	(27)		(36)					
2	3857-3866	9	9	45	Hemipelagic mud	← Late Miocene →	Miocene		
(Drilled)	(3866-3883)	(17)		(62)					
3	3883-3892	9	9	71	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(3892-3954)	(62)		(133)					
4	3954-3963	9	9	142	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(3963-4031)	(68)		210					
5	4031-4040	9	0	219	Hemipelagic mud				
6	4040-4049	9	9	228	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(4049-4127)	(78)		(306)					
7	4127-4136	9	3.6	315	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(4136-4222)	(86)		(401)					
8	4222-4231	9	4	410	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(4231-4316)	(85)		(495)					
9	4316-4325	9	2.5	504	Hemipelagic mud	← Middle Miocene →	Miocene		
(Drilled)	(4325-4436)	(111)		(615)					
10	4436-4438	2	1	617	Hemipelagic mud and ankerite layer	← Middle Miocene →	Miocene		

^aAll intervals are measured by drill pipe from derrick floor which is 10 meters above water surface.

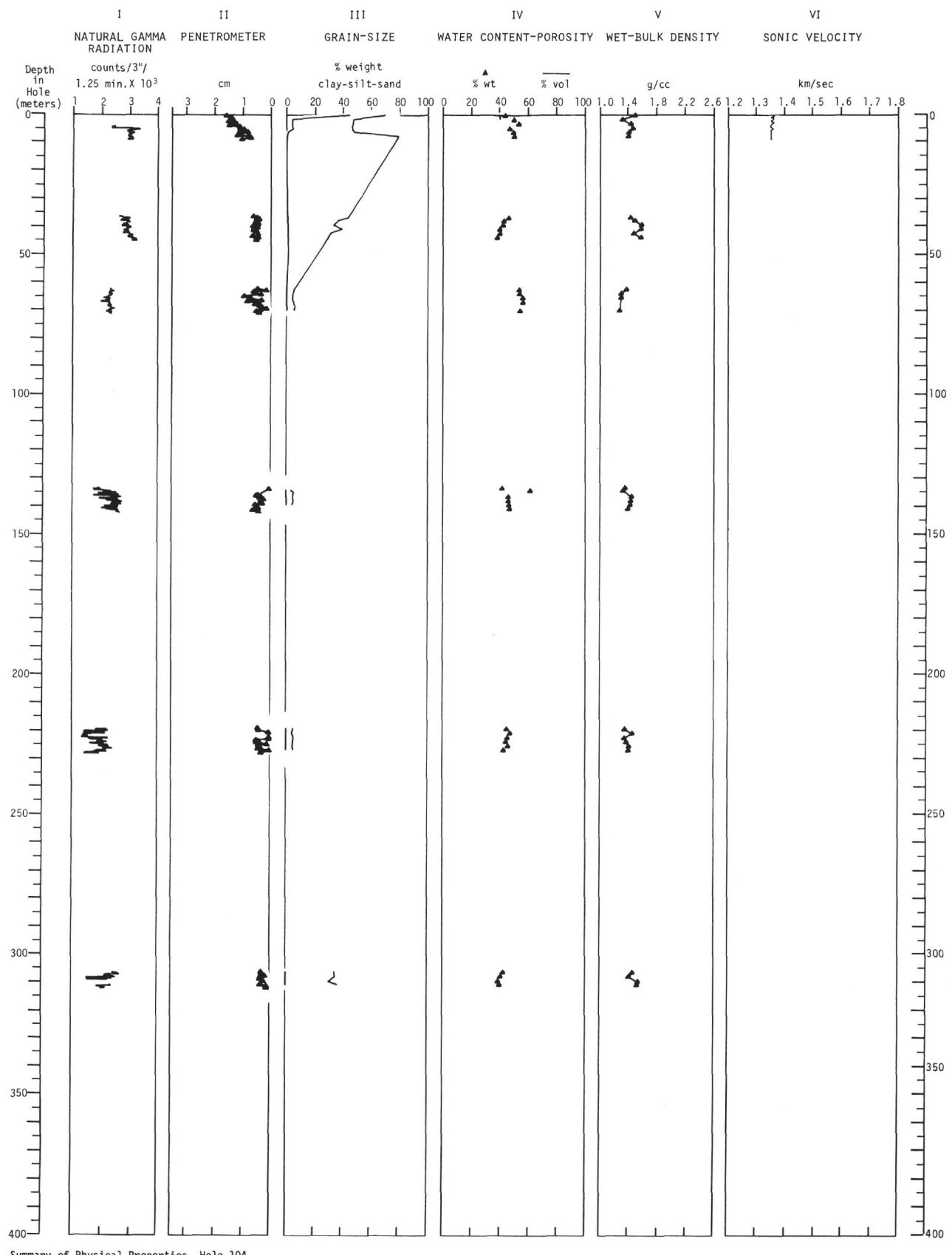
Figure 4c. Core Summary table, Site 104.

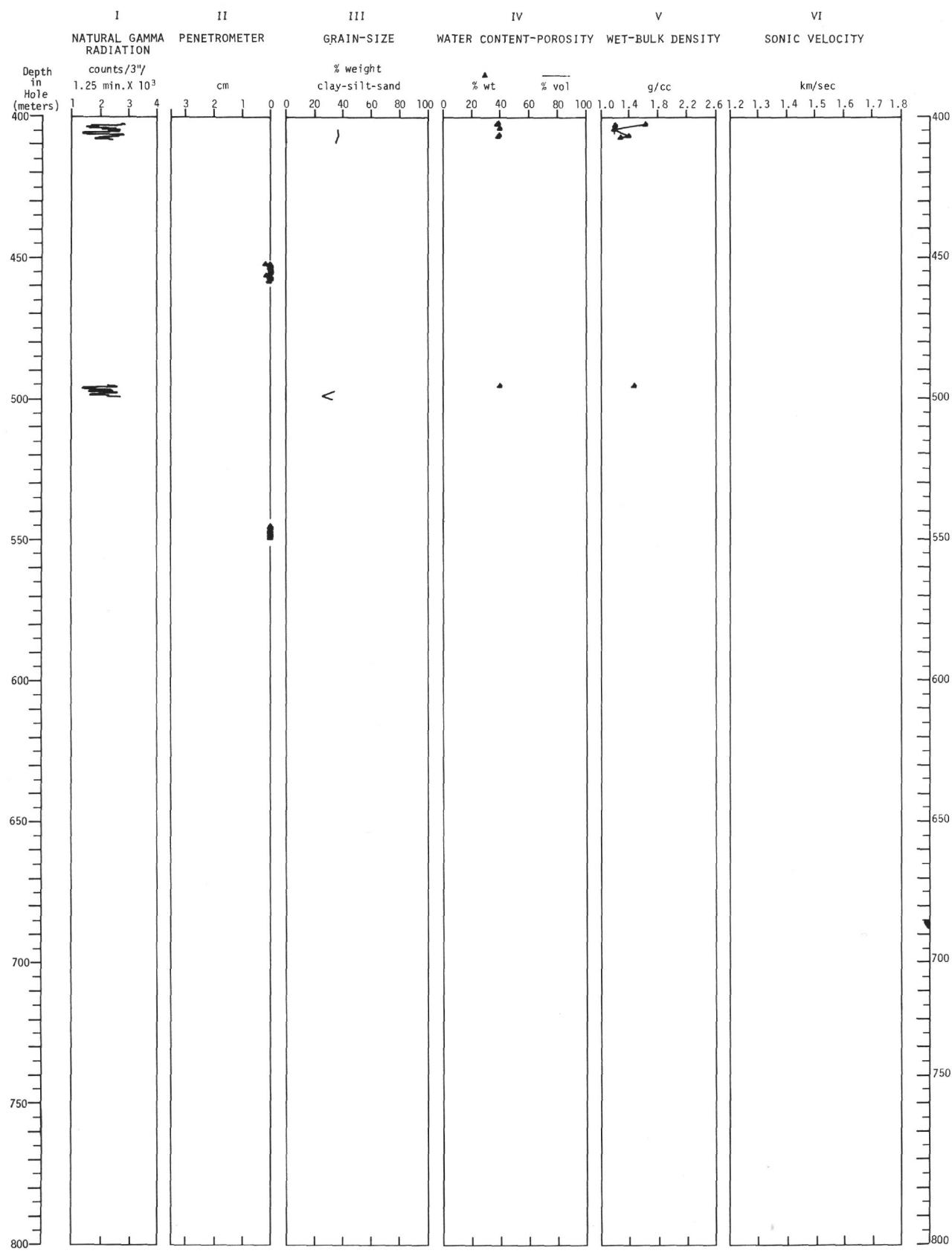


Summary of Physical Properties, Hole 102

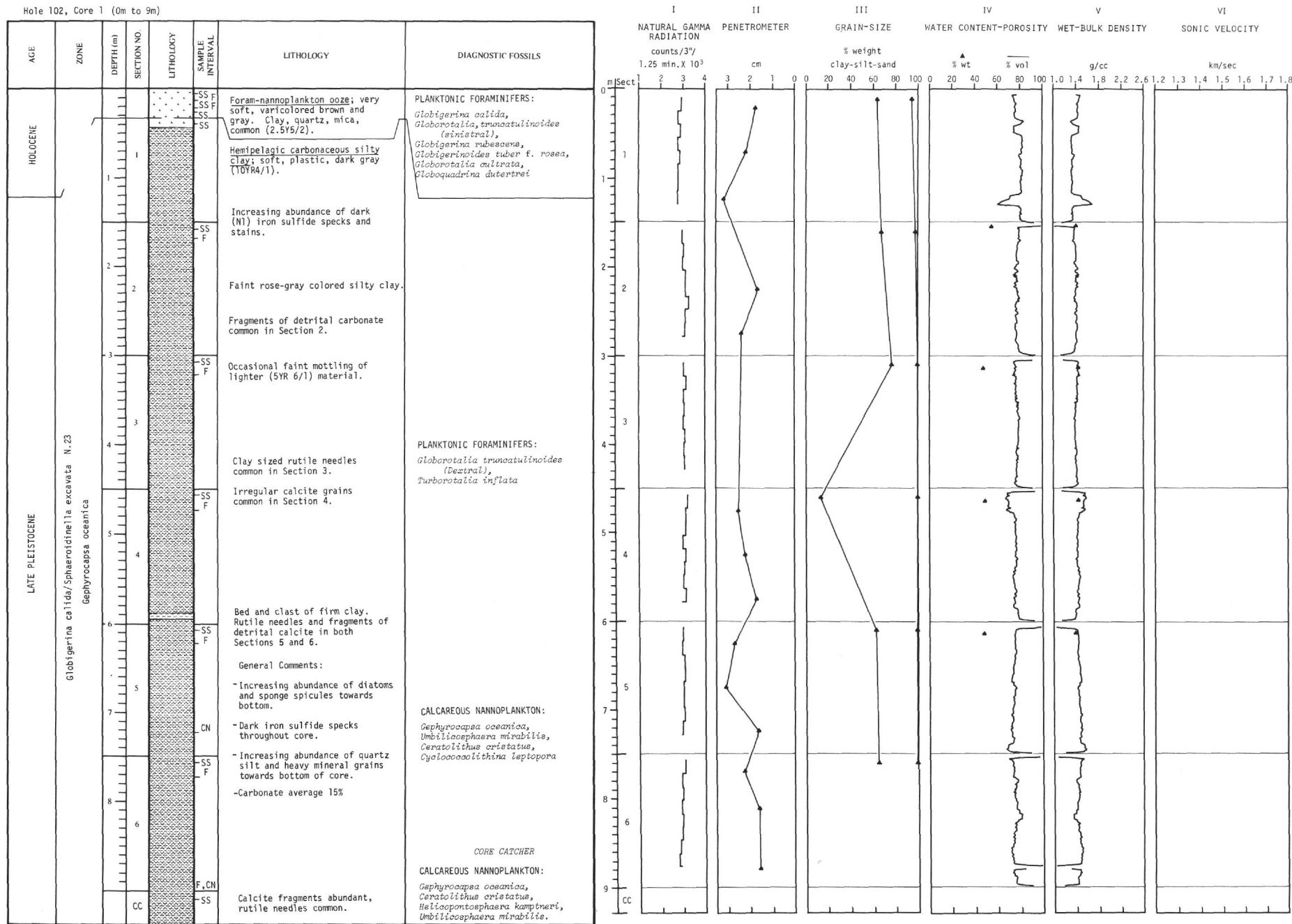


Summary of Physical Properties, Hole 103



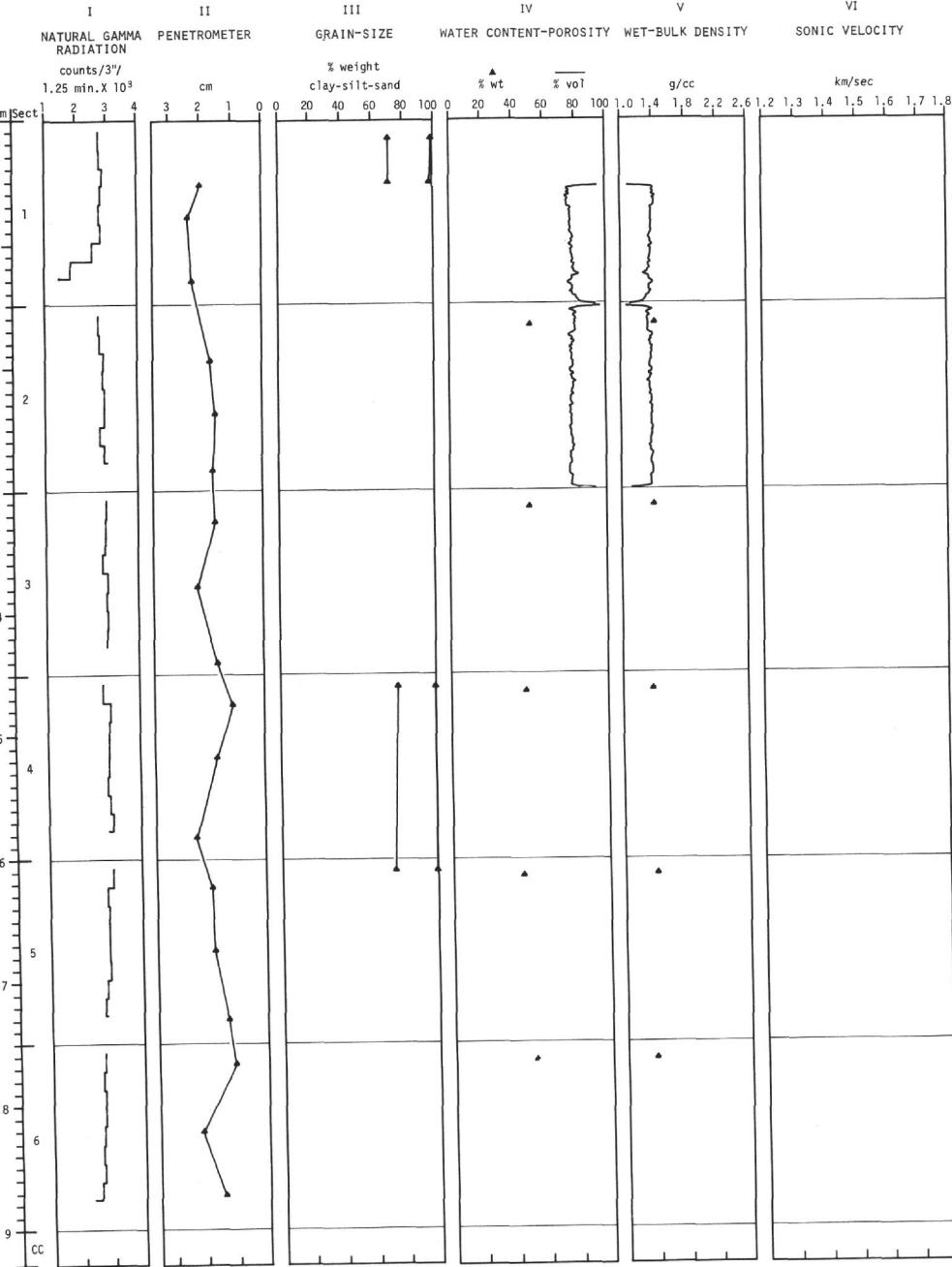


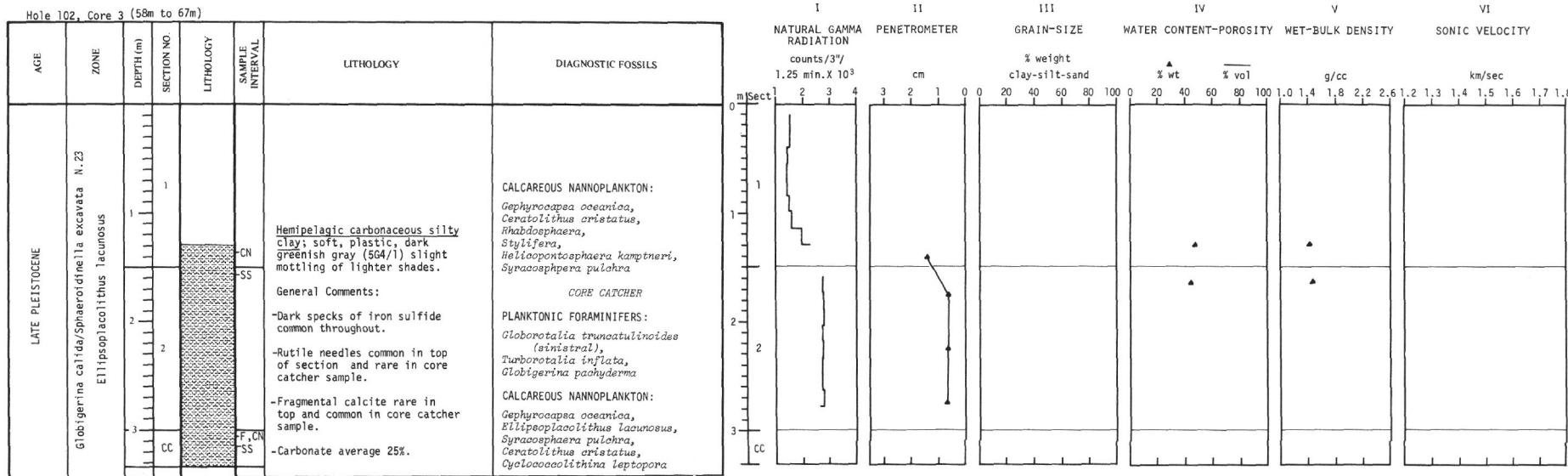
Summary of Physical Properties, Hole 104 (Cont'd)



Hole 102, Core 2 (18m to 27m)

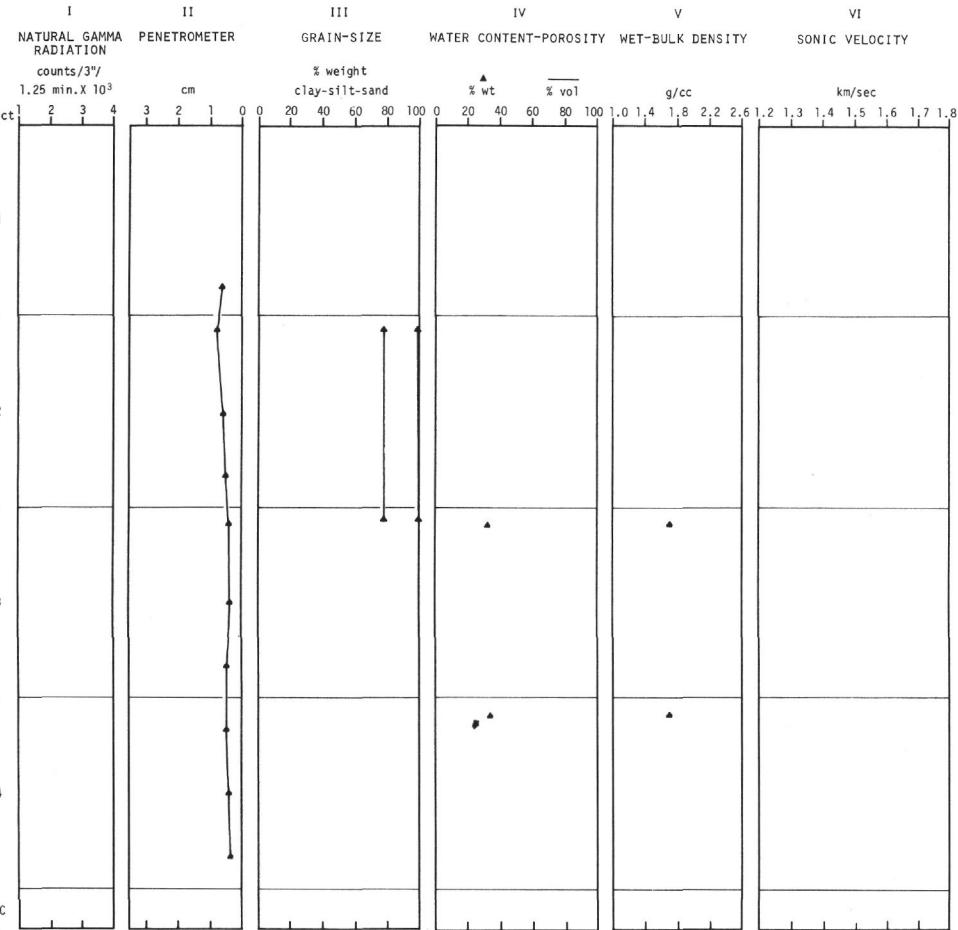
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE PLEISTOCENE	Globigerina calida/Sphaeroidinella excavata N 23 Ellipsiplacolithus lacunosus	1	1	SS		Hemipelagic carbonaceous silty clay; soft, plastic, dark greenish gray (5G4/1) with mottling of greenish gray (5G6/1) and light olive gray (5Y6/1).	
		2	2	SS			CALCAREOUS NANNOPLANKTON: <i>Gephyrocapsa oceanica</i> , <i>Ellipsiplacolithus lacunosus</i> , <i>Ceratolithus cristatus</i> , <i>Syracosphaera pulchra</i> , <i>Cyclococcolithina leptopora</i>
		3	3	SS		Heavy minerals common in Sections 2 and 3. Rutile needles common in Sections 3 and 4.	
		4	4	SS		Radiolarians rare in Sections 4 and 5.	
		5	5	SS		Occasional pyrite nodules.	
		6	6	SS		General Comments: -Sediment broken into foamy chunks due to gas expansion. -Abundant black specks of iron sulfide common throughout. -Abundant radiolarians, diatoms, and sponge spicules. -Common organic matter and calcite fragments throughout. -Increasing amount of sand, nannoplankton and radiolarians towards bottom of core. -Carbonate average 10%.	
		7	7	SS		Gradual change to greenish gray color (5G6/1).	CORE CATCHER PLANKTONIC FORAMINIFERS: <i>Globorotalia truncatulinoides</i> (<i>sinistral</i>), <i>Turborotalita inflata</i>
		8	8	CC	F, CN		CALCAREOUS NANNOPLANKTON: <i>Gephyrocapsa oceanica</i> , <i>Ceratolithus cristatus</i> , <i>Rhabdosphera stylifera</i> , <i>Syracosphaera pulchra</i> , <i>Cyclococcolithina leptopora</i>





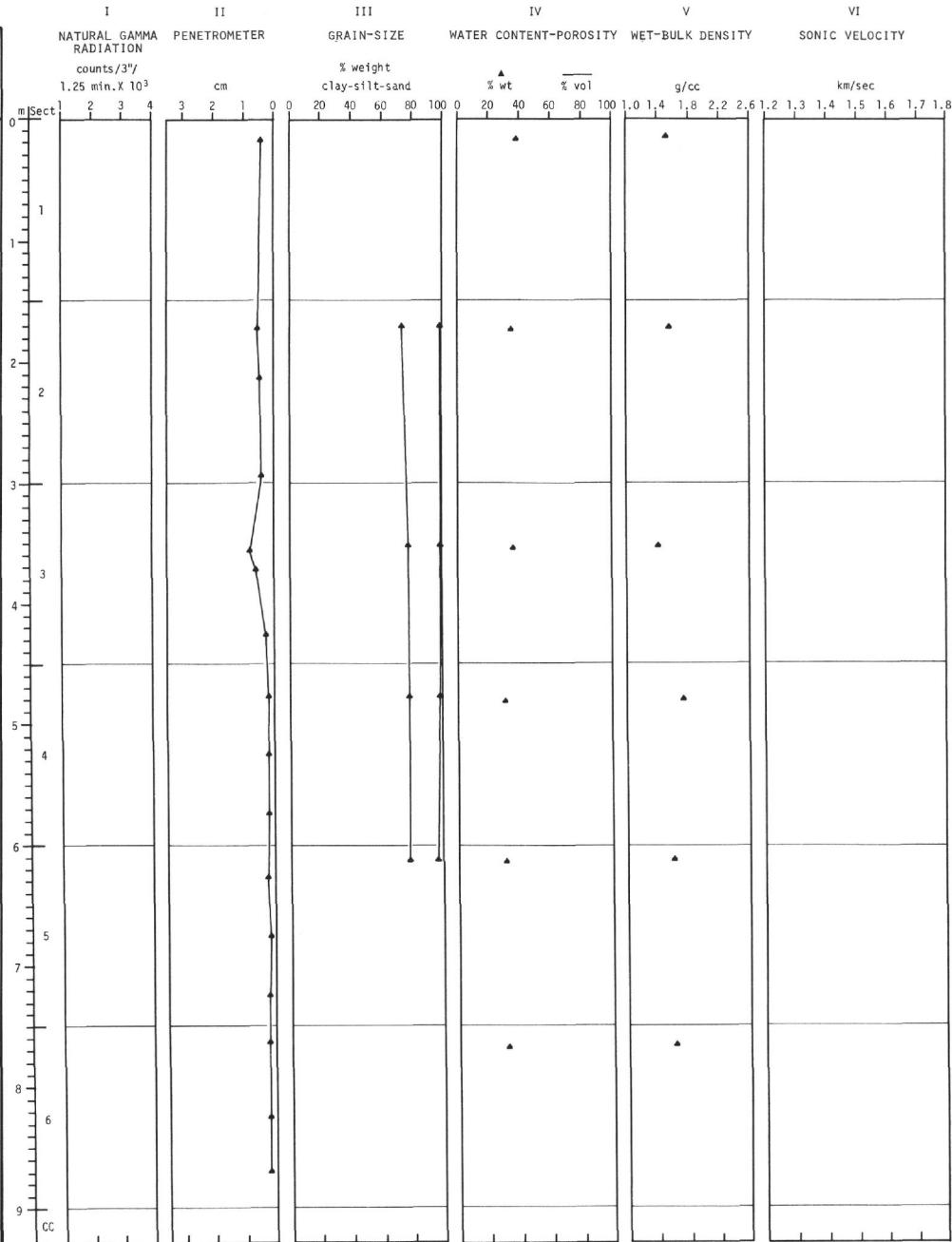
Hole 102, Core 4 (96m to 105m)

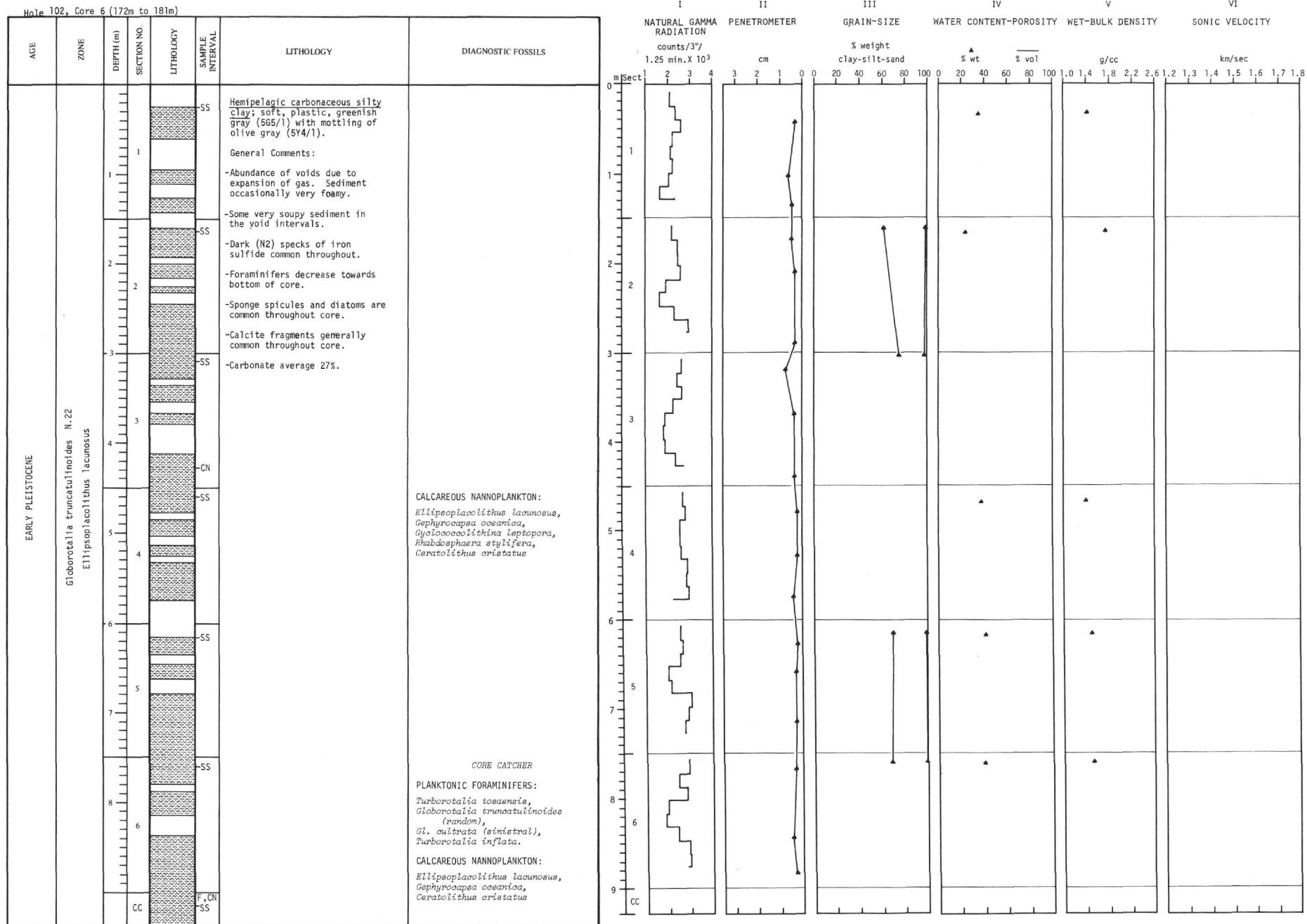
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
EARLY PLEISTOCENE	Globorotalia truncatulinoides N. 22 Ellipsoplaclithus lacunosus	1	1	Hemipelagic carbonaceous silty clay; soft, plastic, greenish gray (566/1) with intense mottling of light greenish gray (567/1).	SS CN		CALCAREOUS NANNOPLANKTON: <i>Ellipsoplaclithus lacunosus</i> , <i>Gephyrocapsa oceanica</i> , <i>Syracosphera pulchra</i> , <i>Rhabdosphera stylifera</i> , <i>Cyclococcolithina leptopora</i> , <i>Ceratolithus cristatus</i>
		2	2	Dark greenish gray (56 4/1) with mottling of brownish gray (5YR4/1) and greenish gray (566/1).	SS		
		3	3	Olive gray (5Y4/1) with moderate mottling of dark greenish gray (56Y4/1).	SS		
		4	4	General Comments: - Radiolarians, diatoms and sponge spicules rare throughout. - Quartz silt abundant in upper portions of core. - Nannoplankton increase towards bottom of core. - Voids due to gas expansion. - Carbonate average 8%. - Black specks of iron sulfide common throughout core. - Calcite fragments abundant and rutile needles rare throughout.	F, CN SS	CORE CATCHER PLANKTONIC FORAMINIFERS: <i>Globorotalia truncatulinoides</i> (Dextral), <i>Globorotalia ultrata</i> (sinistral), <i>Turborotalita inflata</i> . CALCAREOUS NANNOPLANKTON: <i>Gephyrocapsa oceanica</i> , <i>Ellipsoplaclithus lacunosus</i> , <i>Rhabdosphera stylifera</i> , <i>Ceratolithus cristatus</i>	
		5					
		6					
		7					

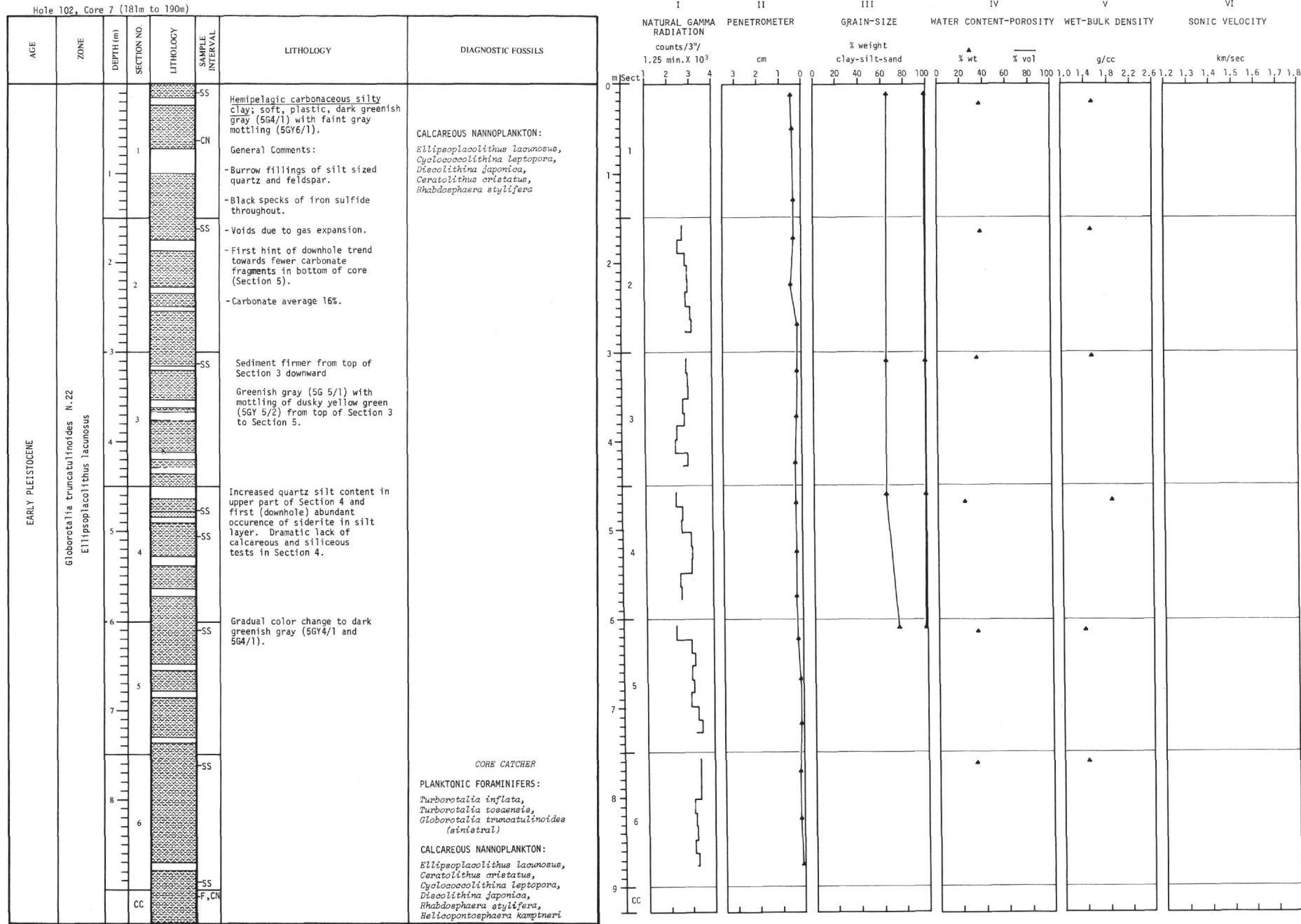


Hole 102, Core 5 (133m to 142m)

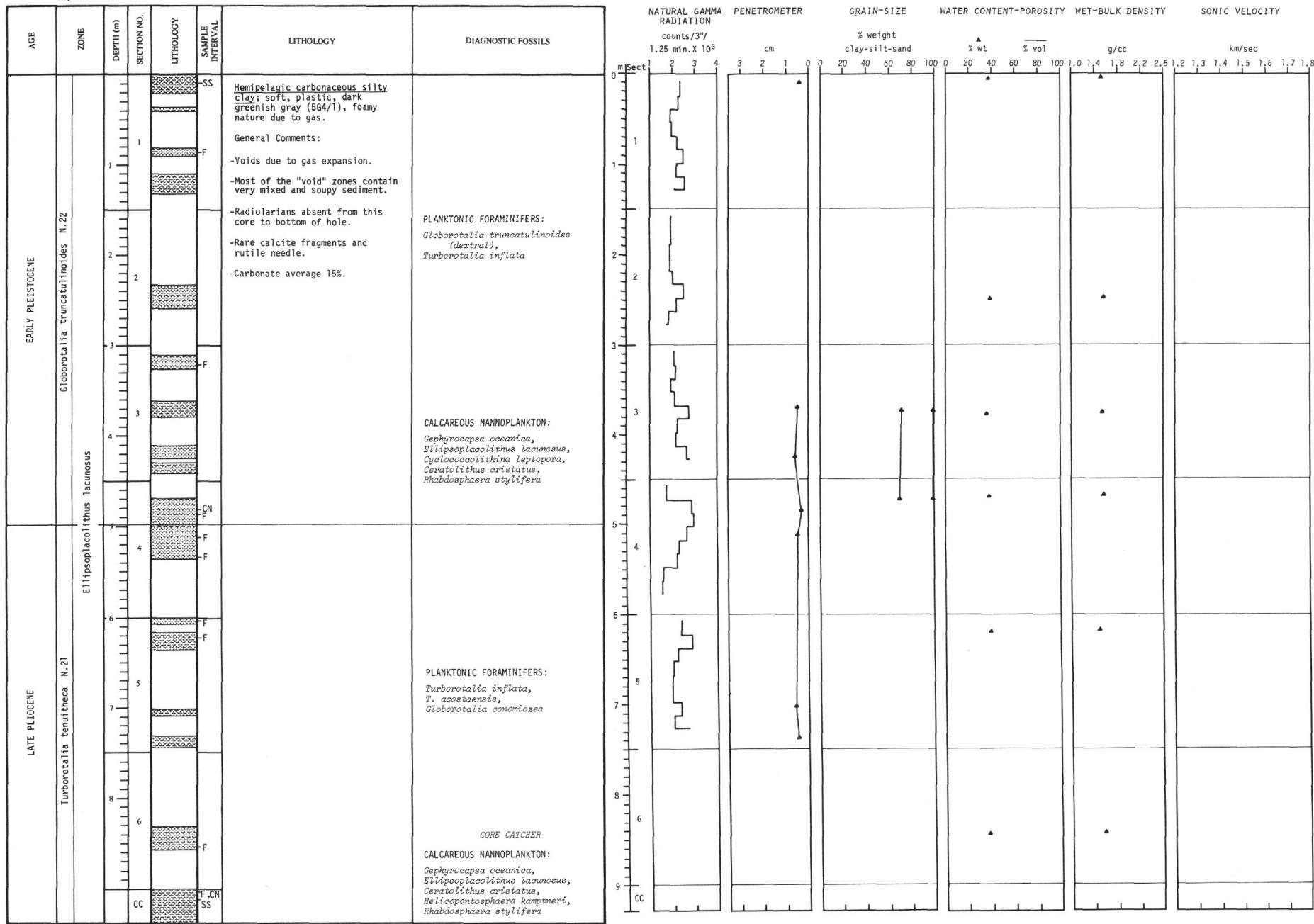
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
EARLY PLEISTOCENE	Globorotalia truncatulinoides N. 22 Ellipsoplaocolithus lacunosus	1	1	SS		Hemipelagic carbonaceous silty clay; soft plastic, greenish gray (5G5/1), slight mottling of olive gray (5Y4/1) and brownish gray (5YR4/1).	General Comments: -Clay mineral content decreases towards the lower half of the core. -Foraminifers increase in abundance towards lower part of the core. -Siliceous forms are rare in core. -Core highly disturbed throughout. -Voids due to gas expansion. -Black specks of iron sulfide common. -Calcite fragments generally common, rutile needles generally rare. Glauconite common in Section 2.
		2	2	SS			Pyrite nodule.
		3	3	SS			CALCAREOUS NANNOPLANKTON: <i>Ellipsoplaocolithus lacunosus</i> , <i>Gephyrocapsa oceanica</i> , <i>Rhabdosphaera stylifera</i> , <i>Discosphaera japonica</i> , <i>Cyclococcolithina leptopora</i>
		4	4	-SS CN			Pyrite nodules.
		5		-SS			
		6	6	-SS			
		7		SS			
		8		SS			
		CC		F, CN			
				SS			





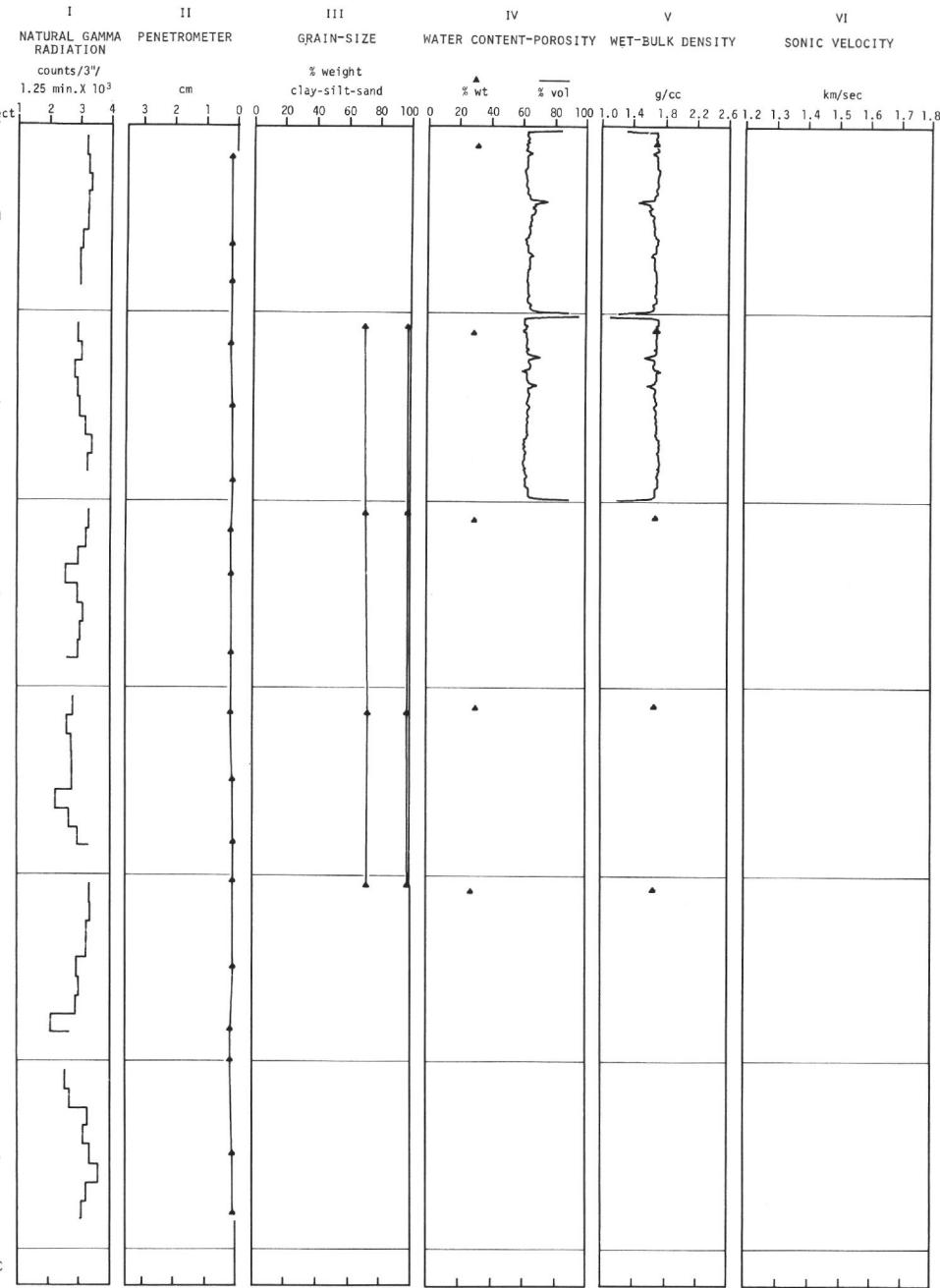


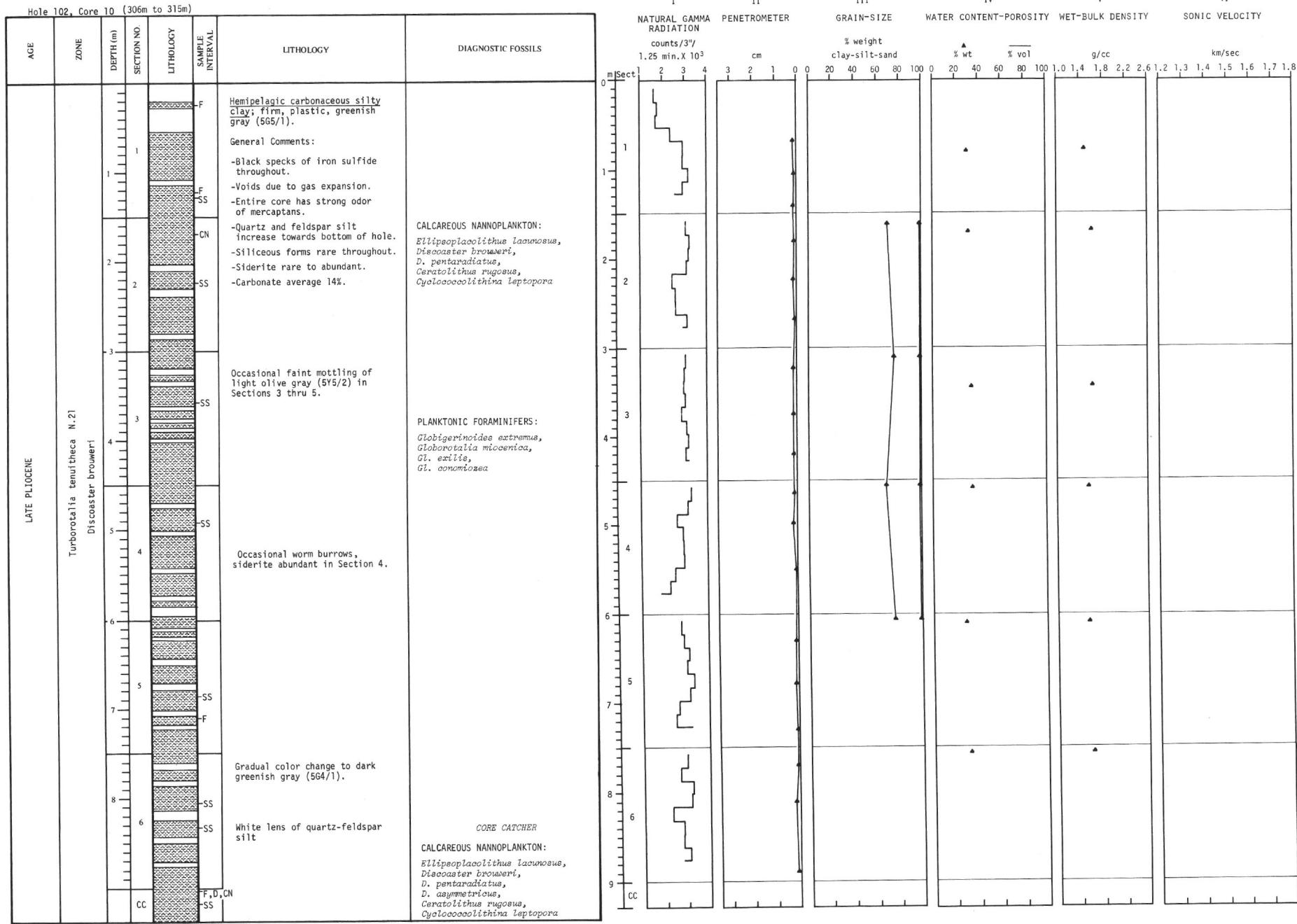
Hole 102, Core 8 (219m to 228m)



Hole 102, Core 9 (266m to 275m)

AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE PLIOCENE	Turborotalita tenuitheca N. [2] Discocaster brouweri						
		1	1	SS, F	CN	Hemipelagic carbonaceous silty clay; slightly indurated plastic, greenish gray (5G5/1) with mottling of olive gray (5Y4/1).	
		2	2	SS		Pyrite nodule.	CALCAREOUS NANNOPLANKTON: <i>Ellipsoplaclithus lacunosus</i> , <i>Discocaster brouweri</i> , <i>D. pentaradiatus</i> , <i>Cyclococcolithina leptopora</i> , <i>Helicopontosphaera kampfneri</i> , <i>Ceratolithus cristatus</i>
		3	3	F	SS	General Comments: -Black specks of iron sulfide common throughout. -Finely bedded nature noticed for first time in the hole. -Voids due to gas expansion. -Silt content and nannoplankton increase toward bottom of core. -Siliceous forms conspicuously absent. -Siderite rhombs rare to common throughout, first consistent occurrence. -Carbonate average 26%.	PLANKTONIC FORAMINIFERS: <i>Turborotalita inflata</i> , <i>Globorotalita conomicoae</i> , <i>Turborotalita costataensis</i> , <i>Globigerinoides extremus</i> , <i>Globorotalita vindobona</i> , <i>Turborotalita toeaensis</i>
		4	4	SS	F	Lens of pyrite silt. Pyrite nodule.	
		5	5	F			
		6	6	SS			
		7					
		8	6				
		CC		F, D, CN	SS		CORE CATCHER DINOFLAGELLATES: <i>Achomosphaera rotulifera</i> CALCAREOUS NANNOPLANKTON: <i>Discocaster brouweri</i> , <i>D. pentaradiatus</i> , <i>Cyclococcolithina leptopora</i> , <i>Ellipsoplaclithus lacunosus</i>





Hole 102, Core 11 (353m to 357m)						
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY
LATE PLIOCENE	Turborotalia tenuitheca N.21 Reticulofenestra pseudounbilica	1	1	SS CN		Hemipelagic carbonaceous silty clay; firm, plastic, dark greenish gray (564/1). General Comments: -Siderite common throughout. -White specks and glass rare to common. -Silt, pyrite, foraminifers and sponge spicules increase towards bottom of core. -Unusually abundant sponge spicules. -Voids due to gas expansion. -Carbonate average 18%.
		2	2	SS		
		3	3	F		
		4	CC	SS, F CN		

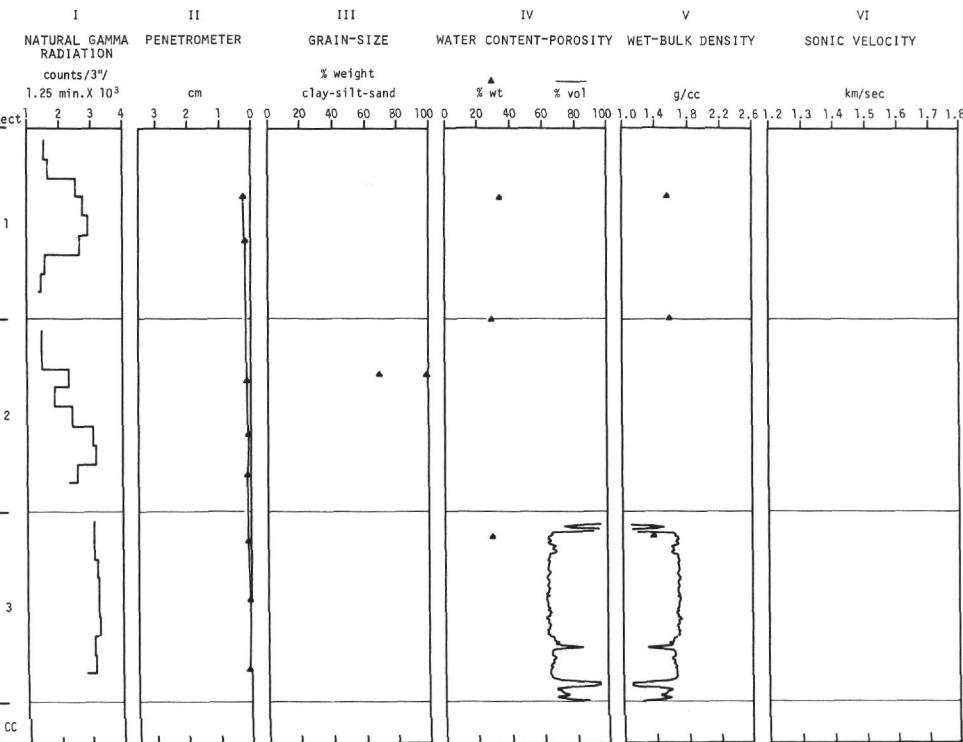
DIAGNOSTIC FOSSILS

CALCAREOUS NANOPLANKTON:
Reticulofenestra pseudounbilica,
Ellipsoplaocolithus lacanoeus,
Cyclococcolithina mactintyrei,
Discaster brouweri,
D. variabilis,
D. pentaradiatus,
D. asymmetricus,
Helicopontosphaera sellii,
H. kampfneri

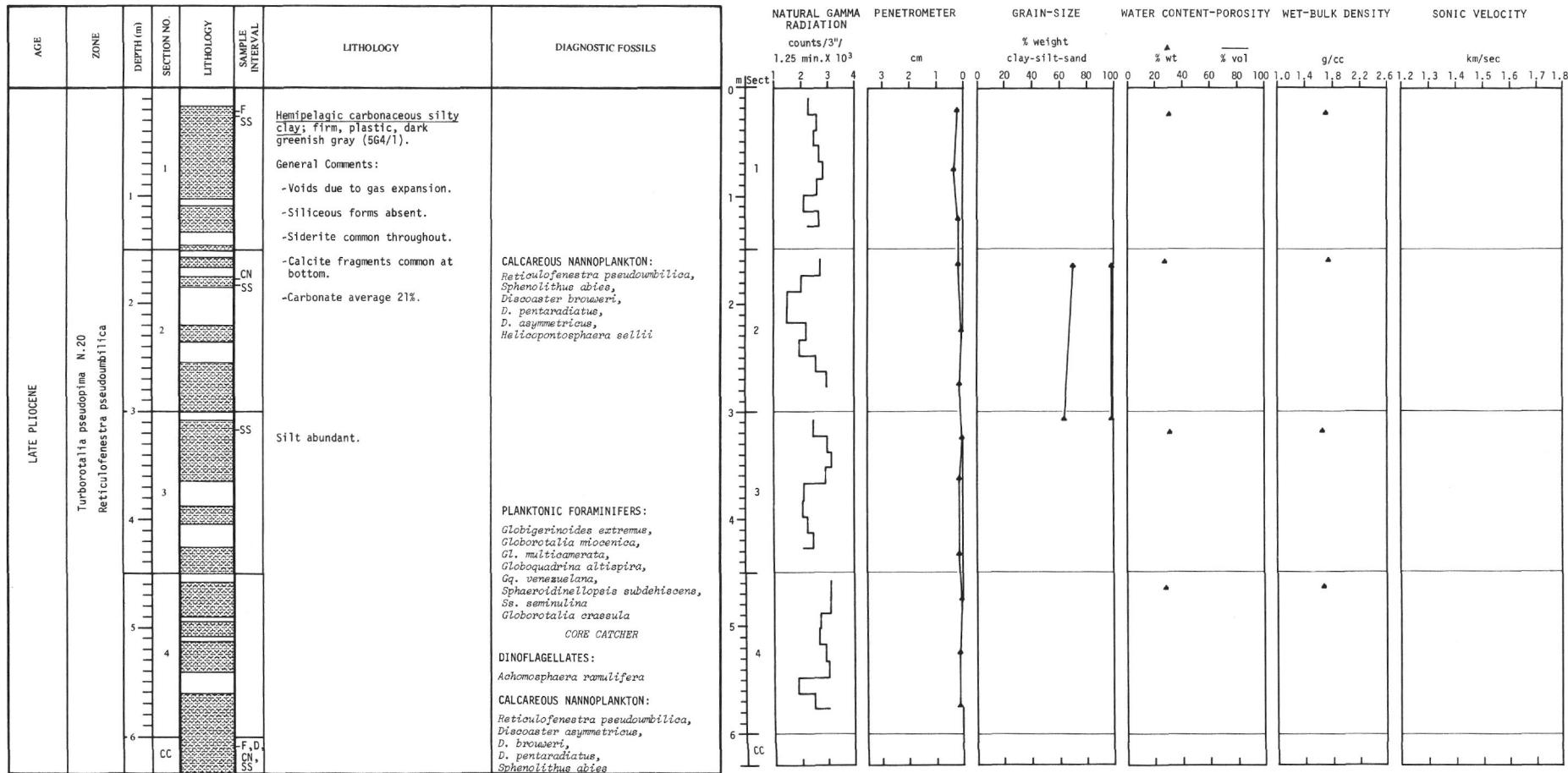
PLANKTONIC FORAMINIFERS:
Globigerinoides extremus,
Globorotalia miccenica,
Gl. multicamerata,
Gl. exilis,
Globoquadrina atlispira,
Sphaeroidinellopsis subdehisces,
Globoquadrina venezuelana,
Globorotalia conomicoae

CORE CATCHER

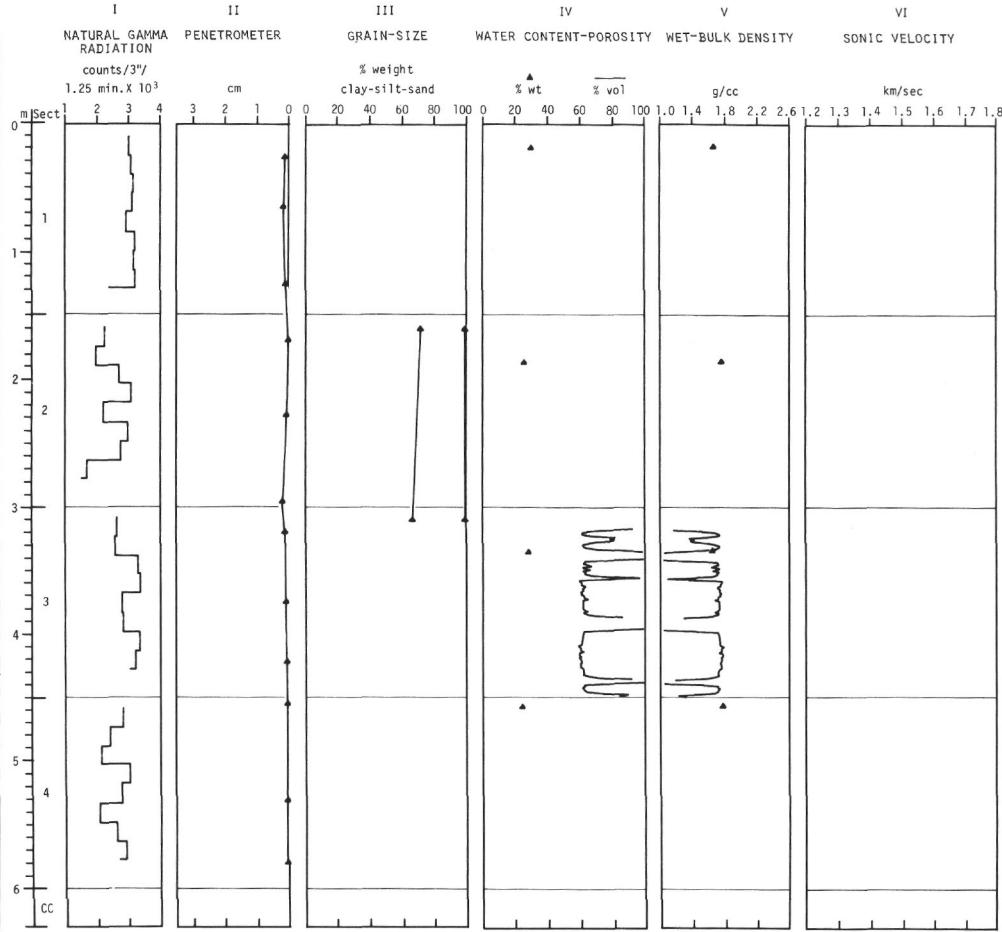
CALCAREOUS NANOPLANKTON:
Reticulofenestra pseudounbilica,
Discaster brouweri,
D. pentaradiatus,
D. asymmetricus,
Sphenolithus abies



Hole 102, Core 12 (419m to 423m)

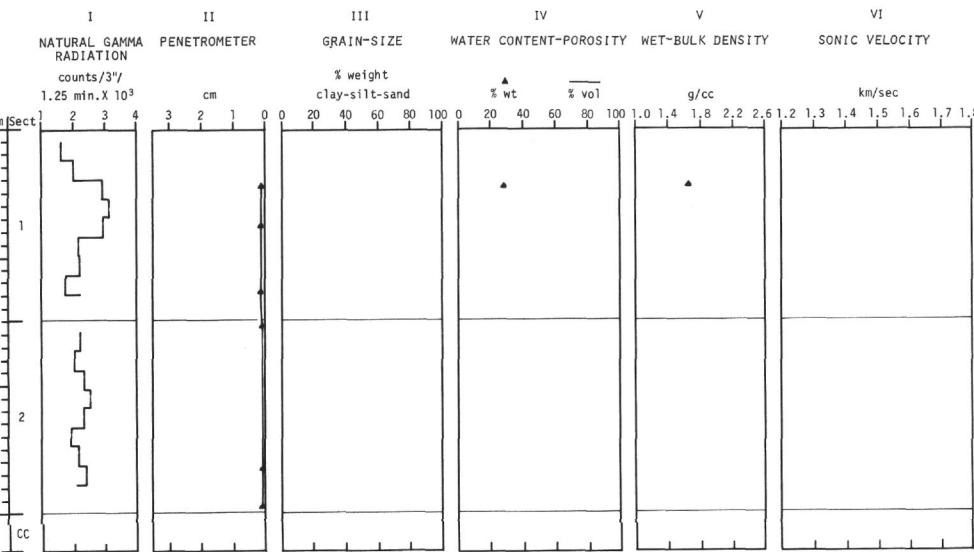


Hole 102, Core 13 (473m to 476m)						
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY
EARLY PLIOCENE	Sphaeroidinella dehisces/Globiquadrina altispira N.19 Discocaster asymmetricus	1	1	SS F		Hemipelagic carbonaceous silty clay; slightly indurated, firm dark greenish gray (564/1). General Comments: -Voids due to gas expansion. -Siderite common throughout. -Siliceous forms absent. -Carbonate average 16%.
		2	2	SS		Abundant sand and silt grains of heavy minerals, quartz, pyrite and glauconite.
		3	3	SS		PLANKTONIC FORAMINIFERS: <i>Globorotalia margaritae</i> , <i>Pulleytina primalis</i> , <i>Globorotalia cibaeensis</i> , <i>Globoquadrina altispira</i> , <i>Globigerinoides extremus</i>
		4	3	F		CALCAREOUS NANOPLANKTON: <i>Discocaster asymmetricus</i> , <i>D. surculus</i> , <i>D. pentaradiatus</i> , <i>Sphenolithus abies</i> , <i>Reticulofenestra pseudowulicina</i> , <i>Helicopontosphaera kampferi</i> , <i>H. sellii</i> CORE CATCHER
		5	4	CN		DINOFLAGELLATES: <i>Achomosphaera ramiifera</i>
		6	CC	F,D,CN,SS		CALCAREOUS NANOPLANKTON: <i>Discocaster asymmetricus</i> , <i>D. surculus</i> , <i>Reticulofenestra pseudowulicina</i> , <i>Sphenolithus abies</i> , <i>Helicopontosphaera sellii</i>



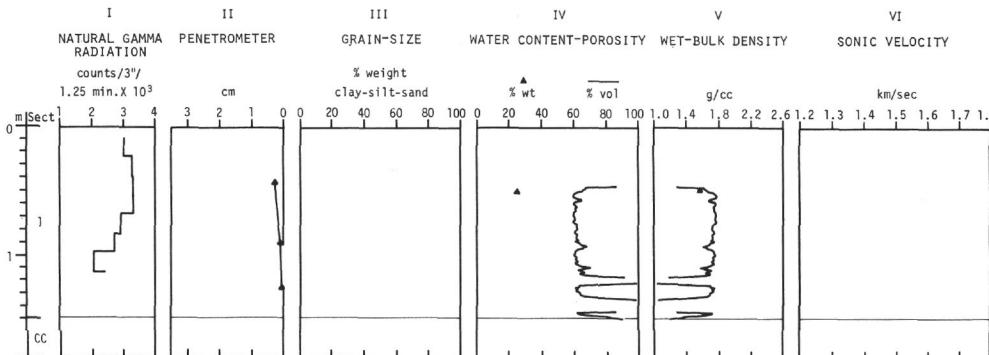
Hole 102, Core 14 (512m to 513m)

AGE	ZONE	LITHOLOGY	DIAGNOSTIC FOSSILS		
EARLY PLIOCENE	Sphaeroidinella dehiscens/Globogaudrina Discocaster asymmetricus				
		1		SS	
		2		CN	
		3		CC	F, D, CN, SS
			Hemipelagic carbonaceous silty clay; slightly indurated, firm, dark greenish gray (564/1). General Comments: -Voids due to gas expansion. -Siderite common throughout. -Siliceous forms absent. -Carbonate average 19%.		
			PLANKTONIC FORAMINIFERS: <i>Globorotalia margaritae</i> , <i>Pulvinatina primalis</i> , <i>Globogaudrina altispira</i> , <i>Globigerinoides extremus</i> , <i>Globorotalia cibicenesis</i>		
			DINOFLAGELLATES: <i>Achomosphaera ramulifera</i>		
			CALCAREOUS NANOPLANKTON: <i>Reticulofenestra pseudoumbilica</i> , <i>Discocaster asymmetricus</i> , <i>D. surculus</i> , <i>D. browseri</i> , <i>Ceratolithus rugosus</i> , <i>Cyclococcolithina macintyreai</i> CORE CATCHER		
			CALCAREOUS NANOPLANKTON: <i>Reticulofenestra pseudoumbilica</i> , <i>Discocaster asymmetricus</i> , <i>D. surculus</i> , <i>D. browseri</i> , <i>Ceratolithus rugosus</i> , <i>C. tricorniculatus</i> , <i>Cyclococcolithina macintyreai</i>		



Hole 102, Core 15 (548m to 549m)

AGE	ZONE	LITHOLOGY	DIAGNOSTIC FOSSILS		
EARLY PLIOCENE	Sphaeroidinella dehiscens/Globogaudrina altispira N. 19 Discocaster asymmetricus				
		1		CN	
		CC		F, D, CN, SS	
			Hemipelagic carbonaceous silty clay; firm, dark greenish gray (564/1) with mottling to grayish olive (104/2). General Comments: -Voids due to gas expansion. -Siderite common. -Carbonate average 20%.		
			PLANKTONIC FORAMINIFERS: <i>Globigerina nepenthes</i> , <i>Globorotalia limbata</i> , <i>Gl. margaritae</i> , <i>Globogaudrina altispira</i> , <i>Globigerinoides extremus</i>		
			CALCAREOUS NANOPLANKTON: <i>Discocaster asymmetricus</i> , <i>D. browseri</i> , <i>D. surculus</i> , <i>D. pentaradiatus</i> , <i>Sphenolithus abies</i> , <i>Ceratolithus rugosus</i> , <i>Reticulofenestra pseudoumbilica</i> CORE CATCHER		
			DINOFLAGELLATES: <i>Achomosphaera ramulifera</i>		
			CALCAREOUS NANOPLANKTON: <i>Discocaster asymmetricus</i> , <i>D. browseri</i> , <i>D. pentaradiatus</i> , <i>D. surculus</i> , <i>Ceratolithus rugosus</i> , <i>Reticulofenestra pseudoumbilica</i>		



Hole 102, Core 16 (584m to 585m)						
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY
EARLY PLIOCENE				F, CN SS		Hemipelagic, carbonaceous, silty clay; indurated, dark greenish gray (564/1).
LATE MIocene	Globorotalia tumida/Sphaeroïdinellopsis paenedeniscens Discaster asymmetricus		1	SS		General Comments: -Concentrations of quartz and feldspar silt in burrows. -Siderite grains common. -White fluorescent material in core catcher sample. -Carbonate average 14%.

DIAGNOSTIC FOSSILS

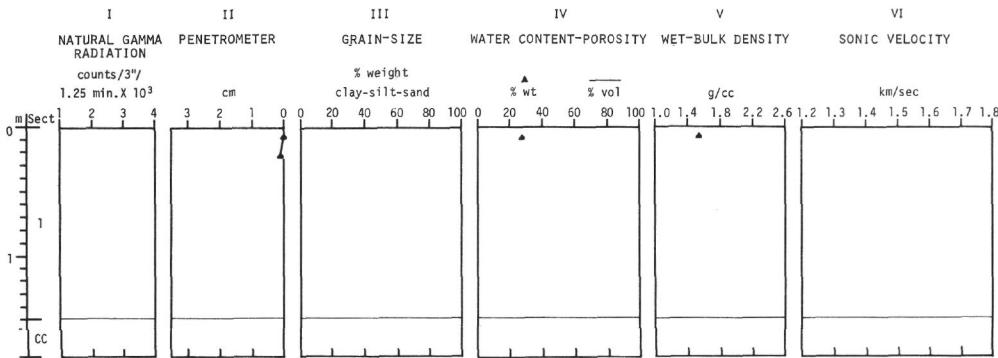
CALCAREOUS NANOPLANKTON:
Reticulofenestra pseudoumbilica,
Sphenolithus abies,
Discaster asymmetricus,
D. browneri,
D. surculatus,
Ceratolithus tricorniculatus

PLANKTONIC FORAMINIFERS:
Globorotalia plesiotumida,
Turborotalita multiloba,
Globigerina repentina,
Globorotalia limbata (dextral),
Globorotalia margaritae

CORE CATCHER

PLANKTONIC FORAMINIFERS:
Globorotalia plesiotumida,
Gl. merotumida,
Gl. lembata (dextral),
Gl. margaritae,
Turborotalita multiloba,
Globigerina repentina,
Globoquadrina altispira

CALCAREOUS NANOPLANKTON:
Reticulofenestra pseudoumbilica,
Sphenolithus abies,
Discaster asymmetricus,
D. exilis,
D. browneri,
D. challengeri,
Ceratolithus tricorniculatus



Hole 102, Core 17 (618m to 619m)						
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY
LATE MIocene	Globorotalia plesiotumida N.17 Ceratolithus tricorniculatus		1	-SS		Hemipelagic, carbonaceous silty clay; indurated, dark greenish gray (564/1).
			CC	-F -SS		General Comments: -Sponge spicules and diatoms common; only occurrence below core 11. -Abundant pyrite spheres in core catcher. -Siderite rare. -Carbonate average 14%.

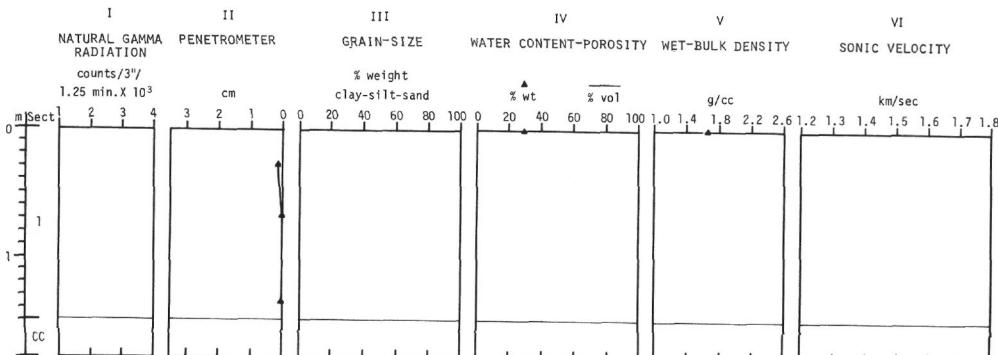
DIAGNOSTIC FOSSILS

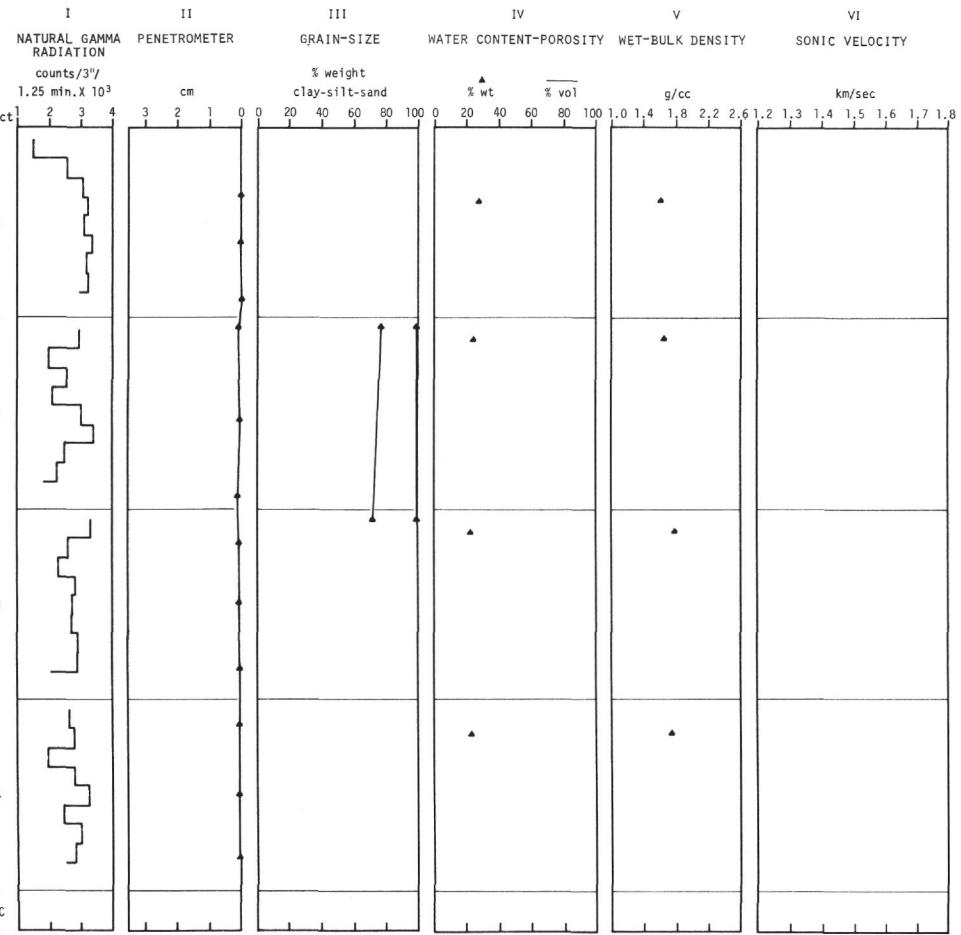
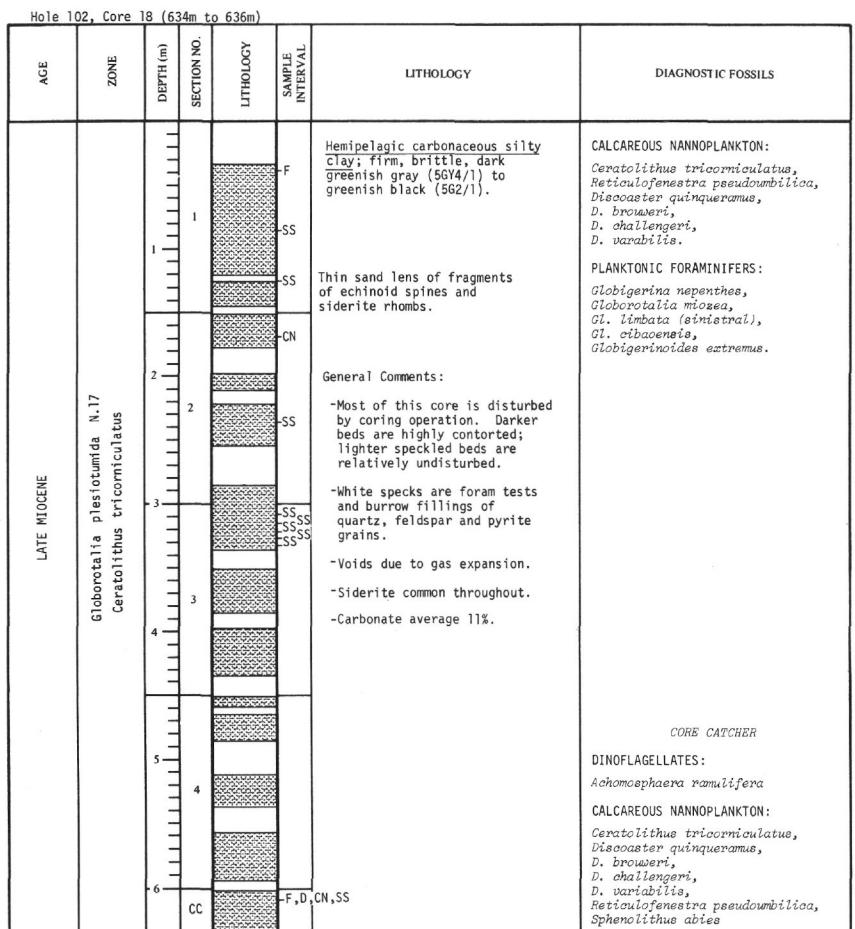
PLANKTONIC FORAMINIFERS:
Globigerinoides extremus,
Globorotalia margaritae,
Gl. limbata (sinistral),
Globigerina nepenthes,
Turborotalita multiloba

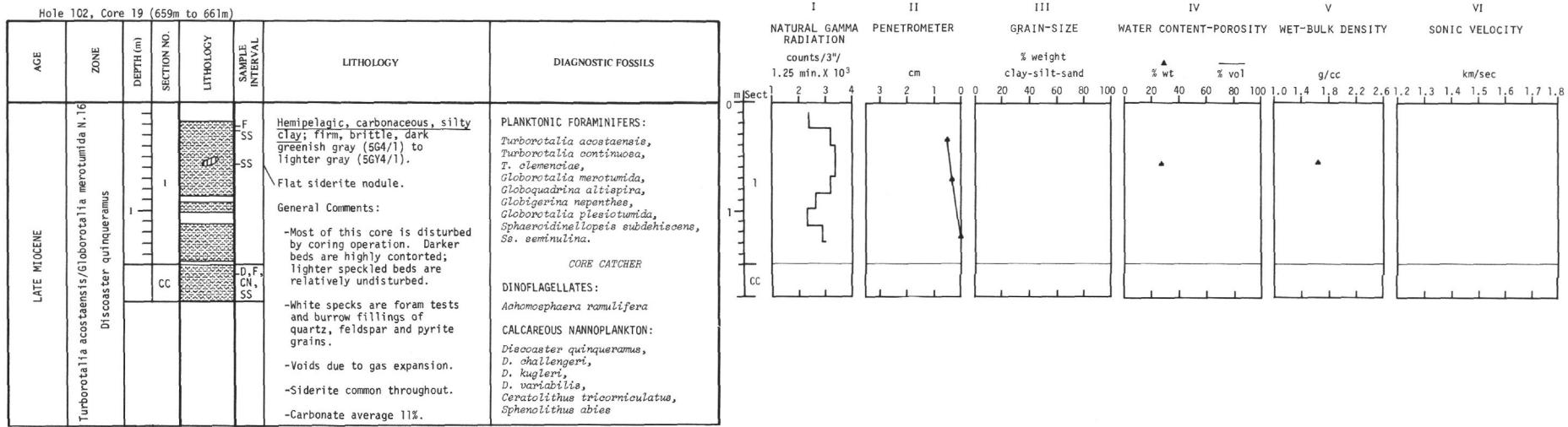
CORE CATCHER

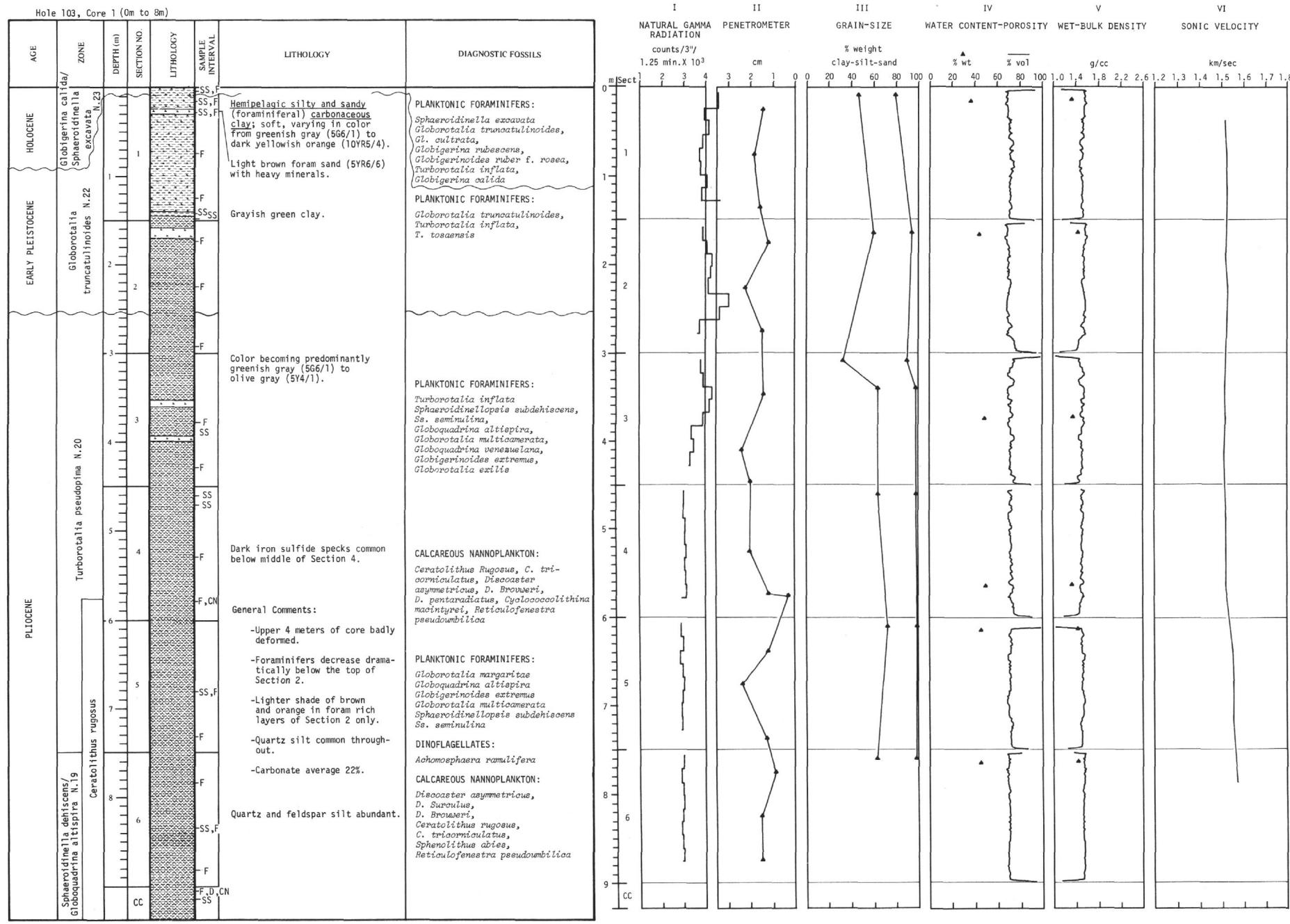
DINOFLAGELLATES:
Achomosphaera ramulifera

CALCAREOUS NANOPLANKTON:
Ceratolithus tricorniculatus,
Discaster challengeri,
D. exilis,
D. variabilis,
D. kugleri,
D. quinquevittatus,
Reticulofenestra pseudoumbilica

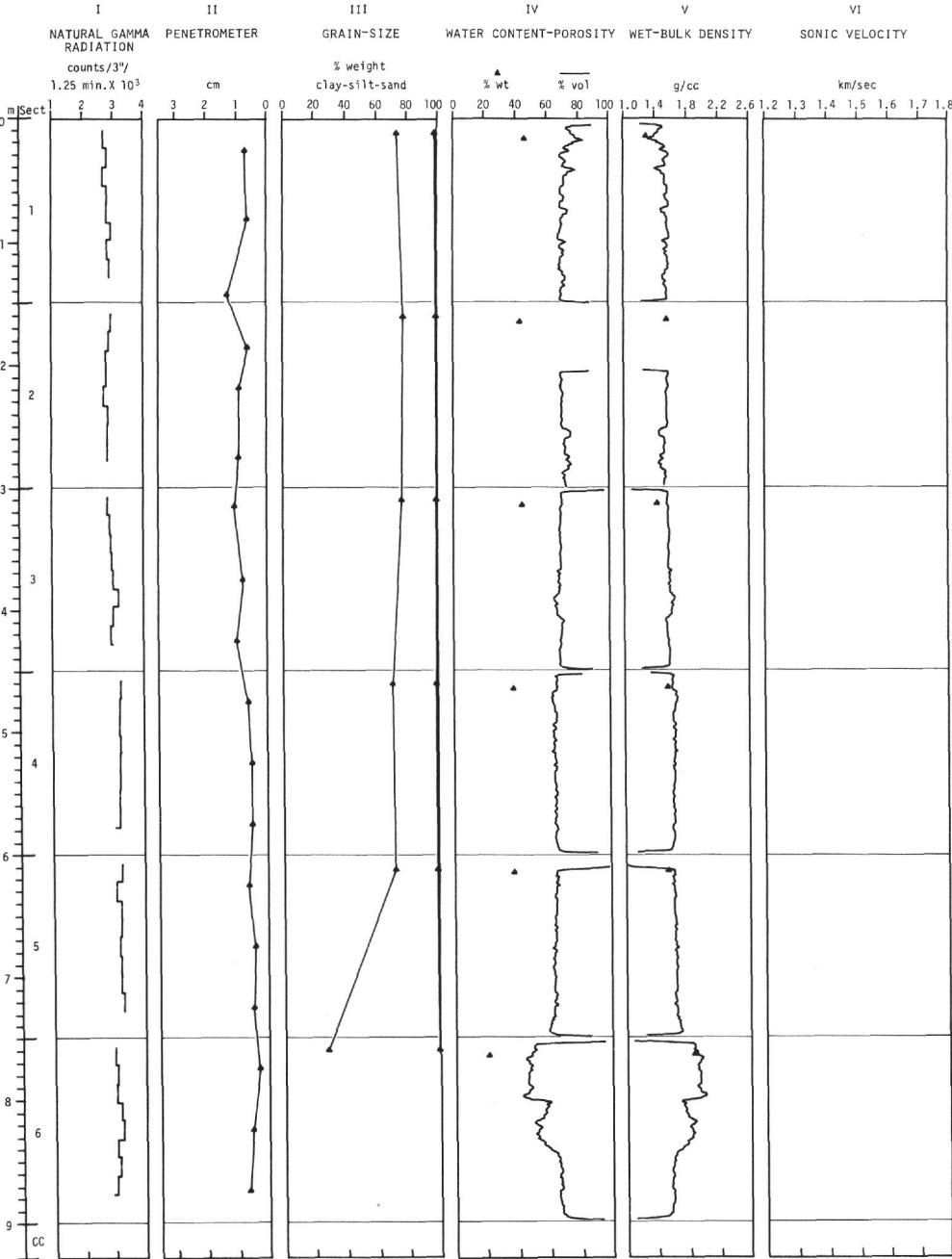






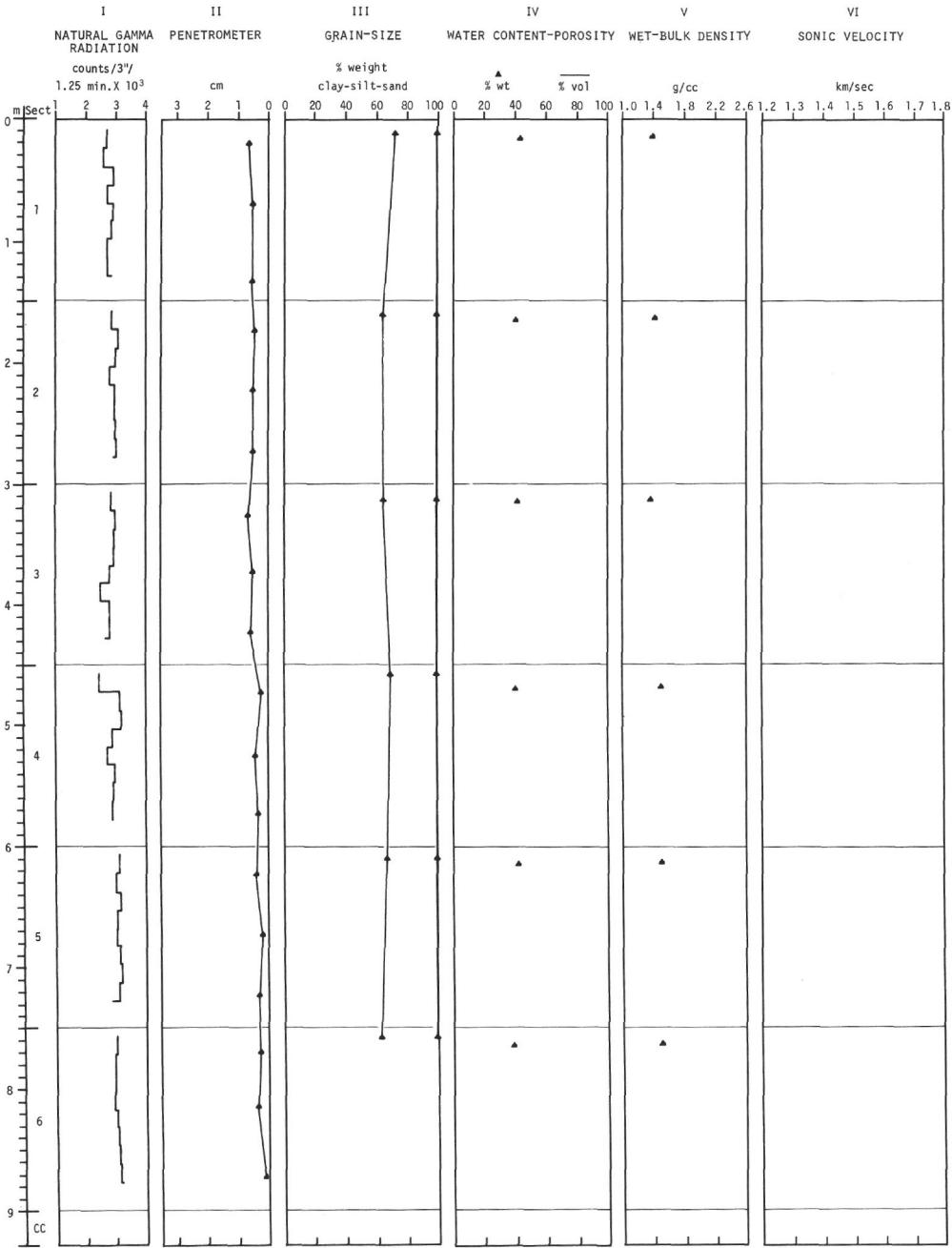


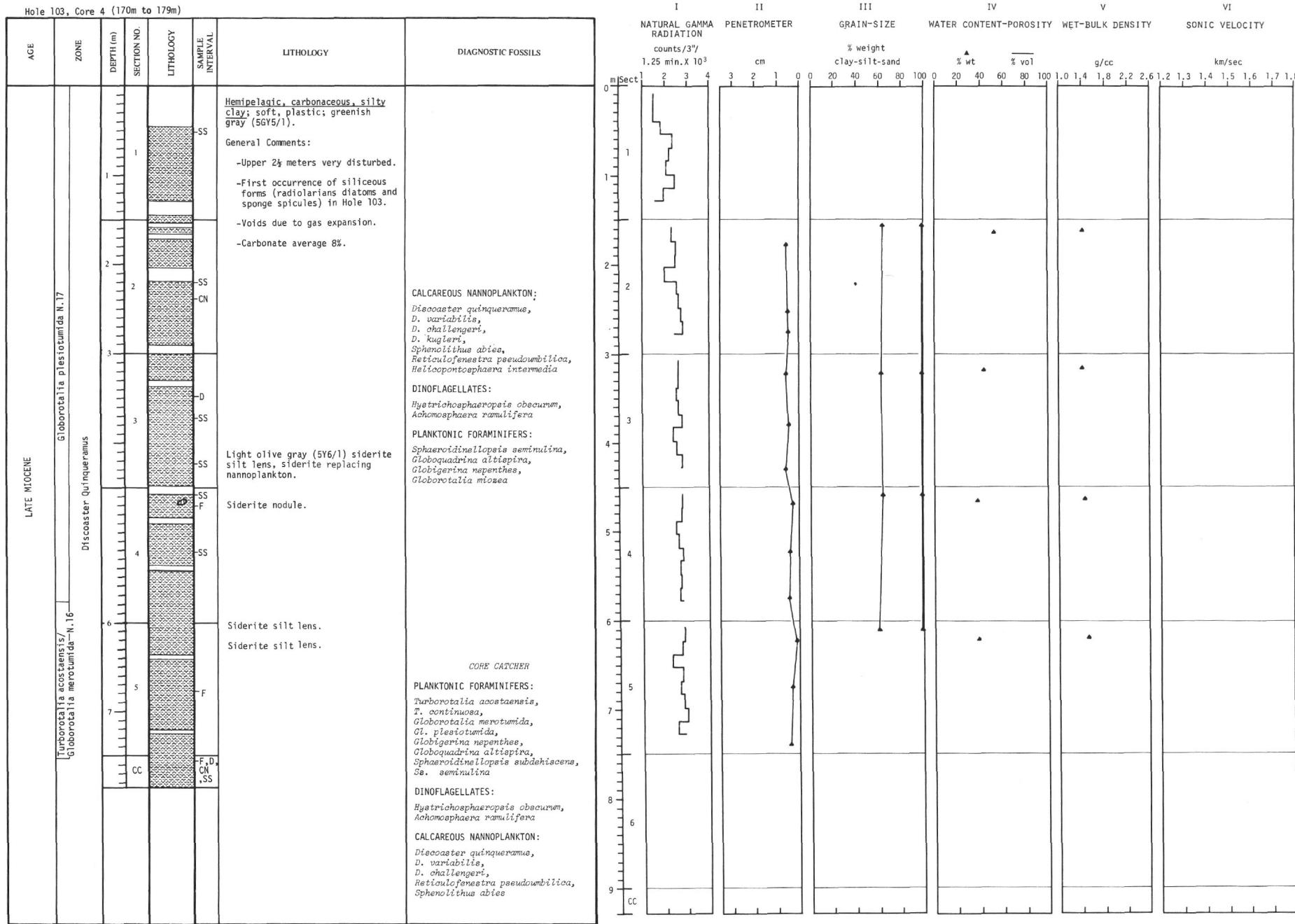
Hole 103, Core 2 (39m to 48m)						
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	DIAGNOSTIC FOSSILS
EARLY PLIOCENE	Ceratolithus tricorniculatus	1	1	SS-F	Hemipelagic carbonaceous silty clay; soft, plastic, medium bluish gray (5B5/1) mottled with yellowish gray (5Y8/1) and greenish gray (5Gy6/1).	General Comments: -Foraminifers rare throughout core. -Core has strong odor of H ₂ S. -Pyrite common to abundant throughout core. -Only concentration of dolomite in Outer Ridge holes - Section 6. -Carbonate average 20%.
LATE MIocene	Globorotalia tumida/Sphaeroidine opsis paeniedensis N. 18	2	2	SS-F,CN		PLANKTONIC FORAMINIFERS: <i>Globigerina nepenthes</i> , <i>Globorotalia cibicenensis</i> , <i>Turborotalia acostensis</i> , <i>Globigerinoides nubila</i> , <i>Globorotalia plectostomida</i> , <i>Sphaeroidine opsis seminudina</i> , <i>S. subdeltoides</i>
		3		SS-F,CN		CALCAREOUS NANNOPLANKTON: <i>Discocaster quinqueramus</i> , <i>D. browseri</i> , <i>D. exilis</i> , <i>D. surculus</i> , <i>Reticulofenestra pseudounbilica</i> , <i>Ceratolithus tricorniculatus</i> , <i>Sphenolithus abies</i>
		4		SS	Small pyrite lined burrows, also quartz and feldspar lined burrows.	
		5		SS		
		6		SS		
		7		SS		
		8		SS	Dolomite zone begins here and continues to end of core. Dolomite in form of rhombs, very abundant, dispersed in the clay.	
		6		SS	Color becoming dark greenish gray (5G4/1) and medium bluish gray (5B5/1) with mottling of yellowish gray (5Y8/1) and greenish gray (5Gy6/1).	CORE CATCHER
		CC		F,D,CN SS		PLANKTONIC FORAMINIFERS: <i>Globorotalia merotumida</i> <i>Globigerinoides nubila</i> <i>Turborotalia acostensis</i> <i>Globigerina nepenthes</i> <i>Globorotalia margaritae</i> <i>Globosphaera altispira</i>
						DINOFLAGELLATES: <i>Achomosphaera ramulifera</i>
						CALCAREOUS NANNOPLANKTON: <i>Ceratolithus tricorniculatus</i> , <i>Discocaster quinqueramus</i> , <i>D. exilis</i> , <i>D. surculus</i> , <i>Sphenolithus abies</i> , <i>Reticulofenestra pseudounbilica</i>

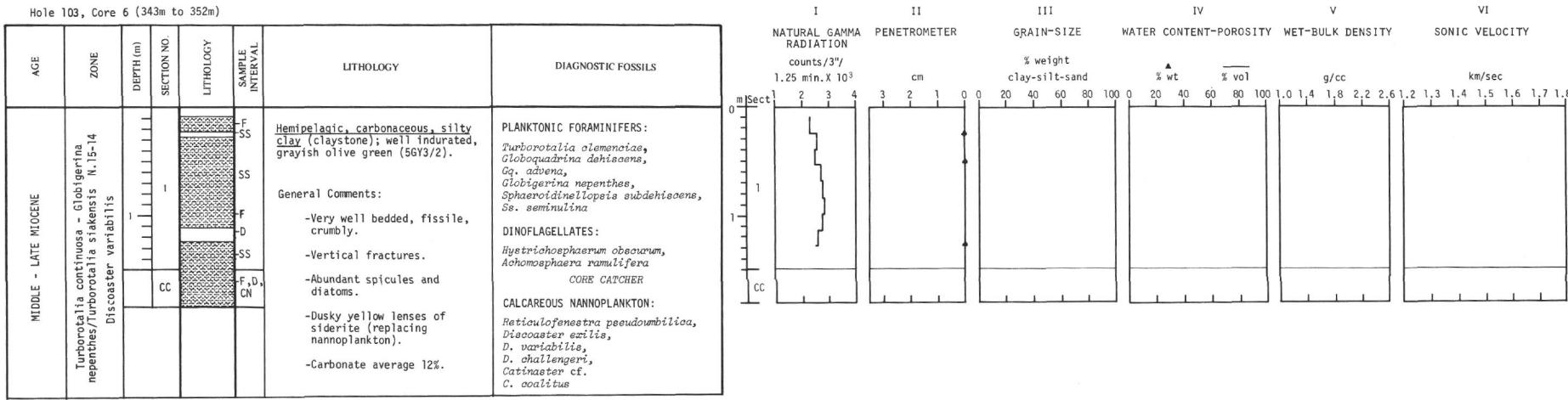
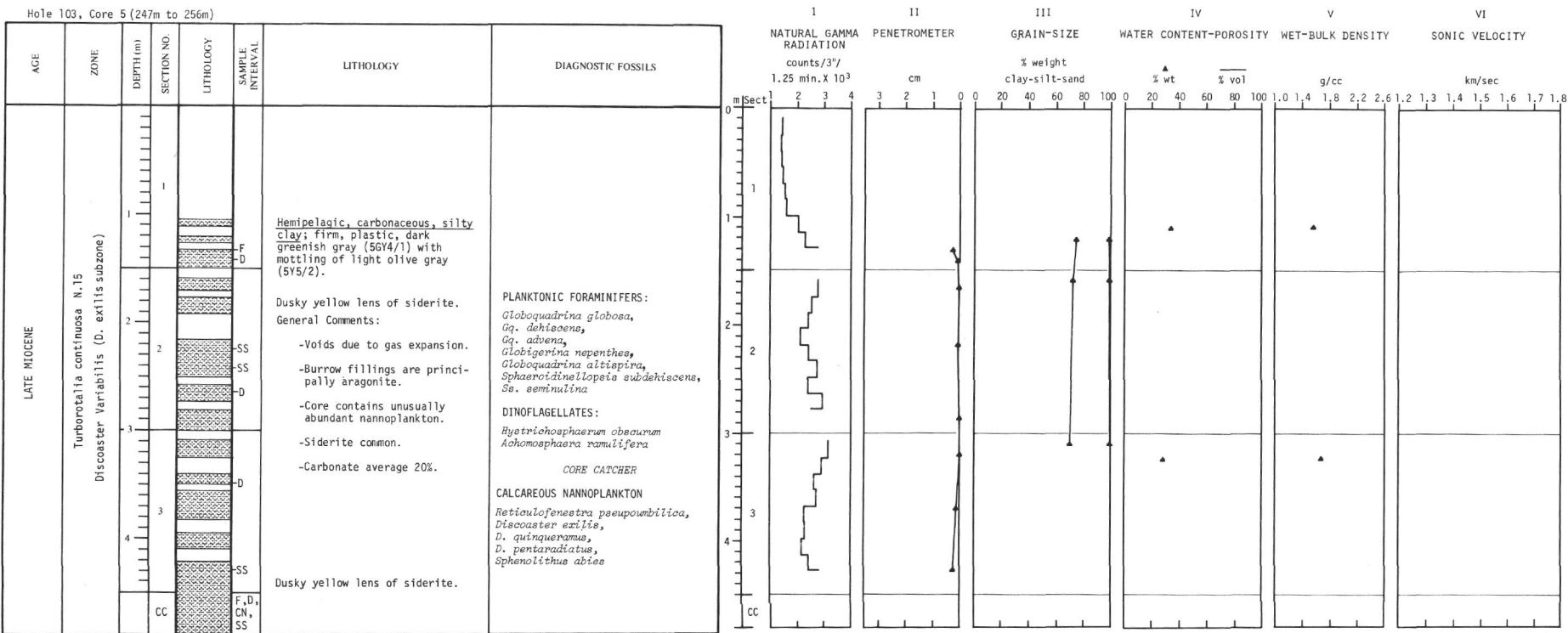


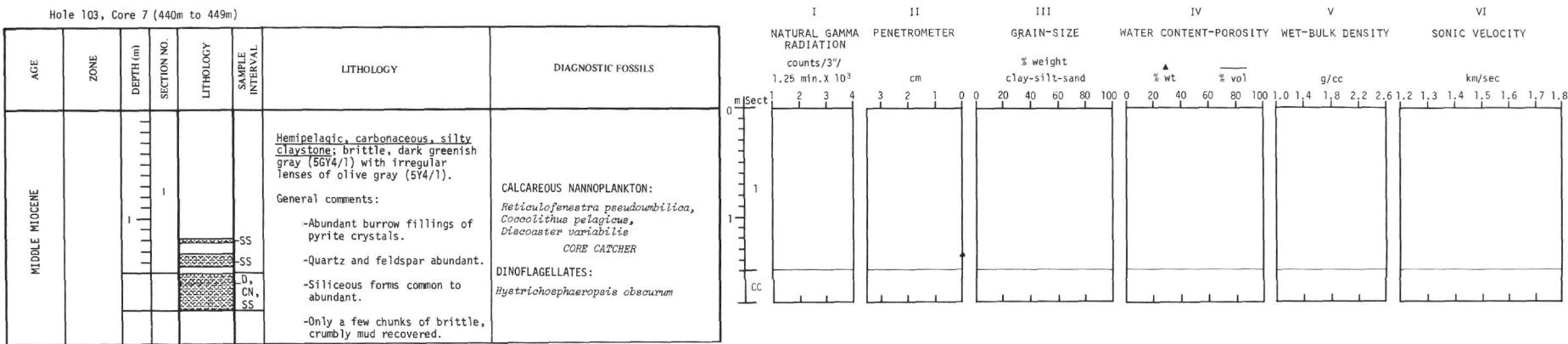
Hole 103, Core 3 (94m to 103m)

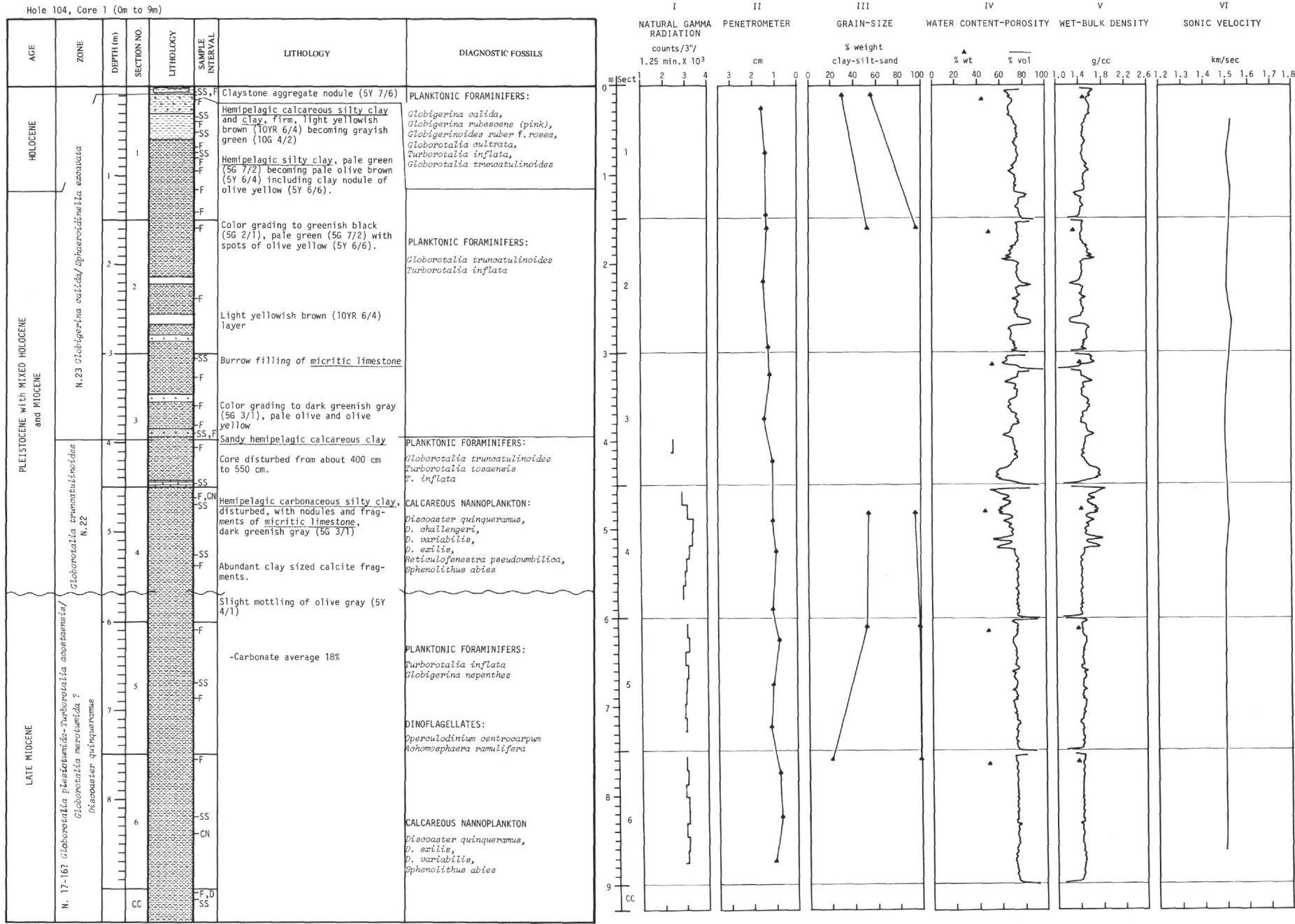
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE Miocene	Globorotalita plesiotumida N.17 Ceratolithus tricorniculatus	1		F -SS -SS		Burrow filled with siderite Hemipelagic, carbonaceous, silty clay; soft, plastic, dark greenish gray (SGY4/1) with mottling of light olive gray (SG5/2).	
		2		-SS		Dusky yellow siderite lens.	
		2		CN SS			
		3				General Comments: -Voids due to gas expansion. -First downhole occurrence of siderite in hole 103. -Carbonate average 4%.	CALCAREOUS NANOPLANKTON: <i>Reticulofenestra pseudumbilica</i> , <i>Cyclococcolithina mactinypret</i> , <i>Discosaster quinqueramus</i> , <i>D. variabilis</i> , <i>D. challengerii</i> , <i>D. exilis</i> , <i>Sphenolithus abies</i>
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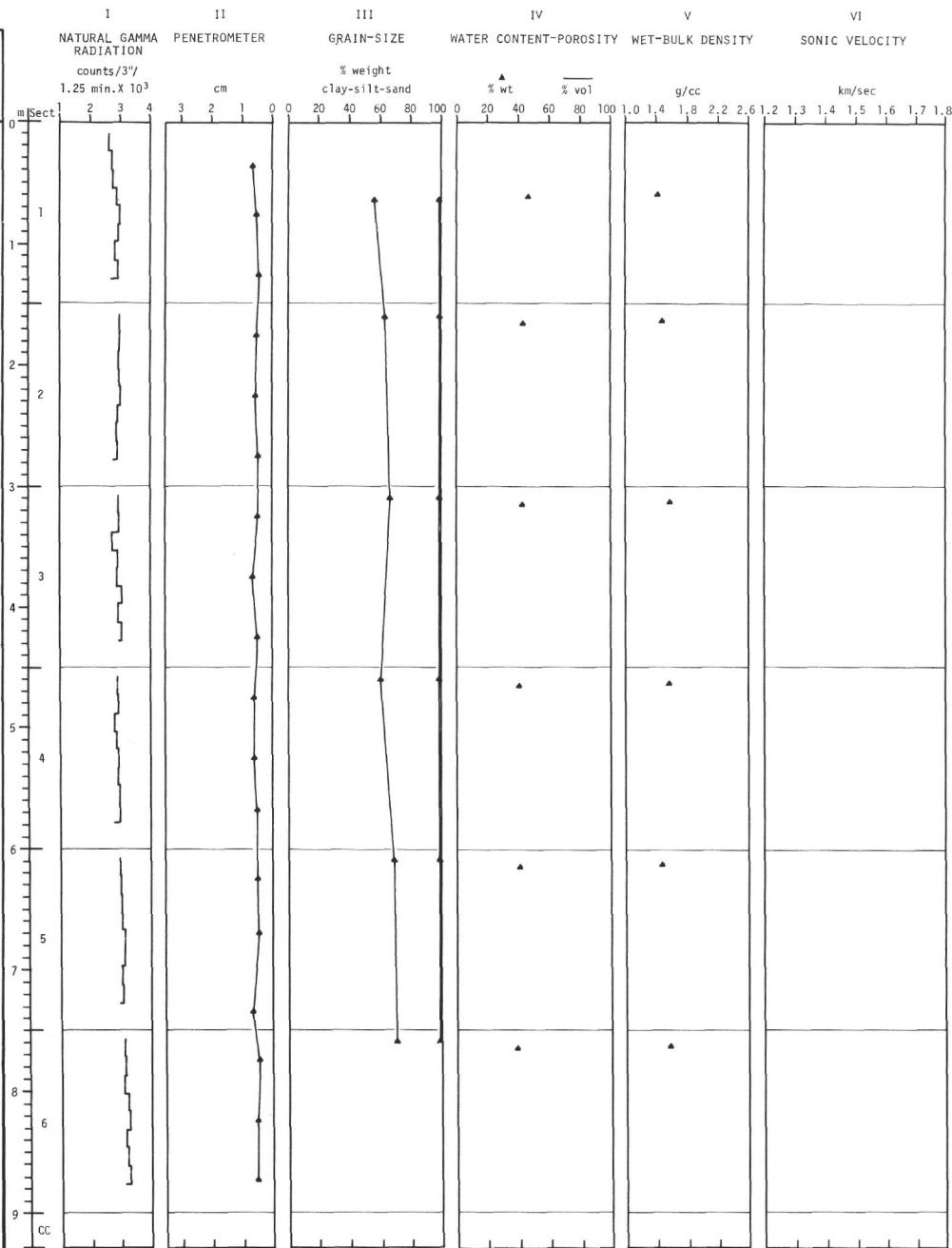




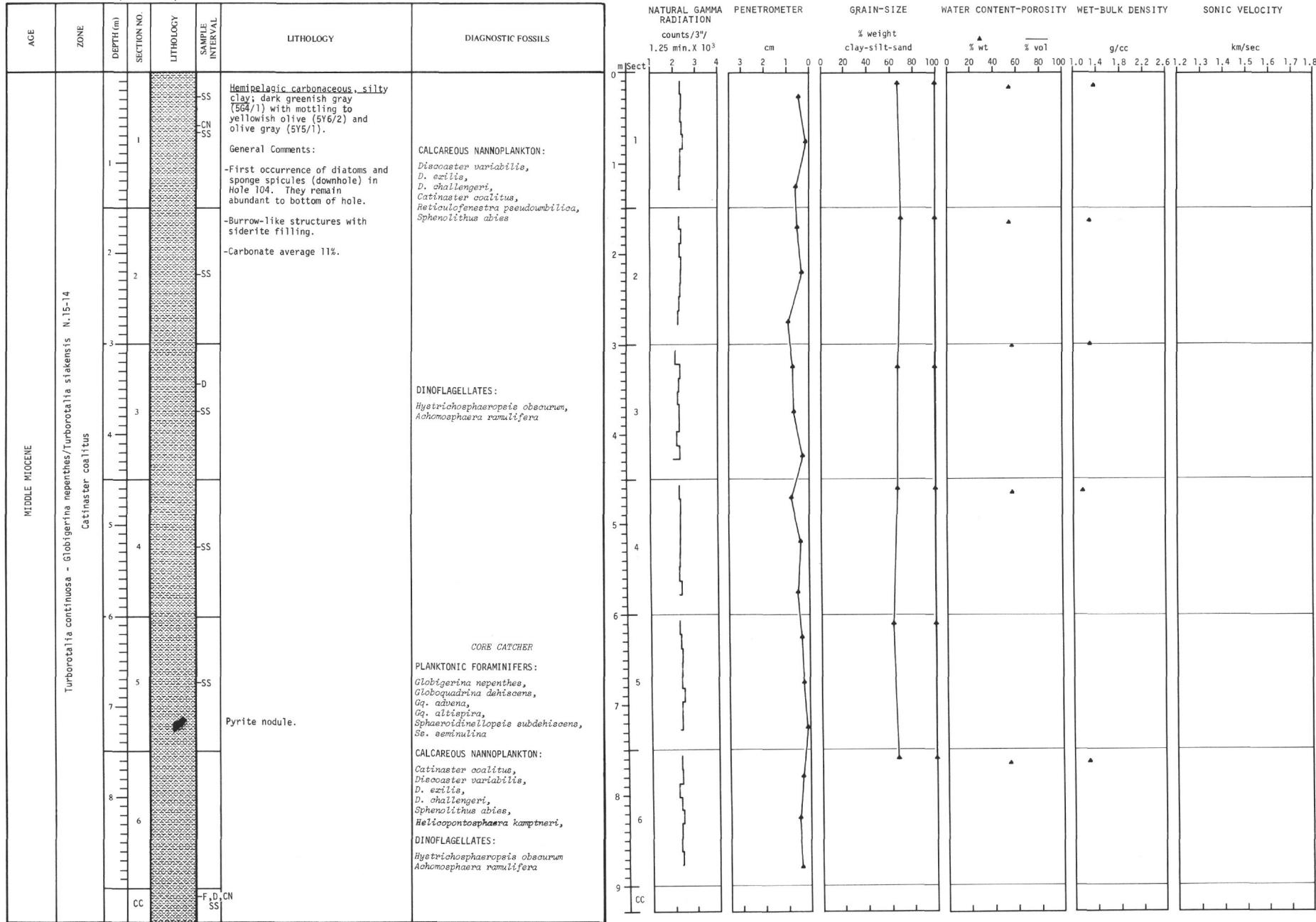


Hole 104, Core 2 (36m to 45m)

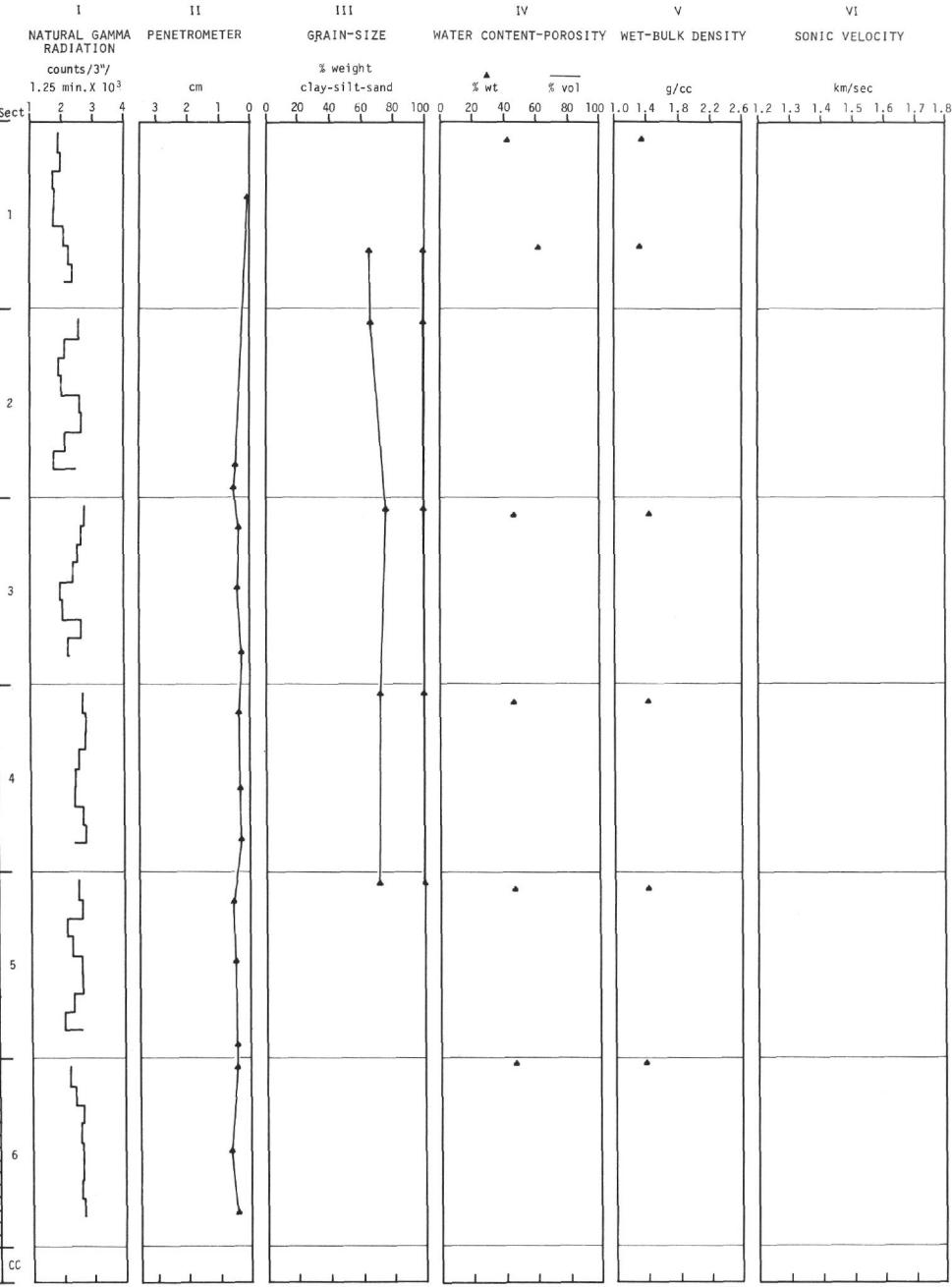
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE MIocene	Turborotalia continuosa N. 15 Discocaster hamatus	1	1	-SS -F	-SS	Hemipelagic carbonaceous silty clay; soft, plastic, dark greenish gray (5G4/1) with grayish olive mottling (10Y4/2). General Comments: -Upper 10cm contains fragments of gray lithographic limestone. -Abundant pyrite burrow fillings. -Strong odor of H ₂ S. -Pyrite common to abundant throughout. -Unusually abundant nannoplankton. -Carbonate average 12%.	PLANKTONIC FORAMINIFERS: <i>Globigerina nepenthes</i> , <i>Globoquadrina dehiscens</i> , <i>Gg. advena</i> , <i>Sphaeroidinellopsis subdehiscens</i> , <i>Ss. seminudina</i> CALCAREOUS NANNOPLANKTON: <i>Reticulofenestra pseudoumbilica</i> , <i>Sphenolithus abies</i> , <i>Discocaster hamatus</i> , <i>D. variabilis</i> , <i>D. brouweri</i> , <i>D. exilis</i> , <i>D. quinqueramus</i> DINOFLAGELLATES: <i>Hystriochospaeropsis obscurum</i> , <i>Achomosphaera ramulifera</i> , <i>A. cf. triangulata</i>
		2	2	-SS	-CN -D		
		3	3	-SS			
		4	4	-SS			
		5				Color becoming medium bluish gray (5B5/1) with moderate mottling of light bluish gray (5B7/1).	
		6		-SS			
		7					
		8					
		9					
		CC			-F,D,CN -SS		



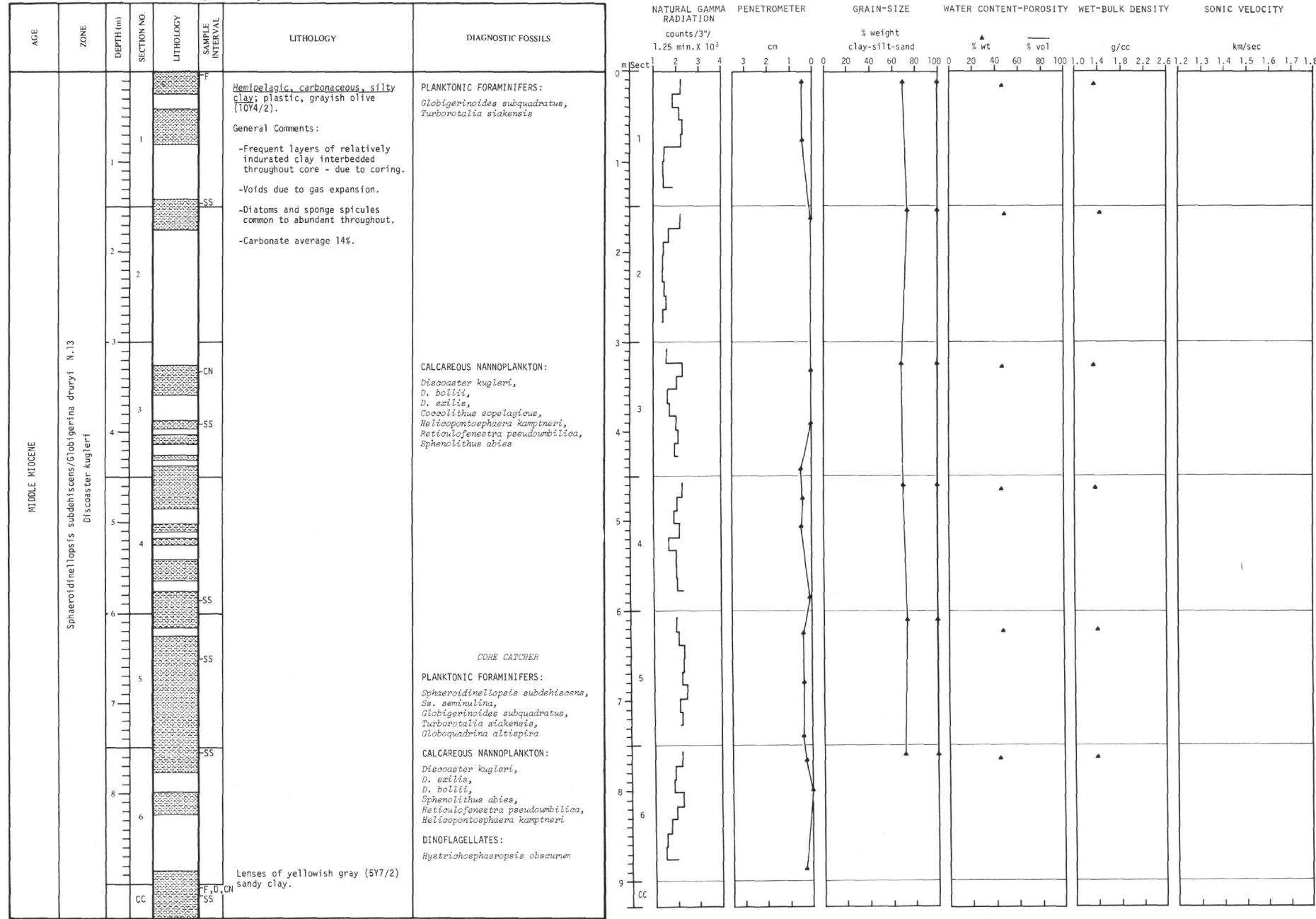
Hole 104, Core 3 (62m to 71m)

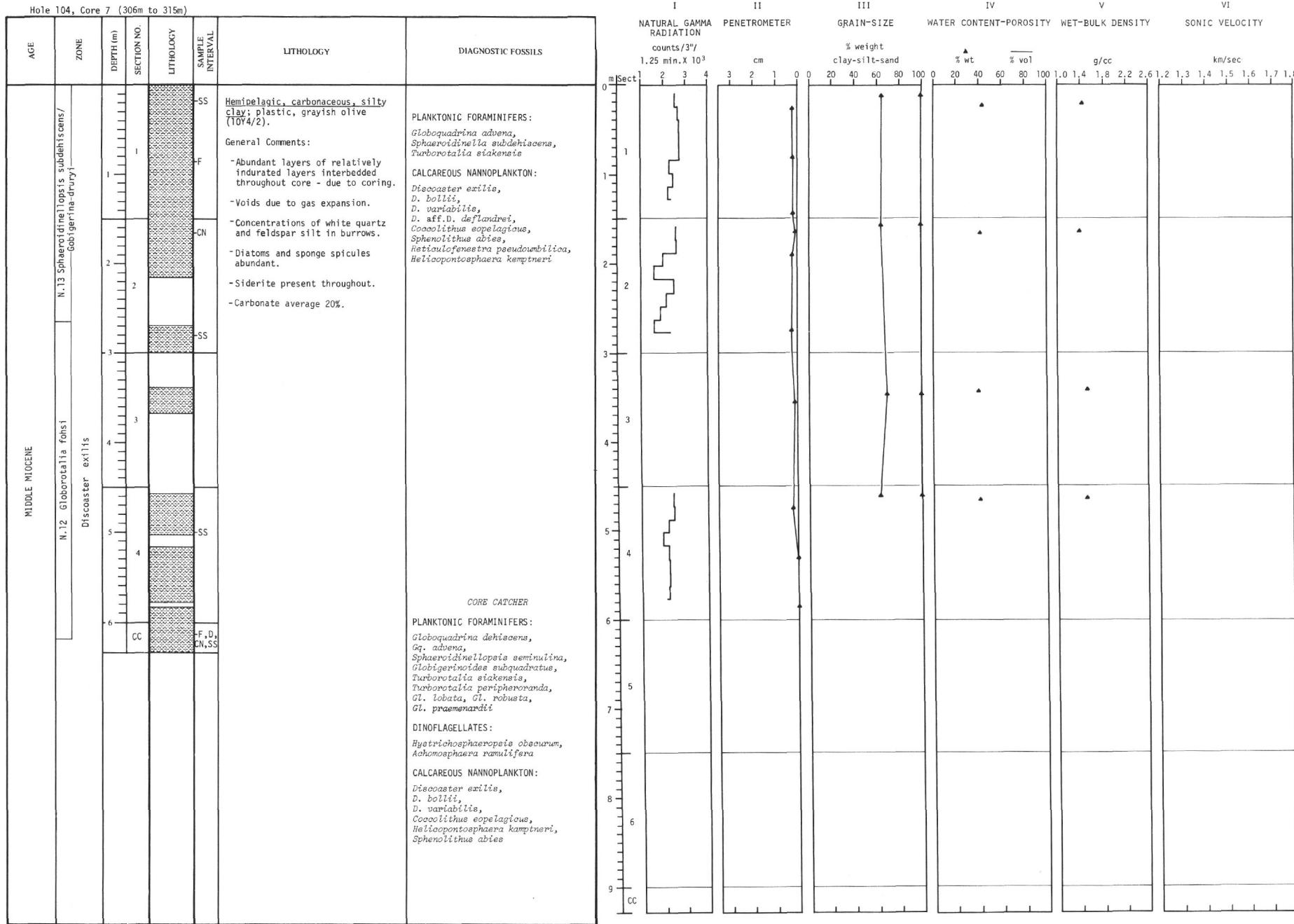


Hole 104, Core 4 (133m to 142m)					
AGE	ZONE	DEPTH (m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL
MIDDLE MIocene	Turborotalia continua - Globigerina nepenthes/Turborotalia stakensis N. 15-14 Discocaster kugleri	133	1	SS	Hemipelagic carbonaceous silty clay; soft, greenish gray (5G5/1) with dusky yellow mottling (5Y6/4).
		1	1	SS	General Comments: -Upper 3 meters badly deformed. -Voids due to gas expansion. -Abundant 5cm thick layers of indurated clay - due to coring. -Siderite common to abundant in yellow material. -Abundant diatoms and sponge spicules. -Carbonate average 7%.
		2	2	SS	
		3	3	CN	
		4	4	SS	CALCAREOUS NANNOPLANKTON: <i>Discocaster bollii</i> , <i>D. kugleri</i> , <i>D. variabilis</i> , <i>Reticulofenestra pseudoumbilica</i> , <i>Sphenolithus abies</i> , <i>Helicopontosphaera kamptneri</i> ,
		5	5	SS	
		6	6	SS	Abundant diatoms in soft material.
		7			CORE CATCHER
		8			PLANKTONIC FORAMINIFERS: <i>Globquadrina advena</i> , <i>Sphaeroidinellopsis seminulina</i> , <i>Ss. subdehiscentiae</i> ,
		6			DINOFLAGELLATES: <i>Hystriocoelsphaeropsis obscurum</i> , <i>Achomosphaera ramulifera</i>
		7			CALCAREOUS NANNOPLANKTON: <i>Discocaster kugleri</i> , <i>D. bollii</i> , <i>D. variabilis</i> , <i>Coccilithus sphaericus</i> , <i>Reticulofenestra pseudoumbilica</i> , <i>Sphenolithus abies</i>
		8			
		9			
		CC			F,D,CN SS

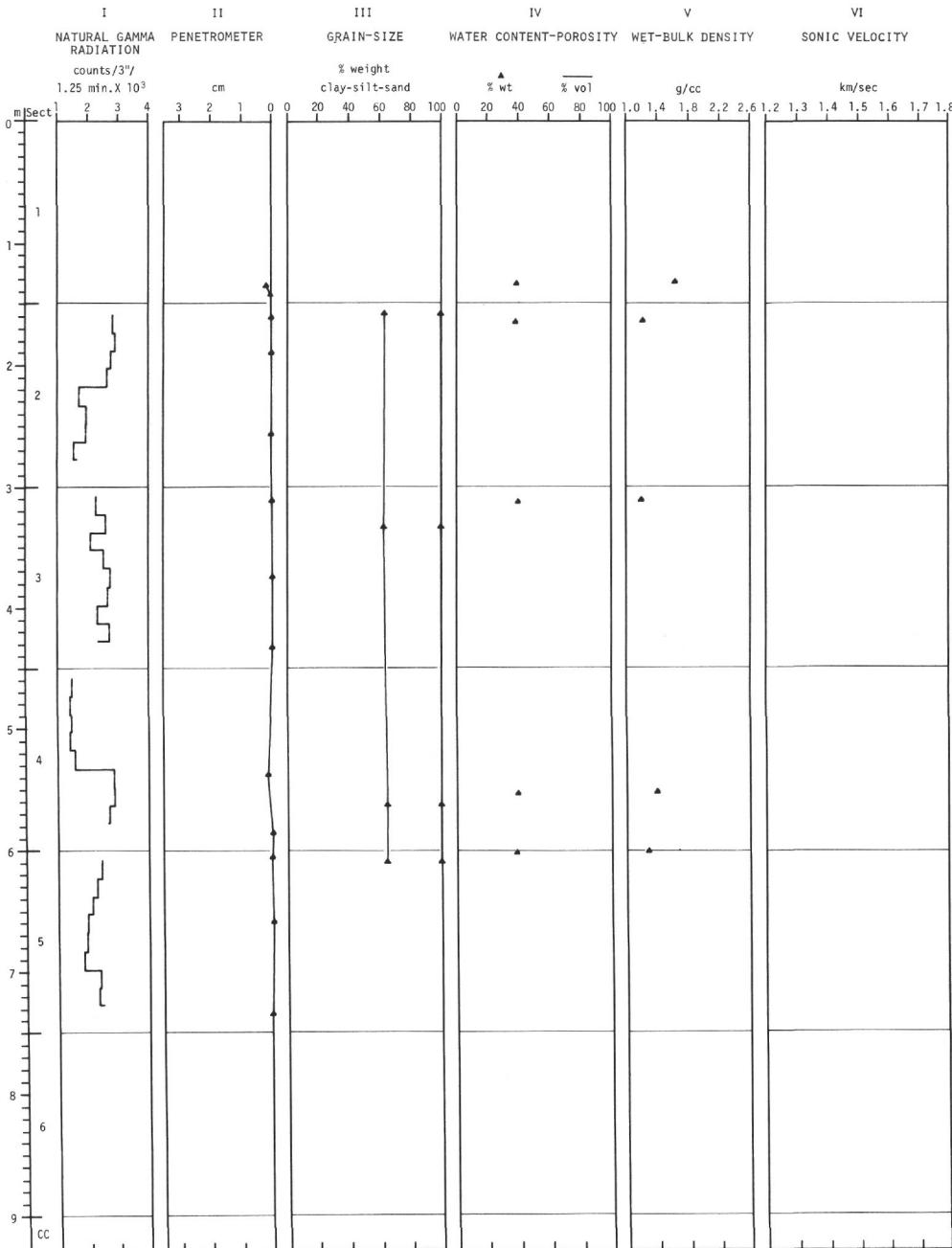


Hole 104, Core 6 (219m to 228m) (no recovery at Core 5)



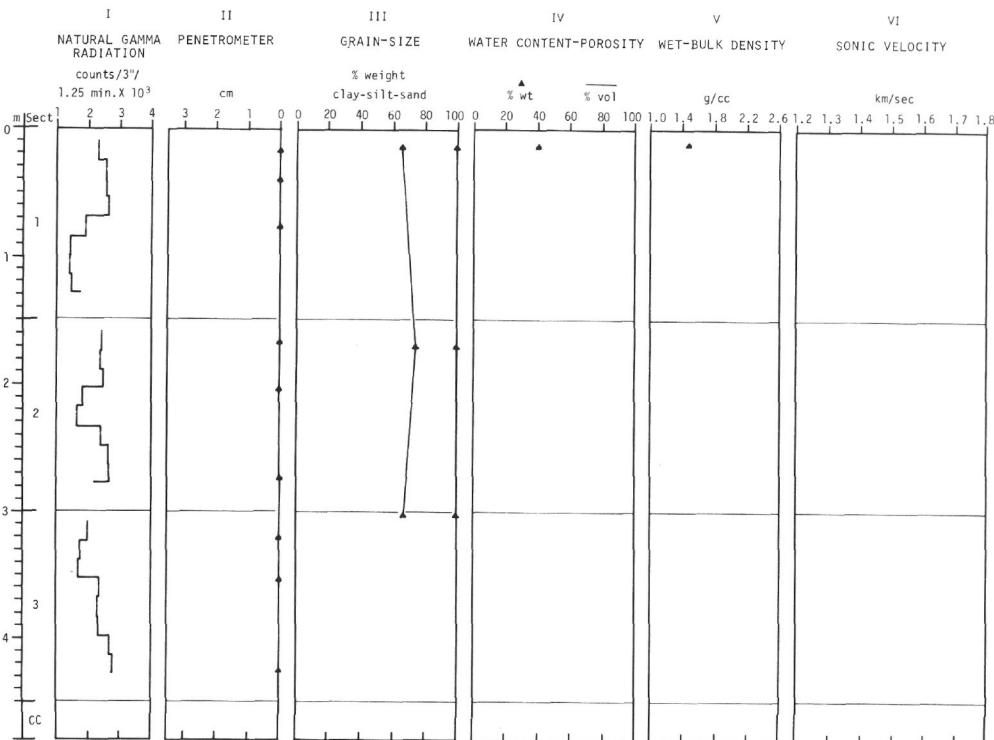


AGE	ZONE	DEPTH(m)	SECTION NO.	LITHOLOGY	SAMPLE INTERVAL	LITHOLOGY		DIAGNOSTIC FOSSILS	
MIDDLE MIocene	Turborotalia praefohsi N. 11 Discocaster exilis	1	1	SS					
		2	2	SS CN		Hemipelagic, carbonaceous, silty clay, firm, grayish olive (10Y4/2).	CALCAREOUS NANNOPLANKTON: <i>Discocaster exilis</i> , <i>D. boliti</i> , <i>D. variabilis</i> , <i>Coccocithus epeLAGicus</i> , <i>Sphaerolithus abies</i> , <i>Helicopontosphaera kampfneri</i> .		
		3	3	SS		General Comments: -Voids due to gas expansion. -Alternate firm and plastic layering due to coring operation. -Vertical fracturing in harder layers. -Diatoms and sponge spicules abundant. -Carbonate average 20%.			
		4	4	SS					
		5							
		6	4	SS					
		7	5	SS					
		CC	CC	F, D, CN					



Hole 104, Core 9 (395m to 504m)

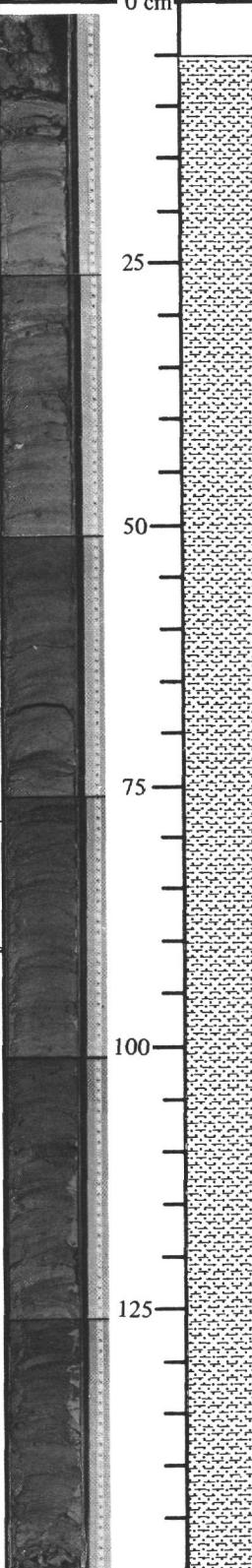
AGE	ZONE	LITHOLOGY	DIAGNOSTIC FOSSILS
DEPTH (m)	SECTION NO.	SAMPLE INTERVAL	
MIDDLE Miocene	Orbulina suturalis/Turborotalia peripheroronda? N. 9?		
	Discaster exilis		
1	1	SS	Hemipelagic carbonaceous silty clay, firm, dry, grayish olive (10Y4/2). General comments: -Voids due to gas expansion. -Thin, lensy bedding apparent. -Diatoms and sponge spicules abundant. -Carbonate average 20%.
2	2	SS	
3	3	SS-F	
4	3	CN	
CC	CC	F, D, CN, SS	



Hole 104, Core 10 (615m to 617m)

AGE	ZONE	LITHOLOGY	DIAGNOSTIC FOSSILS
DEPTH (m)	SECTION NO.	SAMPLE INTERVAL	
MIDDLE Miocene	? Orbulina suturalis/Turborotalia peripheroronda? Sphenolithus heteromorphus		
1	1	SS-D, CN, SS-D	Ankerite, very hard. Hemipelagic silty claystone; greenish gray (5G5/1), gassy. General Comments: -Bedding disturbed by abundant burrows. -Diatoms and sponge spicules - abundant. -Radiolarians - common.
CC	CC	F, D, CN, SS	

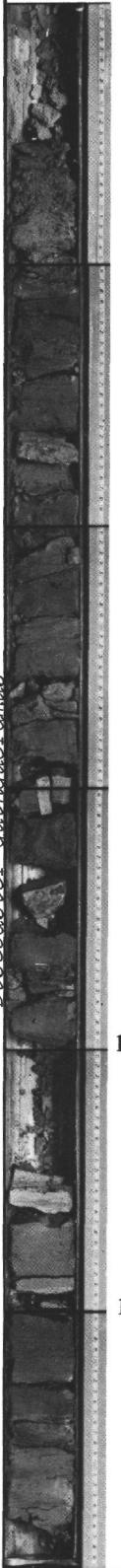
Hole 102, Core 11, Sect. 3

AGE	ZONE	LITHOLOGY	SAMP. INT.	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE PLIOCENE	<i>Turborotalia tenuitheca</i> N. 21 <i>Reticulofenestra pseudounbilica</i>	 <p>0 cm 25 50 75 100 125</p>	CN ZF	<p>Hemipelagic carbonaceous silty clay; dark greenish gray (5G 4/1), moderately indurated but plastic. Siderite, sponge spicules, clay minerals abundant. Quartz, pyrite, plant debris common. Diatoms, mica, glass and glauconite rare.</p> <p>Light colored specks are siderite concentrations.</p> <p>Voids due to gas expansion.</p>	<p>CALCAREOUS NANNOPLANKTON:</p> <p><i>Discoaster brouweri</i>, <i>D. pentaradiatus</i>, <i>D. asymmetricus</i>, <i>Reticulofenestra pseudounbilica</i>, <i>Cyclococcolithina macintyrei</i>, <i>Helicopontosphaera sellii</i></p> <p>PLANKTONIC FORAMINIFERS:</p> <p><i>Globigerinoides extremus</i>, <i>Globorotalia miocenica</i>, <i>Gl. multicamerata</i>, <i>Globoquadrina altispira</i>, <i>Gq. venezuelana</i></p>

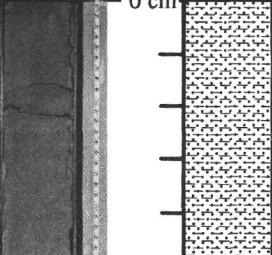
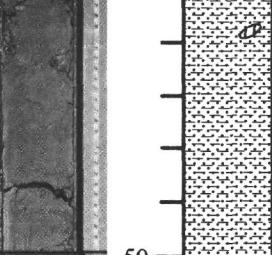
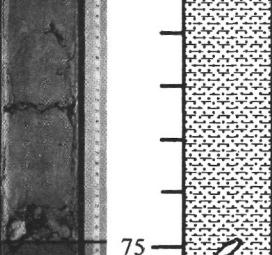
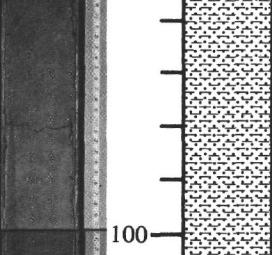
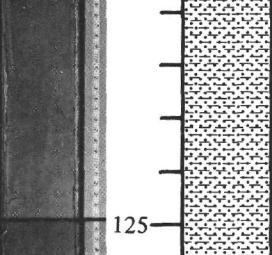
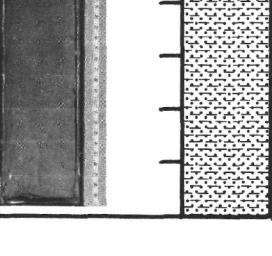
Hole 102, Core 17, Sect. 1

AGE	ZONE	LITHOLOGY	SAMP. INT.	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE MIocene	<i>Globorotalia plesiotamida</i> N. 17 <i>Ceratolithus tricorniculatus</i>		0 cm SS	Hemipelagic, siliceous, silty clay; dark greenish gray (5G 4/1), indurated.	
			25	Clay minerals abundant. Radiolarians, plant debris, quartz, mica glauconite rare.	
			50	Nannoplankton, calcite fragments, diatoms, sponge spicules and pyrite common.	
			75		
			100	Clay minerals abundant, nannoplankton, diatoms, sponge spicules, plant debris, pyrite common. Calcite fragments, siderite quartz, glauconite rare.	CALCAREOUS NANNOPLANKTON: <i>Ceratolithus tricorniculatus</i> , <i>Discoaster quinqueramus</i> , <i>D. exilis</i> , <i>D. challengerii</i> , <i>Reticulofenestra pseudoumbilica</i> , <i>Sphenolithus abies</i> .
			125	Pyrite spheres and clay minerals abundant. Diatoms, sponge spicules, common.	PLANKTONIC FORAMINIFERS: <i>Globoquadrina altispira</i> , <i>Sphaeroidinellopsis seminulina</i> , <i>Globorotalia margaritae</i> , <i>Globigerina nepenthes</i>
				Nannoplankton, quartz, mica, dolomite rhombs, rare.	

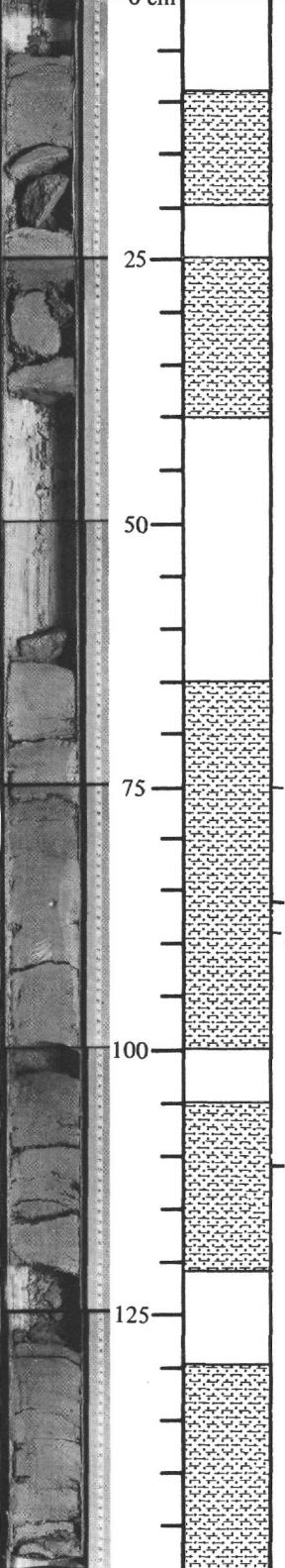
Hole 102, Core 19, Sect. 1

AGE	ZONE	LITHOLOGY	SAMP. INT.	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE MIocene <i>Turborotalia acostaensis/Globorotalia merotumida N.16</i> <i>Discoaster quinqueramus</i>				<p>0 cm</p> <p>Hemipelagic silty carbonaceous clay, varying between dark greenish gray (5G 4/1) and light greenish gray (5GY 4/1).</p> <p>F: Light beds are relatively undisturbed - dark layers reveal distortion caused by coring operation.</p> <p>Voids are due to gas expansion.</p> <p>50</p> <p>50 cm mark (Sidérite nodule).</p> <p>White specks are concentrations of foraminifers or quartz and feldspar silt.</p> <p>75</p> <p>Pyrite occurs as frambooidal grains in burrow fillings.</p> <p>Clay minerals dominant. Quartz, plant debris, and nannoplankton are common.</p> <p>100</p> <p>Foraminifers rare.</p> <p>125</p>	<p>PLANKTONIC FORAMINIFERS:</p> <p><i>Globorotalia miozea</i>, <i>Globigerina nepenthes</i>, <i>Globorotalia merotumida</i>, <i>Turborotalia continuosa</i>, <i>Turborotalia clemenciae</i></p> <p>CALCAREOUS NANNOPLANKTON:</p> <p><i>Reticulofenestra pseudoumbilica</i>, <i>Discoaster quinqueramus</i>, <i>D. kugleri</i>, <i>D. exilis</i>, <i>Ceratolithus tricorniculatus</i>, <i>Helicopontosphaera Kamptneri</i></p>

Hole 103, Core 3, Section 6

AGE	ZONE	LITHOLOGY	SAMP. INT	LITHOLOGY	DIAGNOSTIC FOSSILS
				Hemipelagic silty carbonaceous clay, dark greenish gray (5GY 4/1) mottled with light olive gray (5Y 5/2), soft, plastic	
				Lens of siderite.	
				Clay minerals and quartz silt abundant. Plant debris abundant. Foraminifera & nannoplankton rare.	
LATE MIocene	<i>Globorotalia plesiotumida</i> <i>Ceratolithus tricorniculatus</i>		SS	Dolomite nodule.	
			CN F		CALCAREOUS NANNOPLANKTON: <i>Discoaster challengerii</i> , <i>D. variabilis</i> , <i>D. quinqueramus</i> , <i>D. brouweri</i> , <i>Sphenolithus abies</i> , <i>Reticulofenestra pseudoumbilica</i>
			D		PLANKTONIC FORAMINIFERS: <i>Globorotalia plesiotumida</i> <i>Gl. cibaoensis</i> , <i>Globigerina nepenthes</i> DINOFLAGELLATES: <i>Achomosphaera ramulifera</i>

Hole 103, Core 5, Section 2

AGE	ZONE	LITHOLOGY	SAMP. INT	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE MIocene	<i>L. Turborotalia continuosa</i> <i>Discoaster quinqueramus</i> N.15			<p>0 cm</p> <p>Hemipelagic silty carbonaceous clay; dark greenish gray (5GY 4/1) mottled with light olive gray (5Y 5/2) firm, plastic.</p> <p>25</p> <p>Burrow (?) fillings composed of aragonite. Voids due to gas expansion. Clay minerals and nannoplankton abundant. Plant debris and pyrite common. Dolomite and foraminifers rare.</p> <p>50</p> <p>75 SS</p> <p>N SS</p> <p>100 D</p> <p>125</p>	<p>CALCAREOUS NANOPLANKTON: <i>Discoaster variabilis</i>, <i>D. exilis</i>, <i>D. bollii</i>, <i>D. hamatus</i>, <i>D. quinqueramus</i>, <i>Sphenolithus abies</i>, <i>Reticulofenestra pseudoumbilica</i>.</p> <p>DINOFLAGELLATES: <i>Hystrichosphaeropsis obscurum</i>, <i>Achomosphaera ramulifera</i></p>

Hole 103, Core 6, Section 1

AGE	LITHOLOGY	SAMP. INT	LITHOLOGY	DIAGNOSTIC FOSSILS
ZONE				
LATE MIocene <i>Turborotalia continua</i> <i>Globigerina nepenthes</i> / <i>Turborotalia siakensis</i> <i>Discoaster variabilis</i> N.15-14		0 cm CN F SS 25 50 75 100 D 125 SS SS	Hemipelagic carbonaceous silty clay, grayish olive olive green (5GY 3/2) with lenses of dusky yellow. Evenly bedded. Sediment is very crumbly and fractured. Clay minerals abundant. Quartz, diatoms, sponge spicules, plant debris common.	CALCAREOUS NANNOPLANKTON: <i>Discoaster variabilis</i> , <i>D. exilis</i> , <i>D. challengerii</i> , <i>Helicopontosphaera itamptneri</i> , <i>Reticulofenestra pseudoumbilica</i> , <i>Scyphosphaera amphora</i> PLANKTONIC FORAMINIFERS: <i>Globoquadrina dehiscens</i> , <i>Globigerina nepenthes</i>
				PLANKTONIC FORAMINIFERS: <i>Globigerina nepenthes</i> , <i>Globoquadrina advena</i>
				DINOFLAGELLATES <i>Hystrichosphaeropsis obscurum</i> , <i>Achromosphaera ramulifera</i>

Hole 104, Core 1, Section 6

AGE	LITHOLOGY	SAMP. INT	LITHOLOGY	DIAGNOSTIC FOSSILS
LATE MIocene <i>Globorotalia plesiotumida</i> <i>Turborotalia acostaensis</i> / <i>Turborotalia merotumida</i> ? <i>Discoaster quinqueramus</i> N. 17-16?		F SS CN F,D	<p>0 cm</p> <p>Hemipelagic silty carbonaceous clay, dark greenish gray (5G 4/1) to olive gray (5Y 4/1), soft and plastic.</p> <p>Clay minerals abundant. Quartz, feldspar and organic matter common.</p>	<p>PLANKTONIC FORAMINIFERS: <i>Turborotalia inflata</i></p> <p>CALCAREOUS NANNOPLANKTON: <i>Discoaster quinqueramus</i>, <i>D. exilis</i>, <i>D. variabilis</i>, <i>Sphenolithus abies</i></p> <p>DINOFLAGELLATES: <i>Operculodinium centracarpum</i>, <i>Achomosphaera ramulifera</i></p> <p>PLANKTONIC FORAMINIFERS: <i>Turborotalia inflata</i>, <i>Globigerina nepenthes</i></p>

Hole 104, Core 10, Section 1

AGE	ZONE	LITHOLOGY	SAMP. INT.	LITHOLOGY	DIAGNOSTIC FOSSILS
MIDDLE MIocene	<i>Orbulina suturalis</i> / <i>Turborotalia peripheroronda</i> ? <i>Sphenolithus heteromorphus</i> N.9?	 <p>0 cm 25 50 75 100 125</p>	SS CN D SS D SS F,D,N	<p>Ankeritic limestone, greenish gray (5GY 5/1)</p> <p>Hemipelagic, siliceous, silty mudstone. Clay minerals, diatoms, sponge spicules abundant.</p> <p>Siderite, radiolarians</p> <p>Irregular bedding throughout.</p>	<p>CALCAREOUS NANNOPLANKTON: <i>Sphenolithus heteromorphus</i>, <i>Discoaster bollii</i>, <i>D. exilis</i> <i>Cyclococcolithina neogammation</i>, <i>Coccolithus eopelagicus</i>, <i>Helicopontosphaera kampfneri</i></p> <p>DINOFLAGELLATES: <i>Achomosphaera ramulifera</i>, <i>Svalbardella</i> sp. <i>Hystrichosphaeropsis obscurum</i></p> <p>CALCAREOUS NANNOPLANKTON: <i>Sphenolithus heteromorphus</i>, <i>Helicopontosphaera kampfneri</i> <i>Discoaster exilis</i>, <i>D. bollii</i>, <i>Cyclococcolithina neogammation</i>, <i>Reticulofenestra pseudoumbilica</i></p> <p>DINOFLAGELLATES: <i>Achomosphaera ramulitera</i>, <i>Svalbardella</i> sp., <i>Spiniferites</i> sp. A</p> <p>PLANKTONIC FORAMINIFERS: <i>Orbulina suturalis</i>, <i>Turborotalia siakensis</i>, <i>T. peripheroronda</i>, <i>Globotruncana dehiscens</i></p>

